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- 1 Ocean biogeochemical reconstructions to estimate historical ocean
- 2 CO2 uptake
- 3 Raffaele Bernardello, Valentina Sicardi, Vladimir Lapin, Pablo Ortega, Yohan Ruprich-
- 4 Robert, Etienne Tourigny, and Eric Ferrer
- 5
- 6 Below we provide point-by-point answers to reviewers comments, together with
- 7 indications of what was changed in the new version of the paper and where to
- 8 locate these changes. Reviewers comments are in **bold** and are followed by an
- 9 "Answer" and a "Changes" sections.
- 10
- 11 Reviewer #1
- 12
- 13 <u>Comment:</u> The manuscript "Ocean biogeochemical reconstructions to estimate
- 14 historical ocean CO2 uptake" by Bernardello et al is a very useful comparison of
- 15 different methods of estimating ocean carbon uptake from ocean-only models
- 16 forced by reanalysis versus also including 3D temperature and salinity data
- 17 assimilation with direct relevance to the current gap between ocean inverse
- 18 estimates and "OBGC" or "OMIP" estimates used by the Global Carbon Project.
- 19 My major criticism is that the authors include only analysis of AMOC and MLD
- 20 changes and ignore impacts on the thermocline structure, pCO2, and
- 21 anthropogenic CO2 (GLODAP) observational constraints and impact on ideal age
- 22 and transient tracers. It is not enough to casually correlate the AMOC increase to
- 23 the anthropogenic CO2 increase in the data assimilation: the authors should at
- least look at the pattern differences in CO2 uptake between the various model
- runs to see where the extra CO2 is accumulating. Below I provide specific places
- 26 where I think such a quick analysis would substantively improve the manuscript.
- 27 <u>Answer:</u> We would like to thank reviewer #1 for this constructive criticism of the paper.
- 28 We agree that our analysis fell short of explaining the reasons for the observed changes
- in CO2 uptake and in general ocean biogeochemistry. We welcome the reviewer's
- 30 suggestions to provide such an analysis and below we respond point by point explaining
- 31 how we have addressed the issues raised.
- 32 <u>Changes:</u> We have introduced 7 new figures representing changes in the meridional
- 33 overturning circulation, changes in the ideal age tracer distribution, changes in the
- anthropogenic carbon accumulation in the interior of the ocean and a comparison
- 35 between model's estimates of anthropogenic carbon distribution and the observation-
- 36 based estimate provided by GLODAP. These new figures are described and discussed
- in the text. Details are given below in response to reviewer #1's specific comments.
- 38
- 39 <u>Comment:</u> 19-20 In the sentence "This becomes particularly important in the
- 40 context of a future decline of global CO2 emissions and the UN Framework
- 41 Convention on Climate Change stocktaking activities" it is not clear why "a future
- 42 decline of global CO2 emissions" makes carbon uptake more important than

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- 43 under scenarios of future increase. The authors may be intending to call out the
- 44 2015 Paris Agreement that seeks climate stability/sustainability and net zero
- 45 emissions at particular temperature thresholds, but the connection should be
- 46 **explicit**.
- 47 <u>Answer:</u> We agree that this sentence was not clear at all. We meant that, in a context of
- 48 declining CO2 emissions, the relative importance of variability in ocean CO2 uptake
- 49 (induced by natural climate variability) increases with respect to the net uptake of the
- 50 anthropogenic fraction of CO2. This means that being able to quantify the natural
- 51 variability becomes relatively more important for the detection and attribution of a
- 52 changing trend in ocean CO2 uptake, when CO2 emissions decline. We modified the53 text to make this point clear.
- 54 <u>Changes:</u> We have modified the text in the introduction to better explain this point (see 55 lines 19-22).
- 56

57 <u>Comment:</u> 139 – remove "we"

- 58 Answer: We have made extensive use of the active form throughout the paper. After
- 59 careful consideration, we decided that changing to passive form would decrease the
- 60 readability of the text.
- 61 <u>Changes:</u> We haven't substantially changed the use of the active form.
- 62

63 <u>Comment:</u> 209-210 – To answer the question "it is hard to pinpoint a single cause for the improvements we see in biogeochemical variables when we apply data 64 assimilation of temperature and salinity." The classical means of doing so is to 65 66 look at changes to ideal age and transient tracers like CFC's and SF6. My expectation is that the OMIP version of the model is overly stratified and that the 67 68 thermocline/warm water sphere is deeper in the assimilation case. While the 69 assimilation increasing AMOC certainly goes in the right direction, I expect it is 70 the enhancement of the shallow gyre circulation of AMOC (rather than the deeper, 71 thermohaline aspect) that is driving the improvement as it applies to all the gyres,

- 72 not just the North Atlantic. It should be easy to see where the changes in DIC
- 73 accumulate whether it is just in the Atlantic below 1000 m (in support of the
- thermohaline mechanism), or throughout the ocean above 1000 m (in support of
- 75 the general thermocline ventilation mechanism). These two comparisons should
- 76 be very easy for the authors to conduct.
- 77 Answer: Unfortunately we didn't include CFCs and SF6 in these runs but we do have the
- 78 ideal age tracer. We have used this tracer to look into the changes in ventilation that
- 79 occur when applying data assimilation, as suggested. At the same time, we compared
- 80 simulations to see where the most marked changes in anthropogenic DIC distribution
- 81 happen within the water column in different regions. Results of this analysis, point to a
- 82 general deeper penetration of anthropogenic DIC, up to 3000m, in the Atlantic as well as
- a deeper penetration of DIC in the Southern Ocean at depths compatible with an
- 84 enhanced formation of Antarctic Intermediate Waters. In the Pacific ocean the
- accumulation of DIC is shallower, mainly above 1000m depth.

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<u>Changes:</u> We have introduced figures 6 and 7 representing the meridional overturning
 circulation for the Atlantic, Pacific and Southern oceans. These are described in the

- result section at lines 183-189 and in the discussion at lines 264-267 and 273-281.
- 89 We have introduced figure 8 representing changes in the distribution of the ideal age
- 90 tracer when applying data assimilation with respect to the omip simulations. This figure 91 is described in the results section at lines 191-196 and in the discussion at lines 264-
- 92 267.
- 93 We have introduced figures 10 and 11 representing changes in the accumulation of
- 94 anthropogenic carbon when applying data assimilation with respect to omip simulations.
- 95 These figures are described in the results section at lines 206-214 and in the discussion
- 96 at lines 269-281.
- 97

98 <u>Comment:</u> 228 – furthering the need to look at the patterns of DIC inventory 99 increase above, more detail on the "ameliorated density profile" is necessary 100 here. For example, it would be helpful if the MLD analysis indicated the direction 101 of the improvement – i.e. it looks like the biases being ameliorated were a deep 102 bias in the northern gyre extension regions and a shallow bias in the Southern 103 Ocean... suggesting that it may be increased ventilation in the Southern Ocean 104 that is the most important. Looking at MLD is certainly a big part of the story, but the relation with the overall ventilated thermocline depth is more relevant to net 105 106 anthropogenic CO2 uptake.

- 107 <u>Answer:</u> The reviewer's interpretation of Figure 7 is correct. A deep bias is reduced in
- 108 the North Atlantic for ERA5 while for JRA55 the bias reduction goes in the opposite
- direction, towards correcting a shallow bias. In the Southern Ocean a shallow bias isclearly reduced in the Pacific and Atlantic sectors for ERA5, while the changes for
- 111 JRA55 are less evident. These changes result in a deeper penetration of DICant in the
- 112 Southern Ocean while in the North Atlantic, because the omip simulations differ
- substantially one from the other (ERA5-omip and JRA55-omip), when data assimilation
- 114 is applied important differences are observed between ERA5 and JRA55.
- 115 <u>Changes:</u> We have modified the description of the MLD figure (now fig. 9) to better
- 116 highlight the direction of the bias reductions (Results, at lines 206-212). We have
- 117 modified the discussion to link the changes in MLD induced by data assimilation with the
- 118 changes observed in both ventilation and anthropogenic carbon distribution. These
- 119 changes are in the discussion at lines 269-281.
- 120
- 121 Comment: 235 I don't find the degradation in nutrients and chlorophyll
- surprising at all as this was the foundational problem in the Park et al., 2018 study
- 123 the authors cite for their decision to reduce nudging near the equator and is
- 124 consistent with what I suspect is increasing ventilation under data assimilation
- 125 increasing surface nutrients and chlorophyll from a baseline configuration in
- 126 which the BGC parameterizations for phytoplankton physiology and nitrogen and
- 127 iron limitation were tuned to match observations of high nutrient/low chlorophyll
- 128 patterns.

- 129 <u>Answer:</u> There is no degradation in nutrients. On the contrary, there is an overall
- 130 improvement in their distribution. This is one of the main results that was showcased in
- 131 Figure 4. The reason the reviewer suggests for the insensitivity (or degradation) of
- 132 chlorophyll to these improved nutrient fields coincides with one of the reasons we
- 133 propose in the paper (e.g. tuning of the BGC model), together with unchanged iron
- 134 availability.
- 135 <u>Changes:</u> We opted for not changing this part of the discussion
- 136

<u>Comment:</u> 255 - I disagree with the assertion that "their direct validation is not
 straightforward" as it seems very straightforward to compare against the surface
 ocean pCO2 product of Landschützer et al (2017):

- 140 Landschützer, P., Gruber, N., & Bakker, D. (2017). An updated observation-based
- 141 global monthly gridded sea surface pCO2 and air-sea CO2 flux product from 1982
- 142 through 2015 and its monthly climatology (NCEI Accession 0160558), edited,
- 143 NOAA National Centers for Environmental Information.
- 144 And anthropogenic CO2 inventories of GLOPAPv2:
- 145 Lauvset, S. K., Lange, N., Tanhua, T., Bittig, H. C., Olsen, A., Kozyr, A., ... & Key, R.
- 146 M. (2022). GLODAPv2. 2022: the latest version of the global interior ocean
- 147 biogeochemical data product. Earth System Science Data Discussions, 2022, 1-37.
- 148 Answer: With respect to CO2 fluxes, with this statement we meant that observation-
- 149 based air-sea CO2 flux products also suffer from large uncertainties derived from the
- 150 limited coverage (spatial and temporal) of the original pCO2 dataset (SOCATv3) and
- 151 from different techniques of interpolation. As a consequence, these products are used as
- an additional line of evidence rather than a benchmark in the Global Carbon Budget (e.g.
- 153 Friedlingstein et al., 2022). We tried to convey this message by including in Figure 1 the
- 154 single members of both the model estimate and the obs-based products. The spread
- around the mean is similar for the two estimates. Additionally, in Figure 2 we show a
- correlation matrix between the 7 obs-based products from GCB2022 (including
- 157 Landschutzer et al., 2017) and the model estimates from GCB2022 (besides our
- simulations). For any model, there is a considerable variability in the value of the
- correlation coefficient while moving across obs-based products (horizontally), pointing to
- a large variability among these. Analogously to the practice adopted by the GCB
- exercise, we decided to evaluate the improvements in surface pCO2 when applying DA
- by comparing directly the original point values from the SOCATv3 dataset with the
- 163 correspondent values in the models, co-located in space and time (Hauck et al. 2020).
- 164 This comparison is presented in Figure 3.
- 165 With respect to DIC, GLODAPv2.2022 is the dataset we have used to calculate the
- 166 reduction in RMSE when applying data assimilation of temperature and salinity (see
- 167 Table 2 and Fig. 4). Total DIC is the variable we have used for Fig. 4. In a previous
- release of GLODAPv2 (Lauvset et al., 2016), a mapped estimate of the accumulated
- anthropogenic carbon in the year 2002 is provided. We will use this estimate as a
- 170 reference when analyzing the changes in the distribution of DIC between omip and DA
- 171 simulations. However, similarly to observation-based products for CO2 fluxes, the

- 172 estimates of anthropogenic DIC distribution suffer from considerable uncertainties linked
- to the diversity of methods used to infer them (e.g. Khatiwala et al., 2013).
- 174 <u>Changes:</u> We have introduced Figures 12 and 13 showing zonal averages for the
- 175 Atlantic, Pacific and Indian oceans of the differences in anthropogenic carbon
- 176 distribution between each simulation and the GLODAP estimate. These figures are
- 177 described in the result section at lines 216-219 and in the discussion at lines 294-299.
- 178
- 179 <u>References</u>
- 180 Friedlingstein, P., O'sullivan, M., Jones, M. W., Andrew, R. M., Gregor, L., Hauck, J., Le
- 181 Quéré, C., Luijkx, I. T., Olsen, A., Peters, G. P., et al.: Global carbon budget 2022, Earth 182 System Science Data Discussions, 2022, 1–159, 2022. <u>https://doi.org/10.5194/essd-14-</u>
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- Hauck, J., Zeising, M., Le Quéré, C., Gruber, N., Bakker, D. C., Bopp, L., Chau, T. T.,
- 185 Gürses, Ö., Ilyina, T., Landschützer, P., et al.: Consistency and challenges in the ocean
- carbon sink estimate for the global carbon budget, Frontiers in Marine Science, 7, 571
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- 188 Khatiwala, S., Tanhua, T., Mikaloff Fletcher, S., Gerber, M., Doney, S. C., Graven, H. D.,
- 189 Gruber, N., McKinley, G. A., Murata, A., Ríos, A. F., and Sabine, C. L.: Global ocean
- 190 storage of anthropogenic carbon, Biogeosciences, 10, 2169–2191,
- 191 https://doi.org/10.5194/bg-10-2169-2013, 2013.
- 192 Lauvset, S. K., Key, R. M., Olsen, A., van Heuven, S., Velo, A., Lin, X., Schirnick, C.,
- 193 Kozyr, A., Tanhua, T., Hoppema, M., Jutterström, S., Steinfeldt, R., Jeansson, E., Ishii,
- 194 M., Perez, F. F., Suzuki, T., and Watelet, S.: A new global interior ocean mapped
- climatology: the $1^{\circ} \times 1^{\circ}$ GLODAP version 2, Earth Syst. Sci. Data, 8, 325–340, https://doi.org/10.5104/ossd.8.225.2016.2016
- 196 https://doi.org/10.5194/essd-8-325-2016, 2016.
- 197
- 198
- 199 **Reviewer #2**
- 200
- 201 Comment: Bernardello et al. provide model-based experiments showing that constraining ocean physics towards observed temperature and salinity results in 202 a better representation of global biogeochemistry. Two sets of simulations are 203 204 done: the standard GCB approach, which uses prescribing boundary conditions 205 from atmospheric reanalyses, and the additional assimilation of observed ocean 206 physical variables. Thus the paper shows nicely that the estimate of the ocean 207 CO2 uptake can be more reliable if considering the physical changes of the ocean. I find the paper clearly written and very convincing. I only have some minor 208
- 209 suggestions on this work.
- 210 <u>Answer:</u> We thank Reviewer#2 for the time dedicated to reviewing our paper and we
- 211 detail below the changes we introduced to address the suggestions.
- 212
- 213 <u>Comment:</u> Although using EN4 gridded data is ok, the authors might be aware of
- the problems in EN4 data (e.g., some instrumental biases are not corrected, the
- 215 gridded fields are shifted to the climatology in data-sparse regions, etc.). Good et
- al. 2013 actually explicitly stated that the data should be used with caution when

- 217 dealing with long-term changes. Thus, it would be worthwhile to have some
- 218 discussions, at least in the conclusion section, that, potentially, using a better
- 219 dataset would have further benefits/improvements. A quite comprehensive
- 220 description of the data issues for ocean temperature can be found in a recent
- 221 online reprint (<u>https://essd.copernicus.org/preprints/essd-2024-42/</u>).
- 222 <u>Answer:</u> Thanks for pointing this out. We have expanded the discussion and conclusion
- sections to highlight these potential issues with EN4. The choice of EN4 comes from a
- long analysis and trial/error attempts to obtain a robust reconstruction that could provide
- 225 initial conditions for near-term climate predictions. Each product comes with its own
- problems and advantages and we agree that these need to be better highlighted in ourcase.
- 228 <u>Changes:</u> We have introduced a paragraph in the discussion (lines 283-292) to
- acknowledge potential limitations of using a specific dataset for interior 3D nudging.
- 230

231 <u>Comment:</u> For the discussion of large-scale circulation, only AMOC is mentioned

and discussed, however, ocean circulation is not only that and the other parts are

also important for carbon uptake. Please expand the discussion related to other

circulation systems such as subtropical gyres,, subpolar gyres, or the water mass

- 235 formation/transformation.
- 236 <u>Answer:</u> We addressed this suggestion with the analysis introduced in response to the
- 237 main criticism of reviewer#1. We agree that we didn't provide enough insight on the
- changes in general circulation that are brought by data assimilation and we introduced 3
- new figures to show the meridional overturning circulation in different regions and the
- changes in ventilation highlighted by changes in the distribution of the ideal age tracer.
- 241 Changes: We have introduced figures 6 and 7 representing the meridional overturning
- 242 circulation for the Atlantic, Pacific and Southern oceans. These are described in the
- result section at lines 183-189 and in the discussion at lines 264-267 and 273-281.
- 244 We have introduced figure 8 representing changes in the distribution of the ideal age
- tracer when applying data assimilation with respect to the omip simulations. This figure
- 246 is described in the results section at lines 191-196 and in the discussion at lines 264-
- 247 267.
- 248

249 <u>Comment:</u> Salinity is rarely discussed, how the improved representation of

- 250 salinity can improve the biogeochemical changes? Any insights would be very
 251 useful.
- 252 <u>Answer:</u> Salinity can play a critical role (even more than temperature) through its
- 253 contribution to the large scale density gradients and deep water mixing. For example, in
- the Labrador Sea, vertical stratification is generally controlled by salinity. So, having the
- 255 correct salinity is critical to improve the MLD. These are indirect impacts on
- 256 biogeochemistry as they manifest through changes in the physical/dynamical state of the
- 257 ocean.

- 258 <u>Changes:</u> We have addressed this aspect with the expansion of the analysis and
- discussion on the changes in circulation and their effect on CO2 uptake and ocean
- 260 biogeochemistry in general.
- 261

262 <u>Comment:</u> Fig.6: probably other SST observation data can be added to assess 263 how large the observational uncertainty is.

- 264 <u>Answer:</u> We agree.
- 265 <u>Changes:</u> We have introduced 2 new SST estimates to figure 14.
- 266
- 267 <u>Comment:</u> Table 1: why a reanalysis data (ORAS5) is used for SSS? It likely
- 268 suffers from spurious shifts due to salinity observation system changes over time
- 269 (i.e., around 2005, from a ship-based system to an Argo-based system).
- 270 <u>Answer:</u> The main reason for this choice is that it allows us to make sure that both the
- 271 SST and SSS fields are physically consistent, which can't be guaranteed when using
- objective analyses such as EN4. We have discussed this point together with the
- 273 expanded discussion on the limitations of EN4.
- 274 <u>Changes:</u> We have introduced this discussion together with the considerations about
- 275 using EN4, at lines 283-292.
- 276