



1 A 20-year satellite-reanalysis-based climatology of extreme 2 precipitation characteristics over the Sinai Peninsula

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17

18 **Abstract**

19 Extreme precipitation events and associated flash floods caused by the synoptic cyclonic-systems have profound
20 impacts on society and the environment particularly in dry regions. This study brings forward a satellite-reanalysis-
21 based approach to quantify the extreme precipitation characteristics over the Sinai Desert in Egypt, from a
22 statistical-synoptic perspective for the period of 2001-2020. Using the satellite remote-sensing precipitation and a
23 set of climate indices, we characterize the spatiotemporal distribution of extreme rainfall climatologies across the
24 Sinai region. Then, using the reanalysis dataset, the synoptic systems responsible for the occurrence of
25 precipitation events along with the major tracks of cyclones during the wet and dry periods are explored. Our
26 results indicate that the temporal changes and spatial patterns of the precipitation events do not show a homogenous
27 tendency, rather lack of spatiotemporal coherence across the Sinai. Northern parts of the Sinai, unlike other areas,
28 exhibit the highest anomalies (approx. ± 45 mm/decade); and the annual rainfall trends indicate a drier-climate in
29 the north at -0.03 mm/decade, while a wetter-climate is observed in the central and southern parts at 0.10 and 0.36
30 mm/decade, respectively. The Mediterranean cyclones accompanied by the Red Sea -and Persian Troughs are
31 responsible for the majority of extreme rainfall events all-round the year. A remarkable spatial relationship
32 between the Sinai's rainfall and the atmospheric variables of sea level pressure, wind direction and vertical velocity
33 is found. Furthermore, the cyclone-tracking analysis indicates that 125 and 31 cyclones (rainfall ≥ 10 mm/day)
34 either formed within -or transferred to the Mediterranean basin and precipitated over the Sinai during wet and dry
35 periods, respectively; while some ($\sim 15\%$ with rainfall > 40 mm/day) being capable of leading to flash flood in the
36 wet period of the region. This study, therefore, sheds new light on the extreme rainfall characteristics and the
37 dominant synoptic mechanisms over the Sinai region in the eastern Mediterranean basin.

38

39 **1 Introduction**

40 Extreme precipitation events can have fundamental impacts on society and human wellbeing by causing mortality
41 (Trenberth *et al.*, 2007; Toreti *et al.*, 2010; Wannous and Velasquez 2017; Charlton-Perez *et al.*, 2019), and by
42 causing property and ecological damages (Zhang *et al.*, 2005; IPCC, 2013; Nastos *et al.*, 2013; Boucek *et al.*,
43 2016). Precipitation extremes are realized as one of the severest natural disasters, among many others (Arnous and
44 Omar 2018). Nevertheless, these events are vital for the water resources of the region especially in the water-
45 limited environments (Peleg *et al.*, 2012; Givati *et al.*, 2019; Levy *et al.*, 2020); however, they also constitute the
46 main trigger of flash floods in arid and hyper-arid areas such as the Sinai Peninsula in Egypt (Ocakoglu *et al.*,
47 2002; David-Novak *et al.*, 2004; El-Magd *et al.* 2010; Farahat *et al.* 2017; Gado, 2020).

48 The eastern Mediterranean is one of the main cyclogenetic regions of the Mediterranean basin (Krichak *et al.*,
49 1997) and globally (Ulbrich *et al.*, 2012; Neu *et al.*, 2013), which in many cases associated with precipitation
50 extremes (Flaounas *et al.*, 2014a, 2014b). As such, most of the heavy precipitation events in this region strongly
51 rely on the presence and frequency of the intense Mediterranean cyclones (Trigo *et al.*, 2002; Kotroni *et al.*, 2006;



52 Pfahl and Wernli 2012; Lionello *et al.*, 2016), accompanied by other precipitation producing-systems at synoptic-
53 scale, sometimes of tropical/sub-tropical origin (Krichak *et al.* 1997; Hochman *et al.*, 2020).
54 A multitude of observational-numerical-synoptic studies has been carried out in relation to the extreme
55 precipitation events over the eastern Mediterranean region to date, such as extreme rainfall analysis (e.g. Alpert *et*
56 *al.*, 2002; Ben David-Novak *et al.*, 2004; Kostopoulou and Jones, 2005; Ben-Zvi, 2009; Mathbout *et al.*, 2018),
57 trends in extreme precipitation (e.g. Yosef *et al.*, 2009; Shohami *et al.*, 2011; Ziv *et al.*, 2013; Ajjur and Riffi,
58 2020), satellite remote-sensing-based analysis of precipitation extremes (e.g. Gabella *et al.*, 2006; Mehta and
59 Yang, 2008; Nastos *et al.*, 2013; Yucel and Onen, 2014), numerical modelling and climate change projections of
60 heavy precipitations (e.g. Tous *et al.*, 2015; Romera *et al.*, 2016; Toros *et al.*, 2018; Zoccatelli *et al.*, 2020; Zittis
61 *et al.*, 2020), flash floods and water resources attributed to extreme rainfall events (e.g. Morin *et al.*, 2007; Samuels
62 *et al.*, 2009; Koutroulis and Tsanis, 2010; Tarolli *et al.*, 2012; Varlas *et al.*, 2018; Zoccatelli *et al.*, 2019; Spyrou
63 *et al.*, 2020; Rinat *et al.*, 2021), synoptic analysis of precipitation extremes and floods (e.g. Dayan *et al.*, 2001,
64 2015; Kahana *et al.*, 2002; Alpert *et al.*, 2004; Tsvieli and Zangvil, 2005; Peleg and Morin, 2012; Raveh-Rubini
65 and Wernli, 2015; Toreti *et al.*, 2016), and cyclogenesis and cyclone tracking (e.g. Alpert and Ziv, 1989; Alpert
66 and Shay-El, 1994; Flocas *et al.*, 2010; Flaounas *et al.*, 2014a, 2014b; Almazroui *et al.*, 2014; Zappa *et al.*, 2014;
67 Ziv *et al.*, 2015).

68 However, literature review of the Sinai region reveals that very limited studies have been carried out so far mainly
69 on the flash floods associated with heavy rainfall events from the ground/satellite-based data analysis approach
70 (e.g. Roushdi *et al.*, 2016; Dadamouny and Schnittler, 2016; Arnous and Omar, 2018; Morsy *et al.*, 2019; Baldi *et*
71 *al.*, 2020) to numerical model experiments (e.g. Cools *et al.*, 2012; El Afandi *et al.*, 2013; Morad, 2016; Prama *et*
72 *al.*, 2020; Omran, 2020; El-Fakharany and Mansour, 2021). In such circumstances, Mohamed and El-Raey (2019)
73 pointed out that limited numbers of extreme precipitation events with high intensities and short durations that
74 typically result in flash floods allegedly are the only sources of the renewable water-resources in the Sinai
75 Peninsula. Thus, it seems necessary to understand, in the first place, the spatiotemporal distribution of extreme
76 precipitation events across the Sinai, and in the second place, to discover the corresponding synoptic-dynamical
77 mechanisms responsible for the occurrence of such events in the region. To our best of knowledge, no study
78 attempted yet to quantify the extreme precipitation characteristics (e.g. anomalies, frequencies and spatial patterns)
79 associated with the synoptic-regional atmospheric circulation and cyclone tracking over the Sinai Peninsula –and
80 even not over the eastern Mediterranean basin, as described and presented in this study. Therefore, to bridge the
81 above-mentioned research-gaps existing in the Sinai, in this study we address the following main research
82 questions:

- 83 i. how are the extreme precipitation climatologies spatiotemporally distributed across the Sinai Peninsula?
- 84 ii. which synoptic large-scale systems are responsible for the occurrence of the Sinai's (extreme)
85 precipitation events?
- 86 iii. what are the major tracks of cyclones and their frequencies over the region?

87 In this research, our data-analysis spans the period from 1st of January 2001 to 31st of December 2020. First, we
88 use satellite remote-sensing precipitation to quantify the anomaly, monthly regime, frequency and spatial patterns
89 of the extreme precipitation events, together with the computation of a set of extreme climate indices. Then, the
90 dominant synoptic atmospheric circulation patterns corresponding to the Sinai extreme precipitation events in the
91 wet –and dry periods are explored using reanalysis data at multiple levels of the atmosphere. Finally, a daily-based
92 frequency of the rainfall producing-systems (cyclone tracking) are tracked and plotted over the region for the wet
93 –and dry periods separately.
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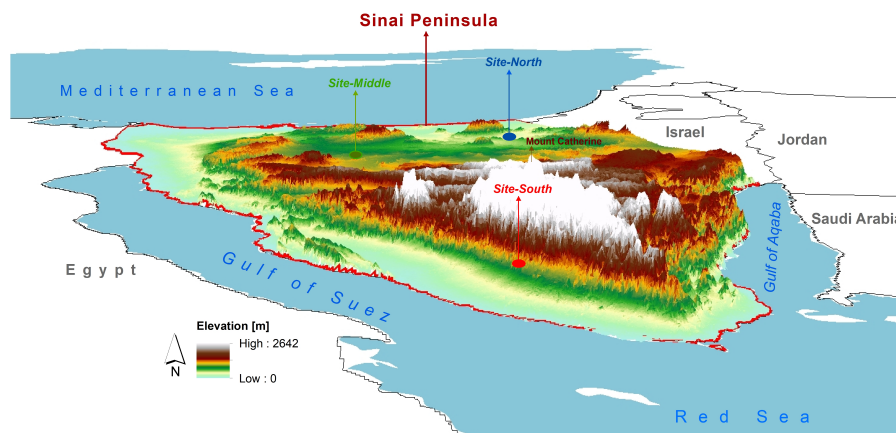
95 **2 Data and methods**

96 **2.1 The description of study area**

97 The Sinai Peninsula is located in the northeast of Egypt with an area of 61,000 km² (Fig. 1) covering about 6% of
98 Egypt's area (Mohamed *et al.*, 2014; Badreldin and Goossens 2013). Sinai lies in an arid to hyper-arid belt of
99 North Africa and belongs to the Saharan-Mediterranean climate classification (Dadamouny and Schnittler, 2016).
100 Nevertheless, it is one of the coldest regions in Egypt due to its high altitudes and mountainous topography, where
101 highest elevations are found toward the southern parts (e.g. Mount Catherine, the highest mountain in Egypt with
102 an elevation of 2642 above ground level (AGL), see Fig. 1). Overall, the Sinai region is characterized by a
103 Mediterranean climate in the north and a semidesert/desert climate in the south (El-Sayed and Habib, 2008).



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105

106 **Figure 1.** The location of the Sinai Peninsula in northeast of Egypt with the underlying three-dimensional
107 topography. Three selected sites in north (Site-North: 30°07'N, 33°09'E), middle (Site-Middle: 30°01'N, 33°50'E)
108 and south (Site-South: 28°50'N, 33°70'E) of the Sinai shown here, used for the site-scale-based calculation of
109 precipitation anomalies.

110

111 2.2 Datasets

112 2.2.1 Satellite Global Precipitation Measurement (GPM)

113 The Global Precipitation Measurement (GPM) is an international satellite mission to provide quasi-global
114 precipitation estimates with a high temporal resolution (30min, daily and monthly) and spatial resolution (0.1°)
115 through the Integrated Multi-satellitE Retrievals (IMERG) product. The GPM mission follows the Tropical
116 Rainfall Measuring Mission (TRMM) program, aiming at improving the satellite-based precipitation observation
117 capability. GPM-IMERG provides different rainfall estimates that are combined from active and passive
118 instruments in the GPM constellation (<https://gpm.nasa.gov/>). Further detailed are given by Huffman *et al.*, (2014).
119 The GPM data has been employed in several studies over the Mediterranean region (e.g. Retalis *et al.*, 2018;
120 Petracca *et al.*, 2018; Caracciolo *et al.*, 2018; Cinzia Marra *et al.*, 2019; Hourngir *et al.*, 2021). In this study, we
121 used the IMERG version 6 GPM-Level 3 final precipitation product (30min/daily) to estimate the extreme
122 precipitation characteristics including rainfall anomalies for the selected sites (see Fig. 1 for the locations), monthly
123 precipitation regime, climate indices (see Table 1), and spatiotemporal variations of the extreme precipitation
124 events for a 20-year period (2001-2020) over the Sinai Peninsula.

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126 2.2.2 NCEP/NCAR reanalysis data

127 To investigate the synoptic-dynamics climatology of the wet-period (October-March) and dry-period (April-
128 September) over the Sinai Peninsula, the required reanalysis variables were obtained from the National Centers
129 for Environmental Prediction and National Center for Atmospheric Research (NCEP/NCAR)
130 <https://psl.noaa.gov/data/gridded/data.ncep.reanalysis.html>. NCEP/NCAR provides the reanalysis data with 6-
131 hours, daily and monthly time-steps at the surface and multiple levels of the atmosphere since 1948 onward with
132 a global 2.5° × 2.5° grid (Kalnay 1996). These datasets have been used in the Mediterranean region in several
133 studies especially with regard to the synoptic analysis of precipitation, blocking systems, storm and cyclone
134 tracking (e.g. Krichak *et al.*, 2002; Trigo *et al.*, 2004; Tolika *et al.*, 2006; Trigo, 2006; Lois, 2009; Barkhordarian
135 *et al.*, 2013; Almazroui and Awad, 2016; Almazroui *et al.*, 2014, 2017; Kotsias *et al.*, 2020). In this study, we used
136 the NCEP/NCAR reanalysis data, among others e.g. ERA-Int with 1.125° spatial grid, due to its coarser resolution.
137 This is mainly because large-scale synoptic systems such as cyclones and anticyclones can be more clearly
138 represented in a coarse resolution; whereas, those will be splitted up into many small-scale systems in the finer
139 resolutions making it hard to diagnose the main controlling patterns. The following NCEP/NCAR reanalysis daily
140 meteorological variables -or derived parameters at multiple atmospheric levels were employed: sea level pressure



141 SLP (hPa), geopotential height HGT (m), relative vorticity RV (10^{-5} S^{-1}), zonal (U) and meridional (V) wind-
 142 components (m s^{-1}), and vertical velocity (ω : Pa s^{-1}).

143

144 2.3 Data analysis approach

145 2.3.1 Calculation of the precipitation anomalies

146 The annual and seasonal changes in total precipitation trend were estimated using the satellite GPM-based daily
 147 precipitation over the period of 2001-2020 for the three selected sites across the Sinai (see Fig. 1 for the locations).
 148 The climatology mean precipitation values and spatial distribution were the main criteria for the selection of the
 149 sites. In this way, each chosen site is representative for its surrounding area in terms of both the precipitation
 150 magnitude and spatial patterns. Thus, the selected sites in north, south and middle parts indicate the max, min and
 151 average amounts of precipitation received across the Sinai region, respectively over a 20-year time-period (see
 152 Fig. 4a). For this analysis, precipitation anomalies (annual and seasonal) are calculated in three steps: i) calculating
 153 the climatology mean of the data, ii) subtracting the mean value from the each year/season values, and ii) drawing
 154 the trend of slopes using the least squares method. Here, the winter includes DJF months (Dec, Jan, Feb) and the
 155 autumn includes SON months (Sep, Oct, Nov). Precipitation anomalies for the spring and summer periods were
 156 found to be close to zero, and therefore excluded.

157

158 2.3.2 Estimate of the extreme precipitation indices

159 The spatiotemporal analysis on the satellite GPM-based daily precipitation timeseries was carried out for the entire
 160 Sinai region. For this, a set of climate functions/indices (see Table 1 for the details) was computed for the period
 161 of 2001-2020 using the Climate Data Operator (CDO) (Schulzweida, 2020), developed in Max-Planck-Institute
 162 for Meteorology (<https://code.mpimet.mpg.de/projects/cdo>).

163

164 **Table 1.** CDO functions and climate indices used in this study (Schulzweida, 2020).

Index	Descriptive name	Definition	Units
<i>monsum</i>	Monthly sum	For every adjacent sequence t_1, \dots, t_n of time steps of the same month it is: $o(t, x) = \sum \{i(t', x), t_1 < t' \leq t_n\}$	mm
<i>yearsum</i>	Yearly sum	For every adjacent sequence t_1, \dots, t_n of time steps of the same year it is: $o(t, x) = \sum \{i(t', x), t_1 < t' \leq t_n\}$	mm
<i>eca_pd</i>	Precipitation days index per time period	Generic ECA operator with daily precipitation $\text{sum} \geq 5 \text{ mm}$.	days
<i>eca_r10mm</i>	Heavy precipitation days index per time period	Specific ECA operator with daily precipitation $\text{sum} \geq 10 \text{ mm}$	days
<i>eca_r20mm</i>	Very heavy precipitation days index per time period	Specific ECA operator with daily precipitation $\text{sum} \geq 20 \text{ mm}$	days
<i>eca_cdd</i>	Consecutive dry days index per time period	Maximum number of dry days with daily precipitation $\text{sum} \geq 1 \text{ mm}$	days
<i>eca_rr1</i>	Wet days index per time period	Number of wet days with daily precipitation $\text{sum} \geq 1 \text{ mm}$	days
<i>eca_sdi</i>	Simple daily intensity index per time period	Average precipitation on wet days with daily precipitation $\text{sum} \geq 1 \text{ mm}$	mm/day

165

166 2.3.3 Synoptic analysis

167 To explore the synoptic and dynamics climatologies at multiple level of the atmosphere responsible for the
 168 occurrence of the extreme precipitation events over the Sinai, the daily reanalysis data obtained from NCEP/NCAR
 169 was investigated. In the first place, the wet-period and dry-period were splitted using the extreme precipitation
 170 indices described above. We realized that, a threshold of $\geq 10 \text{ mm/day}$ is desirable in determining the wet and dry
 171 periods for the Sinai region. Thus, the wettest months were found from October to March, defined as the wet-
 172 period, and the driest months from April to September, defined as the dry period. In the second place, using the



173 above-mentioned satellite-reanalysis variables (see Sect. 2.2.), the dominant synoptic features and regional
174 atmospheric circulation patterns accompanied by the spatial correlations between the Sinai's precipitation and key
175 meteorological variables were computed and analyzed for the wet and dry periods for the climatology period of
176 2001-2020.

177

178 2.3.4 Cyclone tracking

179 In line with the synoptic analysis mentioned above, the daily trajectories of the rainy systems (cyclones)
180 precipitated over the Sinai region (≥ 10 mm) were tracked and plotted for the wet and dry periods (2001-2020).
181 This analysis consists of three major steps as follows: First, daily composites of SLP together with the relative
182 vorticity and streamflow at 850-hPa level were generated i.e. 7305 composites maps at a synoptic-scale; and the
183 same numbers of daily total precipitation maps were produced over the Sinai Peninsula also. Second, all the 20-
184 year daily precipitation events of the Sinai were monitored one-by-one, those with < 10 mm disregarded, and the
185 remaining events were classified into five categories based on rainfall magnitude as follows: category 1 (10-
186 20mm), category 2 (21-30mm), category 3 (31-40mm), category 4 (41-50mm) and category 5 (> 51 mm). For each
187 classified rainfall-event the corresponding cyclone was tracked since it was born (cyclogenesis) until disappeared
188 (cyclolysis) from the region using the above-mentioned composite charts. Third, finally the cyclone tracking charts
189 for the wet and dry periods produced over the study area.

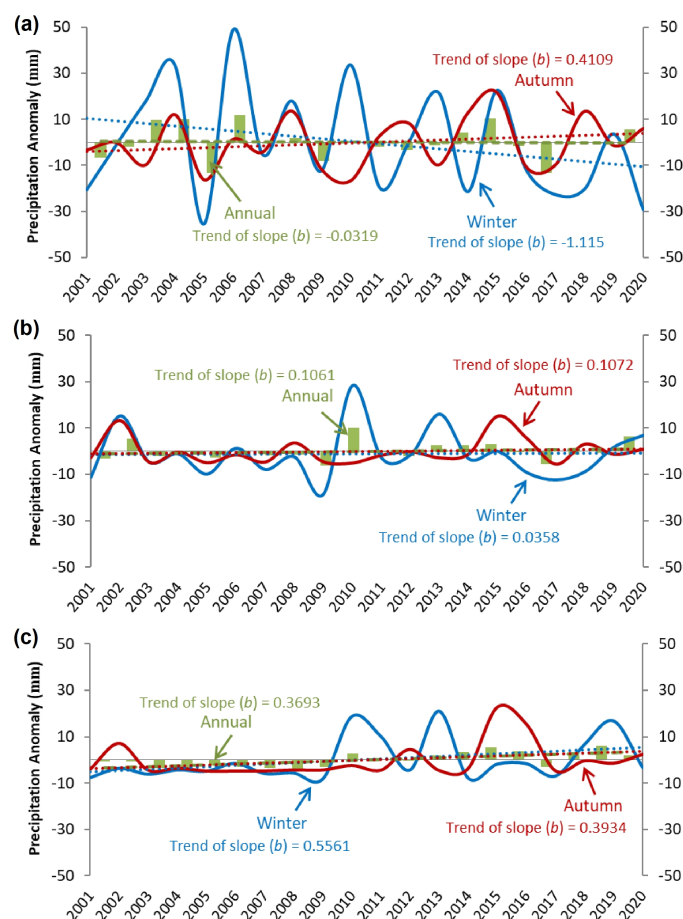
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191 3 Results

192 3.1 Climatology analysis of the precipitation characteristics

193 3.1.1 Annual and seasonal changes in precipitation trend

194 The average changes in the annual and seasonal precipitation trends during the recent period of 2001-2020 across
195 the Sinai Peninsula are shown in Fig. 2. The sign and magnitude of both annual and seasonal precipitation
196 variations do not show spatial continuity Sinai-wide, rather indicating a spatially dependency. This implies that,
197 the site located in north of the Sinai (Fig. 2a, see Fig. 1 for the location) exhibits a greater rate of the annual and
198 interannual changes with a higher anomaly (approx. ± 45 mm/decade) compared to the sites located in the middle
199 (Fig. 2b) and south (Fig. 2c) of the Sinai. The largest annual (inconsiderable) negative trend is found in north of
200 the Sinai (-0.03 mm/decade), while the central (0.10 mm/decade) and southern (0.36 mm/decade) sites exhibit
201 positive trends. This means that, annual precipitation rates over the northeastern parts and middle-towards-
202 southern parts of the Sinai are experiencing a drier and wetter climatic condition, respectively. In terms of
203 interannual trends, all sites show positive changes, expect for the winter precipitation in the north-site expressing
204 a remarkable negative trend (-1.11 mm/decade), as shown in Fig. 2a. Amongst, the maximum and minimum
205 positive interannual changes are found in wintertime precipitations of the south-site (0.55 mm/decade) and the
206 middle-site (0.03 mm/decade), respectively (Figs. 2b and 2c).
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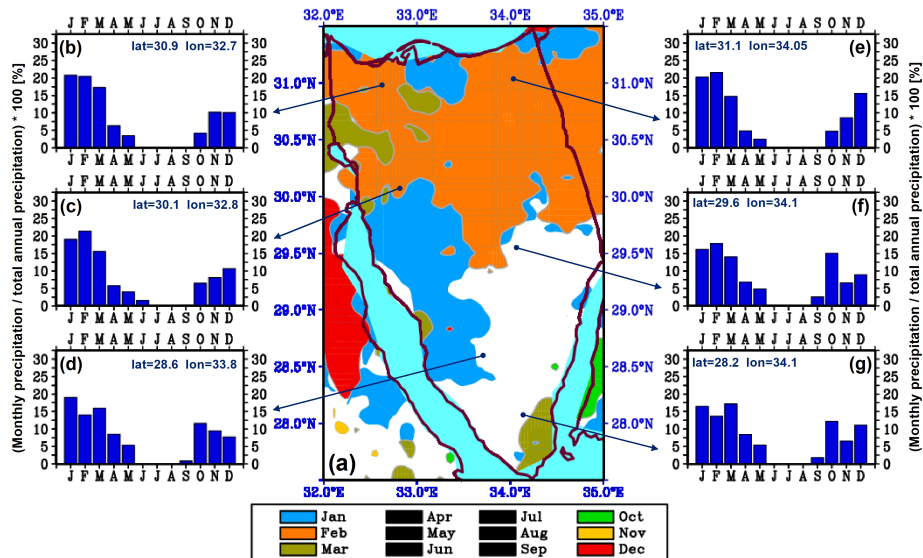
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 209 **Figure 2.** Annual and seasonal changes in the total precipitation (mm) for three selected sites in (a) north, (b)
 210 middle and (c) south of the Sinai Peninsula during 2001-2020. For the locations of the sites, see Figure 1. In the
 211 panels, winter (DJF) is given in blue, autumn (SON) is given in red, and annual in green color.
 212

213 3.1.2 Precipitation regime

214 Figure 3 represents the monthly precipitation regime climatologies (2001-2020) over the Sinai region. High
 215 (>20%) ratios of monthly precipitation over annual precipitation are estimated in the winter months of January
 216 and February, mostly found in the mid-to-north of the Sinai. March indicates some patches of high ratios in the
 217 south and northwest also, as shown in Fig. 3a. However, the period of April to September (colored in black in the
 218 legend) receives less than 20% ratio of the annual precipitation. This implies that spring and summer months
 219 experience longer dry weather periods than the winter season. Considering the autumn months, the areas with 20%
 220 ratio of annual precipitation remain largely out of the Sinai domain, except for a few mini-patches. Therefore,
 221 winter is the rainiest season throughout the Sinai. It is also noted that, full ratios of monthly precipitation to annual
 222 precipitation for individual months of the year are illustrated in Fig. S1, which provides further details on the
 223 monthly precipitation regime in the study area. Furthermore, to compare the monthly ratios of precipitation across
 224 the Sinai, the bar charts for the given sites covering the whole Sinai were plotted (Figs. 3b-3g). The highest and
 225 lowest ratios are found in the winter and summer months, respectively. However, by a closer look it becomes clear
 226 that chosen sites do vary in terms of magnitude and trends in the monthly precipitation ratios. For instance, in most
 227 sites the highest monthly ratio is observed in February, except for sites located in the Sinai southwest (which, is
 228 January – Fig.3d) and southeast (which, is March - Fig.3g). Likewise, an inconsistent seasonal trend is also



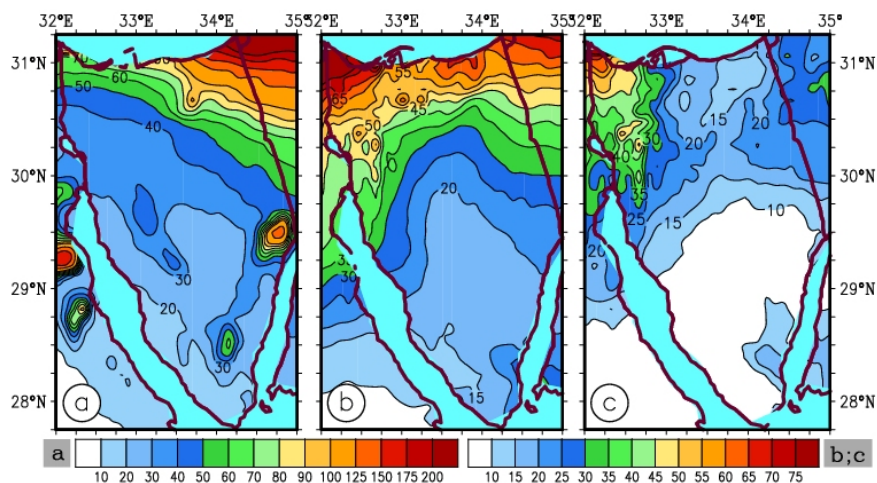
229 remarkable for the autumn months, meaning that the northern sites indicate an positive trend from late summer to
 230 the end of autumn (Figs. 3a, 3b and 3e). The southern sites, however, represent a contrasting pattern with respect
 231 to the monthly precipitation regime (Figs. 3d, 3f and 3g).
 232



233
 234 **Figure 3.** Monthly precipitation regime: (a) ratio of monthly sum precipitation to annual total precipitation (%),
 235 where only ratios >20% are plotted for each month; panels b-g indicate the monthly ratios (January to December)
 236 for the selected sites across the Sinai Peninsula for the climatology period of 2001-2020. It is noted that in the
 237 panel a, the monthly ratios from April to September (colored in black in the legend) are below 20%, thus not plotted
 238 here, but full ratios (%) are illustrated in Fig. S1 in a monthly basis.
 239

240 3.1.3 Spatiotemporal variations of the extreme precipitation events

241 The spatial precipitation patterns in terms of the climatology average, the rainiest month and the wettest day for
 242 the period of 2001-2020 in the Sinai Peninsula are illustrated in Figure 4. The climatology map of precipitation
 243 markedly demonstrate that northeast and southwest parts of the Sinai receive the highest (ranged between 100 and
 244 150 mm/year) and the lowest (ranged between 20 and 30 mm/year) amounts of annual precipitation, respectively
 245 (Fig. 4a). This implies that precipitation unevenly distributed over the Sinai. However, most parts of the region
 246 receive not as high as 40 mm/year, except for the northern areas close to the Mediterranean Sea. With respect to
 247 the occurrence of the precipitation extremes in the Sinai, we discovered that the rainiest month (out of 240 months)
 248 was in March 2020 (Fig. 4b) with a wide range of rainfall values from 15 to 30 mm/month in the south and from
 249 50 to 70 mm/month in the north. Interestingly, the wettest day (out of 7305 days) has been occurred in the same
 250 month/year, that is March 12, 2020 (Fig. 4c), thus it is not surprising to see an analogous spatial pattern when
 251 compared to the rainiest month (Fig. 4b) but with less magnitude.
 252 We also identified the twelve rainiest months out of 240 months (see Fig. S2) and the twelve wettest days out of
 253 7305 days (see Fig. S3). It was found that 9 out of 12 extreme month/day cases occurred in the winter season (Jan,
 254 Feb and Mar) with the highest frequency occurrence in January (5 cases); while only 3 out of 12 cases took place
 255 in autumn (Oct and Dec). In all the extreme precipitation cases, either the rainiest months –or the wettest days, a
 256 similar spatial pattern was captured with respect to the spatial precipitation distribution, meaning that the maxima
 257 were recorded in the north and the minima in south of the Sinai Peninsula.
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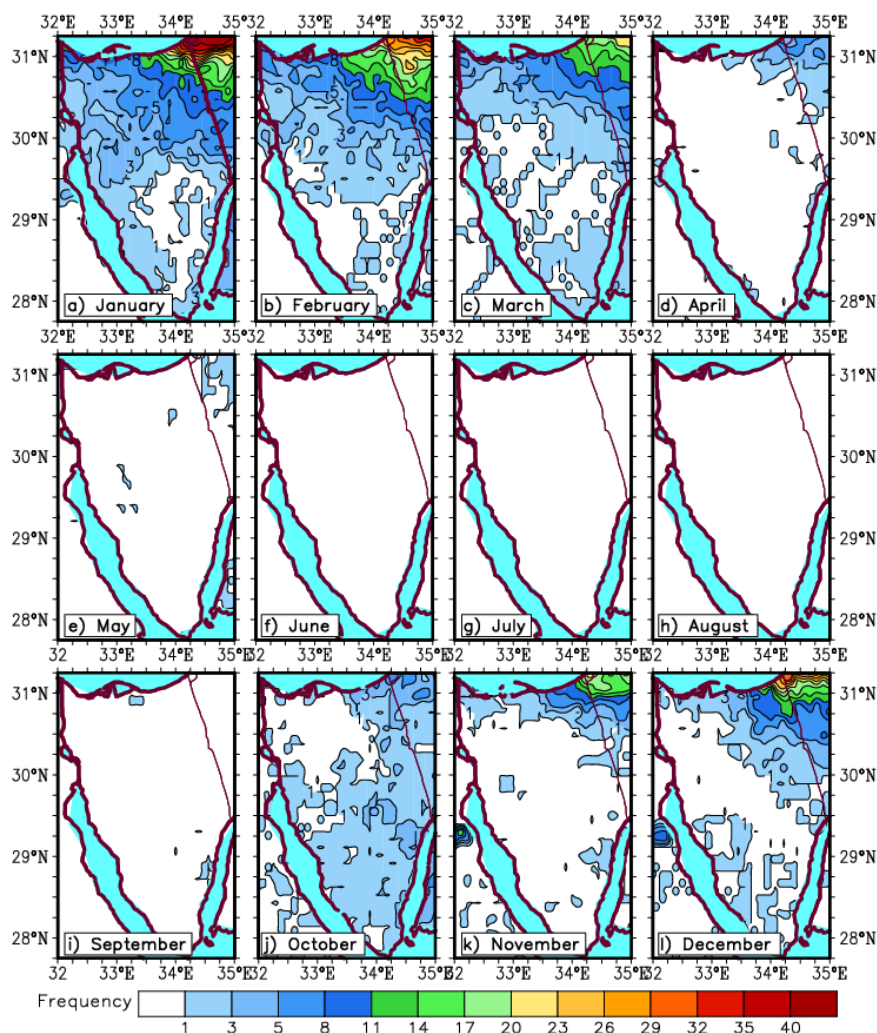


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Figure 4. The precipitation spatial patterns over the Sinai Peninsula: a) climatology map of mean annual precipitation (2001-2020); b) the wettest month i.e. March 2020 (out of 240 months); and c) the wettest day i.e. March 12, 2020 (out of 7305 days). Units are in mm.

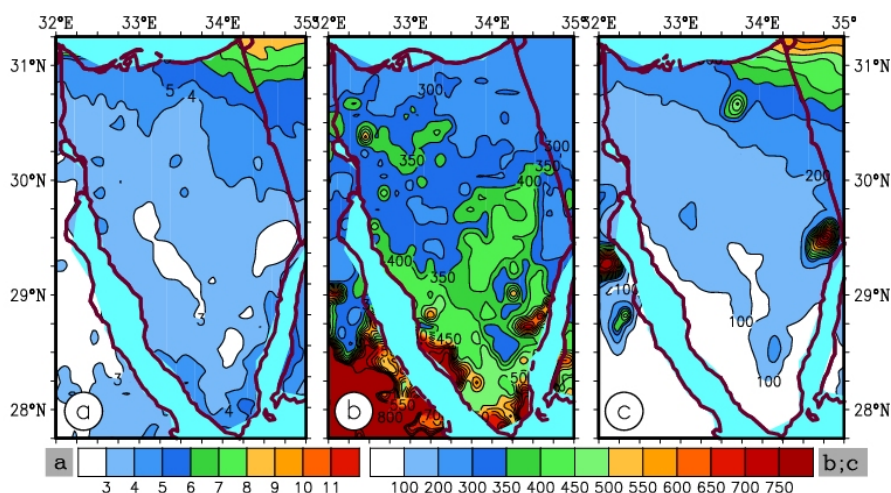
With respect to abundance of extreme rainfall events, frequency occurrence of the Sinai's precipitation extremes with different thresholds was plotted (see Fig. S4). The highest and lowest frequencies with a threshold of ≥ 5 mm/day are occurred in the north (ranged between 100 and 250) and south of the Sinai (ranged between 20 and 40), respectively. Higher precipitation frequency thresholds of ≥ 10 mm/day and ≥ 20 mm/day follow the same spatial pattern as to threshold of ≥ 5 mm/day across the region, but with much lower frequencies. The distribution of the extreme precipitation frequencies, regardless of their thresholds, are in very good agreement with the spatial pattern of the precipitation climatology map (see Fig. 4a).

To find out the precipitation spatial patterns of the Sinai at monthly basis, the frequency of events with ≥ 10 mm/day, as a middle-threshold among others, was selected. According to Figure 5, the occurrence of the precipitation frequencies with ≥ 10 mm/day are limited to the cold periods of winter (ranging from 1 to 40, Figs. 5a-5c) and autumn (ranging from 1 to 30, Figs. 5j-5l). While for the period from April to September (the growing season) almost no precipitation has been recorded (Figs. 5d-5i), suggesting a prolonged dry-period in the region. The winter months comparatively receive higher numbers of rainfall events especially in the northeast (with the maximum frequencies in January (Fig. 5a)), when compared to the autumn months (with the minimum frequencies in November (Fig. 5k)). It is worth mentioning that, monthly spatial precipitation patterns with a threshold frequency of ≥ 20 mm/day were computed also (see Fig. S5). It was found that the frequency occurrence of the extreme events reduced almost by half in comparison with ≥ 10 mm/day precipitation frequency (Fig. 5). Furthermore, it is mostly occurred in the late autumn and early winter episodes, and only limited to a small part of the northeastern Sinai.



284
 285 **Figure 5.** Frequency occurrence of the monthly precipitation climatology events with a threshold of ≥ 10 mm/day
 286 for the period of 2001-2020 (7305 days) over the Sinai Peninsula. Units are in days.
 287

288 The dryness and wetness conditions across the Sinai region were also explored by computing the simple daily
 289 intensity index (SDII), number of consecutive dry days (CDD) and number of wet days (RR1) (Fig. 6). It can be
 290 clearly seen that the highest SDII is observed in the northeast with an intensity of ≥ 6 mm/day. Interestingly, the
 291 lowest SDII is not seen in the south (even though the minimum precipitation magnitude and frequency is located
 292 there – see Fig. 4), but in central parts of the Sinai with ≤ 3 mm/day (Fig. 6a). CDD is remarkable in the south with
 293 ≥ 600 out of 7305 days, indicating that these areas receive less than 1mm rainfall for a long period; however, it is
 294 gradually decreasing northward with ≤ 300 days (Fig. 6b). Unlike CDD, it is not surprising to observe that RR1
 295 is the lowest in the south (≤ 100 days) and innermost parts (≤ 200 days), but rapidly increases towards the northeast
 296 of the region (≥ 350 days), as shown in Figure 6c. These results (Fig. 6) are in good agreement with the above
 297 findings over the Sinai Peninsula.
 298



299
300 **Figure 6.** The extreme daily precipitation indices with a threshold of ≥ 1 mm/day: a) simple daily intensity index
301 (SDII, mm/day), b) consecutive dry days (CDD, days), and c) wet days index (RR1, days) for the period of 2001-
302 2020 (7305 days) over the Sinai Peninsula. Units are: (a) in mm and (b-c) in days.

303
304 Lastly, we plotted the monthly precipitation climatologies (2001-2020) together with ranks of 12 months (out of
305 240) with the highest amount of precipitation received in the Sinai Peninsula (Fig. S6a). The most extreme
306 precipitation event occurred in March 2020 over the past two decades, followed by February 2019 and January
307 2013. Also, timeseries of areal-averaged daily precipitation in the year 2020 was illustrated over the Sinai (Fig.
308 S6b). The precipitation distribution is noticeably limited to the cold months (October-March), which is consistent
309 with Figure 5, with the severest storm recorded during 11-13 March (Fig. S6c). Allegedly, the peak rainfall hours
310 (>30 mm) occurred in the afternoon of the March 12, 2020, as represented by onset and termination of the most
311 powerful rainy system in hourly intervals of the subplot in Fig. S6c. It may worth mentioning that, the exceptional
312 storm event of 11-13 March 2020 over the Sinai is deeply investigated via a data-analysis and simulation-
313 experiment approach in a follow-up research.

314 3.2 Synoptic analysis of the wet and dry periods

316 Spatial distribution of the monthly mean precipitation amounts and magnitudes indicated a remarkable difference
317 between the wet period (ranged 5-70 mm/month) and dry period (ranged 1-3 mm/month) for the climatology
318 period of 2001-2020 over the Sinai Peninsula (see Fig. S7). However, despite a large dissimilarity in precipitation
319 values of the wet and dry periods, their spatial pattern climatologies largely resemble. This implies, precipitation
320 amounts in both periods are notably increased from southern parts towards northeast of the Sinai. In the follow-up
321 sub-sections, therefore, the atmospheric large-scale systems corresponding for the occurrence of the wet and dry
322 periods are synoptically investigated.

324 3.2.1 Synoptic patterns and atmospheric circulation structure

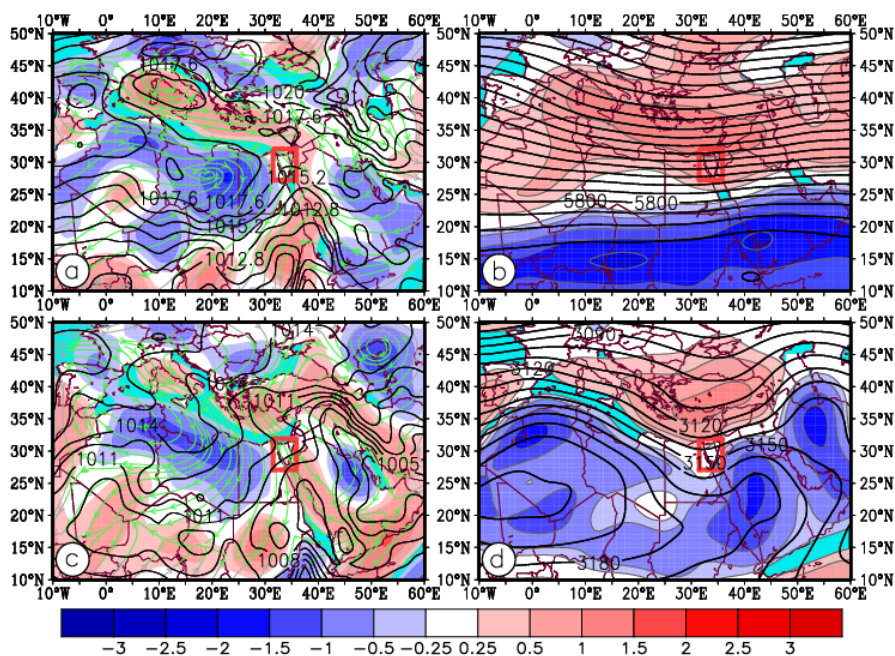
325 Figure 7 represents the climatology (2001-2020) of the synoptic patterns and cyclogenesis at surface and 500-hPa
326 atmospheric levels during the wet and dry periods over the Mediterranean basin including the Sinai Peninsula
327 (marked by a red box). In the wet period at surface level (Fig. 7a), two major sources of strong cyclonic activities
328 (cyclogenesis) are observed over the Mediterranean western part (at the lee of Alps Mountains over Gulf of Genoa)
329 and eastern part (at the lee of Taurus Mountains over Cyprus) – see Fig. 12 for the locations. These areas are found
330 by the closed SLP contours along with strong positive vorticity over the west and east parts of the Mediterranean
331 Sea, respectively. Cyprus low allegedly is responsible for the occurrence of majority of rainfall events over the
332 eastern Mediterranean including the Sinai. For instance, about 80% of the rainfall in the cold period of Israel are
333 associated with the Cyprus cyclone systems, as pointed out by Saaroni *et al.*, (2010). In wet period, the Red Sea



334 trough as a lower-level system, is another significant synoptic system that influencing the eastern Mediterranean
335 region, but mostly in the autumn (Ziv *et al.*, 2021). As shown in Fig. 7a, this trough is developed as a result of the
336 coexistence of the eastern African cyclone namely Sudan's Low and Saudi Arabia's anticyclone. Its high impact
337 on the eastern Mediterranean area is depending on the position of the Red Sea trough axis, that is, the eastern
338 position, as pointed out by e.g. Saaroni *et al.* (1998), and Tsvieli and Zangvil, (2005). However, the impact of the
339 Red Sea trough on the Sinai's rainfalls is limited compared to northeastern parts of the Mediterranean basin, mostly
340 due to the geographical location of the Sinai. In line with the lower levels, the pressure pattern at 500-hPa level
341 shows a synoptic-scale trough (of the persistent low center) with high positive vorticities providing a suitable
342 condition for occurrence of rainfall events over the Mediterranean region extending towards the Middle-East areas
343 (Fig. 7b).

344 In contrast to the wet period, surface level pattern of the dry period differs strongly over the region (Fig. 7c). In
345 the dry period, hardly ever cyclones are produced in the western Mediterranean as dominated by the high-pressure
346 systems extending from the north Atlantic Ocean and north of Africa. Limited low-pressure systems however are
347 typically developed over the eastern Mediterranean. This is due to the formation of a trough extending from the
348 Persian Gulf (which, develops as the result of the topographic impact of Zagros Mountains in western Iran) via
349 Taurus Mountains in the southern Turkey into the eastern Mediterranean basin (see Fig. 12 for the locations). The
350 Sinai region locating in the southeastern Mediterranean basin, as shown in Fig. 7c, is highly influenced by the
351 ridge of the north Africa so-called Azores anticyclone, rather than the Persian Trough that impacting mostly the
352 northeastern parts of the Mediterranean. Thus, at midlevel of 500-hPa geopotential height, the eastern
353 Mediterranean is mostly subjected to persistent air subsidence, and only a limited trough is formed with relatively
354 high positive vorticity over the eastern Mediterranean (Fig. 7d). This results in preventing rainfall to a large extent
355 over the region during the dry period. Therefore, the Sinai area receives much less amount of precipitation in terms
356 of magnitude and frequency, compared to those received over northeastern parts (such as Israel) of the
357 Mediterranean basin. These results are in good agreement with the findings reported by Alpert *et al.* (1990) and
358 Saaroni and Ziv (2000).

359



360

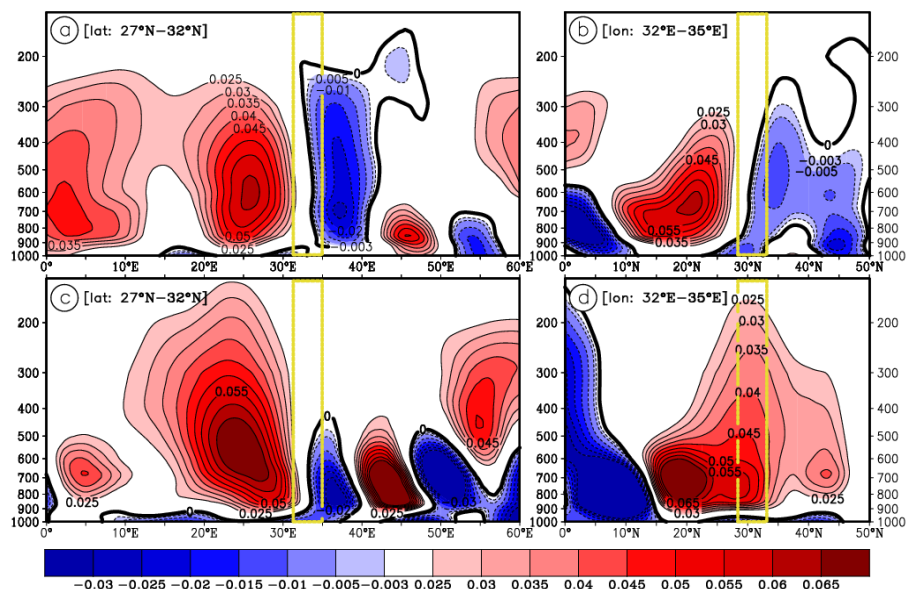
361 **Figure 7.** Climatology synoptic condition during the wet-period from October to March (a and b) and dry-period
362 from April to September (c and d) during the period of 2001-2020 over the Sinai Peninsula (red box in each panel);



363 a) composites of sea level pressure (*black contours*, hPa), 925-hPa relative vorticity (*shading*, 10^{-5} S^{-1}) and
364 streamflow (green streamline); b) 500-hPa composite of geopotential height (*isolines*, m) and relative vorticity
365 (*shading*, 10^{-5} S^{-1}); c and d same as in *a* and *b* panels respectively, but for the dry period.
366

367 Besides the synoptic pressure-systems described above, vertical velocity motions (ω) could further reveal
368 discrepancies between the wet and dry periods, from a dynamical perspective. Increase in the synoptic precipitatin
369 events over the wet period is inevitably attributed to the existence and duration of strong rising parcels of air and
370 upward vertical streams over the Sinai -and in the nearby regions. The ω cross-section along the longitude
371 (Fig. 8a) represents a maximum core with negative value of -0.03 Pa s^{-1} occurs at 800-700 hPa levels (at above
372 36°E) extending up to 250 hPa. It also indicates that, unlike to the western parts, eastern parts of the Sinai
373 experiencing a relatively strong rising condition at multiple levels of the atmosphere during the wet period.

374



375

376 **Figure 8.** Vertical velocity cross-section (ω : Pa s^{-1}) for: the wet period of October to March (a and b), and
377 dry period of April to September (c and d) over the period of 2001-2020. Omega values averaged for the latitudes
378 of 27°N - 32°N across the longitude (a and c panels), and for the longitudes of 32°E - 35°E across the latitude (b and
379 d panels). Yellow box in each panel indicates the location of Sinai Peninsula along longitude and latitude.
380

381 A similar pattern analogous to the longitude cross-section is observed also along the latitude (Fig. 8b). This means
382 that, the maximum core of vertical velocity with the value of -0.006 Pa s^{-1} is seen towards the northeast of the
383 Siani (at below 35°N) in particular at higher levels. However, when it comes to the dry period, a much weaker
384 negative omega is observed, mostly limited to lower levels of the atmosphere along the longitude (Fig. 8c), and it
385 is allegedly positive (sinking) in particular on the southern parts of the Sinia region along the latitude (Fig. 8d). In
386 such circumstances, the rising of air is strictly restricted. These results (Fig. 8) therefore further clarify, among
387 many others, why the northeast-parts of the Sinai receive higher (intense) amount of precipitation compared to rest
388 of the Sinai Peninsula, that is, partially due to the stronger vertical velocity motions in both the dry -and especially
389 wet periods.
390

391

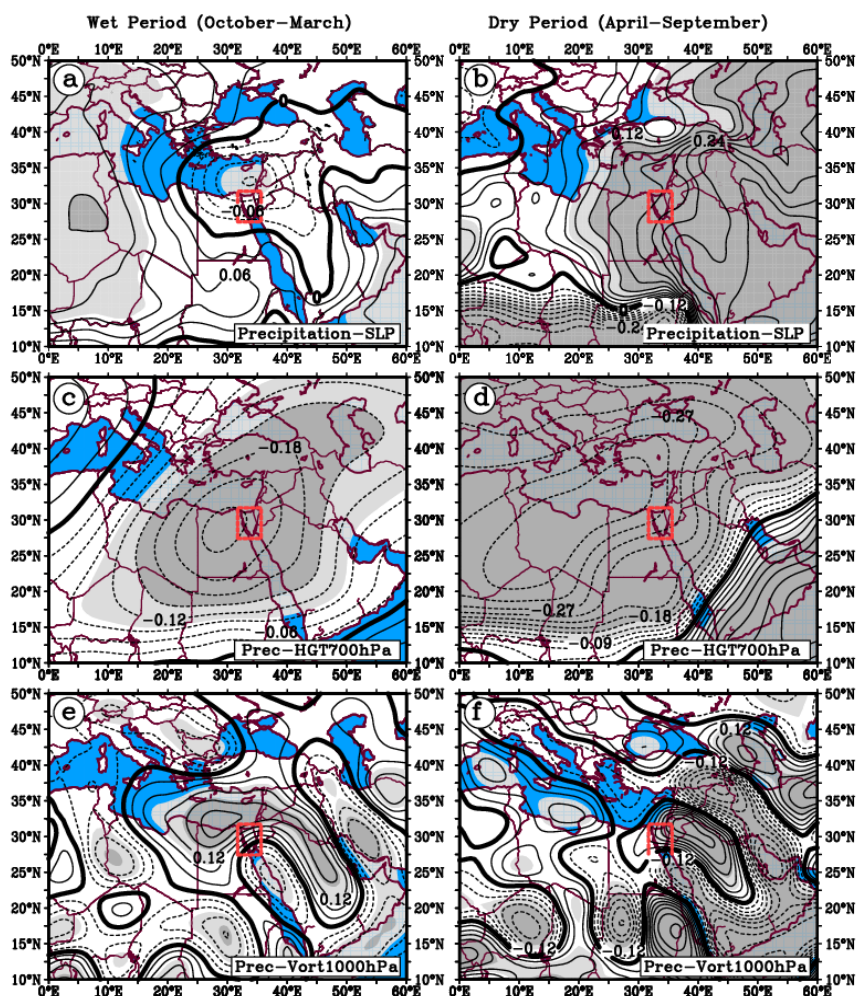
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393 3.2.2 Spatial correlation analysis

394 In this section, daily-scale relationships of the Sinai's rainfalls associated with the regional atmospheric variations
395 responsible for the occurrence of wet and dry periods are explored. Figures 9a and 9b show the spatial correlation
396 patterns between the Sinai's precipitation and regional sea level pressure (SLP) during the wet-period and dry-
397 period, respectively. A meaningful negative correlation ($r = -0.1$) is seen over the Sinai region. This indicates that
398 a strong association is realized between higher rainfall events (magnitude and frequency) and lower surface
399 pressure fields over eastern Mediterranean including the Sinai in wet period (Fig. 9a). Contrariwise, a positive
400 correlation ($r = 0.25$) is found between the rainfall and SLP over the Sinai (Fig. 9b), highlighting the dominance
401 of high-pressure fields over the region that restrict rising of the air during the dry period. The spatial patterns at
402 midlevel of 700-hPa also represent a meaningful negative correlation ($r = -0.24$) between the Sinai's rainfall and
403 geopotential height (HGT) during wet period (Fig. 9c). A similar spatial pattern with a higher correlation
404 coefficient ($r = -0.3$) is observed in the dry period also. However, a significant decrease in the region's precipitation
405 could be justified by the predominance of subtropical high-pressure centers and increase of HGT during the dry
406 period; thus, a meaningful relationship is formed between the two (Fig. 9d). The potential vorticity (PV) at the
407 low-level of 1000-hPa correlates positively with the rainfalls in both wet and dry periods, indicating a cyclonic
408 circulation in lower atmosphere over the Sinai region. However, positive PV ($r = 0.12$) has been dominated over
409 the eastern Mediterranean including the Sinai during the wet period (Fig. 9e); whereas, its impact remarkably
410 diminished over the region in dry period (Fig. 9f), resulting in decrease of precipitation in the eastern
411 Mediterranean basin.

412

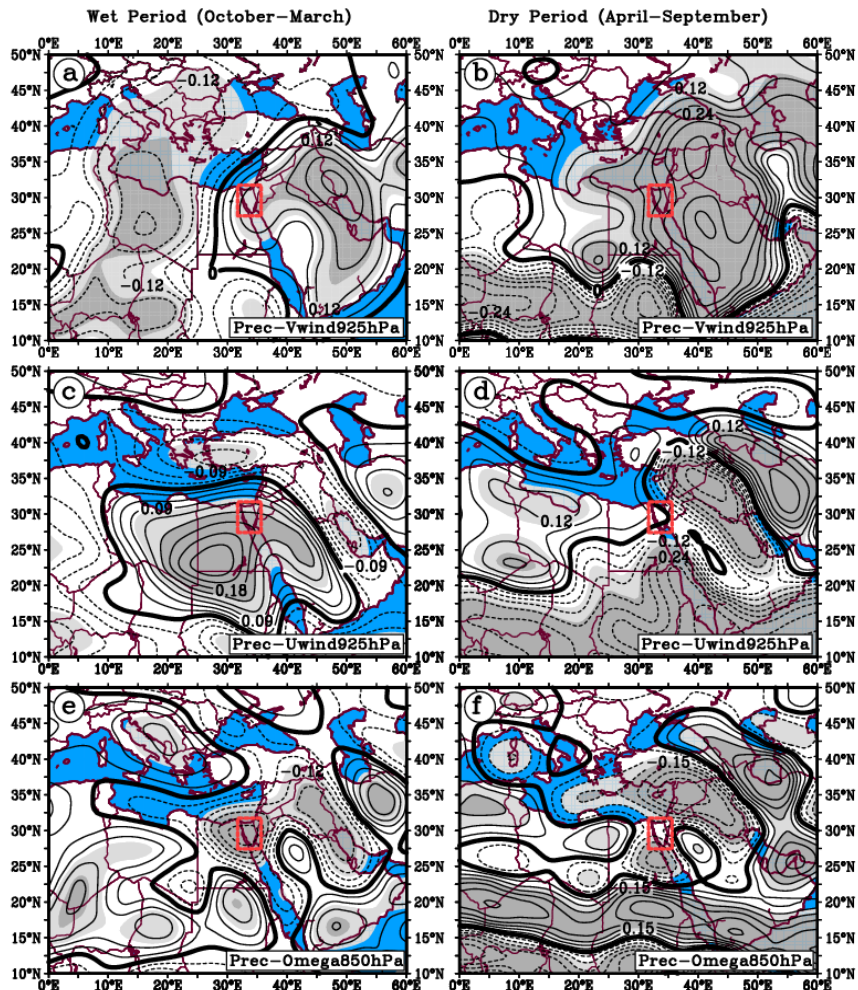


413

414 **Figure 9.** Spatial correlation patterns between the daily precipitation amounts averaged over the Sinai Peninsula
 415 (red box in each panels) and key atmospheric variables for the wet-period of October to March (left panels) and
 416 dry-period of April to September (right panels) over the period of 2001-2020. In each panels, correlations are
 417 between precipitation and: (a and b) SLP; (c and d) geopotential height HGT at 700-hPa; and (e and f) relative
 418 vorticity RV at 1000-hPa. The 95% and 99% meaningful correlations are shown in light-gray and dark-gray colors,
 419 respectively.
 420

421 A coupling-correlation-pattern, as shown in Fig. 10, is observed with respect to the precipitation and the meridional
 422 wind (V-wind) at 925-hPa level over the Sinai region during the wet period (Fig. 10a). This indicates, the Sinai's
 423 rainfall positively correlated ($r = 0.12$) with southerlies found across the Middle-East with a core on the
 424 Mesopotamia, see Fig. 12 for the locations), but negatively correlated ($r = -0.15$) with northerlies found over
 425 central/eastern Mediterranean and north of Africa. This provides a suitable condition for moisture transport from
 426 the Red Sea (by the southerlies) and the Mediterranean Sea (by the northerlies) into the study area.

427



428

429 **Figure 10.** Same as in Fig. 9, but for the correlations between precipitation and: (a and b) meridional wind (V-
430 wind) at 925-hPa; (c and d) zonal wind (U-wind) at 925-hPa; and (e and f) vertical velocity (omega) at 850-hPa.

431

432 In contrast, the region is dominated by the southerly winds during the dry period (Fig. 10b), which limits the role
433 of Mediterranean to feed the region with abundant moisture, thus rain events are largely reduced. Interestingly,
434 likewise the V-wind, a similar coupling-pattern is also observed between precipitation and zonal wind (U-wind)
435 at 925-hPa level over the area during wet period (Fig. 10c). In such circumstances, the Sinai's rainfall positively
436 correlates ($r = 0.15$) with westerlies over the eastern Mediterranean basin. However, in the dry period (Fig. 10d),
437 the Sinai's precipitation is largely associated with the negative predominant westerlies over the Mesopotamia and
438 north of Saudi Arabia. Finally, the Sinai's wet period rainfall correlates negatively ($r = -0.18$) with the omega at
439 lower atmosphere (at 850-hPa, Fig. 10e) over the eastern Mediterranean basin indicating a strong vertical velocity.
440 The relationships of the Sinai's rainfall and vertical velocity are largely weakened ($r = -0.08$) during the dry period
441 (Fig. 10f), thus limits the rising of air to a large extent.

442

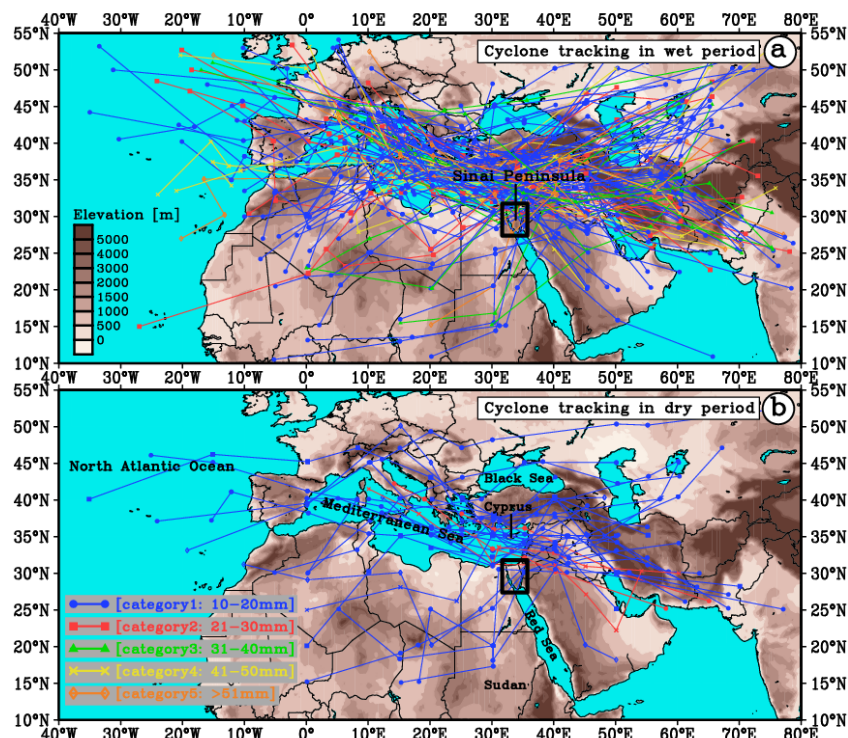
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444



445 **3.3 Cyclone tracking in the wet and dry periods**

446 Figure 11 illustrates the daily tracks of cyclones (with a rainfall $\geq 10\text{mm/day}$) over the Sinai region in the wet and
 447 dry periods for the climatology period of 2001-2020. Total numbers of cyclones during the wet and dry periods
 448 were found to be 125 and 31 cases, respectively. The cyclones of each period were classified into five categories
 449 (see Table 2) based on the total rainfall received across the Sinai Peninsula. During the wet period, large majority
 450 of the cyclone systems (75%) occur within the categories of 1 and 2 (rainfall ranged 10-30mm/day). This implies
 451 that less significant storms have struck the Sinai during wet period. Yet, about 15% of the cyclones (with a rainfall
 452 $>40\text{mm/day}$) are potentially able to produce torrential rainfalls, which may lead to flash floods over the region.
 453



454
 455 **Figure 11.** Daily track of cyclones that precipitated ($\geq 10\text{mm/day}$) over the Sinai Peninsula during: (a) wet-period
 456 from October to March, and (b) dry-period from April to September for the climatology period of 2001-2020 (7305
 457 days). Details of all cyclones (156) classified into five categories are given in Table 2.
 458

459 **Table 2.** Cyclone tracking characteristics over the Sinai Peninsula for the period of 2001-2020.

Cyclone classification	Total precipitation range	Frequency and percentage of cyclones	
		Wet period	Dry period
Category 1	10-20 mm	77 (61.2%)	27 (87%)
Category 2	21-30 mm	17 (13.8%)	4 (13%)
Category 3	31-40 mm	12 (9.7%)	-
Category 4	41- 50 mm	10 (8.1%)	-
Category 5	> 51 mm	9 (7.2%)	-
-	-	125 (100%)	31 (100%)

460



461 Concerning the cyclogenesis, the Mediterranean Sea plays a significant role on either cyclogenesis -or
462 strengthening the cyclones passing through the region (Alpert and Shay, 1994; Flocas *et al.*, 2010; Almazroui *et al.*,
463 2014); this point becomes clear by looking at Fig. 11a. However, considerable numbers of the cyclonic systems
464 are also generated either in the North Atlantic Ocean (then, transferred into the region via passenger cyclones) -or
465 as the result of the Red Sea Trough (Krichak *et al.*, 1997; de Vries *et al.*, 2013; Hochman *et al.*, 2020). Figure 11b
466 also shows the daily tracks of 31 cyclones passed through the Sinai region during the dry period. Unlike to the wet
467 period (Fig. 11a), not only the number of cyclones reduced significantly, but also their magnitudes. The highest
468 frequency of cyclones, according to Table 2, occurs in category 1 with 27 cyclones (87%); and followed by only
469 4 cyclones (13%) in the category of 2, which have been formed within the Mediterranean (unlike category 2 of the
470 wet period) and then moved eastwards. Interestingly, no cyclonic systems (with a rainfall >10mm/day) taken place
471 within the past twenty-years during the dry period over the Sinai region.
472

473 **4 Discussion**

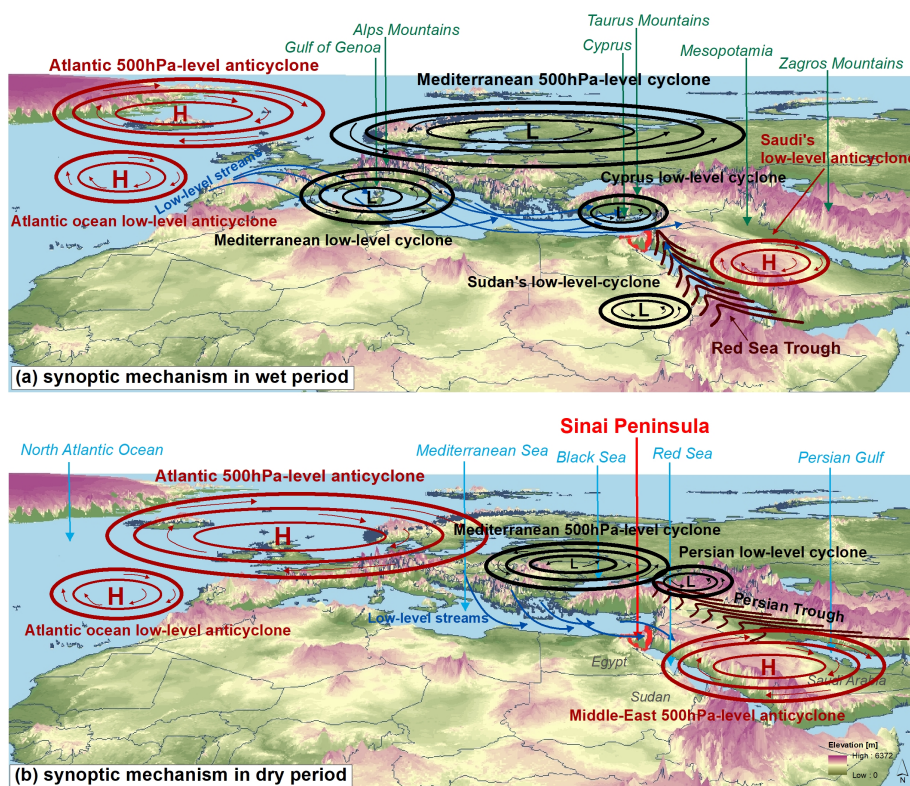
474 The main focus of this study remains on quantifying the extreme precipitation events from a statistical and synoptic
475 perspective over the Sinai Peninsula in the eastern Mediterranean basin over the past two-decades. The literature
476 of the Sinai area is poor; meaning that, although several (relevant) studies have been conducted over the eastern
477 Mediterranean (e.g. Krichak *et al.*, 1997; Alpert *et al.*, 2002; Gabella *et al.*, 2006; Nastos *et al.*, 2013; Mathbout
478 *et al.*, 2018; Rinat *et al.*, 2021); minimal studies however are available over the area, yet mostly focused on heavy
479 rainfall-related flash floods (El Afandi *et al.*, 2013; Dadamouny and Schnittler, 2016; Arnous and Omar, 2018;
480 Baldi *et al.*, 2020; El-Fakharany and Mansour, 2021). Thus, the novelty of this research is a combination of the
481 satellite-reanalysis approach for a long-term data analysis. This enabled us to quantify precipitation characteristics
482 (i.e. annual/inter-annual anomalies, monthly regime, magnitude, frequency and spatial patterns), and to discover
483 the major synoptic systems (cyclogenesis, regional circulation pattern, spatial correlation and cyclone tracking)
484 attributed to the occurrence of such heavy rainfalls across the Sinai region.

485 Our statistical analysis, as one of the first analyses over the Sinai, revealed that distribution of the rainfall events
486 significantly varies in time and space over the region. Temporally, the annual and interannual precipitation trends
487 showed mostly positive tendencies in most parts, especially towards south of the area (Fig. 2). This result is
488 consistent with the previous findings achieved over the surrounding areas in the eastern Mediterranean basin such
489 as in Israel and Gaza-Strip, as pointed out by Yosef *et al.* (2009), Ziv *et al.* (2013) and Ajjur and Riffi (2020). This
490 positive trend, however, may contribute to increase the occurrence of flash floods especially over the southern
491 Sinai, where the trends were found to be more robust along with a higher elevation gradient (see Fig. 1). Besides,
492 with respect to the temporal rainfall regime of the Sinai, as shown in Figure 3, it was found that the highest ratios
493 of the monthly precipitation from the total annual, occur in early winter; yet mostly limited to the northern Sinai.
494 This, denotes that the remaining months could experience a mild-to-severe prolonged dry-weather condition
495 throughout the year. From a spatial perspective, we also found that precipitation climatologies are quite unevenly
496 distributed across the Sinai. So that, the north/northeastern parts receive the highest (>100 mm/year) and
497 south/southwest receive the lowest (<30 mm/year) amounts of precipitation (Fig. 4a). Using the monthly frequency
498 of rainfalls in the Sinai (Fig. 5), we determined the wet-period (October-March) and dry-period (April-September)
499 in the Sinai. Although, a profound dissimilarity was found in the monthly precipitation values during the wet and
500 dry periods (ranging from 5-70mm/month to 1-3mm/month, respectively); however, their spatial patterns were
501 largely resemblance, meaning that the rainfall amounts are notably increased from the south towards northeast of
502 the Sinai region during both periods (Fig. S7).

503 Our synoptic analysis (Fig. 7) was conducted to explore the association of the synoptic systems to the precipitation
504 occurrence over the Sinai during the wet -and dry periods (2001-2020). Basically, majority of the cyclones (rainy
505 systems) affecting the study area are generated within the Mediterranean basin itself and the nearby regions, which
506 spatiotemporally are smaller and have shorter lifetimes compared to those of the north Atlantic systems; a similar
507 result was also reported by Trigo *et al.* (1999) and Buzzi *et al.* (2005). Yet, they are capable of inducing extreme
508 precipitation events and floods in some cases (Homar *et al.*, 2007). Accordingly, we also found that during wet
509 period (Fig. 12a), two major sources of cyclonic activities (cyclogenesis) are responsible for majority of the rainfall
510 events over the region; these are located in the western part (at the lee of Alps Mountains over Gulf of Genoa) and
511 eastern part (at the lee of Taurus Mountains over Cyprus) of the Mediterranean Sea. The cyclones formed over the
512 Cyprus allegedly play a significant role in the occurrence of rainfalls over the eastern Mediterranean (Saaroni *et al.*,
513 2010). Besides, another key synoptic system that plays a secondary role in the eastern Mediterranean's rainfall



514 during the wet period is the Red Sea Trough, which is developed as a result of the coexistence of the Sudan's Low
 515 and Saudi Arabia's Anticyclone (see Fig. 12a). However, allegedly the Red Sea Trough has a limited contribution
 516 to the Sinai's rainfall compared to the northeastern parts of the Mediterranean basin such as over Israel (Saaroni
 517 *et al.*, 1998, and Tsvieli and Zangvil, 2005). However, during the dry period (Fig. 12b), the number of the
 518 Mediterranean's cyclones are significantly reduced (see Fig. 11b) due to the predominance of the high-pressure
 519 systems extending from the north Atlantic and north of Africa. This situation largely prevents the rising of air and,
 520 in turn, condensation, which all limit rainfall genesis over the region during the dry period. However, as the result
 521 of the northwestwards extension of the Persian Trough into the eastern Mediterranean basin, limited cyclones
 522 develop and produce rainfalls over the eastern Mediterranean (Alpert *et al.*, 1990; Saaroni and Ziv 2000) including
 523 the Sinai region, as shown in Fig. 12b.
 524



525 **Figure 12.** Schematic representation of the dominant synoptic systems corresponding for the Sinai precipitation
 526 events (-and also the eastern Mediterranean region) in: (a) wet-period from October to March and (b) dry-period
 527 from April to September for the climatology period of 2001-2020.
 528
 529

530 With respect to the relationships of the Sinai's rainfall against key regional atmospheric variables (Figs. 9 and 10),
 531 we found meaningful correlations amongst, but varied remarkably during the wet and dry periods. In this context,
 532 a special coupling-correlation pattern was observed between the Sinai's rain against U-wind and V-wind
 533 components in wet period (Figs. 10a and 10c). However, despite a strong association between rainfall and
 534 atmospheric variables, their correlation coefficients were found to be relatively low ($< \pm 0.3$). A couple of major
 535 controlling factors, among others, could explain these low r -values. First, a long timeseries of the variables in each
 536 episode (> 3600 days), and second, a very low rate of annual rainfall over the Sinai (on average 10-100 mm/year).
 537 Regarding the former, for instance, we did examine with fewer timeseries (e.g. 100 days), then r -values doubled
 538 (-or tripled in some cases). Therefore, seemingly with a longer timeseries, more smoothed correlation coefficients
 539 are expected. It is also noted that we found that the magnitude of correlations in the dry period are notably high.
 540 This could be explained by a semi-stationary structure of the pressure systems over the region, which despite a



541 low numbers of rainy-days, play a crucial role in increase of r-values of the dry period compared to those of in the
542 wet period. This implies that, allegedly presence of the low-pressure patterns at low atmospheric levels over the
543 eastern Mediterranean during the dry period of the year are well associated with lower precipitation.

544 Finally, a daily track of cyclones precipitated over the Sinai region ($\geq 10\text{mm/day}$) was drawn separately for the
545 wet period (125 cyclones, Fig. 11a) and the dry period (31 cyclones, Fig. 11b). All cyclones were classified into
546 five categories (see Table 2) based on the total rainfall received across the Sinai. Basically, the occurrence and
547 frequency of precipitation events in the eastern Mediterranean (including the Sinai) are largely associated with the
548 passage of cyclonic systems (Ulbrich *et al.*, 2012), of which most of the cyclones are generated within the
549 Mediterranean basin in particular during the winter season (Campins *et al.*, 2000; Nissen *et al.*, 2010) (see Fig.
550 11a). Amongst, some cyclones are capable of inducing extreme precipitation and floods in the region (Buzzi *et al.*,
551 2005; Homar *et al.*, 2007). We found that about 15% of the cyclones (rainfall $>40\text{mm/day}$) in the wet period are
552 potentially able to produce torrential rainfalls leading to flash floods over the Sinai. Unlike the wet period (Fig.
553 11a), both number of cyclones (from 125 to 31) and their magnitudes (from 5 to 2 categories) reduced significantly
554 in dry period (Fig. 11b). Considering the monthly frequency of the cyclones passing through the region, during
555 the wet period (Fig. 11a) February receives the highest numbers of cyclones with 26 out of 125 (20.8%), and
556 followed by January (No. 25, 20%), December (No. 24, 19.2%), March (No. 21, 16.8%), November (No. 16,
557 12.8%), and finally October with the lowest number of 13 (10.4%). We also found that the cyclone of March 12,
558 2020 was the most significant torrential rainy-system ($>70\text{mm/day}$) ever occurred in the Sinai region (and perhaps
559 in the surrounding areas) over the past two decades. This is followed by the second extremist one occurred on
560 December 27, 2006 with more than 62mm/day rainfall over the Sinai. The monthly frequency of cyclones during
561 the dry period also showed that April was by far the first with a total number of 20 out of 31 (65%), followed by
562 May (No. 9, 29%), September (No. 2, 6%), and with a zero number for the rest of months (i.e. June, July and
563 August). Amongst, cyclones of April 5, 2006 (27mm/day) and September 30, 2012 (24mm/day) were found to be
564 the extreme ones, respectively.

566 5 Conclusions

567 The satellite remote-sensing precipitation (GPM) and reanalysis data (NCEP/NCAR) were employed in the present
568 research to explore the extreme precipitation characteristics over the Sinai Peninsula for the climatology period of
569 2001-2020. This was achieved by i) quantifying the anomaly, monthly regime, frequency and spatial patterns of
570 the extreme rainfall events, ii) investigating the synoptic-scale systems responsible for the occurrence of rainfalls,
571 and iii) determining the major tracks of cyclones during the wet and dry periods. The key findings are therefore
572 summarized into three major pillars:

- 573 i. *Spatiotemporal characteristics of rainfall*: annual/interannual precipitation trends did not show a spatial
574 coherency across the Sinai; rather indicated a spatial dependency. The precipitation regimes also
575 revealed that high ratios of annual rainfalls are mostly received in the early winter. Further, precipitation
576 climatology of the Sinai indicated northeast and southwest of the region receive in order the highest
577 ($>100\text{mm/year}$) and the lowest ($<30\text{mm/year}$) annual rainfall. Also, the distribution of extreme rainfall
578 frequencies, regardless of their thresholds, resembled. This means that highest and lowest frequencies
579 occur from January to March (wet period), and from April to September (dry period), respectively.
- 580 ii. *Synoptic atmospheric systems*: majority of cyclones precipitating over the Sinai are generated within the
581 Mediterranean basin (at leeward of the Alps and Taurus Mountains over Gulf of Genoa and Cyprus,
582 respectively), accompanied by the Red Sea Trough at lower-levels during the wet period. These systems
583 either are absent- or weakened significantly during dry period of the region; however, limited low-
584 systems are developed as the result of the Persian Trough extending northwestwards. Further, the spatial
585 correlations of the Sinai's rainfall against key regional variables at multiple levels of the atmosphere
586 revealed meaningful correlation patterns, varied largely in wet and dry periods. Amongst, the relationship
587 of Sinai's rainfall against SLP, U-V winds and vertical velocity was remarkable.
- 588 iii. *Cyclone tracking*: A total number of 125 and 31 cyclones (rainfall $\geq 10\text{mm/day}$) was tracked during wet
589 and dry periods, respectively. Amongst, 75% of the cyclones produced precipitation ranged 10-
590 30mm/day ; while about 15% generated torrential rainfall with $>40\text{mm/day}$, being capable of leading to
591 flash floods in the wet period of the Sinai. However, both frequency (from 125 to 31 cyclones) and
592 magnitude (from 5 to 2 classes) of the cyclones reduced during the dry period as compared to those
593 occurring during the wet season.



594 **Code and data availability.** A set of climate functions/indices computed using the Climate Data Operator (CDO)
595 developed at the Max-Planck-Institute for Meteorology are available at: <https://code.mpimet.mpg.de/projects/cdo>.
596 The satellite GPM and NCEP/NCAR reanalysis datasets used are publicly available at: <https://gpm.nasa.gov/> and
597 <https://psl.noaa.gov/data/gridded/data.ncep.reanalysis.html>, respectively.

598
599 **Supplement.** The supplement related to this article is available online at: (*will be added by the journal*)

600
601 **Author contributions.** MS, BH, AM, SCD, TH, JA and PL designed the study. MS, AM, SCD and PL developed
602 the research goals, and MS wrote the initial manuscript. MS and AM designed and produced the figures and tables.
603 All authors contributed to the interpretation of results and improvement of the manuscript.

604
605 **Acknowledgements.** This research was financially supported by the European centre of excellence for sustainable
606 water technology (Wetsus). The authors would like to acknowledge the Max-Planck-Institute for Meteorology for
607 developing the CDO-tool's functions used in this study to estimate a set of climate indices. A special thanks goes
608 to the NASA/Goddard Space Flight Center for providing the GPM-IMERG (V06B) satellite rainfall data. We also
609 gratefully appreciate the NOAA-NCEP/NCAR reanalysis dataset used in this research.

610 611 **References**

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