



# 1 **Exploration of a novel geoengineering solution: lighting up tropical forests at** 2 **night**

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## 8 **Abstract**

9 Plants primarily conduct photosynthesis in the daytime, offering an opportunity to increase  
10 photosynthesis and carbon sink by providing light at night. We used a fully coupled Earth  
11 System Model to quantify the carbon sequestration and climate effects of a novel carbon removal  
12 proposal: lighting up tropical forests at night via lamp networks above the forest canopy.  
13 Simulation results show that additional light increased tropical forest carbon sink by  $10.4 \pm 0.05$   
14 petagrams of carbon per year during a 16-year lighting experiment, resulting in a decrease in  
15 atmospheric CO<sub>2</sub> and suppression of global warming. In addition, local temperature and  
16 precipitation increased. The energy requirement for capturing one ton of carbon is lower than  
17 that of Direct Air Carbon Capture. When the lighting experiment was terminated, tropical forests  
18 started to release carbon slowly. This study suggests that lighting up tropical forests at night  
19 could be an emergency solution to climate change, and carbon removal actions focused on  
20 enhancing ecosystem productivity by altering environmental factors in the short term could  
21 induce post-action CO<sub>2</sub> outgassing.

## 22 **Short summary**

23 Numerical experiments with a coupled Earth System Model show that large-scale nighttime  
24 artificial lighting in tropical forests will significantly increase carbon sink, local temperature, and  
25 precipitation, and requires less energy than Direct Air Carbon Capture for capturing 1 ton  
26 carbon, suggesting that it could be a powerful climate mitigation option. Side effects include the  
27 CO<sub>2</sub> outgassing after the termination of the nighttime lighting and the impacts on local wildlife.

28 **Keywords:** climate change; Earth system model; geoengineering; carbon cycle; tropical forests

29



## 30 1. Introduction

31 Anthropogenic greenhouse gas (GHG) emissions have led the global mean temperature to  
32 increase by approximately 1.1 degree Celsius since the industrial revolution(IPCC, 2013, 2018;  
33 IPCC AR6 WGI, 2021). Changes in climate have caused impacts on natural ecosystems and  
34 human societies, such as mass ice sheet melt(Jevrejeva et al., 2016), devastating heat  
35 waves(Dosio et al., 2018), and increase in extreme climate events(Kirchmeier-Young and Zhang,  
36 2020), exposing natural and human systems to uncertainties and the risks of unsustainable  
37 development(Gao et al., 2019, 2020). Despite the scientific consensus on climate change,  
38 emission-reduction efforts have made slow or little progress with global GHG emissions  
39 continuing to rise(IPCC AR6 WGI, 2021). In this context, geoengineering options are  
40 increasingly being considered as means of deliberately intervening in Earth's climate system in  
41 the second half of the 21st century(IPCC AR6 WGI, 2021; Moore et al., 2015).

42 Existing geoengineering proposals tend to align with two fundamentally different strategies:  
43 Solar Geoengineering (SG)(Abatayo et al., 2020; Proctor et al., 2018; Robock et al., 2009) and  
44 Carbon Capture and Sequestration (CCS)(IPCC, 2005; Jones, 2008; Leung et al., 2014). SG and  
45 related techniques reduce the amount of incoming radiation from the sun typically via  
46 stratospheric aerosol injection, subsequently affecting the planet's temperature. Although they  
47 may be able to offset temperature increase rapidly, previous studies indicate the potential for  
48 political instability(Abatayo et al., 2020) and negative impacts on human health(Robock et al.,  
49 2009) and agriculture(Proctor et al., 2018). Comparatively, CCS removes carbon from the global  
50 carbon cycle by artificial machines and saves it for long-term storage or for industrial  
51 reutilization(IPCC, 2005). While technically feasible, the environmental risks for the transport  
52 and storage of CO<sub>2</sub>, limited carbon storage capability, and high cost remain large obstacles of  
53 implementing CCS(IPCC, 2005; Jones, 2008; Leung et al., 2014).

54 In this study, the authors propose a novel geoengineering solution: lighting up tropical forests at  
55 night by installing lamp networks above the forest canopy(Graham et al., 2003), which lengthens  
56 photoperiods and leads to greater photosynthesis and carbon sequestration, and helps mitigate  
57 climate change. Contrasting to traditional CCS techniques, this strategy utilizes nature carbon  
58 sink to capture and sequester CO<sub>2</sub> from air and avoids long-distance transport and geological  
59 storage.

60 Structurally intact tropical forests are by far the most efficient carbon-capture method(Mitchard,  
61 2018), and they act as an important carbon sink against rising CO<sub>2</sub> levels(Pan et al., 2011;



62 Sullivan et al., 2020). Although intact tropical forest growth is likely suffering from warming  
63 and moisture stress induced by anthropogenic greenhouse gas emissions(Aguirre-Gutiérrez et al.,  
64 2020; Doughty et al., 2015; Gatti et al., 2021; Hubau et al., 2020), light is still the primary factor  
65 limiting tropical tree growth due to cloud cover, especially during the rainy season(Boisvenue  
66 and Running, 2006; Graham et al., 2003). Previous studies have shown that longer photoperiods  
67 facilitate the bud break and flowering in tropical forests(Borchert et al., 2005; Rivera et al.,  
68 2002). A greenhouse study in 1978 showed that a tropical tree species grown for one year under  
69 a 15-hour photoperiod treatment had an average stem length twice that of the same species  
70 grown under an 8-hour photoperiod treatment(Stubblebine et al., 1978). These studies suggest  
71 that longer photoperiods might have a positive effect on vegetative growth in tropical forests.

72 Earth System Models provides state-of-the-art computer simulations of key processes and  
73 climate states across the Earth(Danabasoglu et al., 2020). In this study the authors used a fully  
74 coupled Earth System Model, Community Earth System Model version 2 (CESM2) developed  
75 by the U.S. National Center for Atmospheric Research(Danabasoglu et al., 2020), to test the  
76 carbon sequestration and climate effects of this geoengineering measure by conducting  
77 numerical lighting experiments. Briefly, we added additional diffuse visible light to tropical  
78 forest canopy at night (see Supplementary Figure 1) assuming that trees will receive light from  
79 multiple directions (e.g., multiple lamps). Tropical forest grids were defined by “Broadleaf  
80 Evergreen Tree Area Percentage” being greater than 60% between 20°N and 20°S. The lighting  
81 experiment started from 12:00 am on January 1<sup>st</sup>, 2015 (UTC time), and the simulation exercise  
82 was conducted across numerous timescales and lighting levels:

- 83 (1) Historical control simulation from 2001 to 2014
- 84 (2) 24-hour lighting experiment with various lighting powers on January 1<sup>st</sup>, 2015
- 85 (3) 16-year lighting experiment with the optimal lighting power from 2015 to 2030
- 86 (4) 20-year simulation after the experiment termination from 2031 to 2050
- 87 (5) Future control simulation from 2015 to 2050

88 Both experiment and control simulations in the future from 2015 to 2050 were on top of the  
89 Shared Socioeconomic Pathways (SSP) 126 scenario(Riahi et al., 2017). Each simulation has a  
90 spatial resolution of 1° and has two members (created from small perturbations to initial  
91 conditions) to provide uncertainty estimation. (see Methods for detailed experimental design)

92



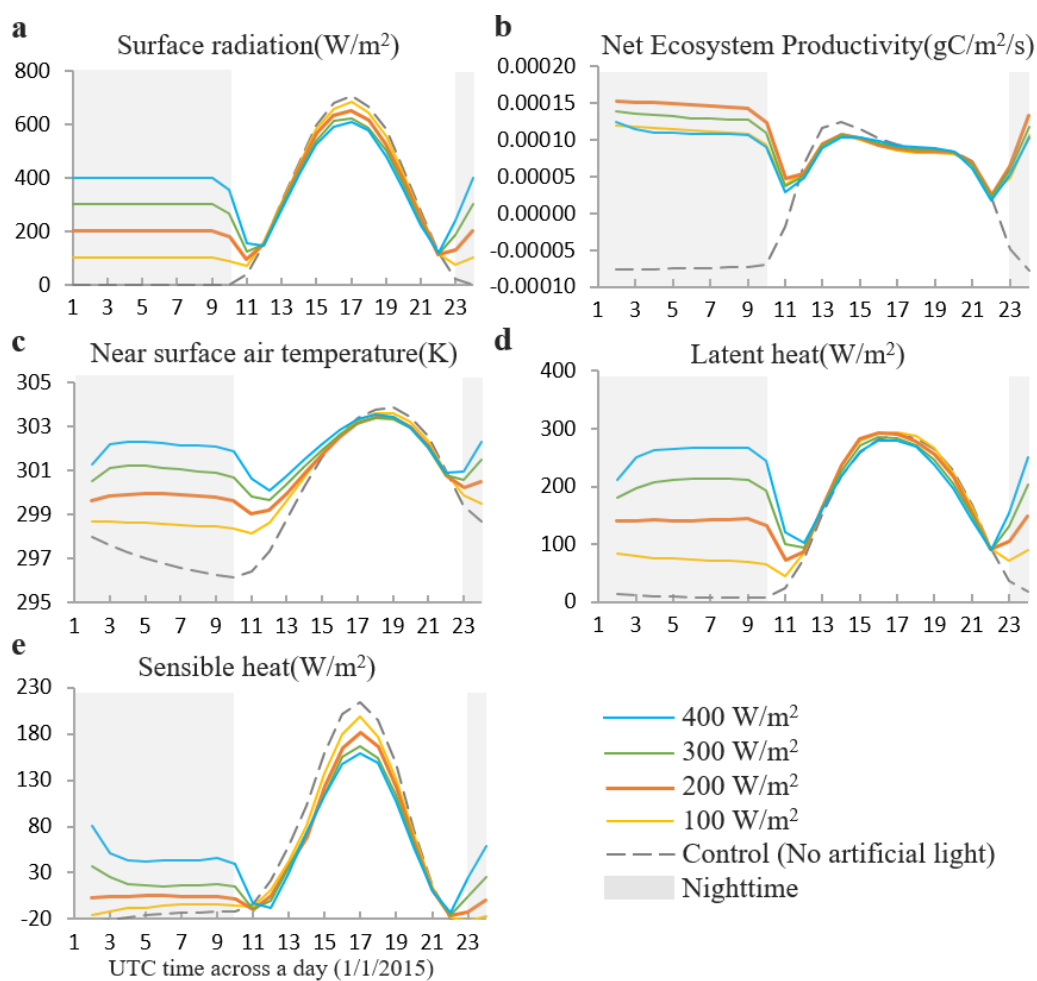
93 **2. Results**

94 2.1 24-hour lighting experiment with various lighting powers on January 1<sup>st</sup> 2015

95 Figure 1 shows the changes in carbon and energy fluxes of Amazonian tropical forests for 24  
96 hours since the start of the nighttime lighting experiment at 12:00 am January 1<sup>st</sup>, 2015 (UTC  
97 time; See Supplementary Figure 2 and 3 for African and Asian tropical forest responses).  
98 Tropical forests had a significant response to nighttime radiation, but the response was different  
99 under 100, 200, 300, and 400 W/m<sup>2</sup> lighting powers. The lighting experiment altered the  
100 nighttime energy balance and increased near surface air temperature, latent heat, and sensible  
101 heat. Higher lighting powers led to greater increases in air temperature, latent heat and sensible  
102 heat. Meanwhile, the additional light activated photosynthesis and increased Net Ecosystem  
103 Productivity (NEP). Nighttime NEP reached the peak at 200W/m<sup>2</sup> and seemed to be suppressed  
104 when the lighting power was higher. Comparison of NEP across lighting powers suggests that  
105 200W/m<sup>2</sup> is optimal in terms of activating additional photosynthesis. African and Asian tropical  
106 forests showed similar responses.  
107



108 **Fig. 1.** Amazonian tropical forest responses for 24 hours since the start of the nighttime lighting  
109 experiment at 12:00 am January 1st, 2015 (UTC time) under various nighttime lighting powers.  
110 Panel (a) refers to surface downward shortwave radiation. Nighttime NEP (b) reached the peak at  
111 200W/m<sup>2</sup>, suggesting that 200W/m<sup>2</sup> is optimal in terms of activating additional photosynthesis.  
112





113 2.2 16-year lighting experiment with the optimal lighting power from 2015 to 2030

114 The yellow lines in Figure 2 show that tropical forest carbon fluxes and climates were  
115 significantly altered by a 16-year continuous lighting experiment at night with a  $200\text{W/m}^2$   
116 power. The annual gross primary production and autotrophic respiration increased by twice near  
117 instantaneously, while the heterotrophic respiration had a slower response and increased  
118 continuously over a longer period. We purport these changes to be due to the increase in local  
119 temperature and the gradual accumulation of organic matter in the soil. Simulation results show  
120 that the lighting experiment also decreased wildfire emissions. However, the expansion of the  
121 coarse woody debris and litter carbon pool could provide more burning materials and increase  
122 wildfire risks. Overall, the net carbon uptake increased to around 25 petagrams of carbon per  
123 year ( $\text{Pg C yr}^{-1}$ ) in the beginning of the lighting experiment, although it decreased with time due  
124 to the continuous increase in heterotrophic respiration. The lighting experiment increased the net  
125 carbon uptake in tropical forests by 15.3 times over the simulation period (from  $0.68 \pm 0.02 \text{ Pg C}$   
126  $\text{yr}^{-1}$  over 2001-2014 to  $11.1 \pm 0.05 \text{ Pg C yr}^{-1}$  over 2015-2030). Among all the absorbed carbon,  
127 75% entered the vegetation carbon pool, 16% entered the coarse woody debris and litter carbon  
128 pool, and 9% entered the soil carbon pool (Figure 3-b).

129 Simulation results show that local climates were also significantly impacted (Figure 2-g,h). The  
130 annual average air temperature increased by around  $1.3^\circ\text{C}$ , and annual precipitation almost  
131 doubled. The temperature and precipitation increase showed no significant seasonal trend  
132 (Supplementary Figure 4). Globally, the atmospheric  $\text{CO}_2$  concentration dropped quickly in the  
133 first several years, while turned flat in the latter of the lighting experiment. As a result, the global  
134 average air temperature increase was suppressed by around  $0.5^\circ\text{C}$ .

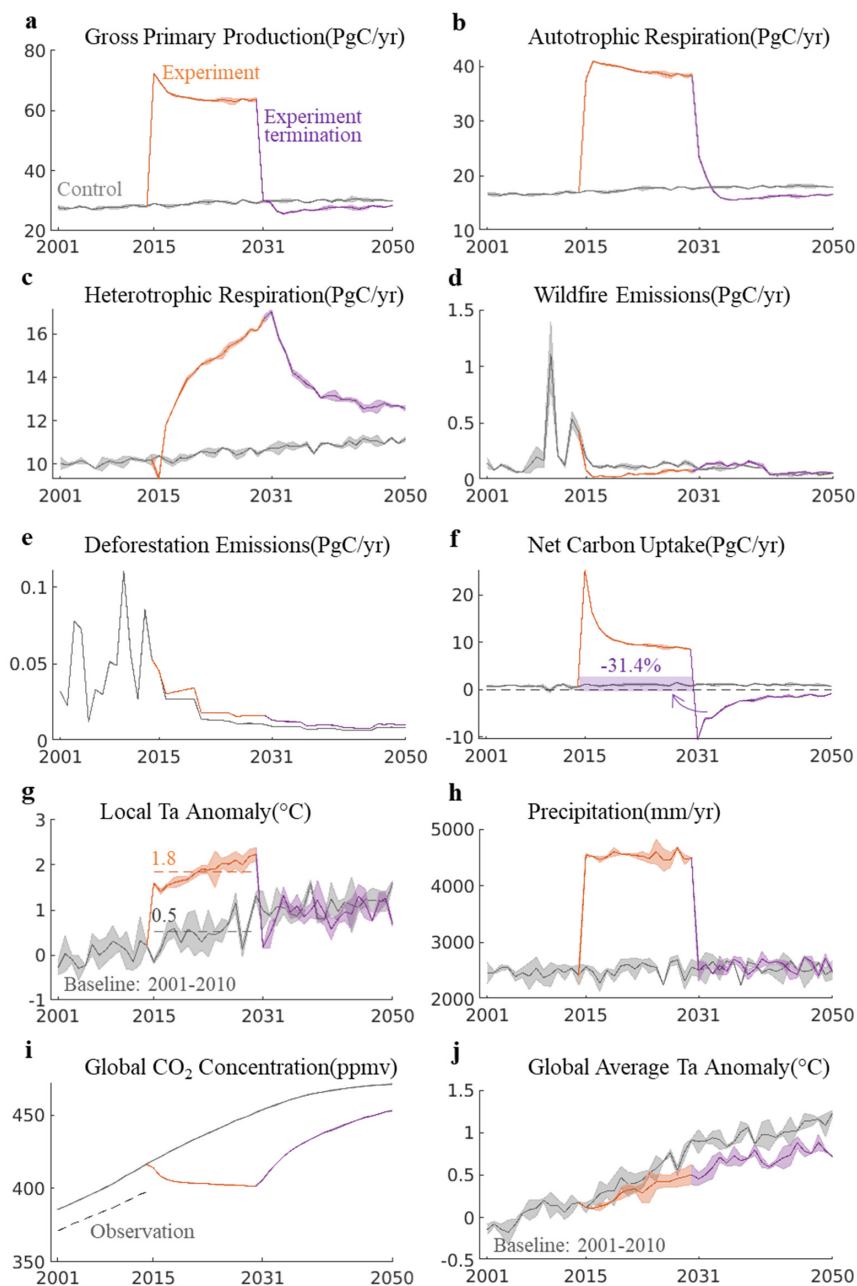
135 Amazonian, African, and Asian tropical forests present different capabilities to offset annual  
136 atmospheric carbon accumulation during the lighting experiment (Figure 4). In the current global  
137 carbon budget (Friedlingstein et al., 2019) (averaged from 2009 to 2018), approximately  $11 \pm 0.5$   
138  $\text{Pg C yr}^{-1}$  was released into atmosphere by anthropogenic activities including fossil fuel  
139 combustion and land use, among which  $2.5 \pm 0.6 \text{ Pg C yr}^{-1}$  was absorbed by ocean,  $3.2 \pm 0.6 \text{ Pg C}$   
140  $\text{yr}^{-1}$  was absorbed by land, and  $4.9 \pm 0.02 \text{ Pg C yr}^{-1}$  was accumulated in atmosphere resulting in  
141 the concerned warming and climate change. The lighting experiment enhanced Amazonian  
142 tropical forest net carbon uptake to  $6.5 \pm 0.04 \text{ Pg C yr}^{-1}$  (averaged during 2015 to 2030),  
143 suggesting that lighting up Amazonian tropical forests along could completely offset  
144 anthropogenic carbon emissions. African and Asian tropical forests showed lower capabilities



145 with the net carbon uptake being approximately  $2.0 \pm 0.002$  and  $2.6 \pm 0.008$  Pg C yr<sup>-1</sup> respectively  
146 (see Supplementary Figure 5, 6, and 7 for Amazonian, African, and Asian tropical forest carbon  
147 flux, carbon amount, and climate responses respectively).  
148



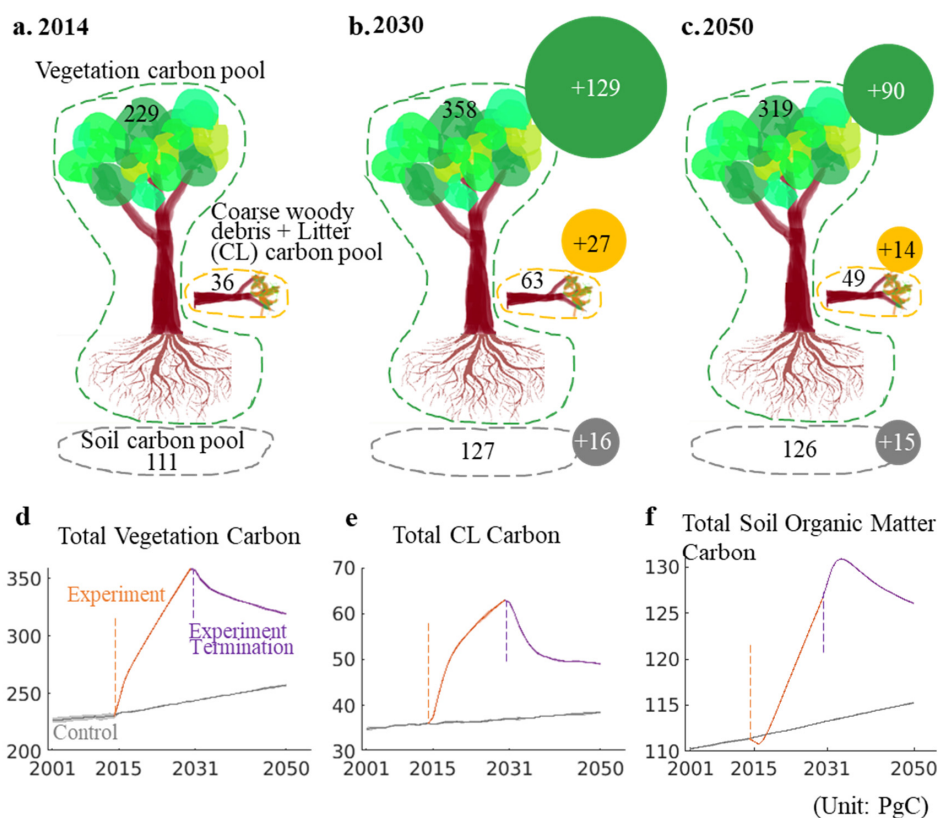
149 **Fig. 2.** Global tropical forest carbon flux and climate responses under and after the lighting  
150 experiment. Ta in panel (g & j): Near surface air temperature. Shaded areas represent  
151 uncertainties, except those in panel (f) which denote carbon released back to atmosphere after the  
152 termination of the lighting experiment.







153 **Fig. 3.** Where Did the Net Absorbed Carbon Go? Global Tropical Forest Carbon Amount  
 154 Responses. Panel (a): the current carbon amount in different carbon pools. Panel (b): carbon  
 155 amount in 2030 after 16-year lighting experiments. Panel (c): carbon amount in 2050 after 20  
 156 years since the termination of the lighting experiments. The solid circles in panel (b) and (c) refer  
 157 to carbon amount changes with respect to panel (a). The numbers in panels (a-c) are based on  
 158 panels (d-f). Tree drawing courtesy of © Ning Zeng.

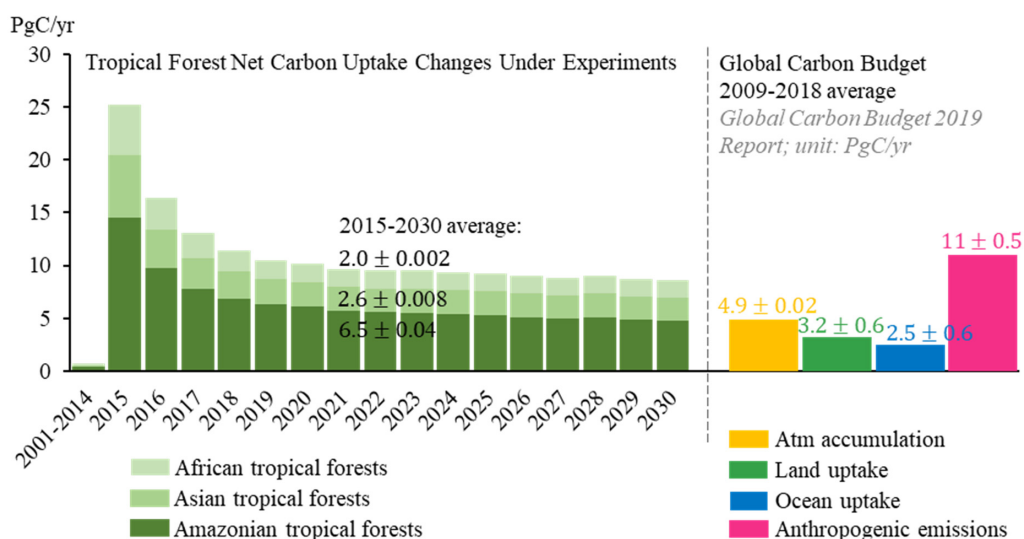


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161 **Fig. 4.** Capabilities of Amazonian, African, and Asian tropical forests to offset annual  
162 atmospheric carbon accumulation.



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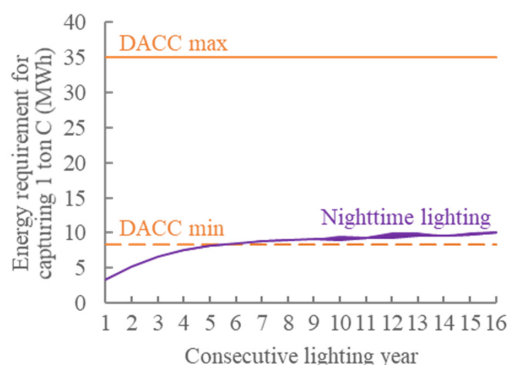
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165 We estimated the energy requirement of this strategy for capturing one ton of carbon, and  
166 compared it to that of Direct Air Carbon Capture (DACC) estimated by recent studies(Chatterjee  
167 and Huang, 2020; Realmonte et al., 2019). As the carbon uptake efficiency of the tropical forest  
168 ecosystem decreases with time when under consecutive nighttime lighting, the energy  
169 requirement for capturing one ton of carbon increases (Figure 5 purple line). Nevertheless, the  
170 energy requirement of this strategy is lower than that of DACC, or is equivalent to the most  
171 optimistic estimation of DACC's energy requirement that excludes the energy costs required for  
172 carbon transport, storage, and utilization.

173



174 **Figure 5.** Energy requirement for DACC and the nighttime lighting strategy



175

176 2.3 20-year simulation after the experiment termination from 2031 to 2050

177 The lighting experiment was terminated at 12:00 am January 1<sup>st</sup>, 2031 (UTC time), and model  
178 simulations continued for 20 years to 2050 (see the purple lines in Figure 2). The annual gross  
179 primary production and autotrophic respiration dropped quickly, ultimately reaching levels that  
180 were even lower than the control period due to a reduction in atmospheric CO<sub>2</sub> (CO<sub>2</sub> has  
181 fertilization effect in the model). Heterotrophic respiration remained high and decreased much  
182 slower at a speed 10 times lower than gross primary production and autotrophic respiration. The  
183 soil organic matter carbon pool continued to expand due to the entering of litter carbon during  
184 the first 2-3 years following the experiment termination (Figure 3-f). The vegetation carbon pool  
185 shrunk as trees produced less leaves (Figure 3-d). As a result, tropical forests turned into a net  
186 carbon source and remained so until the end of the simulation in 2050 (Figure 2-f). 31.4% of the  
187 carbon that had been absorbed during the lighting experiment was released back to the  
188 atmosphere. This number would likely be higher if the simulation continued. As a result, the  
189 global atmospheric CO<sub>2</sub> concentration returned to a level slightly lower than the control scenario.  
190 Local air temperature and precipitation returned to control levels.

191



192 **3. Discussion**

193 Physiological responses of tropical trees to near 24-hour photoperiods at the ecosystem level  
194 remains one of the biggest uncertainties in model simulations. Some field experiments indicate  
195 that higher CO<sub>2</sub> did not increase carbon sequestration of forests without added nutrients(Oren et  
196 al., 2001), suggesting tree growth might be limited by nutrient supply. The simulated local  
197 warming might also suppress tree growth(Gatti et al., 2021). Some observational evidence shows  
198 that intact tropical forest carbon sinks have been negatively influenced by warming and moisture  
199 stress(Doughty et al., 2015; Gatti et al., 2021) and might be reaching saturation(Hubau et al.,  
200 2020). However, the model predicted increase in precipitation, and previous studies have shown  
201 hydro climate plays a key role in deciding the effects of warming on tree growth(Guan et al.,  
202 2015; Reich et al., 2018). No direct evidence exists to verify the simulation results. Ecosystem-  
203 level field experiments are needed to understand how tropical forest ecosystems respond to  
204 longer photoperiods.

205 CESM2 likely overestimated the local air temperature increase in tropical forests for the  
206 omission of chemical energy stored during photosynthesis(Sellers, 1992). In CESM2 and other  
207 modern Earth System Models(Sellers, 1992), the canopy energy equation(Danabasoglu et al.,  
208 2020) uses the solar radiation absorbed by the vegetation to calculate temperature:

209 
$$-\vec{S}_v + \vec{L}_v(T_v) + H_v(T_v) + \lambda E_v(T_v) = 0 \quad (1)$$

210 where  $\vec{S}_v$  is the solar radiation absorbed by the vegetation,  $\vec{L}_v$  is the net longwave radiation  
211 absorbed by vegetation, and  $H_v$  and  $\lambda E_v$  are the sensible and latent heat fluxes from vegetation,  
212 respectively.  $\vec{L}_v, H_v,$  and  $\lambda E_v$  depend on the vegetation temperature  $T_v$ .

213 The chemical energy that is stored during photosynthesis and released by respiration is ignored  
214 as the net chemical energy usually amounts to less than 1% of absorbed insolation (around  
215 0.6%(Trenberth et al., 2009)). In our lighting experiment from 2015 to 2030, however, 17% of  
216 absorbed insolation was fixed in the ecosystem as chemical energy (Figure 2-f) and did not  
217 contribute to local air temperature increase. The model failed to exclude this chemical energy  
218 storage from the energy equation. Therefore, the model overestimated the local temperature  
219 increase. This suggests that the temperature simulation results should be treated carefully when  
220 Earth System Models are used to do extreme scenario experiments associated with  
221 biogeochemistry.



222 Tropical forests experienced significant increase in carbon sink during the lighting experiment,  
223 but ultimately transitioned from a sink to a source after the experiment was terminated (Figure 2-  
224 f). Studies(Koven et al., 2021; Tokarska and Zickfeld, 2015) investigating the effects of  
225 overshoot future scenarios (positive carbon emissions followed by net-negative emissions) on  
226 terrestrial carbon cycle have observed similar phenomenon. During a positive emissions phase,  
227 terrestrial carbon cycles tend to absorb some fraction of added CO<sub>2</sub>; however, during a removal  
228 phase they tend to release CO<sub>2</sub>. The mechanism of these phenomena is the different responding  
229 rates of vegetative primary productivity and heterotrophic respiration to lengthening and  
230 shortening photoperiods, or increasing and decreasing atmospheric CO<sub>2</sub>, with primary  
231 productivity responding much quicker than heterotrophic respiration. It is understandable when  
232 considering the diurnal pattern of forest carbon uptake. In the daytime, forests act as a carbon  
233 sink because photosynthesis is greater than respiration. In the nighttime respiration continues  
234 while photosynthesis abates, making forests a carbon source. Additional light/CO<sub>2</sub> would  
235 increase carbon sink by increasing both photosynthesis and respiration (sometimes referred to as  
236 a fertilization effect). When the additional light/CO<sub>2</sub> is removed, photosynthesis decreases  
237 quickly while respiration remains high, making forests a greater carbon source. It suggests that  
238 carbon removal actions focused on enhancing ecosystem productivity by altering environmental  
239 factors in the short term could induce this post-action CO<sub>2</sub> outgassing.

240 Overall, lighting up tropical forests at night has led to significant increase in carbon uptakes,  
241 decrease in atmospheric CO<sub>2</sub> concentration, and suppression of global warming as simulated by  
242 Earth System Model. However, it has strong side effects after the termination of nighttime  
243 lighting. In addition, local ecosystem changes could have negative impacts on local wildlife.  
244 Practical issues include the large demand for clean energy and the difficulties for  
245 implementation. From a positive standing it might be treated as an emergency climate solution if  
246 the society relies heavily on carbon removal to adjust the Earth's climate in the future. Paris  
247 Agreement set climate goals to limit global warming to well below 2 degree Celsius and  
248 preferably to 1.5 degree Celsius compared to pre-industrial levels(Lawrence et al., 2018). To  
249 accomplish the Paris Agreement's climate goals, different engineering levels (lighting powers,  
250 areas, and periods) might be needed under various anthropogenic emission scenarios, with high-  
251 emission scenarios possibly requiring high engineering levels. This study investigated the highest  
252 engineering level (lighting up global tropical forests at night with the optimal power) under a  
253 low-emission scenario (see Methods). Further research is needed to investigate the relationship



254 between engineering levels and emission scenarios in the context of global climate goals set out  
255 by the Paris Agreement(Lawrence et al., 2018).

256 Current geoengineering studies mainly focus on the evaluation of climate goals that a potential  
257 solution might or might not accomplish; however, the changes in Earth's climate after  
258 terminating a geoengineering measure tend to be overlooked. This study suggests the importance  
259 of post-geoengineering analysis in geoengineering studies.

#### 260 **4. Methods**

261 The CESM2 is an open-source community coupled model consisting of atmosphere, ocean, land,  
262 sea-ice, land-ice, river, and wave models that exchange states and fluxes via a  
263 coupler(Danabasoglu et al., 2020). In this study, we used standard CESM2 configurations and  
264 enabled all modules including the Community Atmosphere Model version 6 (CAM6), the  
265 Parallel Ocean Program version 2 (POP2) with an ocean biogeochemistry component, the  
266 Community Land Model version 5 (CLM5) with a land biogeochemistry component, CICE  
267 version 5.1.2 (CICE5), the Community Ice Sheet Model Version 2.1 (CISM2.1), the Model for  
268 Scale Adaptive River Transport (MOSART), and the NOAA WaveWatch-III ocean surface wave  
269 prediction model version 3.14 (WW3). The CESM2 is part of the Couple Model Intercomparison  
270 Project Phase 6 (CMIP6) core simulations as well as about 20 Model Intercomparison Projects  
271 (MIPs) within CMIP6. Extensive evaluation suggests that the CESM2 simulations exhibit  
272 agreement with satellite era observations of the climate mean state, seasonal cycle, and  
273 interannual variability, which has identified CESM2 as among the most realistic climate models  
274 in the world(Danabasoglu et al., 2020).

##### 275 4.1 Historical Control Simulation from 2001 to 2014

276 CESM2 has published its official historical simulation datasets from 1850-2014 on the Earth  
277 System Grid Federation (ESGF; <https://esgf-node.llnl.gov/search/cmip6>). This study analyzed  
278 the historical simulation datasets of two members from 2001 to 2014 produced by the CESM2  
279 esm-hist-BPRP case.

##### 280 4.2 Future Experiment and Control Simulations from 2015 to 2050

281 The selection of 2015 as the start year of the lighting experiment follows CMIP6 future scenario  
282 simulation rules. The future experiment simulations and control simulations were both based on  
283 the Shared Socioeconomic Pathways (SSP) 126 scenario(Riahi et al., 2017), which is a low-  
284 emission (low fossil fuel combustion and deforestation) scenario. The Earth's climate state under



285 SSP126 is close to the current climate state with respect to high-emission scenarios. Therefore,  
286 the selection of SSP126 controlled variables and allowed us to see how the lighting experiment  
287 along influences tropical forest carbon fluxes and climate. This study ran the CESM2 esm-  
288 SSP126-BPRP case with the official restart files from historical simulations (esm-hist-BPRP  
289 case). Thus, no model spin up was needed. All simulations were forced with specified  
290 greenhouse gas emissions rather than atmospheric greenhouse gas concentrations, so the  
291 atmospheric CO<sub>2</sub> concentration was prognostic and land and ocean carbon cycles feed back on  
292 atmospheric CO<sub>2</sub>. Each simulation has a nominal horizontal resolution of 1° and has two  
293 members created from small perturbations to initial climate states to estimate uncertainties.

#### 294 4.3 The Lighting Experiment Design

295 The authors modified the radiation module (Rapid Radiative Transfer Model for General  
296 circulation models, RRTMG) of CESM2 to add diffuse visible light to tropical forest canopy at  
297 night. CESM2 judges if a grid column is at daytime or nighttime by calculating its cosine (solar  
298 zenith angle) at each time step. If the cosine is positive, the land module calculates and passes  
299 the surface albedo to the atmosphere module and the atmosphere module calculates the radiation  
300 fluxes with the surface albedo and the incoming solar radiation. We made modifications in all  
301 active modules to assign tropical forests' cosine a positive value (could be any number from 0 to  
302 1) when tropical forests were at night. As a result, all modules regarded tropical forests to be at  
303 daytime at every time step.

304 CESM2 divides the solar insolation into four components: direct visible light, diffuse visible  
305 light, direct near infrared light, and diffuse near infrared light. The authors assume that the  
306 artificial light would be provided by a lamp network above the forest canopy and that trees  
307 receive light from multiple directions. Therefore, the artificial light was specified as diffuse  
308 visible light for simplification. In the model, we assigned the diffuse visible light component  
309 with 100, 200, 300, or 400 and other components with 0. The surface albedo was still calculated  
310 by the land module and passed to the atmosphere module. The radiation fluxes were then  
311 calculated by the surface albedo and the specified solar insolation.

312



313 **Code and Data Availability**

314 CESM2 is an open-source community climate model preserved at  
315 <https://doi.org/10.1029/2019MS001916>. All data have been included in the manuscript.

316 **Author contribution**

317 XG designed the study and performed the simulations. XG, SL, DW, YL, BH, and AJ  
318 contributed to the data interpretation. XG drafted the original version of the manuscript. SL and  
319 DW reviewed and edited the manuscript.

320 **Competing interests**

321 Authors declare that they have no competing interests.

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327 **References**

- 328 Abatayo, A. Lou, Bosetti, V., Casari, M., Ghidoni, R. and Tavoni, M.: Solar geoengineering may  
329 lead to excessive cooling and high strategic uncertainty, *Proc. Natl. Acad. Sci. U. S. A.*, 117(24),  
330 doi:10.1073/pnas.1916637117, 2020.
- 331 Aguirre-Gutiérrez, J., Malhi, Y., Lewis, S. L., Fauset, S., Adu-Bredu, S., Affum-Baffoe, K.,  
332 Baker, T. R., Gvozdevaite, A., Hubau, W., Moore, S., Peprah, T., Ziemińska, K., Phillips, O. L.  
333 and Oliveras, I.: Long-term droughts may drive drier tropical forests towards increased  
334 functional, taxonomic and phylogenetic homogeneity, *Nat. Commun.*, 11(1),  
335 doi:10.1038/s41467-020-16973-4, 2020.
- 336 Boisvenue, C. and Running, S. W.: Impacts of climate change on natural forest productivity -  
337 Evidence since the middle of the 20th century, *Glob. Chang. Biol.*, 12(5), doi:10.1111/j.1365-  
338 2486.2006.01134.x, 2006.
- 339 Borchert, R., Renner, S. S., Calle, Z., Havarrete, D., Tye, A., Gautier, L., Spichiger, R. and Von  
340 Hildebrand, P.: Photoperiodic induction of synchronous flowering near the Equator, *Nature*,  
341 433(7026), doi:10.1038/nature03259, 2005.





- 342 Chatterjee, S. and Huang, K. W.: Unrealistic energy and materials requirement for direct air  
343 capture in deep mitigation pathways, *Nat. Commun.*, 11(1), doi:10.1038/s41467-020-17203-7,  
344 2020.
- 345 Danabasoglu, G., Lamarque, J. F., Bacmeister, J., Bailey, D. A., DuVivier, A. K., Edwards, J.,  
346 Emmons, L. K., Fasullo, J., Garcia, R., Gettelman, A., Hannay, C., Holland, M. M., Large, W.  
347 G., Lauritzen, P. H., Lawrence, D. M., Lenaerts, J. T. M., Lindsay, K., Lipscomb, W. H., Mills,  
348 M. J., Neale, R., Oleson, K. W., Otto-Bliesner, B., Phillips, A. S., Sacks, W., Tilmes, S., van  
349 Kampenhout, L., Vertenstein, M., Bertini, A., Dennis, J., Deser, C., Fischer, C., Fox-Kemper, B.,  
350 Kay, J. E., Kinnison, D., Kushner, P. J., Larson, V. E., Long, M. C., Mickelson, S., Moore, J. K.,  
351 Nienhouse, E., Polvani, L., Rasch, P. J. and Strand, W. G.: The Community Earth System Model  
352 Version 2 (CESM2), *J. Adv. Model. Earth Syst.*, 12(2), doi:10.1029/2019MS001916, 2020.
- 353 Dosio, A., Mentaschi, L., Fischer, E. M. and Wyser, K.: Extreme heat waves under 1.5 °c and  
354 2 °c global warming, *Environ. Res. Lett.*, 13(5), doi:10.1088/1748-9326/aab827, 2018.
- 355 Doughty, C. E., Metcalfe, D. B., Girardin, C. A. J., Amézquita, F. F., Cabrera, D. G., Huasco, W.  
356 H., Silva-Espejo, J. E., Araujo-Murakami, A., Da Costa, M. C., Rocha, W., Feldpausch, T. R.,  
357 Mendoza, A. L. M., Da Costa, A. C. L., Meir, P., Phillips, O. L. and Malhi, Y.: Drought impact  
358 on forest carbon dynamics and fluxes in Amazonia, *Nature*, 519(7541),  
359 doi:10.1038/nature14213, 2015.
- 360 Friedlingstein, P., Jones, M. W., O'Sullivan, M., Andrew, R. M., Hauck, J., Peters, G. P., Peters,  
361 W., Pongratz, J., Sitch, S., Le Quéré, C., DBakker, O. C. E., Canadell, J. G., Ciais, P., Jackson,  
362 R. B., Anthoni, P., Barbero, L., Bastos, A., Bastrikov, V., Becker, M., Bopp, L., Buitenhuis, E.,  
363 Chandra, N., Chevallier, F., Chini, L. P., Currie, K. I., Feely, R. A., Gehlen, M., Gilfillan, D.,  
364 Gkritzalis, T., Goll, D. S., Gruber, N., Gutekunst, S., Harris, I., Haverd, V., Houghton, R. A.,  
365 Hurtt, G., Ilyina, T., Jain, A. K., Joetzjer, E., Kaplan, J. O., Kato, E., Goldewijk, K. K.,  
366 Korsbakken, J. I., Landschützer, P., Lauvset, S. K., Lefèvre, N., Lenton, A., Lienert, S.,  
367 Lombardozzi, D., Marland, G., McGuire, P. C., Melton, J. R., Metzl, N., Munro, D. R., Nabel, J.  
368 E. M. S., Nakaoka, S. I., Neill, C., Omar, A. M., Ono, T., Pregon, A., Pierrot, D., Poulter, B.,  
369 Rehder, G., Resplandy, L., Robertson, E., Rödenbeck, C., Séférian, R., Schwinger, J., Smith, N.,  
370 Tans, P. P., Tian, H., Tilbrook, B., Tubiello, F. N., Van Der Werf, G. R., Wiltshire, A. J. and  
371 Zaehle, S.: Global carbon budget 2019, *Earth Syst. Sci. Data*, 11(4), doi:10.5194/essd-11-1783-  
372 2019, 2019.



- 373 Gao, X., Liang, S. and He, B.: Detected global agricultural greening from satellite data, *Agric.*  
374 *For. Meteorol.*, 276–277, doi:10.1016/j.agrformet.2019.107652, 2019.
- 375 Gao, X., Liang, S. and Sauer, J.: Greening Hiatus in Eurasian Boreal Forests Since 1997 Caused  
376 by a Wetting and Cooling Summer Climate, *J. Geophys. Res. Biogeosciences*, 125(9),  
377 doi:10.1029/2020JG005662, 2020.
- 378 Gatti, L. V., Basso, L. S., Miller, J. B., Gloor, M., Gatti Domingues, L., Cassol, H. L. G., Tejada,  
379 G., Aragão, L. E. O. C., Nobre, C., Peters, W., Marani, L., Arai, E., Sanches, A. H., Corrêa, S.  
380 M., Anderson, L., Von Randow, C., Correia, C. S. C., Crispim, S. P. and Neves, R. A. L.:  
381 Amazonia as a carbon source linked to deforestation and climate change, *Nature*, 595(7867),  
382 doi:10.1038/s41586-021-03629-6, 2021.
- 383 Graham, E. A., Mulkey, S. S., Kitajima, K., Phillips, N. G. and Wright, S. J.: Cloud cover limits  
384 net CO<sub>2</sub> uptake and growth of a rainforest tree during tropical rainy seasons, *Proc. Natl. Acad.*  
385 *Sci. U. S. A.*, 100(2), doi:10.1073/pnas.0133045100, 2003.
- 386 Guan, K., Pan, M., Li, H., Wolf, A., Wu, J., Medvigy, D., Caylor, K. K., Sheffield, J., Wood, E.  
387 F., Malhi, Y., Liang, M., Kimball, J. S., Saleska, S. R., Berry, J., Joiner, J. and Lyapustin, A. I.:  
388 Photosynthetic seasonality of global tropical forests constrained by hydroclimate, *Nat. Geosci.*,  
389 8(4), doi:10.1038/ngeo2382, 2015.
- 390 Hubau, W., Lewis, S. L., Phillips, O. L., Affum-Baffoe, K., Beeckman, H., Cuní-Sanchez, A.,  
391 Daniels, A. K., Ewango, C. E. N., Fauset, S., Mukinzi, J. M., Sheil, D., Sonké, B., Sullivan, M. J.  
392 P., Sunderland, T. C. H., Taedoumg, H., Thomas, S. C., White, L. J. T., Abernethy, K. A., Adu-  
393 Bredu, S., Amani, C. A., Baker, T. R., Banin, L. F., Baya, F., Begne, S. K., Bennett, A. C.,  
394 Benedet, F., Bitariho, R., Bocko, Y. E., Boeckx, P., Boundja, P., Brienen, R. J. W., Brncic, T.,  
395 Chezeaux, E., Chuyong, G. B., Clark, C. J., Collins, M., Comiskey, J. A., Coomes, D. A.,  
396 Dargie, G. C., de Haulleville, T., Kamdem, M. N. D., Doucet, J. L., Esquivel-Muelbert, A.,  
397 Feldpausch, T. R., Fofanah, A., Foli, E. G., Gilpin, M., Gloor, E., Gonmadje, C., Gourlet-Fleury,  
398 S., Hall, J. S., Hamilton, A. C., Harris, D. J., Hart, T. B., Hockemba, M. B. N., Hladik, A., Ifo, S.  
399 A., Jeffery, K. J., Jucker, T., Yakusu, E. K., Kearsley, E., Kenfack, D., Koch, A., Leal, M. E.,  
400 Levesley, A., Lindsell, J. A., Lisingo, J., Lopez-Gonzalez, G., Lovett, J. C., Makana, J. R.,  
401 Malhi, Y., Marshall, A. R., Martin, J., Martin, E. H., Mbayu, F. M., Medjibe, V. P., Mihindou,  
402 V., Mitchard, E. T. A., Moore, S., Munishi, P. K. T., Bengone, N. N., Ojo, L., Ondo, F. E., Peh,  
403 K. S. H., Pickavance, G. C., Poulsen, A. D., Poulsen, J. R., Qie, L., Reitsma, J., Rovero, F.,  
404 Swaine, M. D., Talbot, J., Taplin, J., Taylor, D. M., Thomas, D. W., Toirambe, B., Mukendi, J.



- 405 T., Tuagben, D., Umunay, P. M., et al.: Asynchronous carbon sink saturation in African and  
406 Amazonian tropical forests, *Nature*, 579(7797), doi:10.1038/s41586-020-2035-0, 2020.
- 407 IPCC: Special report on carbon dioxide capture and storage, New York., 2005.
- 408 IPCC: Climate Change 2013: The Physical Science Basis, Contribution of Working Group I,  
409 Fifth Assess. Rep. Intergov. Panel Clim. Chang., 2013.
- 410 IPCC: IPCC 2018 Report: Global Warming of 1.5 °C, in Global Warming of 1,5 C Chapter I.,  
411 2018.
- 412 IPCC AR6 WGI: Climate Change 2021 The Physical Science Basis., 2021.
- 413 Jevrejeva, S., Jackson, L. P., Riva, R. E. M., Grinsted, A. and Moore, J. C.: Coastal sea level rise  
414 with warming above 2 °C, *Proc. Natl. Acad. Sci. U. S. A.*, 113(47),  
415 doi:10.1073/pnas.1605312113, 2016.
- 416 Jones, N.: Sucking carbon out of the air, *Nature*, doi:10.1038/news.2008.1319, 2008.
- 417 Kirchmeier-Young, M. C. and Zhang, X.: Human influence has intensified extreme precipitation  
418 in North America, *Proc. Natl. Acad. Sci. U. S. A.*, 117(24), doi:10.1073/pnas.1921628117, 2020.
- 419 Koven, C., Arora, V., Cadule, P., Fisher, R., Jones, C., Lawrence, D., Lewis, J., Lindsey, K.,  
420 Mathesius, S., Meinshausen, M., Mills, M., Nicholls, Z., Sanderson, B., Swart, N., Wieder, W.  
421 and Zickfeld, K.: 23rd Century surprises: Long-term dynamics of the climate and carbon cycle  
422 under both high and net negative emissions scenarios, *Earth Syst. Dyn. Discuss.*,  
423 doi:10.5194/esd-2021-23, 2021.
- 424 Lawrence, M. G., Schäfer, S., Muri, H., Scott, V., Oschlies, A., Vaughan, N. E., Boucher, O.,  
425 Schmidt, H., Haywood, J. and Scheffran, J.: Evaluating climate geoengineering proposals in the  
426 context of the Paris Agreement temperature goals, *Nat. Commun.*, 9(1), doi:10.1038/s41467-  
427 018-05938-3, 2018.
- 428 Leung, D. Y. C., Caramanna, G. and Maroto-Valer, M. M.: An overview of current status of  
429 carbon dioxide capture and storage technologies, *Renew. Sustain. Energy Rev.*, 39,  
430 doi:10.1016/j.rser.2014.07.093, 2014.
- 431 Mitchard, E. T. A.: The tropical forest carbon cycle and climate change, *Nature*, 559(7715),  
432 doi:10.1038/s41586-018-0300-2, 2018.



- 433 Moore, J. C., Grinsted, A., Guo, X., Yu, X., Jevrejeva, S., Rinke, A., Cui, X., Kravitz, B.,  
434 Lenton, A., Watanabe, S. and Ji, D.: Atlantic hurricane surge response to geoengineering, Proc.  
435 Natl. Acad. Sci. U. S. A., 112(45), doi:10.1073/pnas.1510530112, 2015.
- 436 Oren, R., Ellsworth, D. S., Johnsen, K. H., Phillips, N., Ewers, B. E., Maier, C., Schäfer, K. V.  
437 R., McCarthy, H., Hendrey, G., McNulty, S. G. and Katul, G. G.: Soil fertility limits carbon  
438 sequestration by forest ecosystems in a CO<sub>2</sub>-enriched atmosphere, *Nature*, 411(6836),  
439 doi:10.1038/35078064, 2001.
- 440 Pan, Y., Birdsey, R. A., Fang, J., Houghton, R., Kauppi, P. E., Kurz, W. A., Phillips, O. L.,  
441 Shvidenko, A., Lewis, S. L., Canadell, J. G., Ciais, P., Jackson, R. B., Pacala, S. W., McGuire,  
442 A. D., Piao, S., Rautiainen, A., Sitch, S. and Hayes, D.: A large and persistent carbon sink in the  
443 world's forests, *Science* (80-. ), 333(6045), doi:10.1126/science.1201609, 2011.
- 444 Proctor, J., Hsiang, S., Burney, J., Burke, M. and Schlenker, W.: Estimating global agricultural  
445 effects of geoengineering using volcanic eruptions, *Nature*, 560(7719), doi:10.1038/s41586-018-  
446 0417-3, 2018.
- 447 Realmonte, G., Drouet, L., Gambhir, A., Glynn, J., Hawkes, A., Köberle, A. C. and Tavoni, M.:  
448 An inter-model assessment of the role of direct air capture in deep mitigation pathways, *Nat.*  
449 *Commun.*, 10(1), doi:10.1038/s41467-019-10842-5, 2019.
- 450 Reich, P. B., Sendall, K. M., Stefanski, A., Rich, R. L., Hobbie, S. E. and Montgomery, R. A.:  
451 Effects of climate warming on photosynthesis in boreal tree species depend on soil moisture,  
452 *Nature*, 562(7726), doi:10.1038/s41586-018-0582-4, 2018.
- 453 Riahi, K., van Vuuren, D. P., Kriegler, E., Edmonds, J., O'Neill, B. C., Fujimori, S., Bauer, N.,  
454 Calvin, K., Dellink, R., Fricko, O., Lutz, W., Popp, A., Cuaresma, J. C., KC, S., Leimbach, M.,  
455 Jiang, L., Kram, T., Rao, S., Emmerling, J., Ebi, K., Hasegawa, T., Havlik, P., Humpenöder, F.,  
456 Da Silva, L. A., Smith, S., Stehfest, E., Bosetti, V., Eom, J., Gernaat, D., Masui, T., Rogelj, J.,  
457 Strefler, J., Drouet, L., Krey, V., Luderer, G., Harmsen, M., Takahashi, K., Baumstark, L.,  
458 Doelman, J. C., Kainuma, M., Klimont, Z., Marangoni, G., Lotze-Campen, H., Obersteiner, M.,  
459 Tabeau, A. and Tavoni, M.: The Shared Socioeconomic Pathways and their energy, land use, and  
460 greenhouse gas emissions implications: An overview, *Glob. Environ. Chang.*, 42,  
461 doi:10.1016/j.gloenvcha.2016.05.009, 2017.



- 462 Rivera, G., Elliott, S., Caldas, L. S., Nicolossi, G., Coradin, V. T. and Borchert, R.: Increasing  
463 day-length induces spring flushing of tropical dry forest trees in the absence of rain, *Trees -*  
464 *Struct. Funct.*, 16(7), doi:10.1007/s00468-002-0185-3, 2002.
- 465 Robock, A., Marquardt, A., Kravitz, B. and Stenchikov, G.: Benefits, risks, and costs of  
466 stratospheric geoengineering, *Geophys. Res. Lett.*, 36(19), doi:10.1029/2009GL039209, 2009.
- 467 Sellers, P.: Biophysical models of land surface processes. In *Climate System Modelling*,  
468 Trenberth KE (ed.), Cambridge University Press., 1992.
- 469 Stubblebine, W., Langenheim, J. H. and Lincoln, D.: Vegetative Response to Photoperiod in the  
470 Tropical Leguminous Tree *Hymenaea courbaril* L, *Biotropica*, 10(1), doi:10.2307/2388100,  
471 1978.
- 472 Sullivan, M. J. P., Lewis, S. L., Affum-Baffoe, K., Castilho, C., Costa, F., Sanchez, A. C.,  
473 Ewango, C. E. N., Hubau, W., Marimon, B., Monteagudo-Mendoza, A., Qie, L., Sonké, B.,  
474 Martinez, R. V., Baker, T. R., Brienen, R. J. W., Feldpausch, T. R., Galbraith, D., Gloor, M.,  
475 Malhi, Y., Aiba, S. I., Alexiades, M. N., Almeida, E. C., De Oliveira, E. A., Dávila, E. Á.,  
476 Loayza, P. A., Andrade, A., Vieira, S. A., Aragão, L. E. O. C., Araujo-Murakami, A., Arets, E. J.  
477 M. M., Arroyo, L., Ashton, P., Gerardo Aymard, C., Baccaro, F. B., Banin, L. F., Baraloto, C.,  
478 Camargo, P. B., Barlow, J., Barroso, J., Bastin, J. F., Batterman, S. A., Beeckman, H., Begne, S.  
479 K., Bennett, A. C., Berenguer, E., Berry, N., Blanc, L., Boeckx, P., Bogaert, J., Bonal, D.,  
480 Bongers, F., Bradford, M., Brearley, F. Q., Brncic, T., Brown, F., Burban, B., Camargo, J. L.,  
481 Castro, W., Céron, C., Ribeiro, S. C., Moscoso, V. C., Chave, J., Chezeaux, E., Clark, C. J., De  
482 Souza, F. C., Collins, M., Comiskey, J. A., Valverde, F. C., Medina, M. C., Da Costa, L.,  
483 Dančsák, M., Dargie, G. C., Davies, S., Cardozo, N. D., De Haulleville, T., De Medeiros, M. B.,  
484 Del Aguila Pasquel, J., Derroire, G., Di Fiore, A., Doucet, J. L., Dourdain, A., Droissart, V.,  
485 Duque, L. F., Ekoungoulou, R., Elias, F., Erwin, T., Esquivel-Muelbert, A., Fauset, S., Ferreira,  
486 J., Llampazo, G. F., Foli, E., Ford, A., Gilpin, M., Hall, J. S., Hamer, K. C., Hamilton, A. C.,  
487 Harris, D. J., Hart, T. B., Hédli, R., et al.: Long-term thermal sensitivity of earth's tropical forests,  
488 *Science* (80-. ), 368(6493), doi:10.1126/science.aaw7578, 2020.
- 489 Tokarska, K. B. and Zickfeld, K.: The effectiveness of net negative carbon dioxide emissions in  
490 reversing anthropogenic climate change, *Environ. Res. Lett.*, 10(9), doi:10.1088/1748-  
491 9326/10/9/094013, 2015.



- 492 Trenberth, K. E., Fasullo, J. T. and Kiehl, J.: Earth's global energy budget, Bull. Am. Meteorol.  
493 Soc., 90(3), doi:10.1175/2008BAMS2634.1, 2009.