

1 **Exploration of a novel geoengineering solution: lighting up tropical forests at**  
2 **night**

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8 **Abstract**

9 Plants primarily conduct photosynthesis in the daytime, offering an opportunity to increase  
10 photosynthesis and carbon sink by providing light at night. We used a fully coupled Earth  
11 System Model to quantify the carbon sequestration and climate effects of a novel carbon removal  
12 proposal: lighting up tropical forests at night via lamp networks above the forest canopy.  
13 Simulation results show that additional light increased tropical forest carbon sink by  $10.4 \pm 0.05$   
14 petagrams of carbon per year during a 16-year lighting experiment, resulting in a decrease in  
15 atmospheric CO<sub>2</sub> and suppression of global warming. In addition, local temperature and  
16 precipitation increased. The energy requirement for capturing one ton of carbon is lower than  
17 that of Direct Air Carbon Capture. When the lighting experiment was terminated, tropical forests  
18 started to release carbon slowly. This study suggests that lighting up tropical forests at night  
19 could be an emergency solution to climate change, and carbon removal actions focused on  
20 enhancing ecosystem productivity by altering environmental factors in the short term could  
21 induce post-action CO<sub>2</sub> outgassing.

22 **Short summary**

23 Numerical experiments with a coupled Earth System Model show that large-scale nighttime  
24 artificial lighting in tropical forests will significantly increase carbon sink, local temperature, and  
25 precipitation, and requires less energy than Direct Air Carbon Capture for capturing 1 ton  
26 carbon, suggesting that it could be a powerful climate mitigation option. Side effects include the  
27 CO<sub>2</sub> outgassing after the termination of the nighttime lighting and the impacts on local wildlife.

28 **Keywords:** climate change; Earth system model; geoengineering; carbon cycle; tropical forests

29

## 30 **1. Introduction**

31 Anthropogenic greenhouse gas (GHG) emissions have led the global mean temperature to  
32 increase by approximately 1.1 degree Celsius since the industrial revolution(IPCC, 2013, 2018;  
33 IPCC AR6 WGI, 2021). Changes in climate have caused impacts on natural ecosystems and  
34 human societies, such as mass ice sheet melt(Jevrejeva et al., 2016), devastating heat  
35 waves(Dosio et al., 2018), and increase in extreme climate events(Kirchmeier-Young and Zhang,  
36 2020), exposing natural and human systems to uncertainties and the risks of unsustainable  
37 development(Gao et al., 2019, 2020). Despite the scientific consensus on climate change,  
38 emission-reduction efforts have made slow or little progress with global GHG emissions  
39 continuing to rise(IPCC AR6 WGI, 2021). In this context, geoengineering options are  
40 increasingly being considered as means of deliberately intervening in Earth's climate system in  
41 the second half of the 21st century(IPCC AR6 WGI, 2021; Moore et al., 2015).

42 Existing geoengineering proposals tend to align with two fundamentally different strategies:  
43 Solar Geoengineering (SG)(Abatayo et al., 2020; Proctor et al., 2018; Robock et al., 2009) and  
44 Carbon Capture and Sequestration (CCS)(IPCC, 2005; Jones, 2008; Leung et al., 2014). SG and  
45 related techniques reduce the amount of incoming radiation from the sun typically via  
46 stratospheric aerosol injection, subsequently affecting the planet's temperature. Although they  
47 may be able to offset temperature increase rapidly, previous studies indicate the potential for  
48 political instability(Abatayo et al., 2020) and negative impacts on human health(Robock et al.,  
49 2009) and agriculture(Proctor et al., 2018). Comparatively, CCS removes carbon from the global  
50 carbon cycle by artificial machines and saves it for long-term storage or for industrial  
51 reutilization(IPCC, 2005). While technically feasible, the environmental risks for the transport  
52 and storage of CO<sub>2</sub>, limited carbon storage capability, and high cost remain large obstacles of  
53 implementing CCS(IPCC, 2005; Jones, 2008; Leung et al., 2014).

54 In this study, the authors propose a novel geoengineering solution: lighting up tropical forests at  
55 night by installing lamp networks above the forest canopy(Graham et al., 2003), which lengthens  
56 photoperiods and leads to greater photosynthesis and carbon sequestration, and helps mitigate  
57 climate change. Contrasting to traditional CCS techniques, this strategy utilizes nature carbon  
58 sink to capture and sequester CO<sub>2</sub> from air and avoids long-distance transport and geological  
59 storage.

60 Structurally intact tropical forests are by far the most efficient carbon-capture method(Mitchard,  
61 2018), and they act as an important carbon sink against rising CO<sub>2</sub> levels(Pan et al., 2011;

62 Sullivan et al., 2020). Although intact tropical forest growth is likely suffering from warming  
63 and moisture stress induced by anthropogenic greenhouse gas emissions(Aguirre-Gutiérrez et al.,  
64 2020; Doughty et al., 2015; Gatti et al., 2021; Hubau et al., 2020), light is still the primary factor  
65 limiting tropical tree growth due to cloud cover, especially during the rainy season(Boisvenue  
66 and Running, 2006; Graham et al., 2003). Studies on the photoperiodic control of tropical trees’  
67 growth typically fall into two categories: physiological field observations under seasonal  
68 variations of day length(Borchert et al., 2005; Pires et al., 2018; Rivera et al., 2002), and  
69 physiological greenhouse observations under experimental variations of the photoperiod(Dixit  
70 and Singh, 2014; Djerrab et al., 2021; Luo et al., 2021; Stubblebine et al., 1978). Field  
71 observations have shown that longer photoperiods facilitate the bud break and flowering in  
72 tropical forests(Borchert et al., 2005; Pires et al., 2018; Rivera et al., 2002). Greenhouse  
73 experiments either lengthen or shorten photoperiods, and results suggest that short photoperiods  
74 reduce plant growth rate and lead to thinner leaves and lower chlorophyll content(Djerrab et al.,  
75 2021; Luo et al., 2021), while long photoperiods increase stem growth rate and stimulate tree  
76 growth(Dixit and Singh, 2014; Stubblebine et al., 1978). These studies are more focused on  
77 specific tropical plant species and tend to agree that longer photoperiods might have a positive  
78 effect on vegetative growth in tropical forests. Ecosystem-level field experiments are critical for  
79 taking into account key environmental factors that are missing in greenhouse experiments (e.g.  
80 water and nutrition constraints), and for informing model parameterizations, although they are  
81 lacking so far.

82 Earth System Models provides state-of-the-art computer simulations of key processes and  
83 climate states across the Earth(Danabasoglu et al., 2020). In this study the authors used a fully  
84 coupled Earth System Model, Community Earth System Model version 2 (CESM2) developed  
85 by the U.S. National Center for Atmospheric Research(Danabasoglu et al., 2020), to test the  
86 carbon sequestration and climate effects of this geoengineering measure by conducting  
87 numerical lighting experiments. Briefly, we added additional diffuse visible light to tropical  
88 forest canopy at night (see Supplementary Figure 1) assuming that trees will receive light from  
89 multiple directions (e.g., multiple lamps). Tropical forest grids were defined by “Broadleaf  
90 Evergreen Tree Area Percentage” being greater than 60% between 20°N and 20°S. The lighting  
91 experiment started from 12:00 am on January 1<sup>st</sup>, 2015 (UTC time), and the simulation exercise  
92 was conducted across numerous timescales and lighting levels:

- 93 (1) Historical control simulation from 2001 to 2014
- 94 (2) 24-hour lighting experiment with various lighting powers on January 1<sup>st</sup>, 2015

- 95 (3) 16-year lighting experiment with the optimal lighting power from 2015 to 2030  
96 (4) 20-year simulation after the experiment termination from 2031 to 2050  
97 (5) Future control simulation from 2015 to 2050  
98 Both experiment and control simulations in the future from 2015 to 2050 were on top of the  
99 Shared Socioeconomic Pathways (SSP) 126 scenario(Riahi et al., 2017). Each simulation has a  
100 spatial resolution of  $1^\circ$  and has two members (created from small perturbations to initial  
101 conditions) to provide uncertainty estimation. (See Methods for detailed experimental design)

102

## 103 **2. Methods**

104 The CESM2 is an open-source community coupled model consisting of atmosphere, ocean, land,  
105 sea-ice, land-ice, river, and wave models that exchange states and fluxes via a  
106 coupler(Danabasoglu et al., 2020). In this study, we used standard CESM2 configurations and  
107 enabled all modules including the Community Atmosphere Model version 6 (CAM6), the  
108 Parallel Ocean Program version 2 (POP2) with an ocean biogeochemistry component, the  
109 Community Land Model version 5 (CLM5) with a land biogeochemistry component, CICE  
110 version 5.1.2 (CICE5), the Community Ice Sheet Model Version 2.1 (CISM2.1), the Model for  
111 Scale Adaptive River Transport (MOSART), and the NOAA WaveWatch-III ocean surface wave  
112 prediction model version 3.14 (WW3). The CESM2 is part of the Couple Model Intercomparison  
113 Project Phase 6 (CMIP6) core simulations as well as about 20 Model Intercomparison Projects  
114 (MIPs) within CMIP6. Extensive evaluation suggests that the CESM2 simulations exhibit  
115 agreement with satellite era observations of the climate mean state, seasonal cycle, and  
116 interannual variability, which has identified CESM2 as among the most realistic climate models  
117 in the world(Danabasoglu et al., 2020).

### 118 2.1 Historical control simulations from 2001 to 2014

119 CESM2 has published its official historical simulation datasets from 1850-2014 on the Earth  
120 System Grid Federation (ESGF; <https://esgf-node.llnl.gov/search/cmip6>). This study analyzed  
121 the historical simulation datasets of two members from 2001 to 2014 produced by the CESM2  
122 esm-hist-BPRP case.

### 123 2.2 Future experiment and control simulations from 2015 to 2050

124 The selection of 2015 as the start year of the lighting experiment follows CMIP6 future scenario  
125 simulation rules. The future experiment simulations and control simulations were both based on

126 the Shared Socioeconomic Pathways (SSP) 126 scenario (Riahi et al., 2017), which is a low-  
127 emission (low fossil fuel combustion and deforestation) scenario. The Earth's climate state under  
128 SSP126 is close to the current climate state with respect to high-emission scenarios. Therefore,  
129 the selection of SSP126 controlled variables and allowed us to see how the lighting experiment  
130 along influences tropical forest carbon fluxes and climate. This study ran the CESM2 esm-  
131 SSP126-BPRP case with the official restart files from historical simulations (esm-hist-BPRP  
132 case). Thus, no model spin up was needed. All simulations were forced with specified  
133 greenhouse gas emissions rather than atmospheric greenhouse gas concentrations, so the  
134 atmospheric CO<sub>2</sub> concentration was prognostic and land and ocean carbon cycles feed back on  
135 atmospheric CO<sub>2</sub>. Each simulation has a nominal horizontal resolution of 1° and has two  
136 members created from small perturbations to initial climate states to estimate uncertainties.

### 137 2.3 The lighting experiment design

138 The authors modified the radiation module (Rapid Radiative Transfer Model for General  
139 circulation models, RRTMG) of CESM2 to add diffuse visible light to tropical forest canopy at  
140 night. CESM2 determines if a grid column is at daytime or nighttime by calculating its cosine  
141 (solar zenith angle) at each time step. A negative cosine indicates the grid column is at nighttime  
142 and the incoming solar radiation would be assigned with zero. A positive cosine indicates  
143 daytime, and the cosine value would be used to calculate incoming solar radiation. The land  
144 module then calculates and passes the surface albedo to the atmosphere module and the  
145 atmosphere module calculates the radiation fluxes with the surface albedo and the model-  
146 calculated incoming solar radiation. We made modifications in all active modules to switch the  
147 sign of tropical forests' cosine from negative to positive when tropical forests were at night. As a  
148 result, all modules regarded tropical forests to be at daytime at every time step.

149 CESM2 divides the incoming solar radiation into four components: direct visible light, diffuse  
150 visible light, direct near infrared light, and diffuse near infrared light. The authors assume that  
151 the artificial light would be provided by a lamp network above the forest canopy and that trees  
152 receive light from multiple directions. Therefore, the artificial light was specified as diffuse  
153 visible light for simplification. In the model, we assigned the diffuse visible light component of  
154 the incoming solar radiation with 100, 200, 300, or 400 and other components with 0. The  
155 surface albedo was still calculated by the land module and passed to the atmosphere module. The  
156 radiation fluxes were then calculated by the model-calculated surface albedo and manually-  
157 specified solar insolation.

158 2.4 The calculation of the energy requirement for capturing one ton of carbon

159 
$$E = (Power \times Area \times Hours)/Carbon \quad (1)$$

160 where E is the energy requirement for capturing one ton of carbon per year; Power is 200W/m<sup>2</sup>  
161 (nighttime lighting power); Area is the tropical forest area  $10.71 \times 10^{10}m^2$  (CESM2 output);  
162 Hours is the amount of nighttime lighting hours per year:  $365 \times 11$ ; Carbon is the net carbon  
163 uptake per year (Fig. 2-f) simulated by CESM2. There is no assumed data in this calculation.

164

165 **3. Results**

166 3.1 24-hour lighting experiment with various lighting powers on January 1<sup>st</sup> 2015

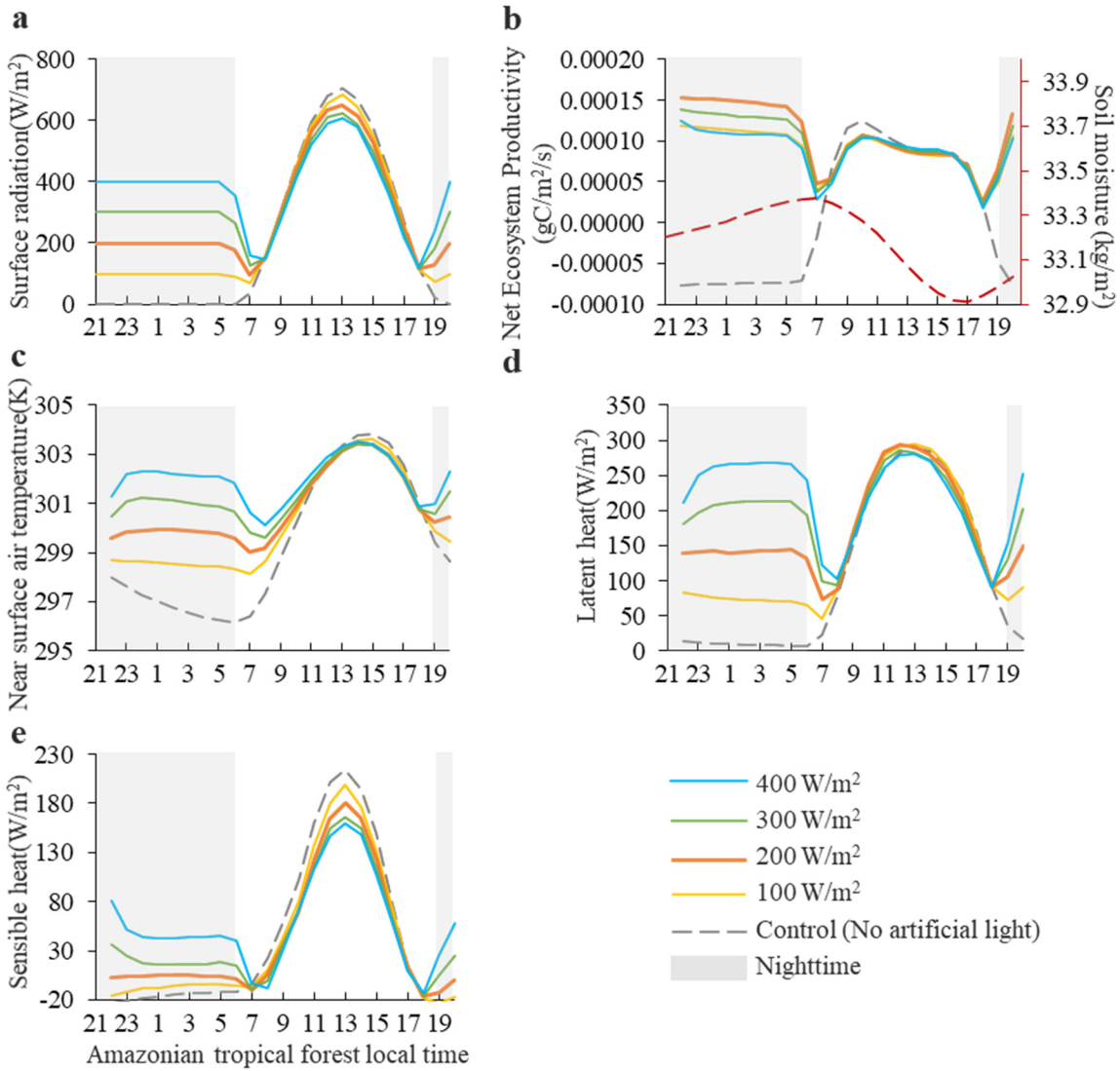
167 Figure 1 shows the changes in carbon and energy fluxes of Amazonian tropical forests for 24  
168 hours since the start of the nighttime lighting experiment at 12:00 am January 1<sup>st</sup>, 2015 (UTC  
169 time; See Supplementary Figure 2 and 3 for African and Asian tropical forest responses).

170 Tropical forests had a significant response to nighttime radiation, but the response was different  
171 under 100, 200, 300, and 400 W/m<sup>2</sup> lighting powers. The lighting experiment altered the  
172 nighttime energy balance and increased near surface air temperature, latent heat, and sensible  
173 heat. Higher lighting powers led to greater increases in air temperature, latent heat and sensible  
174 heat. Meanwhile, the additional light activated photosynthesis and increased Net Ecosystem  
175 Productivity (NEP). Nighttime NEP reached the peak at 200W/m<sup>2</sup> and seemed to be suppressed  
176 when the lighting power was higher. Comparison of NEP across lighting powers suggests that  
177 200W/m<sup>2</sup> is optimal in terms of activating additional photosynthesis. The nighttime NEP is  
178 higher than daytime because nighttime surface radiation is solely diffuse visible light while  
179 daytime surface radiation is composed of direct NIR (~16%), diffuse NIR (~30%), direct visible  
180 light (~15%), and diffuse visible light (~39%). African and Asian tropical forests showed similar  
181 responses.

182 During daytime in the control simulation, the maximum NEP shows up around 9:00-11:00 am  
183 (Fig. 1-b). It is not likely to be due to clouds according to the diurnal pattern of the surface  
184 downward shortwave radiation (Fig.1-a). We examined the diurnal curve of the soil moisture  
185 (the red dash line in Fig. 1-b), and it seems to be due to soil moisture stress. Soil moisture was  
186 consumed quickly in the morning which led to water stress for plant growth in the afternoon. The  
187 soil moisture pattern also explains the biased distribution of daytime surface air temperature  
188 (Fig.1-c), and slightly biased daytime latent heat (Fig.1-d), and daytime sensible heat (Fig.1-e).

189

190 **Fig. 1.** Amazonian tropical forest responses for 24 hours since the start of the nighttime lighting  
 191 experiment at 12:00 am January 1st, 2015 (UTC time) under various nighttime lighting powers.  
 192 Panel (a) refers to surface downward shortwave radiation. Nighttime NEP (b) reached the peak at  
 193  $200\text{W/m}^2$ , suggesting that  $200\text{W/m}^2$  is optimal in terms of activating additional photosynthesis.



194



195 3.2 16-year lighting experiment with the optimal lighting power from 2015 to 2030

196 The yellow lines in Figure 2 show that tropical forest carbon fluxes and climates were  
197 significantly altered by a 16-year continuous lighting experiment at night with a 200W/m<sup>2</sup>  
198 power. The annual gross primary production and autotrophic respiration increased by twice near  
199 instantaneously, while the heterotrophic respiration had a slower response and increased  
200 continuously over a longer period. We purport these changes to be due to the increase in local  
201 temperature and the gradual accumulation of organic matter in the soil. Simulation results show  
202 that the lighting experiment also decreased wildfire emissions as soil moisture increased despite  
203 the expansion of the coarse woody debris and litter carbon pool that provides potential burning  
204 materials. Overall, the net carbon uptake increased to around 25 petagrams of carbon per year  
205 (Pg C yr<sup>-1</sup>) in the beginning of the lighting experiment, although it decreased with time due to  
206 the continuous increase in heterotrophic respiration. The lighting experiment increased the net  
207 carbon uptake in tropical forests by 15.3 times over the simulation period (from 0.68±0.02 Pg C  
208 yr<sup>-1</sup> over 2001-2014 to 11.1±0.05 Pg C yr<sup>-1</sup> over 2015-2030). Among all the absorbed carbon,  
209 75% entered the vegetation carbon pool, 16% entered the coarse woody debris and litter carbon  
210 pool, and 9% entered the soil carbon pool (Figure 3-b).

211 Simulation results show that local climates were also significantly impacted (Figure 2-g,h). The  
212 annual average air temperature increased by around 1.3°C, and annual precipitation almost  
213 doubled. The temperature and precipitation increase showed no significant seasonal trend  
214 (Supplementary Figure 4). Globally, the atmospheric CO<sub>2</sub> concentration dropped quickly in the  
215 first several years, while turned flat in the latter of the lighting experiment. As a result, the global  
216 average air temperature increase was suppressed by around 0.5°C.

217 Amazonian, African, and Asian tropical forests present different capabilities to offset annual  
218 atmospheric carbon accumulation during the lighting experiment (Figure 4). In the current global  
219 carbon budget(Friedlingstein et al., 2019) (averaged from 2009 to 2018), approximately 11±0.5  
220 Pg C yr<sup>-1</sup> was released into atmosphere by anthropogenic activities including fossil fuel  
221 combustion and land use, among which 2.5±0.6 Pg C yr<sup>-1</sup> was absorbed by ocean, 3.2±0.6 Pg C  
222 yr<sup>-1</sup> was absorbed by land, and 4.9±0.02 Pg C yr<sup>-1</sup> was accumulated in atmosphere resulting in  
223 the concerned warming and climate change. The lighting experiment enhanced Amazonian  
224 tropical forest net carbon uptake to 6.5±0.04 Pg C yr<sup>-1</sup> (averaged during 2015 to 2030),  
225 suggesting that lighting up Amazonian tropical forests along could completely offset  
226 anthropogenic carbon emissions. African and Asian tropical forests showed lower capabilities

227 with the net carbon uptake being approximately  $2.0 \pm 0.002$  and  $2.6 \pm 0.008$  Pg C yr<sup>-1</sup> respectively  
228 (see Supplementary Figure 5, 6, and 7 for Amazonian, African, and Asian tropical forest carbon  
229 flux, carbon amount, and climate responses respectively).

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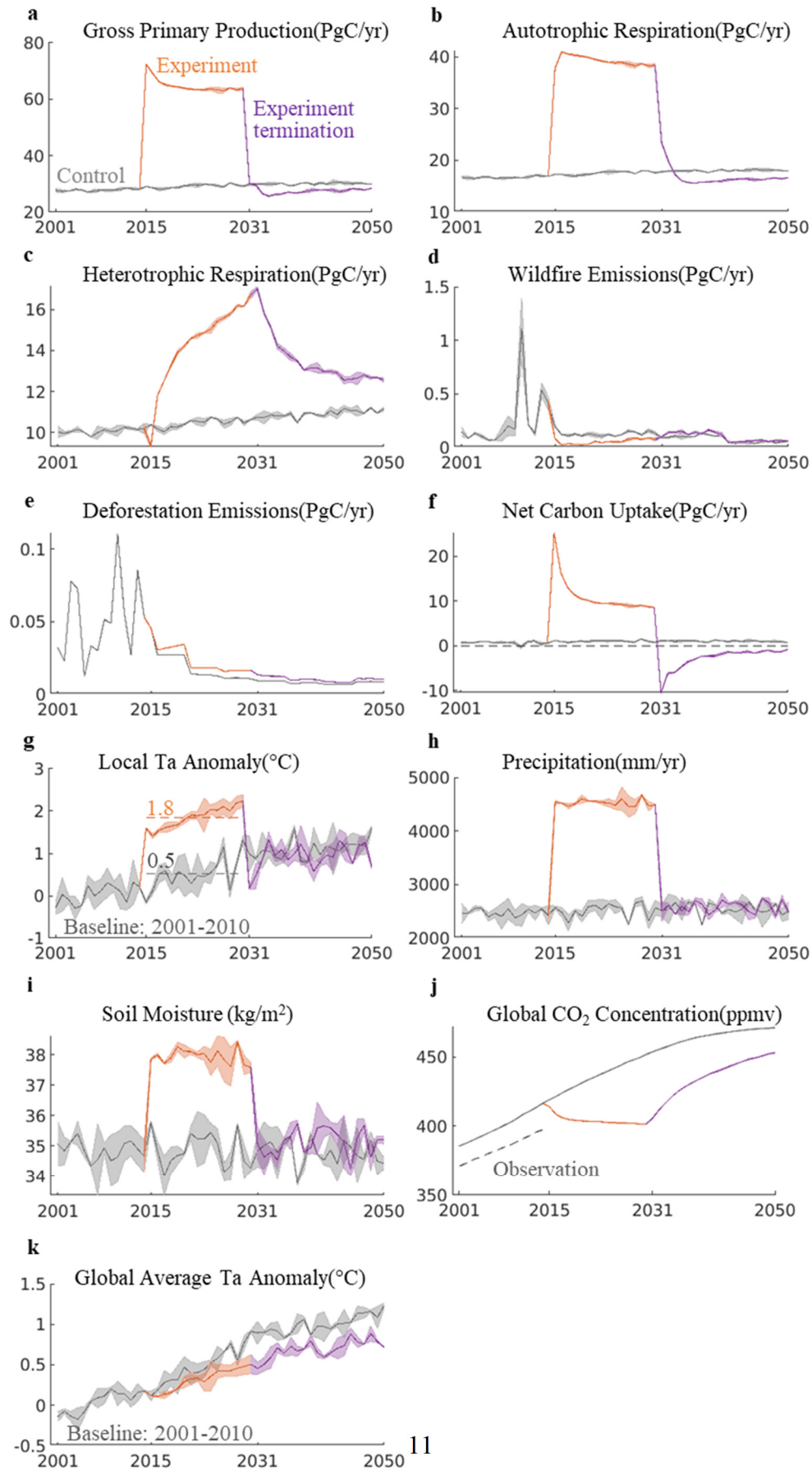
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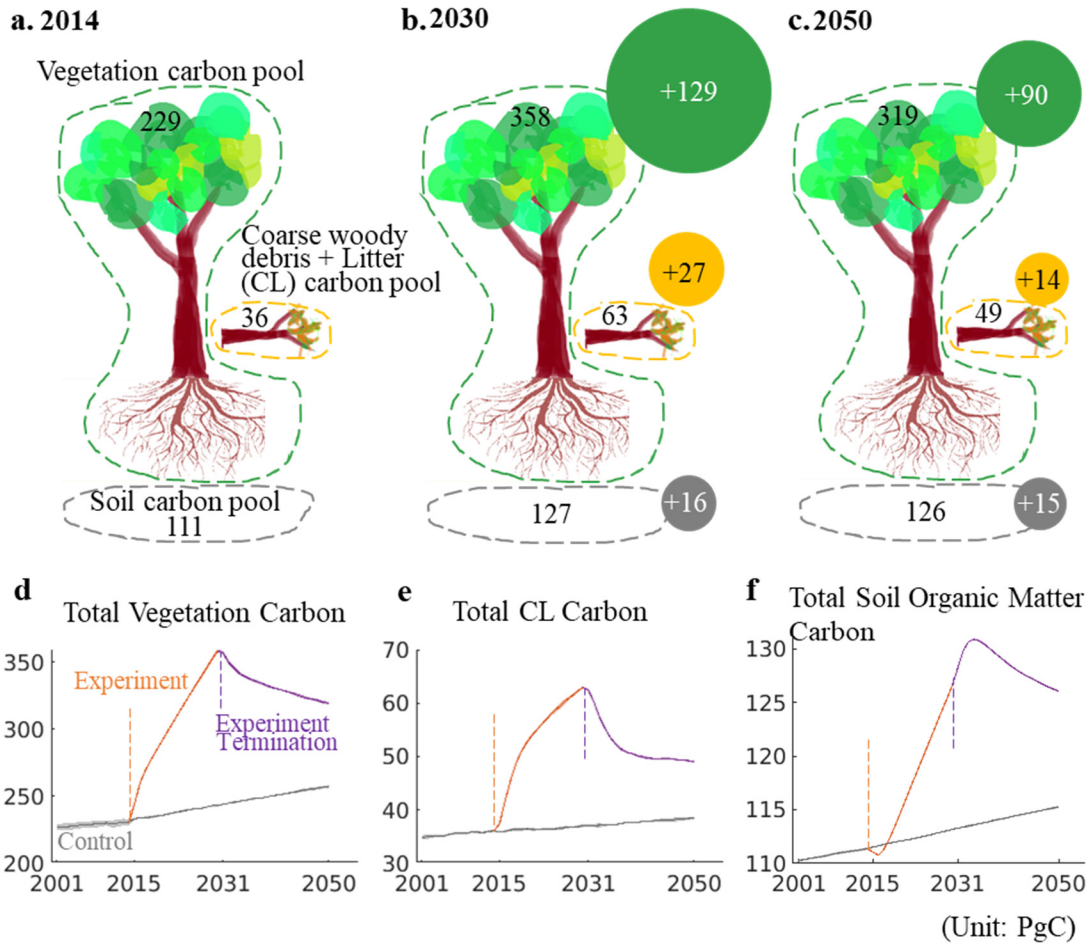
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247

248 **Fig. 2.** Global tropical forest carbon flux and climate responses under and after the lighting  
249 experiment. Ta in panel (g & k): Near surface air temperature. Soil moisture in Panel (i) refers to  
250 the mass of water in the 10cm soil surface. Shaded areas represent uncertainties.



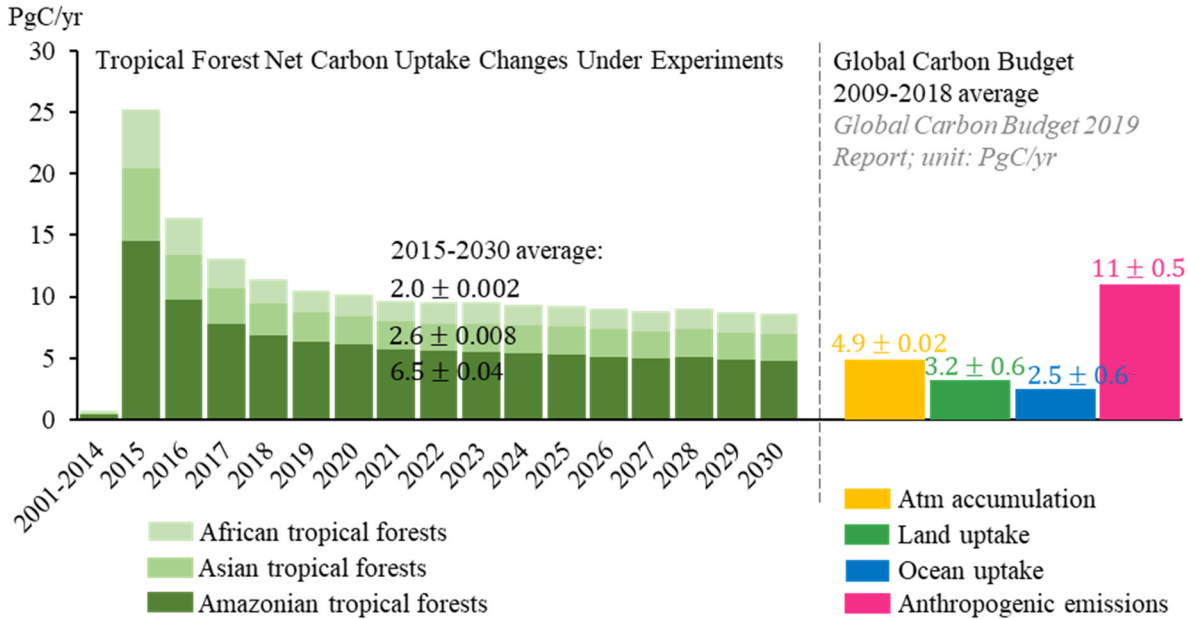
252 **Fig. 3.** Where Did the Net Absorbed Carbon Go? Global Tropical Forest Carbon Amount  
 253 Responses. Panel (a): the current carbon amount in different carbon pools. Panel (b): carbon  
 254 amount in 2030 after 16-year lighting experiments. Panel (c): carbon amount in 2050 after 20  
 255 years since the termination of the lighting experiments. The solid circles in panel (b) and (c) refer  
 256 to carbon amount changes with respect to panel (a). The numbers in panels (a-c) are based on  
 257 panels (d-f). Tree drawing courtesy of © Ning Zeng.



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260 **Fig. 4.** Capabilities of Amazonian, African, and Asian tropical forests to offset annual  
 261 atmospheric carbon accumulation.



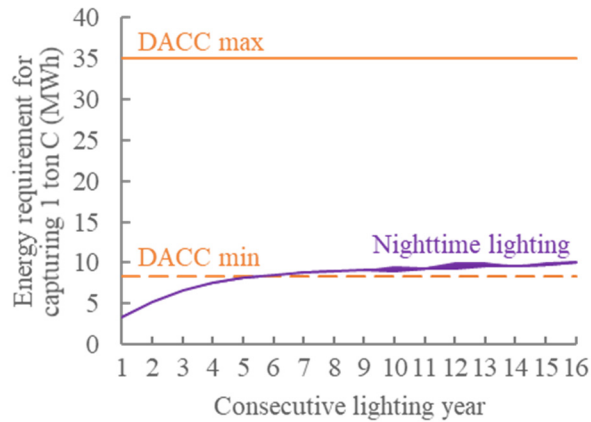
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264 We estimated the energy requirement of this strategy for capturing one ton of carbon (see  
 265 Methods), and compared it to that of Direct Air Carbon Capture (DACC) estimated by recent  
 266 studies(Chatterjee and Huang, 2020; Realmonte et al., 2019). As the carbon uptake efficiency of  
 267 the tropical forest ecosystem decreases with time when under consecutive nighttime lighting, the  
 268 energy requirement for capturing one ton of carbon increases (Figure 5 purple line).  
 269 Nevertheless, the energy requirement of this strategy is lower than that of DACC, or is  
 270 equivalent to the most optimistic estimation of DACC's energy requirement that excludes the  
 271 energy costs required for carbon transport, storage, and utilization. (See Discussion)

272

273 **Figure 5.** Energy requirement for DACC and the nighttime lighting strategy



274

275 3.3 20-year simulation after the experiment termination from 2031 to 2050

276 The lighting experiment was terminated at 12:00 am January 1<sup>st</sup>, 2031 (UTC time), and model  
277 simulations continued for 20 years to 2050 (see the purple lines in Figure 2). The annual gross  
278 primary production and autotrophic respiration dropped quickly, ultimately reaching levels that  
279 were even lower than the control period due to a reduction in atmospheric CO<sub>2</sub> (CO<sub>2</sub> has  
280 fertilization effect in the model). Heterotrophic respiration remained high and decreased much  
281 slower at a speed 10 times lower than gross primary production and autotrophic respiration. The  
282 soil organic matter carbon pool continued to expand due to the entering of litter carbon during  
283 the first 2-3 years following the experiment termination (Figure 3-f). The vegetation carbon pool  
284 shrunk as trees produced less leaves (Figure 3-d). As a result, tropical forests turned into a net  
285 carbon source and remained so until the end of the simulation in 2050 (Figure 2-f). 31.4% of the  
286 carbon that had been absorbed during the lighting experiment was released back to the  
287 atmosphere. This number would likely be higher if the simulation continued. As a result, the  
288 global atmospheric CO<sub>2</sub> concentration returned to a level slightly lower than the control scenario.  
289 Local air temperature and precipitation returned to control levels.

290

#### 291 4. Discussion

292 Physiological responses of tropical trees to longer photoperiods at the ecosystem level remains  
293 one of the biggest uncertainties in model simulations. Some field experiments indicate that  
294 higher CO<sub>2</sub> did not increase carbon sequestration of forests without added nutrients(Oren et al.,  
295 2001), suggesting tree growth might be limited by nutrient supply. The simulated local warming  
296 might also suppress tree growth(Gatti et al., 2021). Some observational evidence shows that  
297 intact tropical forest carbon sinks have been negatively influenced by warming and moisture  
298 stress(Doughty et al., 2015; Gatti et al., 2021) and might be reaching saturation(Hubau et al.,  
299 2020). However, the model predicted increase in precipitation and soil moisture, and previous  
300 studies have shown hydro climate plays a key role in deciding the effects of warming on tree  
301 growth(Guan et al., 2015; Reich et al., 2018). No direct evidence exists to verify the simulation  
302 results. Ecosystem-level field experiments are needed to understand how tropical forest  
303 ecosystems respond to longer photoperiods.

304 CESM2 likely overestimated the local air temperature increase in tropical forests for the  
305 omission of chemical energy stored during photosynthesis(Sellers, 1992). In CESM2 and other  
306 modern Earth System Models(Sellers, 1992), the canopy energy equation(Danabasoglu et al.,  
307 2020) uses the solar radiation absorbed by the vegetation to calculate temperature:

$$308 \quad -\vec{S}_v + \vec{L}_v(T_v) + H_v(T_v) + \lambda E_v(T_v) = 0 \quad (2)$$

309 where  $\vec{S}_v$  is the solar radiation absorbed by the vegetation,  $\vec{L}_v$  is the net longwave radiation  
310 absorbed by vegetation, and  $H_v$  and  $\lambda E_v$  are the sensible and latent heat fluxes from vegetation,  
311 respectively.  $\vec{L}_v$ ,  $H_v$ , and  $\lambda E_v$  depend on the vegetation temperature  $T_v$ .

312 The chemical energy that is stored during photosynthesis and released by respiration is ignored  
313 as the net chemical energy usually amounts to less than 1% of absorbed insolation (around  
314 0.6%(Trenberth et al., 2009)). In our lighting experiment from 2015 to 2030, however, 17% of  
315 absorbed insolation was fixed in the ecosystem as chemical energy (Figure 2-f) and did not  
316 contribute to local air temperature increase. The model failed to exclude this chemical energy  
317 storage from the energy equation. Therefore, the model overestimated the local temperature  
318 increase. This suggests that the temperature simulation results should be treated carefully when  
319 Earth System Models are used to do extreme scenario experiments associated with  
320 biogeochemistry.

321 Tropical forests experienced significant increase in carbon sink during the lighting experiment,  
322 but ultimately transitioned from a sink to a source after the experiment was terminated (Figure 2-  
323 f). Studies(Koven et al., 2021; Tokarska and Zickfeld, 2015) investigating the effects of  
324 overshoot future scenarios (positive carbon emissions followed by net-negative emissions) on  
325 terrestrial carbon cycle have observed similar phenomenon. During a positive emissions phase,  
326 terrestrial carbon cycles tend to absorb some fraction of added CO<sub>2</sub>; however, during a removal  
327 phase they tend to release CO<sub>2</sub>. The mechanism of these phenomena is the different responding  
328 rates of vegetative primary productivity and heterotrophic respiration to lengthening and  
329 shortening photoperiods, or increasing and decreasing atmospheric CO<sub>2</sub>, with primary  
330 productivity responding much quicker than heterotrophic respiration. It is understandable when  
331 considering the diurnal pattern of forest carbon uptake. In the daytime, forests act as a carbon  
332 sink because photosynthesis is greater than respiration. In the nighttime respiration continues  
333 while photosynthesis abates, making forests a carbon source. Additional light/CO<sub>2</sub> would  
334 increase carbon sink by increasing both photosynthesis and respiration (sometimes referred to as  
335 a fertilization effect). When the additional light/CO<sub>2</sub> is removed, photosynthesis decreases  
336 quickly while respiration remains high, making forests a greater carbon source. It suggests that  
337 carbon removal actions focused on enhancing ecosystem productivity by altering environmental  
338 factors in the short term could induce this post-action CO<sub>2</sub> outgassing.

339 The global average surface air temperature remained below the control level after the termination  
340 of lighting experiments due to a reduction in atmospheric CO<sub>2</sub> concentration (Fig.2-k). However,  
341 the local air temperature went back to the control level and seems to be not influenced by CO<sub>2</sub>  
342 reduction (Fig.2-g). We attribute it to two possible reasons. First, different regions tend to have  
343 diverse air temperature responses to global CO<sub>2</sub> changes. Arctic regions show a much larger  
344 temperature increase in response to CO<sub>2</sub> increase, while the temperature increase in tropical  
345 regions is not that significant. Similarly, the CO<sub>2</sub> reduction may exert diverse impacts on  
346 temperature changes in different regions. Second, the temperature change in tropical forests at  
347 the termination of the experiment is controlled by two factors in this study, decreased incoming  
348 shortwave radiation and reduced CO<sub>2</sub>. The former has a much larger impact on the local energy  
349 balance than the latter. Therefore, the influence of CO<sub>2</sub> reduction on local tropical forests is not  
350 as large as on the global scale.

351 Large clean energy requirements have always been a hurdle to large-scale deployment of any  
352 Carbon Dioxide Removal (CDR) techniques, including DAC and the strategy we discuss in this  
353 study. Our estimation suggests that the energy requirement of this strategy for capturing one ton



354 of carbon is less than that of DACC. Specifically, if we give DACC 100 units of energy  
355 (100MWh) per year, DACC could fix 3-12 ton carbon per year. If we give forests 100 units of  
356 extra energy per year, forests could fix around 19.5 ton carbon per year on average (15-year  
357 average: 29 ton carbon in the first year and 10 ton carbon in the 15th year due to an increase of  
358 soil respiration); however, only 17 units of energy are actively used to fix carbon, and the rest 83  
359 units of energy end up as heat which increases local temperature. Therefore, the energy use  
360 efficiency of this strategy is low, which is a major drawback.

361 Other than the direct lighting energy, this strategy requires additional energy associated with  
362 manufacturing and installing lamp networks, constructing electricity transmission devices, so on  
363 and so forth. To make a direct comparison to DACC, we only focus on the energy requirement  
364 specifically for carbon capture. Therefore, we didn't include the energy costs associated with  
365 engineering aspects, as the estimation of DACC's energy requirement does not include the  
366 energy costs required for carbon transport, storage, and utilization. In this study, we also mainly  
367 focus on the physical understanding of tropical forest ecosystem's responses to nighttime  
368 artificial lighting, so we didn't have much discussion on engineering aspects (how such a  
369 network of lamps could be constructed) as well as costs estimates. Nevertheless, the estimation  
370 of additional energy costs and the engineering feasibility are important, and we will attempt to  
371 address these issues in future studies.

372 As to the energy source, we assume this strategy only uses clean energy coming from solar, wind  
373 or nuclear farms to avoid extra carbon emissions when providing light to forests. In terms of  
374 technical analysis, more clean energy can be acquired by deploying more low-carbon energy  
375 generation plants across the globe (e.g., building large-scale solar and wind farms in the Sahara  
376 Desert(Li et al., 2018)). In terms of economic analysis, however, both DACC and this strategy  
377 are energetically and financially costly, and therefore, are unrealistic at present(Chatterjee and  
378 Huang, 2020). Moreover, even if the clean energy generation capacity increases, we cannot  
379 expect the global clean energy supply to only be invested to absorb CO<sub>2</sub>. Nevertheless, if society  
380 has urgency to intervene in Earth's climate by removing CO<sub>2</sub> from the atmosphere in the late half  
381 of the 21th century, and/or an energy revolution realizes and we achieve the status of a  
382 significant surplus of clean energy, CDR would still be a powerful and effective climate  
383 mitigation strategy.

384 Another critical negative impact of this strategy is the potential threat to local wildlife and  
385 biodiversity. Tropical forests are the repository of a large proportion of Earth's biodiversity, and

386 many of the organisms in the tropics are nocturnal or crepuscular, with organisms and  
387 interactions occurring in darkness. An extension of photoperiods could disrupt the habit of  
388 nocturnal creatures and have unexpectedly large impacts on ecosystem biodiversity. In addition,  
389 the disruption, disturbance and habitat fragmentation resulting from installing lights throughout  
390 tropical forests and throughout the forest canopies could exacerbate the negative impacts of this  
391 strategy. Given the potentially inverse relationship between more light at night and ecosystem  
392 health, policy makers may consider extending the photoperiod to an appropriate level to increase  
393 carbon sequestration meanwhile protecting local biodiversity from disastrous impacts. The  
394 tradeoff between nighttime carbon sequestration and biodiversity preservation should be  
395 rigorously evaluated and weight in the decision making process.

396 Overall, lighting up tropical forests at night has led to significant increase in carbon uptakes,  
397 decrease in atmospheric CO<sub>2</sub> concentration, and suppression of global warming as simulated by  
398 Earth System Model. However, it has strong side effects after the termination of nighttime  
399 lighting. In addition, local ecosystem changes could have negative impacts on local wildlife.  
400 Practical issues include the large demand for clean energy and the difficulties for  
401 implementation. From a positive standing it might be treated as an emergency climate solution if  
402 the society relies heavily on carbon removal to adjust the Earth's climate in the future. Paris  
403 Agreement set climate goals to limit global warming to well below 2 degree Celsius and  
404 preferably to 1.5 degree Celsius compared to pre-industrial levels(Lawrence et al., 2018). To  
405 accomplish the Paris Agreement's climate goals, different engineering levels (lighting powers,  
406 areas, and periods) might be needed under various anthropogenic emission scenarios, with high-  
407 emission scenarios possibly requiring high engineering levels. This study investigated the highest  
408 engineering level (lighting up global tropical forests at night with the optimal power) under a  
409 low-emission scenario (see Methods). Further research is needed to investigate the relationship  
410 between engineering levels and emission scenarios in the context of global climate goals set out  
411 by the Paris Agreement(Lawrence et al., 2018).

412 Current geoengineering studies mainly focus on the evaluation of climate goals that a potential  
413 solution might or might not accomplish; however, the changes in Earth's climate after  
414 terminating a geoengineering measure tend to be overlooked. This study suggests the importance  
415 of post-geoengineering analysis in geoengineering studies.

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419 **Code and Data Availability**

420 CESM2 is an open-source community climate model preserved at  
421 <https://doi.org/10.1029/2019MS001916>. All data have been included in the manuscript.

422 **Author contribution**

423 XG designed the study and performed the simulations. XG, SL, DW, YL, BH, and AJ  
424 contributed to the data interpretation. XG drafted the original version of the manuscript. SL and  
425 DW reviewed and edited the manuscript.

426 **Competing interests**

427 Authors declare that they have no competing interests.

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