

1 **Exploration of a novel geoengineering solution: lighting up tropical forests at** 2 **night**

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8 **Abstract**

9 Plants primarily conduct photosynthesis in the daytime, offering an opportunity to increase
10 photosynthesis and carbon sink by providing light at night. We used a fully coupled Earth
11 System Model to quantify the carbon sequestration and climate effects of a novel carbon removal
12 proposal: lighting up tropical forests at night via lamp networks above the forest canopy.
13 Simulation results show that additional light increased tropical forest carbon sink by 10.4 ± 0.05
14 petagrams of carbon per year during a 16-year lighting experiment, resulting in a decrease in
15 atmospheric CO₂ and suppression of global warming. In addition, local temperature and
16 precipitation increased. The energy requirement for capturing one ton of carbon is lower than
17 that of Direct Air Carbon Capture. When the lighting experiment was terminated, tropical forests
18 started to release carbon slowly. This study suggests that lighting up tropical forests at night
19 could be an emergency solution to climate change, and carbon removal actions focused on
20 enhancing ecosystem productivity by altering environmental factors in the short term could
21 induce post-action CO₂ outgassing.

22 **Short summary**

23 Numerical experiments with a coupled Earth System Model show that large-scale nighttime
24 artificial lighting in tropical forests will significantly increase carbon sink, local temperature, and
25 precipitation, and requires less energy than Direct Air Carbon Capture for capturing 1 ton
26 carbon, suggesting that it could be a powerful climate mitigation option. Side effects include the
27 CO₂ outgassing after the termination of the nighttime lighting and the impacts on local wildlife.

28 **Keywords:** climate change; Earth system model; geoengineering; carbon cycle; tropical forests

29

30 **1. Introduction**

31 Anthropogenic greenhouse gas (GHG) emissions have led the global mean temperature to
32 increase by approximately 1.1 degree Celsius since the industrial revolution(IPCC, 2013, 2018;
33 IPCC AR6 WGI, 2021). Changes in climate have caused impacts on natural ecosystems and
34 human societies, such as mass ice sheet melt(Jevrejeva et al., 2016), devastating heat
35 waves(Dosio et al., 2018), and increase in extreme climate events(Kirchmeier-Young and Zhang,
36 2020), exposing natural and human systems to uncertainties and the risks of unsustainable
37 development(Gao et al., 2019, 2020). Despite the scientific consensus on climate change,
38 emission-reduction efforts have made slow or little progress with global GHG emissions
39 continuing to rise(IPCC AR6 WGI, 2021). In this context, geoengineering options are
40 increasingly being considered as means of deliberately intervening in Earth's climate system in
41 the second half of the 21st century(IPCC AR6 WGI, 2021; Moore et al., 2015).

42 Existing geoengineering proposals tend to align with two fundamentally different strategies:
43 Solar Geoengineering (SG)(Abatayo et al., 2020; Proctor et al., 2018; Robock et al., 2009) and
44 Carbon Capture and Sequestration (CCS)(IPCC, 2005; Jones, 2008; Leung et al., 2014). SG and
45 related techniques reduce the amount of incoming radiation from the sun typically via
46 stratospheric aerosol injection, subsequently affecting the planet's temperature. Although they
47 may be able to offset temperature increase rapidly, previous studies indicate the potential for
48 political instability(Abatayo et al., 2020) and negative impacts on human health(Robock et al.,
49 2009) and agriculture(Proctor et al., 2018). Comparatively, CCS removes carbon from the global
50 carbon cycle by artificial machines and saves it for long-term storage or for industrial
51 reutilization(IPCC, 2005). While technically feasible, the environmental risks for the transport
52 and storage of CO₂, limited carbon storage capability, and high cost remain large obstacles of
53 implementing CCS(IPCC, 2005; Jones, 2008; Leung et al., 2014).

54 In this study, the authors propose a novel geoengineering solution: lighting up tropical forests at
55 night by installing lamp networks above the forest canopy(Graham et al., 2003), which lengthens
56 photoperiods and leads to greater photosynthesis and carbon sequestration, and helps mitigate
57 climate change. Contrasting to traditional CCS techniques, this strategy utilizes nature carbon
58 sink to capture and sequester CO₂ from air and avoids long-distance transport and geological
59 storage.

60 Structurally intact tropical forests are by far the most efficient carbon-capture method(Mitchard,
61 2018), and they act as an important carbon sink against rising CO₂ levels(Pan et al., 2011;

62 Sullivan et al., 2020). Although intact tropical forest growth is likely suffering from warming
63 and moisture stress induced by anthropogenic greenhouse gas emissions(Aguirre-Gutiérrez et al.,
64 2020; Doughty et al., 2015; Gatti et al., 2021; Hubau et al., 2020), light is still the primary factor
65 limiting tropical tree growth due to cloud cover, especially during the rainy season(Boisvenue
66 and Running, 2006; Graham et al., 2003). Studies on the photoperiodic control of tropical trees’
67 growth typically fall into two categories: physiological field observations under seasonal
68 variations of day length(Borchert et al., 2005; Pires et al., 2018; Rivera et al., 2002), and
69 physiological greenhouse observations under experimental variations of the photoperiod(Dixit
70 and Singh, 2014; Djerrab et al., 2021; Luo et al., 2021; Stubblebine et al., 1978). Field
71 observations have shown that longer photoperiods facilitate the bud break and flowering in
72 tropical forests(Borchert et al., 2005; Pires et al., 2018; Rivera et al., 2002). Greenhouse
73 experiments either lengthen or shorten photoperiods, and results suggest that short photoperiods
74 reduce plant growth rate and lead to thinner leaves and lower chlorophyll content(Djerrab et al.,
75 2021; Luo et al., 2021), while long photoperiods increase stem growth rate and stimulate tree
76 growth(Dixit and Singh, 2014; Stubblebine et al., 1978). These studies are more focused on
77 specific tropical plant species and tend to agree that longer photoperiods might have a positive
78 effect on vegetative growth in tropical forests. Ecosystem-level field experiments are critical for
79 taking into account key environmental factors that are missing in greenhouse experiments (e.g.
80 water and nutrition constraints), and for informing model parameterizations, although they are
81 lacking so far. Previous studies have shown that longer photoperiods facilitate the bud break and
82 flowering in tropical forests(Borchert et al., 2005; Rivera et al., 2002). A greenhouse study in
83 1978 showed that a tropical tree species grown for one year under a 15-hour photoperiod
84 treatment had an average stem length twice that of the same species grown under an 8-hour
85 photoperiod treatment(Stubblebine et al., 1978). These studies suggest that longer photoperiods
86 might have a positive effect on vegetative growth in tropical forests.

87 Earth System Models provides state-of-the-art computer simulations of key processes and
88 climate states across the Earth(Danabasoglu et al., 2020). In this study the authors used a fully
89 coupled Earth System Model, Community Earth System Model version 2 (CESM2) developed
90 by the U.S. National Center for Atmospheric Research(Danabasoglu et al., 2020), to test the
91 carbon sequestration and climate effects of this geoengineering measure by conducting
92 numerical lighting experiments. Briefly, we added additional diffuse visible light to tropical
93 forest canopy at night (see Supplementary Figure 1) assuming that trees will receive light from
94 multiple directions (e.g., multiple lamps). Tropical forest grids were defined by “Broadleaf

95 Evergreen Tree Area Percentage” being greater than 60% between 20°N and 20°S. The lighting
96 experiment started from 12:00 am on January 1st, 2015 (UTC time), and the simulation exercise
97 was conducted across numerous timescales and lighting levels:

98 (1) Historical control simulation from 2001 to 2014

99 (2) 24-hour lighting experiment with various lighting powers on January 1st, 2015

100 (3) 16-year lighting experiment with the optimal lighting power from 2015 to 2030

101 (4) 20-year simulation after the experiment termination from 2031 to 2050

102 (5) Future control simulation from 2015 to 2050

103 Both experiment and control simulations in the future from 2015 to 2050 were on top of the
104 Shared Socioeconomic Pathways (SSP) 126 scenario(Riahi et al., 2017). Each simulation has a
105 spatial resolution of 1° and has two members (created from small perturbations to initial
106 conditions) to provide uncertainty estimation. (see Methods for detailed experimental design)

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108 **2. Results**

109 2.1 24-hour lighting experiment with various lighting powers on January 1st 2015

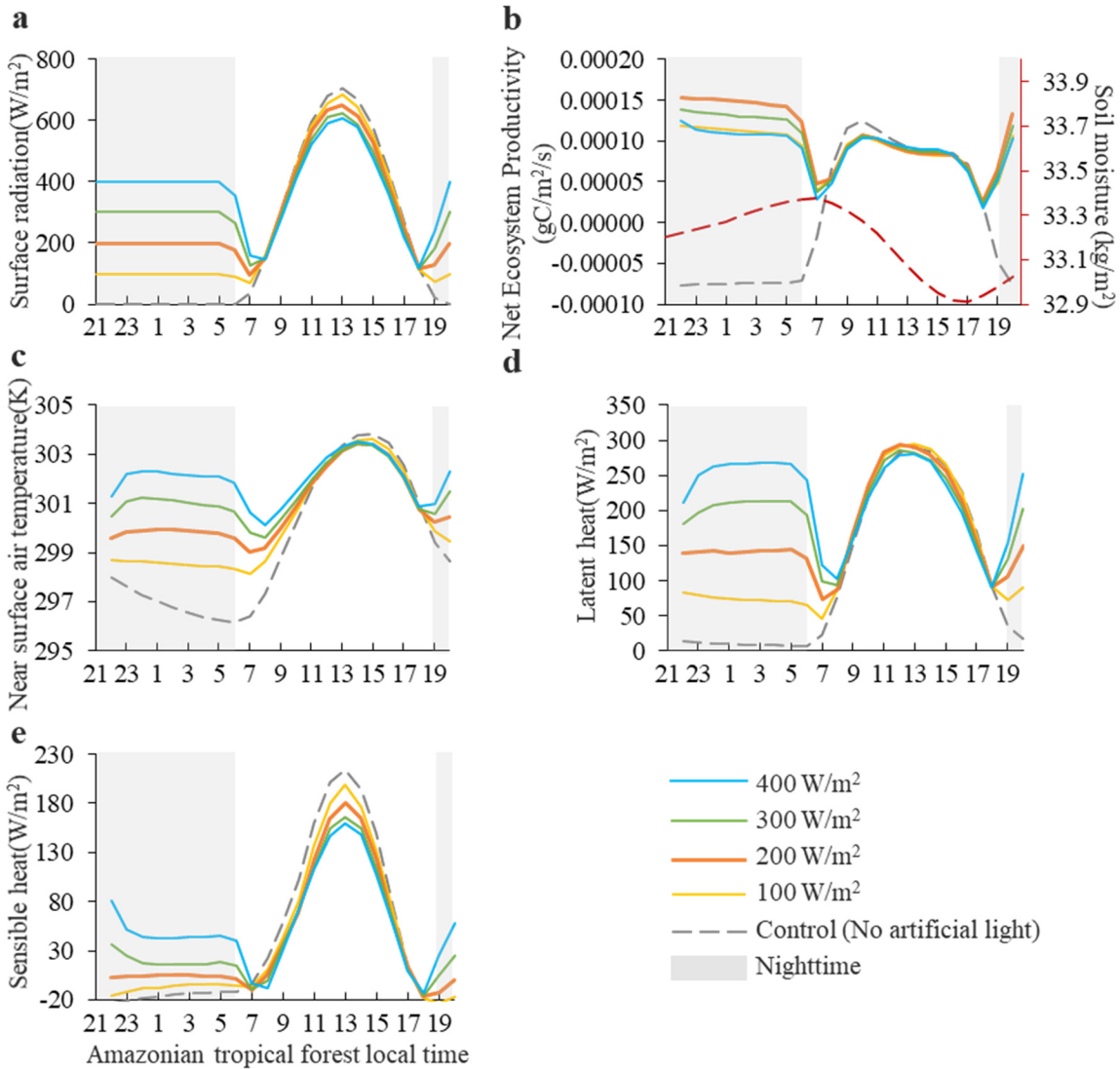
110 Figure 1 shows the changes in carbon and energy fluxes of Amazonian tropical forests for 24
111 hours since the start of the nighttime lighting experiment at 12:00 am January 1st, 2015 (UTC
112 time; See Supplementary Figure 2 and 3 for African and Asian tropical forest responses).

113 Tropical forests had a significant response to nighttime radiation, but the response was different
114 under 100, 200, 300, and 400 W/m² lighting powers. The lighting experiment altered the
115 nighttime energy balance and increased near surface air temperature, latent heat, and sensible
116 heat. Higher lighting powers led to greater increases in air temperature, latent heat and sensible
117 heat. Meanwhile, the additional light activated photosynthesis and increased Net Ecosystem
118 Productivity (NEP). Nighttime NEP reached the peak at 200W/m² and seemed to be suppressed
119 when the lighting power was higher. Comparison of NEP across lighting powers suggests that
120 200W/m² is optimal in terms of activating additional photosynthesis. The nighttime NEP is
121 higher than daytime because nighttime surface radiation is solely diffuse visible light while
122 daytime surface radiation is composed of direct NIR (~16%), diffuse NIR (~30%), direct visible
123 light (~15%), and diffuse visible light (~39%). African and Asian tropical forests showed similar
124 responses.

125 During daytime in the control simulation, the maximum NEP shows up around 9:00-11:00 am
126 (Fig. 1-b). It is not likely to be due to clouds according to the diurnal pattern of the surface
127 downward shortwave radiation (Fig.1-a). We examined the diurnal curve of the soil moisture
128 (the red dash line in Fig. 1-b), and it seems to be due to soil moisture stress. Soil moisture was
129 consumed quickly in the morning which led to water stress for plant growth in the afternoon. The
130 soil moisture pattern also explains the biased distribution of daytime surface air temperature
131 (Fig.1-c), and slightly biased daytime latent heat (Fig.1-d), and daytime sensible heat (Fig.1-e).

132

133 **Fig. 1.** Amazonian tropical forest responses for 24 hours since the start of the nighttime lighting
 134 experiment at 12:00 am January 1st, 2015 (UTC time) under various nighttime lighting powers.
 135 Panel (a) refers to surface downward shortwave radiation. Nighttime NEP (b) reached the peak at
 136 200W/m^2 , suggesting that 200W/m^2 is optimal in terms of activating additional photosynthesis.



137

138 2.2 16-year lighting experiment with the optimal lighting power from 2015 to 2030

139 The yellow lines in Figure 2 show that tropical forest carbon fluxes and climates were
140 significantly altered by a 16-year continuous lighting experiment at night with a 200W/m²
141 power. The annual gross primary production and autotrophic respiration increased by twice near
142 instantaneously, while the heterotrophic respiration had a slower response and increased
143 continuously over a longer period. We purport these changes to be due to the increase in local
144 temperature and the gradual accumulation of organic matter in the soil. Simulation results show
145 that the lighting experiment also decreased wildfire emissions as soil moisture increased despite
146 the expansion of the coarse woody debris and litter carbon pool that provides potential burning
147 materials. However, the expansion of the coarse woody debris and litter carbon pool could
148 provide more burning materials and increase wildfire risks. Overall, the net carbon uptake
149 increased to around 25 petagrams of carbon per year (Pg C yr⁻¹) in the beginning of the lighting
150 experiment, although it decreased with time due to the continuous increase in heterotrophic
151 respiration. The lighting experiment increased the net carbon uptake in tropical forests by 15.3
152 times over the simulation period (from 0.68±0.02 Pg C yr⁻¹ over 2001-2014 to 11.1±0.05 Pg C
153 yr⁻¹ over 2015-2030). Among all the absorbed carbon, 75% entered the vegetation carbon pool,
154 16% entered the coarse woody debris and litter carbon pool, and 9% entered the soil carbon pool
155 (Figure 3-b).

156 Simulation results show that local climates were also significantly impacted (Figure 2-g,h). The
157 annual average air temperature increased by around 1.3°C, and annual precipitation almost
158 doubled. The temperature and precipitation increase showed no significant seasonal trend
159 (Supplementary Figure 4). Globally, the atmospheric CO₂ concentration dropped quickly in the
160 first several years, while turned flat in the latter of the lighting experiment. As a result, the global
161 average air temperature increase was suppressed by around 0.5°C.

162 Amazonian, African, and Asian tropical forests present different capabilities to offset annual
163 atmospheric carbon accumulation during the lighting experiment (Figure 4). In the current global
164 carbon budget (Friedlingstein et al., 2019) (averaged from 2009 to 2018), approximately 11±0.5
165 Pg C yr⁻¹ was released into atmosphere by anthropogenic activities including fossil fuel
166 combustion and land use, among which 2.5±0.6 Pg C yr⁻¹ was absorbed by ocean, 3.2±0.6 Pg C
167 yr⁻¹ was absorbed by land, and 4.9±0.02 Pg C yr⁻¹ was accumulated in atmosphere resulting in
168 the concerned warming and climate change. The lighting experiment enhanced Amazonian
169 tropical forest net carbon uptake to 6.5±0.04 Pg C yr⁻¹ (averaged during 2015 to 2030),

170 suggesting that lighting up Amazonian tropical forests along could completely offset
171 anthropogenic carbon emissions. African and Asian tropical forests showed lower capabilities
172 with the net carbon uptake being approximately 2.0 ± 0.002 and 2.6 ± 0.008 Pg C yr⁻¹ respectively
173 (see Supplementary Figure 5, 6, and 7 for Amazonian, African, and Asian tropical forest carbon
174 flux, carbon amount, and climate responses respectively).

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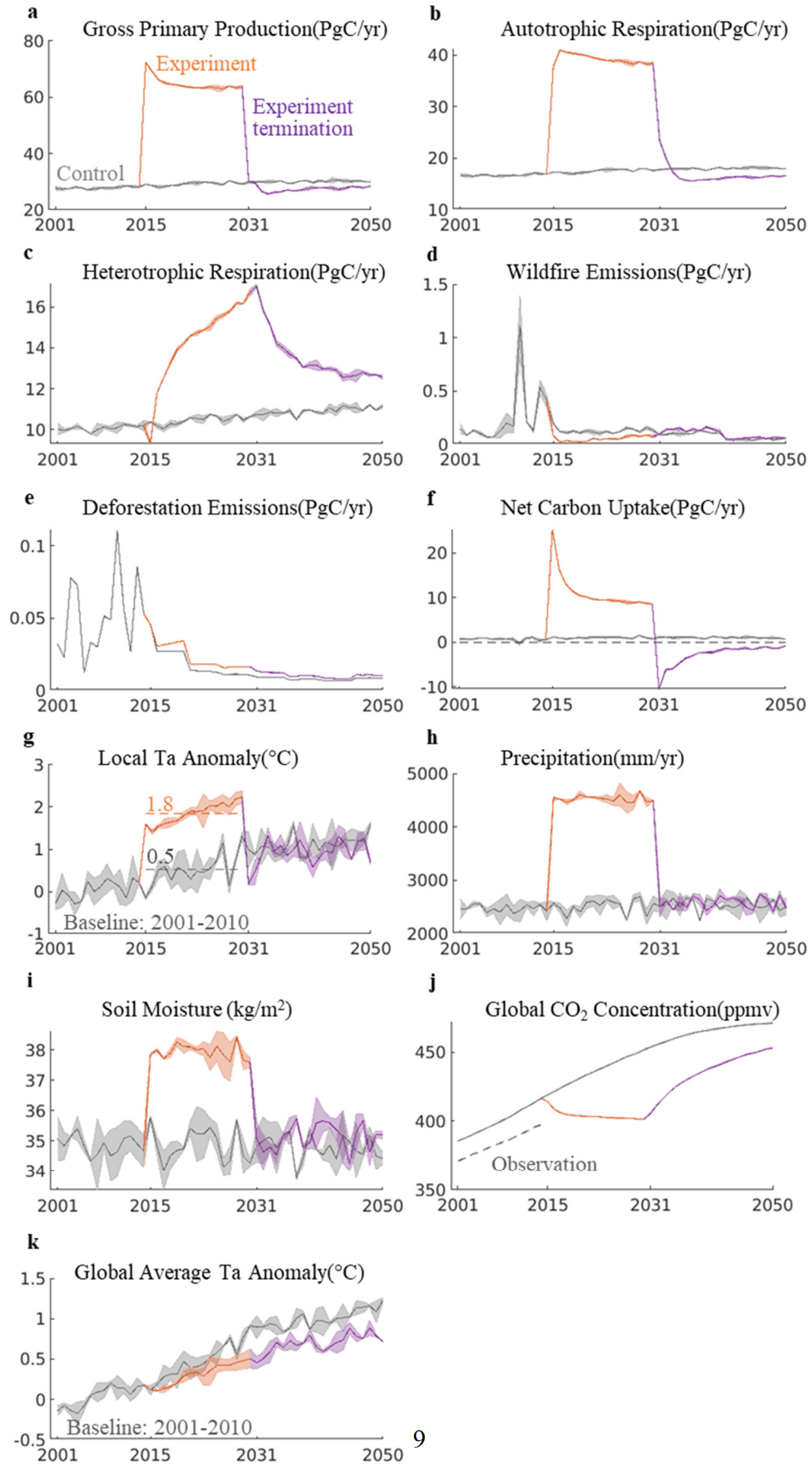
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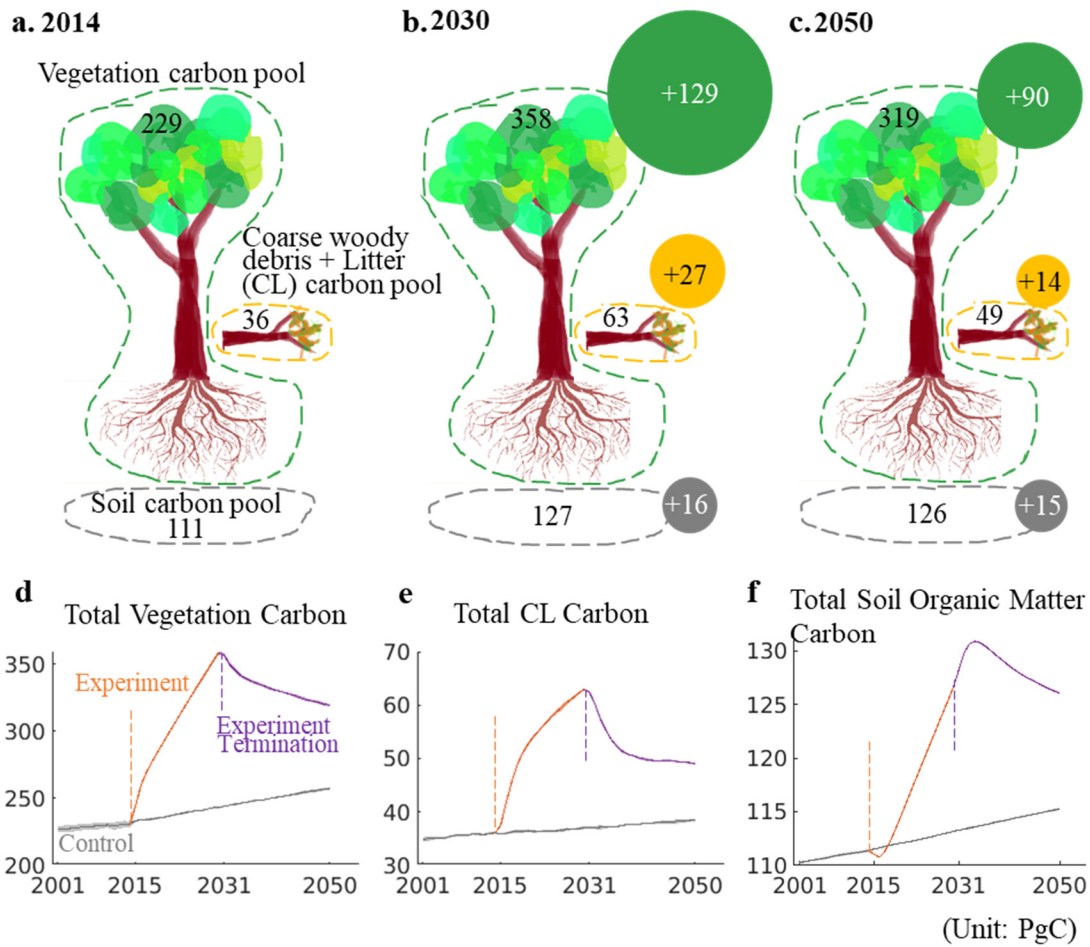
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193 **Fig. 2.** Global tropical forest carbon flux and climate responses under and after the lighting
194 experiment. Ta in panel (g & k): Near surface air temperature. Soil moisture in Panel (i) refers to
195 the mass of water in the 10cm soil surface. Shaded areas represent uncertainties.



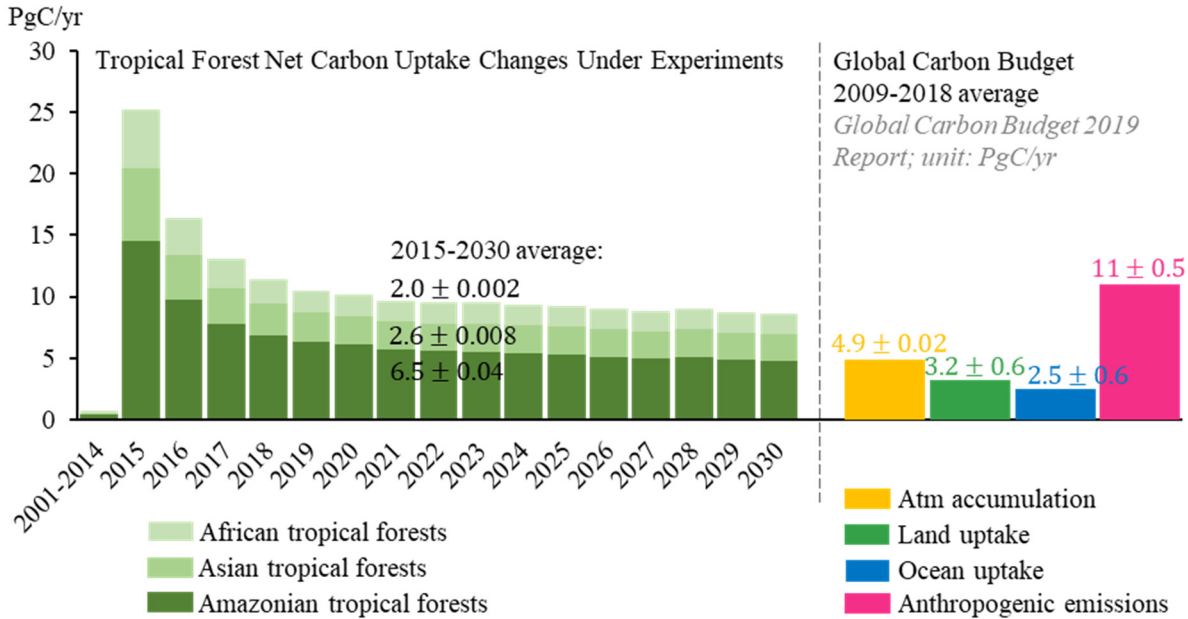
197 **Fig. 3.** Where Did the Net Absorbed Carbon Go? Global Tropical Forest Carbon Amount
 198 Responses. Panel (a): the current carbon amount in different carbon pools. Panel (b): carbon
 199 amount in 2030 after 16-year lighting experiments. Panel (c): carbon amount in 2050 after 20
 200 years since the termination of the lighting experiments. The solid circles in panel (b) and (c) refer
 201 to carbon amount changes with respect to panel (a). The numbers in panels (a-c) are based on
 202 panels (d-f). Tree drawing courtesy of © Ning Zeng.



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205 **Fig. 4.** Capabilities of Amazonian, African, and Asian tropical forests to offset annual
 206 atmospheric carbon accumulation.



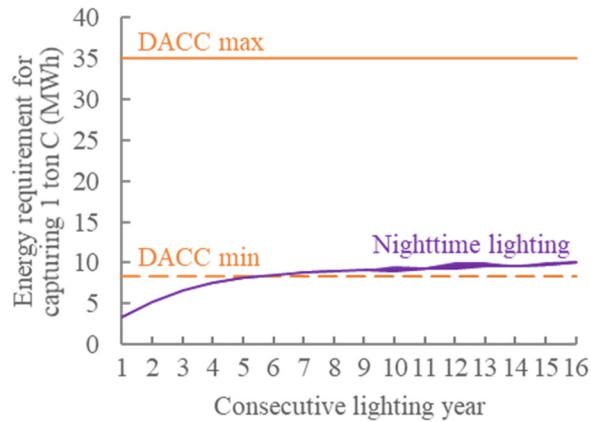
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209 We estimated the energy requirement of this strategy for capturing one ton of carbon ([see](#)
 210 [Methods](#)), and compared it to that of Direct Air Carbon Capture (DACC) estimated by recent
 211 studies(Chatterjee and Huang, 2020; Realmonte et al., 2019). As the carbon uptake efficiency of
 212 the tropical forest ecosystem decreases with time when under consecutive nighttime lighting, the
 213 energy requirement for capturing one ton of carbon increases (Figure 5 purple line).
 214 Nevertheless, the energy requirement of this strategy is lower than that of DACC, or is
 215 equivalent to the most optimistic estimation of DACC's energy requirement that excludes the
 216 energy costs required for carbon transport, storage, and utilization. ([see Discussion](#))

217

218 **Figure 5.** Energy requirement for DACC and the nighttime lighting strategy



219

220 2.3 20-year simulation after the experiment termination from 2031 to 2050

221 The lighting experiment was terminated at 12:00 am January 1st, 2031 (UTC time), and model
222 simulations continued for 20 years to 2050 (see the purple lines in Figure 2). The annual gross
223 primary production and autotrophic respiration dropped quickly, ultimately reaching levels that
224 were even lower than the control period due to a reduction in atmospheric CO₂ (CO₂ has
225 fertilization effect in the model). Heterotrophic respiration remained high and decreased much
226 slower at a speed 10 times lower than gross primary production and autotrophic respiration. The
227 soil organic matter carbon pool continued to expand due to the entering of litter carbon during
228 the first 2-3 years following the experiment termination (Figure 3-f). The vegetation carbon pool
229 shrunk as trees produced less leaves (Figure 3-d). As a result, tropical forests turned into a net
230 carbon source and remained so until the end of the simulation in 2050 (Figure 2-f). 31.4% of the
231 carbon that had been absorbed during the lighting experiment was released back to the
232 atmosphere. This number would likely be higher if the simulation continued. As a result, the
233 global atmospheric CO₂ concentration returned to a level slightly lower than the control scenario.
234 Local air temperature and precipitation returned to control levels.

235

236 3. Discussion

237 Physiological responses of tropical trees to longer photoperiods at the ecosystem level remains
238 one of the biggest uncertainties in model simulations. Some field experiments indicate that
239 higher CO₂ did not increase carbon sequestration of forests without added nutrients(Oren et al.,
240 2001), suggesting tree growth might be limited by nutrient supply. The simulated local warming
241 might also suppress tree growth(Gatti et al., 2021). Some observational evidence shows that
242 intact tropical forest carbon sinks have been negatively influenced by warming and moisture
243 stress(Doughty et al., 2015; Gatti et al., 2021) and might be reaching saturation(Hubau et al.,
244 2020). However, the model predicted increase in precipitation and soil moisture, and previous
245 studies have shown hydro climate plays a key role in deciding the effects of warming on tree
246 growth(Guan et al., 2015; Reich et al., 2018). No direct evidence exists to verify the simulation
247 results. Ecosystem-level field experiments are needed to understand how tropical forest
248 ecosystems respond to longer photoperiods.

249 CESM2 likely overestimated the local air temperature increase in tropical forests for the
250 omission of chemical energy stored during photosynthesis(Sellers, 1992). In CESM2 and other
251 modern Earth System Models(Sellers, 1992), the canopy energy equation(Danabasoglu et al.,
252 2020) uses the solar radiation absorbed by the vegetation to calculate temperature:

$$253 \quad -\vec{S}_v + \vec{L}_v(T_v) + H_v(T_v) + \lambda E_v(T_v) = 0 \quad (1)$$

254 where \vec{S}_v is the solar radiation absorbed by the vegetation, \vec{L}_v is the net longwave radiation
255 absorbed by vegetation, and H_v and λE_v are the sensible and latent heat fluxes from vegetation,
256 respectively. \vec{L}_v , H_v , and λE_v depend on the vegetation temperature T_v .

257 The chemical energy that is stored during photosynthesis and released by respiration is ignored
258 as the net chemical energy usually amounts to less than 1% of absorbed insolation (around
259 0.6%(Trenberth et al., 2009)). In our lighting experiment from 2015 to 2030, however, 17% of
260 absorbed insolation was fixed in the ecosystem as chemical energy (Figure 2-f) and did not
261 contribute to local air temperature increase. The model failed to exclude this chemical energy
262 storage from the energy equation. Therefore, the model overestimated the local temperature
263 increase. This suggests that the temperature simulation results should be treated carefully when
264 Earth System Models are used to do extreme scenario experiments associated with
265 biogeochemistry.

266 Tropical forests experienced significant increase in carbon sink during the lighting experiment,
267 but ultimately transitioned from a sink to a source after the experiment was terminated (Figure 2-
268 f). Studies(Koven et al., 2021; Tokarska and Zickfeld, 2015) investigating the effects of
269 overshoot future scenarios (positive carbon emissions followed by net-negative emissions) on
270 terrestrial carbon cycle have observed similar phenomenon. During a positive emissions phase,
271 terrestrial carbon cycles tend to absorb some fraction of added CO₂; however, during a removal
272 phase they tend to release CO₂. The mechanism of these phenomena is the different responding
273 rates of vegetative primary productivity and heterotrophic respiration to lengthening and
274 shortening photoperiods, or increasing and decreasing atmospheric CO₂, with primary
275 productivity responding much quicker than heterotrophic respiration. It is understandable when
276 considering the diurnal pattern of forest carbon uptake. In the daytime, forests act as a carbon
277 sink because photosynthesis is greater than respiration. In the nighttime respiration continues
278 while photosynthesis abates, making forests a carbon source. Additional light/CO₂ would
279 increase carbon sink by increasing both photosynthesis and respiration (sometimes referred to as
280 a fertilization effect). When the additional light/CO₂ is removed, photosynthesis decreases
281 quickly while respiration remains high, making forests a greater carbon source. It suggests that
282 carbon removal actions focused on enhancing ecosystem productivity by altering environmental
283 factors in the short term could induce this post-action CO₂ outgassing.

284 The global average surface air temperature remained below the control level after the termination
285 of lighting experiments due to a reduction in atmospheric CO₂ concentration (Fig.2-k). However,
286 the local air temperature went back to the control level and seems to be not influenced by CO₂
287 reduction (Fig.2-g). We attribute it to two possible reasons. First, different regions tend to have
288 diverse air temperature responses to global CO₂ changes. Arctic regions show a much larger
289 temperature increase in response to CO₂ increase, while the temperature increase in tropical
290 regions is not that significant. Similarly, the CO₂ reduction may exert diverse impacts on
291 temperature changes in different regions. Second, the temperature change in tropical forests at
292 the termination of the experiment is controlled by two factors in this study, decreased incoming
293 shortwave radiation and reduced CO₂. The former has a much larger impact on the local energy
294 balance than the latter. Therefore, the influence of CO₂ reduction on local tropical forests is not
295 as large as on the global scale.

296 Large clean energy requirements have always been a hurdle to large-scale deployment of any
297 Carbon Dioxide Removal (CDR) techniques, including DAC and the strategy we discuss in this
298 study. Our estimation suggests that the energy requirement of this strategy for capturing one ton

299 of carbon is less than that of DACC. Specifically, if we give DACC 100 units of energy
300 (100MWh) per year, DACC could fix 3-12 ton carbon per year. If we give forests 100 units of
301 extra energy per year, forests could fix around 19.5 ton carbon per year on average (15-year
302 average: 29 ton carbon in the first year and 10 ton carbon in the 15th year due to an increase of
303 soil respiration); however, only 17 units of energy are actively used to fix carbon, and the rest 83
304 units of energy end up as heat which increases local temperature. Therefore, the energy use
305 efficiency of this strategy is low, which is a major drawback.

306 Other than the direct lighting energy, this strategy requires additional energy associated with
307 manufacturing and installing lamp networks, constructing electricity transmission devices, so on
308 and so forth. To make a direct comparison to DACC, we only focus on the energy requirement
309 specifically for carbon capture. Therefore, we didn't include the energy costs associated with
310 engineering aspects, as the estimation of DACC's energy requirement does not include the
311 energy costs required for carbon transport, storage, and utilization. In this study, we also mainly
312 focus on the physical understanding of tropical forest ecosystem's responses to nighttime
313 artificial lighting, so we didn't have much discussion on engineering aspects (how such a
314 network of lamps could be constructed) as well as costs estimates. Nevertheless, the estimation
315 of additional energy costs and the engineering feasibility are important, and we will attempt to
316 address these issues in future studies.

317 As to the energy source, we assume this strategy only uses clean energy coming from solar, wind
318 or nuclear farms to avoid extra carbon emissions when providing light to forests. In terms of
319 technical analysis, more clean energy can be acquired by deploying more low-carbon energy
320 generation plants across the globe (e.g., building large-scale solar and wind farms in the Sahara
321 Desert(Li et al., 2018)). In terms of economic analysis, however, both DACC and this strategy
322 are energetically and financially costly, and therefore, are unrealistic at present(Chatterjee and
323 Huang, 2020). Moreover, even if the clean energy generation capacity increases, we cannot
324 expect the global clean energy supply to only be invested to absorb CO₂. Nevertheless, if society
325 has urgency to intervene in Earth's climate by removing CO₂ from the atmosphere in the late half
326 of the 21th century, and/or an energy revolution realizes and we achieve the status of a
327 significant surplus of clean energy, CDR would still be a powerful and effective climate
328 mitigation strategy.

329 Another critical negative impact of this strategy is the potential threat to local wildlife and
330 biodiversity. Tropical forests are the repository of a large proportion of Earth's biodiversity, and

331 many of the organisms in the tropics are nocturnal or crepuscular, with organisms and
332 interactions occurring in darkness. An extension of photoperiods could disrupt the habit of
333 nocturnal creatures and have unexpectedly large impacts on ecosystem biodiversity. In addition,
334 the disruption, disturbance and habitat fragmentation resulting from installing lights throughout
335 tropical forests and throughout the forest canopies could exacerbate the negative impacts of this
336 strategy. Given the potentially inverse relationship between more light at night and ecosystem
337 health, policy makers may consider extending the photoperiod to an appropriate level to increase
338 carbon sequestration meanwhile protecting local biodiversity from disastrous impacts. The
339 tradeoff between nighttime carbon sequestration and biodiversity preservation should be
340 rigorously evaluated and weight in the decision making process.

341 Overall, lighting up tropical forests at night has led to significant increase in carbon uptakes,
342 decrease in atmospheric CO₂ concentration, and suppression of global warming as simulated by
343 Earth System Model. However, it has strong side effects after the termination of nighttime
344 lighting. In addition, local ecosystem changes could have negative impacts on local wildlife.
345 Practical issues include the large demand for clean energy and the difficulties for
346 implementation. From a positive standing it might be treated as an emergency climate solution if
347 the society relies heavily on carbon removal to adjust the Earth's climate in the future. Paris
348 Agreement set climate goals to limit global warming to well below 2 degree Celsius and
349 preferably to 1.5 degree Celsius compared to pre-industrial levels(Lawrence et al., 2018). To
350 accomplish the Paris Agreement's climate goals, different engineering levels (lighting powers,
351 areas, and periods) might be needed under various anthropogenic emission scenarios, with high-
352 emission scenarios possibly requiring high engineering levels. This study investigated the highest
353 engineering level (lighting up global tropical forests at night with the optimal power) under a
354 low-emission scenario (see Methods). Further research is needed to investigate the relationship
355 between engineering levels and emission scenarios in the context of global climate goals set out
356 by the Paris Agreement(Lawrence et al., 2018).

357 Current geoengineering studies mainly focus on the evaluation of climate goals that a potential
358 solution might or might not accomplish; however, the changes in Earth's climate after
359 terminating a geoengineering measure tend to be overlooked. This study suggests the importance
360 of post-geoengineering analysis in geoengineering studies.

361 **4. Methods**

362 The CESM2 is an open-source community coupled model consisting of atmosphere, ocean, land,
363 sea-ice, land-ice, river, and wave models that exchange states and fluxes via a
364 coupler(Danabasoglu et al., 2020). In this study, we used standard CESM2 configurations and
365 enabled all modules including the Community Atmosphere Model version 6 (CAM6), the
366 Parallel Ocean Program version 2 (POP2) with an ocean biogeochemistry component, the
367 Community Land Model version 5 (CLM5) with a land biogeochemistry component, CICE
368 version 5.1.2 (CICE5), the Community Ice Sheet Model Version 2.1 (CISM2.1), the Model for
369 Scale Adaptive River Transport (MOSART), and the NOAA WaveWatch-III ocean surface wave
370 prediction model version 3.14 (WW3). The CESM2 is part of the Couple Model Intercomparison
371 Project Phase 6 (CMIP6) core simulations as well as about 20 Model Intercomparison Projects
372 (MIPs) within CMIP6. Extensive evaluation suggests that the CESM2 simulations exhibit
373 agreement with satellite era observations of the climate mean state, seasonal cycle, and
374 interannual variability, which has identified CESM2 as among the most realistic climate models
375 in the world(Danabasoglu et al., 2020).

376 4.1 Historical control simulations from 2001 to 2014

377 CESM2 has published its official historical simulation datasets from 1850-2014 on the Earth
378 System Grid Federation (ESGF; <https://esgf-node.llnl.gov/search/cmip6>). This study analyzed
379 the historical simulation datasets of two members from 2001 to 2014 produced by the CESM2
380 esm-hist-BPRP case.

381 4.2 Future experiment and control simulations from 2015 to 2050

382 The selection of 2015 as the start year of the lighting experiment follows CMIP6 future scenario
383 simulation rules. The future experiment simulations and control simulations were both based on
384 the Shared Socioeconomic Pathways (SSP) 126 scenario(Riahi et al., 2017), which is a low-
385 emission (low fossil fuel combustion and deforestation) scenario. The Earth's climate state under
386 SSP126 is close to the current climate state with respect to high-emission scenarios. Therefore,
387 the selection of SSP126 controlled variables and allowed us to see how the lighting experiment
388 along influences tropical forest carbon fluxes and climate. This study ran the CESM2 esm-
389 SSP126-BPRP case with the official restart files from historical simulations (esm-hist-BPRP
390 case). Thus, no model spin up was needed. All simulations were forced with specified
391 greenhouse gas emissions rather than atmospheric greenhouse gas concentrations, so the
392 atmospheric CO₂ concentration was prognostic and land and ocean carbon cycles feed back on

393 atmospheric CO₂. Each simulation has a nominal horizontal resolution of 1° and has two
394 members created from small perturbations to initial climate states to estimate uncertainties.

395 4.3 The lighting experiment design

396 The authors modified the radiation module (Rapid Radiative Transfer Model for General
397 circulation models, RRTMG) of CESM2 to add diffuse visible light to tropical forest canopy at
398 night. CESM2 determines if a grid column is at daytime or nighttime by calculating its cosine
399 (solar zenith angle) at each time step. A negative cosine indicates the grid column is at nighttime
400 and the incoming solar radiation would be assigned with zero. A positive cosine indicates
401 daytime, and the cosine value would be used to calculate incoming solar radiation. If the cosine
402 is positive, the land module then calculates and passes the surface albedo to the atmosphere
403 module and the atmosphere module calculates the radiation fluxes with the surface albedo and
404 the model-calculated incoming solar radiation. We made modifications in all active modules to
405 assign-switch the sign of tropical forests' cosine a positive value (could be any number from 0 to
406 1) from negative to positive when tropical forests were at night. As a result, all modules regarded
407 tropical forests to be at daytime at every time step.

408 CESM2 divides the incoming solar radiation into four components: direct visible light, diffuse
409 visible light, direct near infrared light, and diffuse near infrared light. The authors assume that
410 the artificial light would be provided by a lamp network above the forest canopy and that trees
411 receive light from multiple directions. Therefore, the artificial light was specified as diffuse
412 visible light for simplification. In the model, we assigned the diffuse visible light component of
413 the incoming solar radiation with 100, 200, 300, or 400 and other components with 0. The
414 surface albedo was still calculated by the land module and passed to the atmosphere module. The
415 radiation fluxes were then calculated by the model-calculated surface albedo and the manually-
416 specified solar insolation.

417 4.4 The calculation of the energy requirement for capturing one ton of carbon

$$418 \quad E = (\text{Power} \times \text{Area} \times \text{Hours}) / \text{Carbon} \quad (2)$$

419 where E is the energy requirement for capturing one ton of carbon per year; Power is 200W/m²
420 (nighttime lighting power); Area is the tropical forest area 10.71 × 10¹⁰m² (CESM2 output);
421 Hours is the amount of nighttime lighting hours per year: 365 × 11; Carbon is the net carbon
422 uptake per year (Fig. 2-f) simulated by CESM2. There is no assumed data in this calculation.

423 **Code and Data Availability**

424 CESM2 is an open-source community climate model preserved at
425 <https://doi.org/10.1029/2019MS001916>. All data have been included in the manuscript.

426 **Author contribution**

427 XG designed the study and performed the simulations. XG, SL, DW, YL, BH, and AJ
428 contributed to the data interpretation. XG drafted the original version of the manuscript. SL and
429 DW reviewed and edited the manuscript.

430 **Competing interests**

431 Authors declare that they have no competing interests.

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