Oceanographic regional climate projections for the Baltic Sea until 2100

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Abstract. The Baltic Sea, located in northern Europe, is a semi-enclosed, shallow and tide-less sea with seasonal sea-ice cover in its northern sub-basins. Its long water residence time contributes to oxygen depletion in the bottom water of its southern sub-basins. In this study, recently performed scenario simulations for the Baltic Sea including marine biogeochemistry were analysed and compared with earlier published projections. Specifically, dynamical downscaling using a regionally coupled atmosphere-ocean climate model was used to regionalise four global Earth System Models. However, as the regional climate model does not include components representing terrestrial and marine biogeochemistry, an additional catchment and a coupled physical-biogeochemical model for the Baltic Sea were included. Previous scenario simulations and scenarios taking into account the impact of various water levels were examined. According to the projections, compared to the present climate, higher water temperatures, a shallower mixed layer with a sharper thermocline during summer, less sea-ice cover and greater mixing in the northern Baltic Sea during winter can be expected. Both the frequency and the duration of marine heat waves will increase significantly, in particular in the coastal zone of the southern Baltic Sea (except in regions with frequent upwellings). Nonetheless, due to the uncertainties in the projections regarding regional winds, the water cycle and the global sea level rise, robust and statistically significant salinity changes could not be identified. The impact of a changing climate on biogeochemical cycling is predicted to be considerable but still smaller than that of plausible nutrient input changes. Implementing the proposed Baltic Sea Action Plan, a nutrient input abatement plan for the entire catchment area, would result in a significantly improved ecological status of the Baltic Sea, including reductions in the size of the hypoxic area also in a future climate, which in turn would increase the resilience of the Baltic Sea against anticipated climate change. While our findings regarding changes in heat-cycle variables mainly confirm earlier scenario simulations, they differ substantially from earlier projections of salinity and biogeochemical cycles, due to differences in experimental setups and in input scenarios for bioavailable nutrients.

1 Introduction

The Baltic Sea is a shallow, semi-enclosed sea located in northern Europe (Fig. 1). It has a mean depth of 54 m but due to its strongly varying bottom topography it can be divided into several sub-basins, with limited transport between them (Sjöberg, 1992). In particular, water exchange between the Baltic Sea and the North Sea is hampered

because of two shallow sills located in narrow channels connecting these two water bodies. Thus, large saltwater inflows occur only sporadically, on average once per year, mainly during the winter season but never during summer (Mohrholz, 2018). Furthermore, because the Baltic Sea is embedded within a catchment area that is about four times larger than the Baltic Sea surface, annual freshwater inputs are large relative to the volume of the Baltic Sea (Bergström and Carlsson, 1994). The volume of the Baltic Sea is ~21,700 km^3 (Sjöberg, 1992) and the turnover time of the total freshwater supply (~16,000 $m^3\,s^{-1}$) is 35 years (Meier and Kauker, 2003). These features contribute to strong horizontal and vertical salinity gradients in the Balti Sea (Fonselius and Valderrama, 2003). Moreover, due to its location and physical characteristics, especially the long water-residence time, the Baltic Sea is vulnerable to external pressures, including eutrophication, pollution and global warming (e.g., Jutterström et al., 2014). Ocean circulation modelling has shown that the time scale of the salinity response to changes in atmospheric and hydrological forcing is 20 years (Meier, 2006).

Some 85 million people, in 14 countries, currently live in the catchment area of the Baltic Sea, and anthropogenic pressure on the marine ecosystem is accordingly high (HELCOM, 2018). Insufficiently treated wastewater, pollutant emissions, overfishing, habitat degradation and intensive marine traffic, including oil transport, place a heavy burden on the Baltic Sea ecosystem (Reckermann et al., 2021). One consequence is oxygen depletion of the Baltic Sea's deep waters, such that bottom areas lack higher life forms (e.g., Carstensen et al., 2014; Meier et al., 2018b). In 2018, the area of dead bottom was equal to that of the Republic of Ireland, ~73,000 km^2, which is about one sixth of the sea surface area of the Baltic Sea. Oxygen depletion in the deeper parts of the Baltic Sea arise from the limited ventilation of those waters and the accelerated oxygen consumption that accompanies the remineralisation of organic matter (Meier et al., 2018b). Hence, nutrient input abatement strategies, such as the Baltic Sea Action Plan (BSAP), have been proposed (HELCOM, 2007), with projections of their impact requested by stakeholders such as the Helsinki Commission (HELCOM) or national environmental protection agencies[1].

Projections of the Baltic Sea's climate at the end of the 21st century were among the first to be made for coastal seas worldwide (Meier and Saraiva, 2020). Already at the beginning of the 2000s, the first scenario simulations, were carried out for selected time slices in present and future climates (e.g. Haapala et al., 2001; Meier, 2002a, b; Omstedt et al., 2000). In the dynamical downscaling approach used for those simulations, regional climate models (RCMs) were employed to refine predictions of global climate change to regional and local scales, in this case for the Baltic Sea (e.g. Rummukainen et al., 2004; Döscher et al., 2002). However, those first projections were based on scenarios consisting of a single global climate model (GCM) and a single greenhouse gas (GHG) concentration (150% increase in equivalent CO_2 concentration in the atmosphere in the future vs. the historical climate) and only covered 10-year time slices. These initial attempts were therefore followed by more advanced scenario simulations using mini-ensembles (e.g. Döscher and Meier, 2004; Meier et al., 2004b; Meier et al., 2004a; Räisänen et al., 2004) and centennial-long simulations (e.g. Meier, 2006; Meier et al., 2006; Meier et al., 2011c; Table 1). However, the latter studies considered only monthly mean changes in the future vs. the present climate, applying a so-called delta approach, while neglecting possible changes in inter-annual variability. From these oceanographic studies it was concluded that "mean annual sea surface temperatures (SSTs) could increase by some 2 to 4˚C by the end of the 21st century. Ice extent in the sea would then decrease by some 50 to 80%. The average salinity of

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[1]https://helcom.fi/helcom-at-work/events/events-2021/ccfs-launch/

77 the Baltic Sea could range between present day values and decreases of as much as 45%. However, it should be
78 noted that these oceanographic findings, with the exception of salinity, are based upon only four regional scenario
79 simulations using two emissions scenarios and two global models" (BACC Author Team, 2008).
80
81 For the second assessment of climate change in the Baltic Sea region (BACC II Author Team, 2015), continuously
82 integrated transient simulations from present to future climates became available and even included marine
83 biogeochemical modules (e.g. Eilola et al., 2013; Friedland et al., 2012; Gräwe and Burchard, 2012; Gräwe et al.,
84 2013; Gröger et al., 2019; Gröger et al., 2021b; Holt et al., 2016; Kuznetsov and Neumann, 2013; Meier et al.,
85 2011b; Meier et al., 2011c; Meier et al., 2012a; Meier et al., 2012c; Meier et al., 2012d; Neumann, 2010; Neumann
86 et al., 2012; Omstedt et al., 2012; Pushpadas et al., 2015; Ryabchenko et al., 2016; Skogen et al., 2014) and higher
87 trophic levels (e.g. Bauer et al., 2019; Ehrnsten et al., 2020; Gogina et al., 2020; Holopainen et al., 2016;
88 MacKenzie et al., 2012; Niiranen et al., 2013; Vuorinen et al., 2015; Weigel et al., 2015). The BACC II Author
89 Team (2015) concluded that "recent studies confirm the findings of the first assessment of climate change in the
90 Baltic Sea basin". A key finding of their report was that "No clear tendencies in saltwater transport were found.
91 However, the uncertainty in salinity projections is likely to be large due to biases in atmospheric and hydrological
92 models. Although wind speed is projected to increase over sea, especially over areas with diminishing ice cover,
93 no significant trend was found in potential energy …" (a measure of energy to homogenize the water column). "In
94 accordance with earlier results, it was found that sea-level rise has greater potential to increase surge levels in the
95 Baltic Sea than does increased wind speed. In contrast to the first BACC assessment (BACC Author Team, 2008),
96 the findings reported in this chapter are based on multi-model ensemble scenario simulations using several GHG
97 emissions scenarios and Baltic Sea models. However, it is very likely that estimates of uncertainty caused by biases
98 in GCMs are still underestimated in most studies" (BACC II Author Team, 2015).
99
100 Since the early 21st century, transient simulations for the period 1960–2100 using regional ocean (Holt et al., 2016;
101 Pushpadas et al., 2015) and regionally coupled atmosphere-ocean models, so-called Regional Climate System
102 Models (RCSMs; Bülow et al., 2014; Dieterich et al., 2019; Gröger et al., 2019; Gröger et al., 2021b), have been
103 available for the entire combined Baltic Sea and North Sea system. An overview was given by Schrum et al. (2016)
104 as part of the North Sea Region Climate Change Assessment Report (NOSCCA, Quante and Colijn, 2016) and by
105 Gröger et al. (2021a) within the Baltic Earth Assessment Report (BEAR) project (this thematic issue).
106
107 There is a notable difference in the salinity projections between the first two assessments (BACC Author Team,
108 2008; BACC II Author Team, 2015) and recent scenario simulations (Meier et al., 2021). The first Baltic Sea
109 scenario simulations, driven by nine RCMs and five GCMs, showed a pronounced negative ensemble mean change
110 in salinity because two of the GCMs included a significant increase in the mean west wind component (Meier et
111 al., 2006). These pronounced changes in the large-scale atmospheric circulation were not a feature of later studies
112 (Saraiva et al., 2019a). However, as the natural variability was poorly sampled, this finding may be coincidental.
113
114 The large spread in river discharge did not decrease between the studies, ranging from −8% to +26% (Meier et al.,
115 2006; 2021). Since in more recent assessments the projected rates of global sea level rise (SLR) were revised
116 upwards (e.g. IPCC, 2019a; Bamber et al., 2019), recent scenario simulations for the Baltic Sea also considered a

117   rise in sea level (Meier et al., 2021). As a consequence of compensating effects of the competing drivers of salinity

118   changes, i.e. wind, freshwater input and sea level, future salinity changes were predicted to be small (Table 2).

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120   In the following, we provide an overview of the projections performed since 2013, i.e. after the last assessment of

121   climate change for the Baltic Sea basin, and compare recent results with previous findings by the BACC II Author

122   Team (2015). We focus on projections for the marine environment, from both physical and biogeochemical

123   perspectives. Among the analysed variables are temperature, salinity, oxygen, phosphate, nitrate, phytoplankton,

124   primary production, nitrogen fixation, hypoxic area and Secchi depth (measuring water transparency). An

125   accompanying study by Christensen et al. (2021) investigated atmospheric projections in the Baltic Sea region.

126   For an overview of the development of RCSMs and their applications, the reader is referred to Gröger et al.

127   (2021a). In our comparisons of the various scenario simulations, we analyse only published data (Table 1), with a

128   focus on two recently generated sets of scenario simulations: BalticAPP and CLIMSEA (Table 1, see Saraiva et

129   al., 2019a, b; Meier et al., 2019a; 2021). These are compared with the previous ECOSUPPORT scenario

130   simulations (Meier et al., 2014) assessed by the BACC II Author Team (2015). Investigations of the impact of

131   climate change on primary production in the Baltic Sea that did not utilise a RCM (Holt et al., 2016; Pushpadas et

132   al., 2015) are not addressed herein, nor are nutrient input reduction scenarios under present climate, e.g. as

133   described by Friedland et al. (2021). To our knowledge, further coordinated experiments aimed at projections for

134   the coupled physical-biogeochemical system of the Baltic Sea after 2013 have not been published. Uncoordinated

135   scenario simulations performed prior to 2013 (including Ryabchenko et al., 2016) and their uncertainties were

136   previously discussed by Meier et al. (2018a; 2019b).

137

138   The paper is organised as follows. In Section 2, the dynamical downscaling method, the catchment and Baltic Sea

139   models, the experimental setup and the analytical strategy are introduced. In Section 3, the historical and future

140   climates results of the three scenario simulations, ECOSUPPORT, BalticAPP and CLIMSEA, are compared.

141   Tables 1, 3 and 4 provide an overview of these (Tables 3 and 4) and other (Table 1) scenario simulations from the

142   literature. A consideration of knowledge gaps and a summary of our findings conclude the study. Acronyms used

143   in this study are defined in Table 5.

2 Methods

2.1 Regionalisation of a changing climate

146   Dieterich et al. (2019) produced an ensemble of scenario simulations with a coupled RCSM, called RCA4-NEMO,

147   which was introduced by Wang et al. (2015). Gröger et al. (2019; 2021b) and Dieterich et al. (2019) validated and

148   analysed the different aspects of the RCA4-NEMO ensemble discussed herein. The atmospheric component,

149   RCA4 (Rossby Centre Atmosphere model Version 4), was run at a resolution of $0.22°$ and 40 vertical levels in the

150   EURO-CORDEX domain (Jacob et al., 2014), and the coupled North Sea-Baltic Sea model NEMO (Nucleus for

151   European Modelling of the Ocean) at a resolution of two nautical miles (3.7 km) and 56 levels. The two

152   components of the RCSM are coupled by sending sea surface data of sea level pressure, energy, mass and

153   momentum fluxes every 3 h from the atmosphere to the ocean model. Conversely, the atmosphere model receives

154   data of sea and ice surface temperatures and the sea-ice fraction and albedo at the same frequency.

155

This RCSM was applied to downscale eight different Earth System Models (ESMs), each one driven by three Representative Concentration Pathways (RCPs). For the Baltic Sea projections, four ESMs (MPI-ESM-LR, EC-Earth, IPSL-CM5A-MR, HadGEM2-ES; see Gröger et al. (2019) and references for the ESMs therein) and the GHG concentration scenarios RCP4.5 and RCP8.5 were selected (Table 3). The four ESMs were part of the Fifth Coupled Model Intercomparison Project (CMIP5; Taylor et al., 2012) and their results were assessed in the Fifth IPCC Assessment Report (AR5; IPCC, 2013).

Surface variables of the atmospheric component were saved at hourly to 6-hourly frequencies to allow for an analysis of means and extremes in present and future climates. As RCA4-NEMO does not contain model components for terrestrial and marine biogeochemistry, two additional models forced with the atmospheric surface fields of RCA4-NEMO, i.e. a catchment and a marine ecosystem model, were employed (Fig. 2).

For the ECOSUPPORT scenario simulations, dynamical downscaling was performed with the regional Rossby Centre Atmosphere Ocean (RCAO) model (Döscher et al., 2002). RCAO consists of the atmospheric component RCA3 (Samuelsson et al., 2011) and the oceanic component RCO (Meier et al., 2003; Meier, 2007), with horizontal grid resolutions of 25 km and six nautical miles (11.1 km) respectively. In the vertical, the ocean model has 41 levels with layer thicknesses ranging between 3 m close to the surface and 12 m at 250 m depth. The latter was the maximum depth in the model.

2.2 Catchment models

In BalticAPP/CLIMSEA and ECOSUPPORT, the catchment model E-HYPE (Hydrological Predictions for the Environment, http://hypeweb.smhi.se), a process-based, high-resolution multi-basin model applied for Europe (Hundecha et al., 2016; Donnelly et al., 2017), and a statistical hydrological model STAT (Meier et al., 2012c), were respectively applied to calculate river runoff and nutrient inputs under changing climate but without considering land surface changes. The statistical model calculates river runoff as precipitation minus evaporation over the catchment area; river-borne nutrient inputs are estimated as the product of a given nutrient concentration and the statistically derived volume flow (Gustafsson et al., 2011; Meier et al., 2012c).

In CLIMSEA, two nutrient input scenarios, defining plausible future pathways of nutrient inputs from rivers, point sources and atmospheric deposition, i.e. the BSAP and reference (REF) scenarios (Saraiva et al., 2019a; b), are used (Fig. 3). In BalticAPP, nutrient input scenarios follow BSAP, REF and Worst Case (WORST) scenarios (Saraiva et al., 2019a; b; Pihlainen et al., 2020). Finally, in ECOSUPPORT, instead of WORST a business-as-usual (BAU) scenario is applied (Gustafsson et al., 2011; Meier et al., 2011b).

In the BSAP scenario in CLIMSEA and BalticAPP, nutrient inputs linearly decrease from the actual values in 2012 (i.e., the average for 2010–2012) to the maximum allowable input in 2020 defined by the mitigation plan. Thereafter, nutrient inputs remain constant until the end of the century. A similar temporal evolution is defined in ECOSUPPORT but with a reference period of 1997–2003 (Gustafsson et al., 2011; their Fig. 3.1).

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194   In the REF scenario, in CLIMSEA and BalticAPP, nutrient inputs are calculated using E-HYPE, which considers
195   the impact of changing river flow on nutrient inputs but neglects any changes in land use or socioeconomic
196   development. These inputs correspond on average to the observed mean inputs during the period 2010–2012.

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198   The two additional, above-mentioned scenarios on future projections, BAU and WORST, are not compared
199   because the corresponding input assumptions differ (see Meier et al., 2018a). However, both are characterised by
200   population growth and intensified agricultural practices such as land cover changes and fertiliser use (HELCOM,
201   2007; Zandersen et al., 2019; Pihlainen et al., 2020). In this study they are discussed only for the sake of
202   completeness.

203

204   A comparison of the historical (1980–2005) and future (2072–2097) periods reveals that the reductions in nutrient
205   inputs under the BSAP scenario are smaller in ECOSUPPORT than in BalticAPP and CLIMSEA (Meier et al.,
206   2018a; their Fig. 3). In ECOSUPPORT and BalticAPP/CLIMSEA using the same physical-biogeochemical model
207   RCO-SCOBI, input changes of bioavailable phosphorus amount to −11 ktons (Model A in Meier et al., 2018a) and
208   −34 ktons (Model C in Meier et al., 2018a) respectively (Table 6). Corresponding input changes in bioavailable
209   nitrogen are −230 and −269 ktons. Table 6 also lists the calculated changes for the other two biogeochemical
210   models in ECOSUPPORT, BALTSEM (Model F in Meier et al., 2018a) and MOM-ERGOM (Model D in Meier
211   et al., 2018a), and for the REF scenarios. A comparison confirms the considerable differences between
212   ECOSUPPORT and BalticAPP/CLIMSEA scenario simulations. In the next section, the Baltic Sea models are
213   introduced.

214   **2.3 Baltic Sea models**

215   This study used data from three different Baltic Sea models. The Swedish Coastal and Ocean Biogeochemical
216   model coupled to the Rossby Centre Ocean model (RCO-SCOBI) is driven by the atmospheric surface field data
217   calculated by either RCAO or RCA4-NEMO and by the river runoff and nutrient input scenarios derived from
218   either STAT or E-HYPE projections and atmospheric deposition (Fig. 2). Atmospheric depositions are assumed
219   to be constant or reduced as in the BSAP (Fig. 3). RCO is a Bryan-Cox-Semtner-type ocean circulation model
220   with horizontal and vertical grid resolutions of 3.7 km and 3 m respectively (Meier et al., 1999; 2003; Meier, 2001;
221   2007). SCOBI is a biogeochemical module of the nutrient-phytoplankton-zooplankton-detritus (NPZD) type; it
222   considers state variables such as phosphate, nitrate, ammonium, oxygen concentration, the phytoplankton
223   concentrations of three algal types (diatoms, flagellates and others, cyanobacteria) and detritus (Eilola et al., 2009;
224   Almroth-Rosell et al., 2011; 2015). RCO-SCOBI has been used in many Baltic Sea climate applications (for an
225   overview see Meier and Saraiva, 2020), evaluated with respect to measurements and compared with other Baltic
226   Sea models (Eilola et al., 2011; Placke et al., 2018; Meier et al., 2018a).

227

228   The Ecological ReGional Ocean Model (ERGOM, see www.ergom.net) is a marine biogeochemical model
229   coupled with an ocean general circulation model and a Hibler-type sea-ice model (MOM, Griffies, 2004); its
230   complexity is roughly the same as that of the RCO-SCOBI model. The horizontal resolution of the model is ~5.6
231   km and thus somewhat coarser than that of RCO-SCOBI but in the surface layer its vertical resolution is higher,
232   i.e. 1.5 m in the upper 30 m and below that depth gradually increasing to as high as 5 m (Neumann et al., 2012).

234 The BAltic sea Long-Term large-Scale Eutrophication Model (BALTSEM) spatially resolves the Baltic Sea into

235 13 dynamically interconnected and horizontally averaged sub-basins with high vertical resolution (Gustafsson et

236 al., 2012). For further details of these and other available Baltic Sea ecosystem models the reader is referred to

237 Meier et al. (2018a).

2.4 Scenario simulations

239 In CLIMSEA, we analysed the ensemble of 48 RCO-SCOBI scenario simulations for the period 1976–2098 (Table

240 3) that was produced following the dynamical downscaling approach described in Sections 2.1–2.3 (Fig. 2) and

241 presented in Meier et al. (2021). Unlike in previous studies (Meier et al., 2011b; Saraiva et al., 2019a), the

242 CLIMSEA scenario simulations also consider various scenarios of global SLR. In the three SLR scenarios starting

243 from the year 2005 that were applied by Meier et al. (2021), the mean sea level changes relative to the seabed

244 projected by the year 2100 are: (scenario 1) 0 m, (scenario 2) the ensemble mean of RCP4.5 (0.54 m) and RCP8.5

245 (0.90 m) IPCC projections (IPCC, 2019b; Hieronymus and Kalén, 2020) and (scenario 3) the 95th percentiles of

246 the lowest case (1.26 m, here combined with RCP4.5) and highest case (2.34 m, here combined with RCP8.5)

247 scenarios following Bamber et al. (2019; Table 3). A deepening of the water depth at all grid points every 10 years

248 increases the relative sea level linearly. The spatially varying land uplift was not considered. For details, the reader

249 is referred to Meier et al. (2021).

251 The CLIMSEA ensemble simulations are compared with earlier ensemble scenario simulations by Meier et al.

252 (2011b; 2012c) and Neumann et al. (2012) (ECOSUPPORT), and by Saraiva et al. (2019a, b) and Meier et al.

253 (2019a) (BalticAPP). Both ECOSUPPORT and BalticAPP rely on a downscaling approach similar to that used in

254 the CLIMSEA projections (Fig. 2). However, the scenario simulations of ECOSUPPORT are based upon different

255 global and regional climate models, three coupled physical-biogeochemical models for the Baltic Sea and previous

256 GHG emission scenarios as detailed by the Fourth IPCC Assessment Report (AR4; Table 1). Compared to

257 BalticAPP, the CLIMSEA ensemble is enlarged by three SLR scenarios (Table 3) whereas previous projections

258 assumed no change in the mean sea level relative to the seabed. The inclusion of SLR scenarios followed the

259 finding that the relative sea level above the sills in the entrance area limits transport and controls salinity in the

260 entire Baltic Sea (Meier et al., 2017). As the relative SLR during the period 1915–2014 was estimated to be 0–1

261 mm year^{-1}, resulting from the net effect of past eustatic SLR and land uplift (Madsen et al., 2019), a lowest-case

262 scenario for the future would be a water level above the sills that is relatively unchanged (Meier et al., 2021). In

263 CLIMSEA, mean and highest-case scenarios follow the median values of the RCP4.5 and RCP8.5 ensembles

264 reported by Oppenheimer et al. (2019) and the 95th percentiles of the lowest- and highest-case scenarios of Bamber

265 et al. (2019; Table 3).

2.5 Analysis

267 *Evaluation of the historical period*

268 In this study, the model results of the BalticAPP and CLIMSEA scenario simulations during the historical period

269 were evaluated by calculating the annual and seasonal mean biases during the historical period obtained with RCO-

270 SCOBI simulations and reanalysis data (Liu et al., 2017). Liu et al. (2017) utilised the Ensemble Optimal

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271   Interpolation (EnOI) method to integrate profiles of temperature, salinity and the concentrations of oxygen,
272   ammonium, nitrate and phosphate determined by the Swedish environmental monitoring program into the RCO-
273   SCOBI model. As reanalysis data are available for the period 1971–1999, we limited our bias calculations to 1976–
274   1999, the overlap period between the historical period of the scenario simulations and the reanalysis data. Model
275   data of historical periods of BalticAPP and ECOSUPPORT scenario simulations were evaluated by Saraiva et al.
276   (2019a, b) and Meier et al. (2011b; 2012c, d) respectively.

277

278   *Mixed-layer depth*
279   The mixed-layer depth (MLD) was calculated following de Boyer Montégut et al. (2004), using a threshold value
280   for the difference between the near-surface water temperature at 10 m depth and the temperature at the MLD of
281    $\Delta T = 0.2°C$.

282

283   *Secchi depth*
284   Secchi depth (SD) is a measure of water transparency and is calculated from SD = 1.7/k(PAR), where k(PAR) is
285   the coefficient of underwater attenuation of the photosynthetically available radiation (Kratzer et al., 2003). Factors
286   controlling k(PAR) in the RCO-SCOBI model are the concentrations of phytoplankton and detritus. In addition,
287   salinity is used in one of the other Baltic Sea models (MOM-ERGOM) of the ECOSUPPORT scenario simulations
288   as a proxy of the spatio-temporal dynamics of coloured dissolved organic matter or yellow substances.

289

290   *Trends*
291   First, the monthly average of SST was computed from the model output every 48 h. The linear trend was then
292   calculated using the Theil-Sen estimator (Theil, 1950; Sen, 1968). The trend computed with this method was the
293   median of the slopes determined by all pairs of sample points. The advantage of this computationally expensive
294   method is that it is much less sensitive to outliers. The significance of the SST trends was evaluated from a Mann-
295   Kendall non-parametric test with a threshold of 95%. The SST trends were computed by season and annually. In
296   the latter case, the annual cycle was removed before the linear trend was computed.

297

298   Following Kniebusch et al. (2019), we performed a ranking analysis to identify the atmospheric drivers others than
299   air temperature that are most important for the monthly variability of SST in each ESM forcing of the CLIMSEA
300   data set and in the RCP scenarios RCP4.5 and RCP8.5. The SST trend is dominated by the trend in air temperature.
301   Thus, to eliminate the air temperature effect on SST, the difference between the SSTs and a linear regression
302   between the SSTs and surface air temperatures (SATs) was calculated. This was followed by applying a cross-
303   correlation analysis of the residual SSTs to determine the main factor driving the SST trend. For each grid point
304   and variable (i.e. cloudiness, latent heat flux and u-v wind components), the explained variance was calculated and
305   the variable explaining the most variance was identified.

306

307   *Marine heat waves*
308   During recent decades, the Baltic Sea region has warmed faster than either the global mean warming (Rutgersson
309   et al., 2015; Kniebusch et al., 2019) or any other coastal sea (Belkin, 2009), making it prone to marine heat waves
310   (MHWs). Indeed, short periods of abnormally high water temperatures have been documented for the Baltic Sea
311   (Suursaar, 2020). MHWs can be defined with reference to the mean climatology (e.g. the 90th, 95th, 98th percentile

temperature) or by temperatures exceeding absolute temperature thresholds, defined with respect to the end-user application (Hobday et al., 2018). In most cases, MHWs are defined by the number of periods, their intensity, their duration and the specific purpose (Hobday et al., 2018). In this study, the focus was on the general impact of climate change and the sensitivity of ecosystem dynamics. Hence, MHWs are defined herein as periods of SST ≥ 20°C lasting for at least 10 consecutive days. For comparison, we showed also MHWs defined as periods of SST exceeding the 95th percentile of the SST distribution also lasting for at least 10 consecutive days.

3 Results

3.1 Historical period

3.1.1 Water temperature

The climate of the Baltic Sea region varies considerably, due to maritime and continental weather regimes. For the period 1970–1999, the annual mean SST was ~7.8°C (Fig. 4). The mean seasonal cycle of the SST is pronounced. Thus, every winter, the northern Baltic Sea, including the Bothnian Bay, Bothnian Sea and the eastern Gulf of Finland, is typically covered by sea ice (not shown). Due to its large latitudinal extension, the Baltic Sea is characterised throughout the year by a distinct SST difference between the colder northern and warmer southern sub-basins (Fig. 4). In the southern Baltic Sea, there is also a pronounced west–east temperature gradient, mainly during summer and autumn, which reflects the large-scale cyclonic circulation that transports warmer, more saline southern waters along the eastern coast and colder, less saline northern waters along the western side (see Gröger et al., 2019, their Suppl. Mat. S1; Fig. 4).

On average, during the period 1976–2005, the climate in the CLIMSEA simulations is warmer than the climate according to the reanalysis data (Fig. 4). During spring and summer, the shallow coastal zone of the northern and eastern Baltic Sea is too warm. The spatially averaged biases during winter, spring, summer and autumn and in the annual mean are 0.8, 0.9, 0.8, 1.0 and 0.9°C. The reason for the warm bias is likely a bias of the RCSM. If driven by the reanalysis data ERA-40 (Uppala et al., 2005), RCA4-NEMO systematically overestimates water temperatures and underestimates sea-ice cover in the Baltic Sea for the period 1976–2005 (Gröger et al., 2019; their Suppl. Mat. S1).

In the ECOSUPPORT scenario simulations, there is also a systematic warm bias of the RCAO driven by GCMs at the lateral boundaries, such that winter water temperatures are too warm and sea-ice cover is too low (Meier et al., 2011d, d; 2012c, d). While these biases occur in all three applied Baltic Sea models (Table 3) forced with RCSM atmospheric surface fields, in the simulations driven by regionalised reanalysis data (ERA-40) the mean biases are smaller (Eilola et al., 2011).

3.1.2 Mixed-layer depth

Figure 5 shows the seasonal MLD cycle calculated after de Boyer Montégut et al. (2004). A deeper MLD with pronounced west-east gradients characterises the open ocean. This is related to the predominant southwesterly wind regime, with the larger wind fetches and higher significant wave heights in the eastern Gotland Basin causing

348   wave-induced vertical mixing. Furthermore, a positive sea-atmosphere temperature contrast favours higher wind

349   speeds ('positive winter thermal feedback loop'; Gröger et al., 2015; 2021b). In spring, a weakening wind regime,

350   which reduces heat exchange (with a shift from heat loss to heat gain), together with the increased solar irradiance

351   leads to a thinner MLD in the southern Baltic Sea while melting sea ice and subsequent thermal convection and

352   wind-induced mixing maintain a MLD > 50 m in the sea's northern part. During summer, when atmosphere-ocean

353   dynamics are weakest, a pronounced thermocline develops and MLDs are shallowest (the 'summer thermal short

354   circuit'; Gröger et al., 2021b). During autumn, the atmosphere cools faster than the Earth's surface, and land

355   masses faster than open sea areas. These increased thermal differences result in a stronger large-scale wind regime

356   with a positive feedback on the MLD.

357

358   The ensemble model mean in CLIMSEA reproduces these dynamics and the spatial pattern relatively well. During

359   the cold season, however, the MLD is somewhat shallower in the simulation than in the reanalysis data of Liu et

360   al. (2017). This may be the result of air-sea coupling. Gröger et al. (2015, 2021b) demonstrated that the complex

361   thermal air-sea feedbacks in winter are less well represented by stand-alone ocean models than by fully coupled

362   ocean-atmosphere GCMs. This can result in SST biases and too-shallow MLDs (Gröger et al., 2015; Fig. 7a

363   therein; Gröger et al., 2021b). However, the real reasons for the underestimated winter MLD are unknown.

364

365   In the literature, MLDs in the ECOSUPPORT scenario simulations have not been analysed.

366   **3.1.3 Marine heat waves**

367   Baltic Sea MHWs are defined herein as periods of >10 days duration during which 1) the SST is > 20°C and 2)

368   the SST exceeds the 95th percentile temperature. The CLIMSEA climate model ensemble mean and the reanalysis

369   data set generated by the same model are compared in Figure 6 (Liu et al., 2017).

370

371   The first MHW index uses a fixed threshold that emphasises the environmental impact of the heat waves. In

372   particular, diazotrophic nitrogen fixation becomes effective at higher temperatures. The spatial pattern of MHWs

373   is strongly related to the simulated SST. Figure 6a shows that MHWs are mostly absent in the open sea of the

374   Baltic proper and further north in the Gulf of Bothnia, but they are highly abundant in shallow marginal bays such

375   as the Gulf of Finland and Gulf of Riga as well as along the coasts. The MHWs produced by the RCO ensemble

376   mean are generally more frequent and of longer duration than those of the reanalysis data set. Furthermore, the

377   coastal signature of high abundance extends further offshore (Fig. 6a). For the Belt Sea and Bay of Lübeck, this

378   leads to considerable deviations from the reanalysis data set.

379

380   The second index is based on a reference climatology, here defined as that of 1976–1999. The number of MHWs

381   (Fig. 6c) correlates negatively with their average duration (Fig. 6d). This is somewhat more pronounced in the

382   reanalysis data set. In general, the patterns obtained with the reanalysis data and the RCO are similar but the

383   amplitude of spatial variance is higher in the former (Fig. 6c), as it includes small-scale regional observations. In

384   the RCO (Fig. 6d), MHWs in the open sea are of the longest duration, with their interruption likely due to the

385   vertical mixing induced by wind events.

386

387 Since MHWs in the Baltic Sea are predominantly a summer phenomenon, the stability of the seasonal thermocline

388 is likely a key element in their dynamics such that processes related to vertical mixing can be considered a

389 benchmark in their simulation by the models. Given that mixing is highly parameterised in current ocean models,

390 the RCO reproduces the spatial patterns of the number and average duration of MHW reasonably well.

391

392 In the literature, MHWs in the ECOSUPPORT scenario simulations have not been analysed.

393 **3.1.4 Salinity**

394 The annual mean sea surface salinity (SSS) distribution shows a large north–south gradient mirroring both the

395 input of freshwater from rivers, mostly located in the northern catchment area, and saltwater inflows from the

396 North Sea (Fig. 7). The SSS drops from about 20 g kg^{-1} in the Kattegat to < 2 g kg^{-1} in the northern Bothnian Bay

397 and eastern Gulf of Finland. For the period 1970–1999, the annual mean SSS of the Baltic Sea including the

398 Kattegat was ~7.3 g kg^{-1}. Large inflows of heavy saltwater from the Kattegat occasionally ventilate the bottom

399 water of the Baltic Sea, filling its deeper regions (Fig. 7). As tides are almost absent, mixing is limited such that

400 the water column is characterised by a pronounced vertical gradient in salinity, and consequently also in density,

401 between the sea surface and the bottom.

402

403 Probably due to differences in the data of the hydrological model (E-HYPE) compared to observations, SSS in the

404 coastal zone and the Kattegat is on average lower in the CLIMSEA climate models than in the reanalysis data of

405 Liu et al. (2017) (Fig. 7). The spatially averaged, annual mean bias is −0.4 g kg^{-1}. Bottom salinities in the Belt

406 Sea, Great Belt area and the Gotland Basin (especially in the northwestern part) are considerably higher and in the

407 Bornholm Basin considerably lower in the climate models than in the reanalysis data (Fig. 7). The spatially

408 averaged, annual mean bias is +0.3 g kg^{-1}. Hence, vertical stratification in the Belt Sea, Great Belt area and the

409 Gotland Basin is also larger in the climate models than in the reanalysis data, because the difference between

410 surface and bottom salinities is a good proxy for vertical stratification.

411

412 In the ECOSUPPORT scenario simulations, SSS is overestimated in the entire Baltic Sea, in particular in its

413 northern and eastern regions (Meier et al., 2011c; 2012c). In both, the ensemble mean bottom salinity and vertical

414 stratification are also overestimated while the bottom salinity in the eastern Gotland Basin is well reproduced

415 (Meier et al., 2012c).

416 **3.1.5 Sea level**

417 Due to the seasonal cycle in wind speed, with wind directions predominantly from the southwest, the sea level in

418 the Baltic Sea varies considerably throughout the year, with the highest levels (~40 cm), measured relative to the

419 Kattegat, occurring during winter at the northern coasts of the Bothnian Bay and at the eastern coasts of the Gulf

420 of Finland (Fig. 7). For the period 1976–1999, the annual mean sea level in the Nordic height system 1960 (NH60)

421 as determined by Ekman and Mäkinen (1996) was ~16 cm, with a horizontal north–south difference of ~35 cm

422 (not shown). This sea level slope was explained by the lighter brackish water in the northeastern Baltic Sea than

423 in the Kattegat and by wind coming from the southwesterly direction, which pushes the water to the north and east

424 (Meier et al., 2004a).

425

426 The differences in the mean sea level between the CLIMSEA climate models and the reanalysis data are small

427 (Fig. 7) and the spatially averaged, winter mean bias is only +0.6 cm. Sea levels in some parts of the coastal zone

428 such as the western Bothnian Sea are higher in the climate models than in the reanalysis data, probably due to

429 lower salinities. The negative sea level bias in the eastern Gotland Basin suggests an intensified, basin-wide

430 cyclonic gyre. The seasonal cycle of the ensemble mean sea level is relatively well simulated, but with an

431 overestimated sea level in early spring and an underestimated sea level in summer at all investigated tide gauge

432 locations compared to observations and to a hindcast simulation driven by regionalised ERA40 data (Fig. 8).

433

434 In the ECOSUPPORT scenario simulations, sea levels were not systematically analysed. In one of the three models

435 (RCO-SCOBI), seasonal mean biases were comparable to the biases in the CLIMSEA scenario simulations (Meier

436 et al., 2011a).

437 **3.1.6 Oxygen concentration and hypoxic area**

438 Since the 1950s, nutrient inputs into the Baltic Sea have increased due to population growth and intensified

439 fertiliser use in agriculture (Gustafsson et al., 2012; Fig. 3). Nutrient inputs reached their peak in the 1980s but

440 have steadily declined following the implementation of nutrient input abatement strategies. Nonetheless, since the

441 1960s, the bottom water of the Baltic Sea below the permanent halocline has been characterised by oxygen

442 depletion and large-scale hypoxia (Fig. 9).

443

444 Consistent with the stratification biases in the deeper sub-basins of the Baltic Sea, summer bottom oxygen

445 concentrations in the Bornholm Basin are higher and those in the Gotland Basin lower in the CLIMSEA/BalticAPP

446 climate simulations than in the reanalysis data of Liu et al. (2017) (Fig. 9). The stronger vertical stratification,

447 especially at the halocline depth, hampers vertical fluxes of oxygen, causing prolonged residence times and lower

448 bottom oxygen concentrations. Spatially averaged biases during winter, spring, summer and autumn and in the

449 annual mean are small but systematic: -0.6, -0.7, -0.7, -0.5 and -0.6 mL L^{-1} respectively.

450

451 In the ECOSUPPORT scenarios, the ensemble mean deep-water oxygen concentration in the eastern Gotland

452 Basin is slightly higher (but within the range of natural variability) and that in the Gulf of Finland significantly

453 lower than determined from observations (Meier et al., 2011b; 2012d).

454 **3.1.7 Nutrient concentrations**

455 Nutrient (i.e. phosphorus and nitrogen) content in the surface layer during winter is a good indicator of the intensity

456 of the following spring bloom. Sea-surface mean winter concentrations of phosphate and nitrate are highest in the

457 coastal zone, in particular close to the mouths of the large rivers in the southern Baltic Sea that transport elevated

458 inputs of nutrients into the sea (Fig. 9).

459

460 For the historical period of 1976–1999, winter surface phosphate concentrations according to the climate

461 simulations are relatively close to those of the reanalysis data (Fig. 9). The concentrations differ substantially only

462 in those coastal regions influenced by large rivers, such as those affected by discharges of the Odra, Vistula and

463 Pärnu rivers. Spatially averaged biases are largest during summer and autumn, with an average bias in summer of
464 +0.2 mmol P m^{-3}.

465

466 Likewise, winter surface nitrate concentrations in the simulations are close to those in the reanalysis data but in
467 coastal regions they differ due to differences in the inputs from large rivers (Fig. 9). This is exemplified by the
468 Gulf of Riga and the eastern Gulf of Finland, where the large differences between them are due to inputs from the
469 Neva River. Spatially averaged biases during winter, spring, summer, autumn and in the annual mean are rather
470 small but systematic: −1.1, −1.3, −0.5, −0.7 and −0.9 mmol N m^{-3} respectively.

471

472 In the ECOSUPPORT scenario simulations, the simulated profiles of phosphate, nitrate and ammonium are within
473 the range of observations for 1978–2007, except in the case of phosphate in the Gulf of Finland (Meier et al.,
474 2012d). According to hindcast simulations, the biases in the coupled physical-biogeochemical models of the Baltic
475 Sea relative to the standard deviations of observations are larger for the northern Baltic Sea than for the Baltic
476 proper (Eilola et al., 2011).

477 **3.1.8 Phytoplankton concentrations**

478 During the period 1976–1999, dense phytoplankton blooms were confined to the coastal zone, i.e. the area with
479 the highest nutrient concentrations (Fig. 10). Water transparency, measured by Secchi depth, is lower in the Baltic
480 Sea than in the open ocean (Fleming-Lehtinen and Laamanen, 2012), and for the period 1970–1999 the annual
481 mean Secchi depth averaged for the entire Baltic Sea, including the Kattegat, was only ~6.6 m. The Secchi depth
482 is also much smaller in the coastal zone than in the open Baltic Sea (Fig. 10), and in the northern Baltic Sea than
483 in the Gotland Basin, attributable to yellow substances originating from land (Fleming-Lehtinen and Laamanen,
484 2012).

485

486 Due to nutrient concentration biases, the annual mean surface phytoplankton concentrations of the simulations are
487 close to those of the reanalysis data of Liu et al. (2017) but they deviate in coastal regions (Fig. 10). Spatially
488 averaged biases during winter, spring, summer, and autumn and in the annual mean are relatively small: +0.02,
489 −0.1, −0.009, +0.06 and −0.008 mg chlorophyll (Chl) m^{-3} respectively. Note that the reanalysis data of Liu et al.
490 (2017) assimilate nutrient and oxygen concentrations but not chlorophyll data.

491

492 Similar results are found for the mean biases in the simulated Secchi depths (Fig. 10). In climate simulations,
493 Secchi depths are systematically deeper in the regions south of Gotland and at the entrance to the Gulf of Finland
494 (northeastern Gotland Basin) than elsewhere in the Baltic Sea. Spatially averaged biases during winter, spring,
495 summer and autumn and in the annual mean are +0.2, +0.4, +0.06, +0.1 and +0.2 m respectively.

496

497 Compared to the Secchi depth data from HELCOM (HELCOM, 2013; their Table 4.3) and Savchuk et al. (2006;
498 their Table 3), the CLIMSEA climate simulations under- and overestimate the Secchi depth in the southwestern
499 and northern Baltic Sea respectively, while in the Gotland Basin the model results well fit the observations (Meier
500 et al., 2019a).

501

502   In the ECOSUPPORT scenario simulations, Secchi depth was not compared with observations.

503   **3.1.9 Biogeochemical fluxes**

504   An evaluation of biogeochemical fluxes, such as primary production and nitrogen fixation, is difficult because

505   observations are lacking. An exception is the study by Hieronymus et al. (2021), in which historical simulations

506   with RCO-SCOBI were compared with in situ observations of nitrogen fixation. The RCO-SCOBI model includes

507   a cyanobacteria life cycle (CLC) model (Hense and Beckmann, 2006; 2010) driven by reconstructed atmospheric

508   and hydrological data. The authors found a satisfactory agreement, with the results mainly within the uncertainty

509   range of the observations. However, simulated monthly mean nitrogen fixation during 1999–2008 showed a

510   prolonged peak period in July and August whereas according to observations the peak was mostly confined to

511   July. It should be noted that the RCO-SCOBI version used in the scenario simulations discussed here (e.g., Saraiva

512   et al., 2019a) does not contain a CLC model.

513   **3.2 Future period**

514   **3.2.1 Water temperature**

515   *Annual and seasonal mean changes*

516   In Figures 11 and 12 and Table 7, annual and seasonal mean SST changes between 1976–2005 and 2069–2098 in

517   RCO-SCOBI are depicted and quantified respectively. The maximum seasonal warming signal propagates

518   between winter and summer from the Gulf of Finland via the Bothnian Sea into the Bothnian Bay (Fig. 11).

519   Maximum warming occurs during summer in the Bothnian Sea and Bothnian Bay. The seasonal patterns of RCP4.5

520   and RCP8.5 are similar although warming is greater in the latter. As SLR has almost no impact on SST changes,

521   BalticAPP and CLIMSEA scenario simulations yield similar results (not shown). The warming level according to

522   ECOSUPPORT is between that predicted by CLIMSEA/BalticAPP RCP4.5 and RCP8.5 because the GHG

523   emissions of the A1B scenario, which forces the ECOSUPPORT ensemble[2], are between those of the RCP4.5 and

524   RCP8.5 scenarios.

525

526   In the CLIMSEA/BalticAPP RCSM projections, the annual mean SST changes in the Baltic Sea driven by four

527   ESMs, i.e. MPI-ESM-LR, EC-EARTH, IPSL-CM5A-MR, HadGEM2-ES, under the RCP8.5 scenario are +2.3,

528   +3.7, +3.5 and +4.7°C respectively (Gröger et al., 2019). Thus, the ensemble mean change is +3.5°C. The

529   corresponding ensemble mean change in the RCO-SCOBI scenario simulations is smaller, +2.9°C. Different

530   MLDs, vertical stratification and sea-ice cover in the two ocean models, RCO-SCOBI and NEMO, may explain

531   the different responses. Indeed, a comparison of the MLD between the two models reveals a shallower MLD in

532   the RCSM than in RCO-SCOBI (not shown), which argues for a higher sensitivity of the RCSM to climate

533   warming.

534

---

[2]One of the scenario simulations of ECOSUPPORT is driven by the A2 scenario, which due to higher GHG emissions is generally warmer than the A1B scenario. However, this particular simulation of the ECHAM5–MPIOM GCM is exceptional and at the end of the 21st century the temperature is not much warmer than that obtained with the corresponding run based on the same model under the A1B scenario.

535 While the spatial patterns of the SST changes in the scenario simulations of ECOSUPPORT (e.g., Meier et al.,
536 2012c) and CLIMSEA (e.g., Saraiva et al., 2019b) are similar, the uncertainties due to the applied global (Meier
537 et al., 2011a) or regional (Meier et al., 2012b) model are in some cases considerable. Of note is the summer
538 ensemble range of the various GCMs (Meier et al., 2011a). The differences in the magnitude of warming are
539 explained by the GHG concentration scenarios shown in Fig. 12.

540

541 *Trends*

542 Since SLR and nutrient input scenarios have a negligible impact on SST changes, only the RCP4.5 and RCP8.5
543 scenarios in CLIMSEA/BalticAPP are compared. The multi-model mean of the annual mean SST trends averaged
544 over the Baltic Sea is ~0.18 °C decade^{-1} and ~0.35 °C decade^{-1} in the RCP4.5 and RCP8.5 scenarios respectively
545 (Fig. 13a, f). At the Baltic Sea scale, seasonal SST trends based on annual values vary only slightly (±0.01 °C
546 decade^{-1} in both scenarios). However at the sub-basin scale, seasonal variations are much stronger, reaching ±0.05
547 °C decade^{-1} in the northern Baltic Sea, with a maximum in summer (Fig. 13). This summer maximum can be
548 explained by the projected decline in sea-ice cover in summer, as occurred during the period 1850–2008
549 (Kniebusch et al., 2019).

550

551 As seen in Figure 14, the relative SST trends indicate faster warming of the northern than the southern Baltic Sea
552 (0.02 °C decade^{-1} and 0.04 °C decade^{-1} in the RCP4.5 and RCP8.5 scenarios respectively), with the largest trends,
553 calculated over the entire period 2006–2099, reaching ~0.24 °C decade^{-1} and ~0.45 °C decade^{-1} in RCP4.5 and
554 RCP8.5 respectively. However, a calculation of the SST trends by 30-year slice periods every 10 years over the
555 entire period shows that annual SST trends are variable over time (not shown). The natural variability appears to
556 modulate these trends, with successive periods of increasing and decreasing SST trends over a period of about 30
557 years. For example, in the RCP8.5 scenario, SST trends gradually increase over the first 50 years of the period,
558 reaching a maximum of 0.5 °C decade^{-1} between 2046 and 2075, before declining slightly from 2060 onwards. As
559 in the RCP4.5 scenario, this is a result of the pronounced natural variability in this scenario. Despite the robustness
560 of the spatial pattern of the SST trends ($p < 0.05$ everywhere), an analysis of SST trends for the four ESM forcings
561 reveals an important dependency of those trends on atmospheric forcings, with a spread of ±0.06 °C decade^{-1} from
562 the multi-model mean of both scenarios (not shown).

563

564 At an annual timescale, the variability in the air temperature, through the sensible heat fluxes, is the main driver
565 of the Baltic Sea's SST (Kniebusch et al., 2019), illustrated here by the high variance of SST explained by air
566 temperature (between 0.85 and 0.95, Fig. 15). The minimum of variance explained is located in the Bothnian Bay,
567 where the sea-ice cover isolates seawater from the air in winter.

568

569 The processes responsible for the SST trends were analysed using a rank analysis of atmospheric variables (i.e.
570 latent heat fluxes, cloud cover and u-v wind components) following Kniebusch et al. (2019; Fig. 16). The second
571 parameter (after SAT) explaining the variability in the SST differs according to the location and ESM.
572 Nevertheless, in all ESMs and in both RCP scenarios, zonal and meridional wind components are the variables
573 that best correlate with SST along most of the coastal areas, probably because of upwelling. In the open sea of the
574 Baltic proper and in the Bothnian Bay, the second most important variable is cloudiness. This is also the case in
575 the Bothnian Sea under the RCP4.5 scenario. However, in RCP8.5 the second most important variable at this

576 location is the latent heat flux. The difference is perhaps due to the absence of sea ice, and therefore the amplified
577 air-sea exchange, under RCP8.5.
578
579 In the vertical, temperature trends are larger in the surface layer than in the winter water of the Baltic Sea above
580 the halocline, thus causing a more intense seasonal thermocline (see Section 3.2.2) in which the trends are largest
581 in spring and summer (not shown). Elevated trends also characterise deep water, due to the influence of saltwater
582 inflows that will be warmer in a future climate because they originate from the shallow entrance area and occur
583 mainly in winter. Hence, in sub-basins that are sporadically ventilated by lateral saltwater inflows, such as the
584 Bornholm Basin and the Gotland Basin, the deep water below the halocline will warm more than the overlaying
585 intermediate layer water.
586
587 In the literature, trends in ECOSUPPORT scenario simulations have not been analysed.

588 **3.2.2 Mixed-layer depth**

589 Figure 17 shows the changes in the MLD. During winter, reduced sea-ice cover in the Bothnian Sea and Bothnian
590 Bay, likely caused by wind-induced mixing, favours a widespread deepening of their MLDs. In spring, the most
591 pronounced feature is a strong shallowing of the MLD in the Bothnian Sea, probably attributable to the radiative
592 fluxes that warm the surface layer and to less thermal convection (Hordoir and Meier, 2012). During the historical
593 period, water temperatures in this area were between 2.0 and 3.0°C (Fig. 4). Thus, in the future, surface water
594 warming between 1.6 and 2.4°C (Fig. 11) may hamper thermal convective mixing.
595
596 The changes during summer are less pronounced. In contrast to winter, there is an overall shallowing in the entire
597 Baltic Sea. This is in agreement with a shallower, more intense thermocline in warming scenarios, as suggested
598 by Gröger et al. (2019), and it is a common feature among the projections, because the changes in wind speed are
599 small (Christensen et al., 2021). Autumn is primarily characterised by a prolongation of the thermal stratification,
600 leading to an overall shallower MLD than during the historical period.
601
602 While Hordoir et al. (2018; 2019) speculated that these changes in thermocline depth during summer will impact
603 the vertical overturning circulation, the meridional overturning circulation in the Baltic proper does not show a
604 clear signal but rather a northward expansion of the main overturning cell (Gröger et al., 2019). Indeed, the effect
605 is expected to be small (Placke et al., 2021).

606 **3.2.3 Marine heat waves**

607 The number of MHWs within climatological 30-year time slices is shown in Figure 18. Under historical climate
608 conditions, MHWs are virtually absent in open ocean areas. They are most frequent in shallow regions and more
609 abundant along the eastern (Baltic States) than the western (Swedish) coasts, which may reflect the greater
610 frequency of coastal upwelling events along the western than the eastern coasts of the Baltic Sea. Even under the
611 RCP4.5 scenario, wide areas of the Baltic proper are affected by MHWs roughly once a year. The strongest
612 response is projected for the high-emission RCP8.5 scenario, and specifically in marginal basins such as the Gulf
613 of Riga and the Gulf of Finland, where in the future MHWs will occur two or three times per year. Not only the

frequency but also the average duration of the MHWs will increase with climate warming. Under RCP8.5, MHWs of ~20 days duration will occur even in the open Gulf of Bothnia (Fig. 18). This increase in MHWs is likewise linked to an increased frequency of tropical nights in the Baltic Sea (Meier et al., 2019a; Gröger et al., 2021b).

MHWs can also be analysed by calculating them with respect to the 95th percentile temperature of the historical reference climate (Fig. 19). For the historical climate, the average duration of MHWs in most regions is < 20–30 days, although in the southern Baltic Sea, especially west of the Baltic proper, MHWs are more frequent. However, the climate change signal is characterised by MHWs that are both more frequent and of longer duration. In RCP4.5, MHWs in the Baltic Sea occur at least every year. The strongest increase in frequency is near the coasts, but the average duration increases less than in the open sea (Fig. 19). This is probably related to repeated cold-water entrainments from the open sea that interrupt warm periods because of the larger variability in the coastal zone than in the open sea. In addition, with their lower heat storage capacity, shallow areas are more sensitive to cold weather events and the associated oceanic heat loss.

3.2.4 Salinity

In the CLIMSEA ensemble, salinity changes are not robust, i.e. the ensemble spread is larger than the signal (Meier et al., 2021). Under both RCP4.5 and RCP8.5, the ensemble mean salinity change is small because the impact on salinity of the projected increase in total river runoff from the entire catchment (Fig. 3) is approximately compensated by the impact of larger saltwater inflows due to the projected SLR (Table 8). Hence, compared to previous studies such as those by Meier et al. (2011b; ECOSUPPORT) and Saraiva et al. (2019a; BalticAPP; Fig. 12), the ensemble mean salinity changes in CLIMSEA are much smaller (Table 8). In idealised sensitivity experiments performed with the RCO-SCOBI model for the period 1850–2008 (Meier et al., 2017; 2019d), the change in the average Baltic Sea salinity (1988–2007) increased linearly with SLR and at a rate of ~1.4 g kg^{-1} m^{-1} (Table 9).

3.2.5 Sea level

Following global sea level changes, SLR in the Baltic Sea will accelerate (Hünicke et al., 2015; Church et al., 2013; Bamber et al., 2019; Oppenheimer et al., 2019; Weisse and Hünicke, 2019), albeit at a somewhat slower rate than the global mean because of the remote impact of the melting Antarctic ice sheet (Grinsted, 2015). Changes in SLR in the North Atlantic (and the Baltic Sea) will be larger in response to the melting of the Antarctic ice sheet than to the melting of Greenland, due to gravitational effects. For a mid-range scenario, SLR in the Baltic Sea is projected to be ~87% of the global mean (Pellikka et al., 2020). Land uplift will partly compensate for the eustatic SLR, in particular in the northern Baltic Sea (e.g. Hill et al., 2010). In RCP2.6 and RCP8.5, the global mean sea level in 2100 is 43 cm and 84 cm higher than during the period 1986–2005 (Oppenheimer et al., 2019). For these two scenarios, likely ranges are 29–59 cm and 61–110 cm respectively. Bamber et al. (2019) assessed ice sheet dynamics in detail and subsequently estimated global-median SLRs in 2100 of 69 cm and 111 cm for low- and high-case scenarios respectively. Likely ranges according to the authors were 49–98 cm and 79–174 cm and very likely ranges 36–126 cm and 62–238 cm.

651 In BalticAPP and CLIMSEA scenario simulations, sea level changes are small (Fig. 12, Table 8) whereas in
652 ECOSUPPORT scenario simulations they are larger, particularly in spring, because one member of the multi-
653 model ensemble considers Archimedes' principle (not shown). Note that the sea level changes shown in Figure 12
654 consider only changing river runoff, changing wind, and melting sea ice as affecting the sea level according to
655 Archimedes' principle (only in the ECOSUPPORT ensemble); as neither the global mean SLR nor land uplift is
656 included, they have to be added (e.g. Meier, 2006; Meier et al., 2004a).

658 In CLIMSEA, there are no statistically significant seasonal changes in the SLR (Fig. 20). In both GHG
659 concentration scenarios, the largest changes are only about ±5 cm. According to these results, systematic changes
660 in the regional wind field (Christensen et al., 2021) and nonlinear effects are negligible. Instead, in the projections,
661 the mean absolute sea level in the Baltic Sea simply follows the mean sea level in the North Atlantic. However,
662 the spatially inhomogeneous isostatic adjustment will considerably alter patterns of sea level changes relative to
663 the sea floor.

665 In response to the global mean SLR, the sea level extremes in the Baltic Sea that are rare today will become more
666 common in the future (e.g. Hieronymus and Kalén, 2020). However, changes in sea level extremes relative to the
667 mean sea level will not be statistically significant because wind velocities are projected to remain unchanged
668 (Christensen et al., 2021). The exceptions are areas with sea-ice decline since they are linked to a decrease in
669 atmospheric stability accompanied by increased wind velocities, the result of increases in temperature and
670 turbulent fluxes (Meier et al., 2011c). These increases will mostly translate as changes from calm to light wind
671 conditions as the stable atmospheric boundary layer becomes less stable. For stronger wind conditions related to
672 high sea level extremes, the impact of stratification effects on mixing is small. In addition, open water areas after
673 sea-ice loss have a smaller surface roughness than ice-covered areas, with the reduced surface friction leading to
674 an increase in wind velocities.

676 As sea level extremes also depend on the path of low-pressure systems (Lehmann et al., 2011; Suursaar and Sooäär,
677 2007), which in a future climate do not show systematic changes (Christensen et al., 2021), changes in sea level
678 extremes relative to the mean sea level are not expected. In addition, a large internal variability at low frequencies
679 prevents the detection of climate-warming-related changes in sea level extremes (Lang and Mikolajewicz, 2019).

680 **3.2.6 Oxygen concentration and hypoxic area**

681 *Bottom oxygen concentration*
682 Projected changes in bottom oxygen concentrations differ considerably between ECOSUPPORT and
683 BalticAPP/CLIMSEA scenario simulations, as illustrated for summer (Fig. 21, Table 10), whereas the differences
684 between BalticAPP (SLR = 0 cm) and CLIMSEA (SLR > 0 cm) scenarios are smaller (Meier et al., 2021). The
685 differences between the ECOSUPPORT and BalticAPP ensembles mainly reflect the different experimental setups
686 of the simulations and the different nutrient input scenarios (Meier et al., 2018a). While in shallow regions without
687 a pronounced halocline future bottom oxygen concentrations decrease in all scenario simulations, due to the lower
688 oxygen saturation concentrations, in deeper offshore regions with a halocline, changes in bottom oxygen
689 concentration depend largely on the nutrient input scenario (Fig. 21). In ECOSUPPORT scenario simulations, the

future bottom oxygen concentration decreases significantly in all scenarios except the BSAP, where in deeper regions it changes only slightly on average (see Meier et al., 2011b). By contrast, in the BalticAPP projections, under the BSAP bottom oxygen concentrations in deeper regions increase considerably, regardless of the degree of warming (see Saraiva et al., 2019a; Meier et al., 2011b). Under RCP4.5, bottom oxygen concentrations increase even under the nutrient inputs of REF and WORST whereas RCP8.5 predicts slight reductions in the Bothnian Sea and southwestern Baltic Sea, in particular under WORST. Similar results were calculated for the CLIMSEA ensemble (Meier et al., 2021).

The differences in the oxygen concentration changes between the ECOSUPPORT and BalticAPP/CLIMSEA ensembles can be explained as follows. In ECOSUPPORT, changes in nutrient input relative to the historical period 1961–2006, including the observed nutrient inputs averaged from the period 1995–2002, were applied (Gustafsson et al., 2011; Meier et al., 2011b). For the historical period 1980–2002, these inputs were lower than in BalticAPP/CLIMSEA scenario simulations because in the latter the observed monthly nutrient inputs, including the pronounced decline from the peak in the 1980s until the much lower recent values, were used as the forcing (Meier et al., 2018a). Furthermore, in ECOSUPPORT, future nutrient inputs under the BSAP scenario were calculated as relative changes, resulting in higher future inputs than in BalticAPP/CLIMSEA, in which absolute values of the BSAP were applied.

Hence, the reductions between future and historical nutrient inputs are smaller in ECOSUPPORT under the BSAP than in BalticAPP/CLIMSEA (Table 6) and result in a smaller response of biogeochemical cycling. We argue that the more realistic historical simulation, including a spin-up since 1850, based on observed or reconstructed nutrient inputs as used in the BalticAPP and CLIMSEA ensembles result in a model response that is more realistic than that of the ECOSUPPORT scenario simulations.

Hypoxic area

In ECOSUPPORT, the hypoxic area is projected to increase under REF and BAU nutrient input scenarios (Meier et al., 2011b). Only under BSAP is there a slight decrease compared to the early 2000s.

In CLIMSEA under REF, the hypoxic area is projected to decrease slightly until about 2050, as a delayed response to nutrient input reductions, and then increase again towards the end of the century, presumably in response to increased nutrient inputs and warming (Fig. 22). Larger hypoxic areas are calculated under RCP8.5 than under RCP4.5. Under BSAP, the hypoxic area is projected to considerably decrease. At the end of the century, the size of the hypoxic area is expected to be 22–78% smaller than the average size during the period 1976–2005. This range represents the results of the various scenario simulations.

In accordance with previous studies, such as Saraiva et al. (2019b) and Meier et al. (2021), the impact of warming (reduced oxygen solubility, increased internal nutrient cycling, increased riverine inputs) and of increasing stratification (decreased ventilation) will be an amplified depletion of oxygen that enlarges the hypoxia area in the Baltic Sea and partially counteracts nutrient input abatement strategies such as the BSAP. However, in all available scenarios the impact of climate change is smaller than the impact of nutrient input changes.

3.2.7 Nutrient concentrations

While in ECOSUPPORT scenario simulations of future climate the projected surface phosphate concentrations in winter increase under all three nutrient input scenarios (except in the Gulf of Finland in BSAP), in BalticAPP projections the surface phosphate concentrations in winter decrease almost everywhere (except in the Odra Bight and adjacent areas in REF and WORST) (not shown). In contrast to the nearly ubiquitous changes in the surface phosphate concentration, larger nitrate concentration changes are usually confined to the coastal zone and differ in their signs. In ECOSUPPORT projections, the increases in winter surface nitrate concentrations in REF and BAU are largest in the Gulf of Riga, the eastern Gulf of Finland and along the eastern coasts of the Baltic proper (not shown). In BalticAPP projections, the increases in winter surface nitrate concentrations in REF and WORST are largest in the Bothnian Bay and the Odra Bight while in the Gulf of Riga and the Vistula lagoon nitrate concentrations decrease. Overall, the differences in surface nutrient concentrations between the two sets of scenario simulations are considerable (not shown) and can be explained by the large differences in nutrient inputs from land. Thus, while the projected changes in inputs in ECOSUPPORT refer to the average inputs during 1995–2002, in BalticAPP scenario simulations the observed historical changes include a decline in nutrient inputs since the 1980s (Meier et al., 2018a).

3.2.8 Phytoplankton concentrations

Annual mean changes in surface phytoplankton concentration (expressed as chlorophyll concentration) follow the changes in nutrient concentrations and are confined to the productive zone along the coasts (Fig. 23). In ECOSUPPORT projections, annual mean Secchi depths decrease in all scenario simulations (see Fig. 24 and Table 11). In the BalticAPP projections, the area-averaged Secchi depths generally increase, except in the combined RCP8.5 and BAU scenario (Table 11), indicating a general improvement of the water quality in future compared to the present climate. The most striking changes occur in the BSAP scenario, in which the Secchi depth increases by up to 2 m in the coastal zone of the eastern Baltic proper. Changes in stratification (illustrated by the differences between BalticAPP and CLIMSEA ensembles and between the CLIMSEA ensemble mean and high SLR scenarios) have only a minor impact on water transparency (Table 11). The overwhelming driver of the changes in the Secchi depth are nutrient input scenarios (illustrated by the differences between ECOSUPPORT and BalticAPP/CLIMSEA ensembles and highlighted by, in some cases, contradictory signs in the changes).

3.2.9 Biogeochemical fluxes

In CLIMSEA under the BSAP, primary production and nitrogen fixation are projected to considerably decrease in a future climate (Fig. 22). According to this scenario, the interannual variability declines. Under REF, nitrogen fixation is projected to slightly decrease until ~2050, as a delayed response to nutrient input reductions, and then to increase towards the end of the century, likely in response to increased nutrient inputs and warming. At the end of the century, primary production and nitrogen fixation will be at the same level as under current conditions. The impact of warming is larger under high than under low nutrient conditions (Saraiva et al., 2019b).

3.2.10 Relation to large-scale atmospheric circulation

The dominant large-scale atmospheric pattern controlling the climate in the Baltic Sea region during winter is the North Atlantic Oscillation (NAO; Hurrell, 1995). However, its influence is not stationary but depends on other modes of variability, such as the Atlantic Multidecadal Oscillation (AMO; Börgel et al., 2020). During the past climate, the relationship between the NAO index and regional climate variables in the Baltic Sea region, such as SST, changed over time (Vihma and Haapala, 2009; Omstedt and Chen, 2001; Hünicke and Zorita, 2006; Chen and Hellström, 1999; Meier and Kauker, 2002; Beranová and Huth, 2008).

Figure 25 shows the calculated ensemble mean winter (December–February) NAO index for the period 2006–2100. For the RCP4.5 emission scenario, the NAO shows high interannual variability. Following a wavelet analysis, the calculated NAO index exhibits decadal variability, which differs for every model (not shown). A comparison of RCP4.5 with the high-emission scenario RCP8.5 shows that the spread of the ensemble increases with increasing GHG concentrations. Figure 25 also depicts the running correlation between the NAO index and the area-averaged SST. The correlation remains positive but it is not constant in time. Also evident from a comparison of RCP4.5 and RCP8.5 is that there are no systematic changes in the two emission scenarios, although for RCP8.5 the ensemble spread is slightly larger.

4 Knowledge gaps

In the largest set of scenario simulations of this study, the CLIMSEA ensemble, only four ESMs were regionalised using only one RCSM; consequently, this ensemble is still too small to estimate the uncertainties caused by ESM and RCSM differences. While nine ESMs with the same RCSM were recently regionalised, they did not include running modules for terrestrial and marine biogeochemistry Gröger et al., 2021b), such that these simulations were not considered in our assessment. The uncertainties related to unresolved physical and biogeochemical processes in the Baltic Sea and on land were also not considered, because only one Baltic Sea and one catchment model were used. Although the CLIMSEA ensemble is larger than the ensembles in previous studies, it is still too small to estimate all sources of uncertainty.

In addition to the uncertainties related to global and regional climate and impact models, pathways of GHG and nutrient emissions are thus far unknown and the role of natural variability versus anthropogenic forcing is not well understood (Meier et al., 2018a; 2019b; 2021). Recent studies suggest that the impact of natural variability, such as the low-frequency AMO, is larger than hitherto estimated. For instance, in paleoclimate simulations the AMO affected Baltic Sea salinity at time scales of 60–180 years (Börgel et al., 2018), which is longer than the simulation periods of available scenario simulations. Furthermore, the AMO may also influence the centres of action of the NAO (Börgel et al., 2020). Lateral tilting of the positions of the Icelandic Low and Azores High explains the changes in the correlations between the NAO and regional variables such as water temperature, sea-ice cover and river runoff in the Baltic Sea region (Börgel et al., 2020). Despite indications that the AMO is affected by climate states such as the Medieval Climate Anomaly and Little Ice Age (Wang et al., 2017; Börgel et al., 2018), how future warming would affect these modes of climate variability is unclear.

802 Changes in sea-ice cover were not analysed in this study because in the recent scenario simulations of the
803 CLIMSEA ensemble sea-ice cover is systematically underestimated. However, we found that future sea-ice cover
804 is projected to be considerably reduced, with an on-average ice-free Bothnian Sea and western Gulf of Finland.
805 Recent results by Höglund et al. (2017) confirmed earlier results by Meier (2002b) and Meier et al. (2011d; 2014),
806 see BACC Author Team (2008).

807

808 The various scenario simulation sets have in common that plausible nutrient input changes have a bigger impact
809 on changes in biogeochemical variables, such as nutrient, phytoplankton and oxygen concentrations, than of either
810 the projected changes in climate, such as warming, or changes in vertical stratification. The latter would be caused
811 by increased freshwater inputs, SLR or changes in regional wind fields, assuming RCP4.5 or RCP8.5 scenarios.
812 Long-term simulations of past climate support these results. Although historical warming had an impact on the
813 size of the present-day hypoxic area, model results suggest that hypoxia in the Baltic Sea is best explained by the
814 increases in nutrient inputs due to population growth and intensified agriculture since 1950 (Gustafsson et al.,
815 2012; Carstensen et al., 2014; Meier et al., 2012a; 2019c, d). Hypoxia is also a feature of the Medieval Climate
816 Anomaly (Zillén and Conley, 2010). However, a preliminary attempt to simulate the past 1000 years could not
817 explain the low-oxygen conditions without substantial increases in nutrient inputs (Schimanke et al., 2012). Thus,
818 the sensitivity of state-of-the-art physical-biogeochemical models to various drivers can be questioned and it is
819 clear that the models do not reproduce all important processes.

820

821 As outlined in previous assessments, current and future bioavailable nutrient inputs from land and atmosphere are
822 unknown and were consequently classified as one of the largest uncertainties (Meier et al., 2019b). For a more
823 detailed discussion of uncertainties in Baltic Sea projections, the reader is referred to Meier et al. (2018a; 2019b;
824 2021).

825 **5 Summary**

826 As shown in Section 3, the latest published scenario simulations confirm the findings of the first and second
827 assessments of climate change in the Baltic Sea region (BACC Author Team, 2008; BACC II Author Team, 2015),
828 namely that, in all projections driven by RCP4.5 and RCP8.5 and by four selected ESMs of CMIP5, water
829 temperature is projected to increase and sea-ice cover to decrease significantly. In the two RCP scenarios, the
830 ensemble mean annual changes in SST between 1978–2007 and 2069–2098 are 2°C and 3°C respectively.
831 Warming would enhance the stability across the seasonal thermocline and cause a shallower MLD during summer.
832 During winter, however, the mixed layer in the northern Baltic Sea would be deeper, probably because of the
833 declining sea-ice cover and the associated intensification of wind speed, waves and vertical mixing. Both the
834 frequency and the duration of MHWs would increase significantly, in particular south of 60ºN and in the coastal
835 zone (except in regions with frequent upwellings).

836

837 The spatial patterns of seasonal SST trends projected for 2006–2099 are similar to those of historical
838 reconstructions of the period 1850–2008, although in most regions the magnitude of the trends is larger. The largest
839 trends are those in summer in the northern Baltic Sea (Bothnian Sea and Bothnian Bay) and thus in regions where
840 under a warmer climate sea ice would melt earlier or disappear completely due to the ice-albedo feedback. This

841   implies that, with increasing warming, SST trends in the northern Baltic Sea will become larger than those in the

842   southern Baltic Sea. Accordingly, in contrast to the present climate, in which mean SSTs considerably decline

843   from south to north, in a future climate the north-south temperature gradient will weaken.

844

845   In contrast to previous scenario simulations, recent scenario simulations considered the impact of the global mean

846   SLR on Baltic Sea salinity, which for the ensemble mean salinity would more or less completely compensate for

847   the effects of the projected increasing river runoff. However, as future changes in all three drivers of salinity (wind,

848   runoff and SLR) are highly uncertain, the spread in the salinity projections of the various ESMs is larger than any

849   signal.

850

851   In agreement with earlier assessments, we conclude that SLR has a greater potential to increase surge levels in the

852   Baltic Sea than does changing wind speed or direction. For the latter, there have been no statistically significant

853   changes during the 21st century thus far.

854

855   In agreement with earlier studies, changes in nutrient input according to the BSAP or REF scenarios will have a

856   larger impact on biogeochemical cycling in the Baltic Sea than will a changing climate driven by RCP4.5 or

857   RCP8.5 scenarios. Furthermore, the impact of climate change will be more pronounced under higher than under

858   lower nutrient conditions. Hence, without further nutrient input reductions, as suggested by the BSAP,

859   eutrophication and oxygen depletion will worsen. However, the response determined in recent studies differs

860   considerably from the responses reported in previous studies, because of more plausible assumptions regarding

861   historical and future nutrient inputs. In some cases this has led to opposite signs in the response of bottom oxygen

862   concentrations. The new scenarios suggest that implementation of the BSAP would lead to a significant

863   improvement in the ecological status of the Baltic Sea regardless of the applied RCP scenario.

864

865   However, recent studies identified SLR as a new global driver. Depending on the combination of SLR and RCP

866   scenarios, the impact on the bottom oxygen concentration may be significant. A higher mean sea level relative to

867   the seabed at the sills would cause increased saltwater inflows, a stronger vertical stratification in the Baltic Sea

868   and a larger hypoxic area. The relationship between vertical stratification and the size of the hypoxic area was

869   confirmed in historical measurements. Nevertheless, recent studies suggest that the difference in future nutrient

870   emissions between the BSAP and REF scenarios is a more important driver than the projected changes in climate

871   with respect to changes in hypoxic area, phytoplankton concentration, water transparency (Secchi depth), primary

872   production and nitrogen fixation.

873

874   The currently available ensembles of scenario simulations are larger than in previous studies. Consequently, the

875   uncertainty range covered by the assessed ESMs and, in turn, the spread of the results are also larger. However,

876   the ensemble size might still be too small and model uncertainty is very likely underestimated. Moreover, natural

877   variability might be a more important source of uncertainty than previously considered for applications in the

878   Baltic Sea.

879

880   In the present climate, the climate variability of the Baltic Sea region during winter is dominated by the impact of

881   the NAO. However, in the past the correlation between the NAO and regional variables such as water temperature

882   or sea ice varied in time. The low-frequency changes in this correlation are projected to continue. Furthermore,
883   systematic changes in the influence of the large-scale atmospheric circulation on regional climate and on the NAO
884   itself could not be detected. While a northward shift in the mean summer position of the westerlies at the end of
885   the 21st century compared to the 20th century was reported (Gröger et al., 2019), it was based upon a limited set
886   of simulations with a few ESMs.

905

906 **Figures**

907

908 **Figure 1:** Bottom topography of the Baltic Sea (depth in m). The Baltic proper comprises the Arkona Basin,

909 Bornholm Basin and Gotland Basin. The border of the analysed domain of the Baltic Sea models is shown as a

910 black line in the northern Kattegat. The tide gauges Klagshamn (55.522ºN, 12.894ºE), Landsort (58.742ºN,

911 17.865ºE), Hamina (60.563ºN, 27.179ºE), and Kalix (65.697ºN, 23.096ºE) are also depicted.

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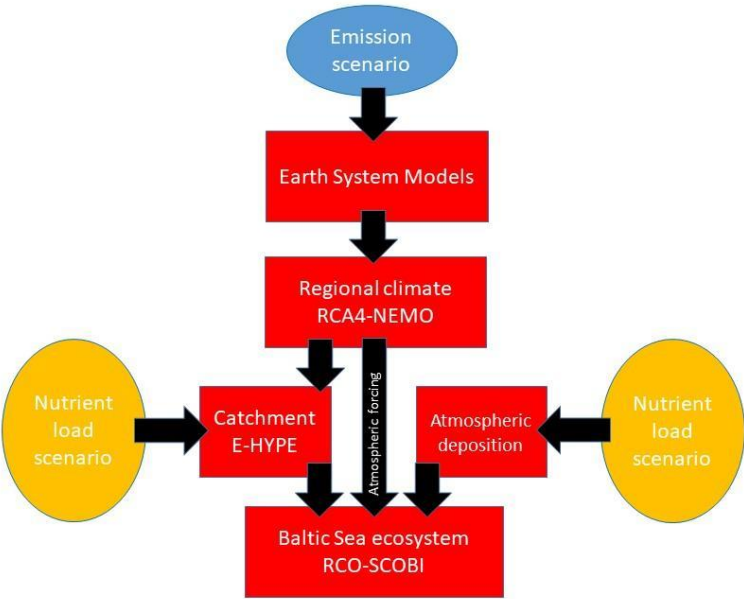
Figure 2. Dynamical downscaling approach for the Baltic Sea region. The models for the various components of the Earth System are explained in Section 2. (Source: Meier et al., 2021)

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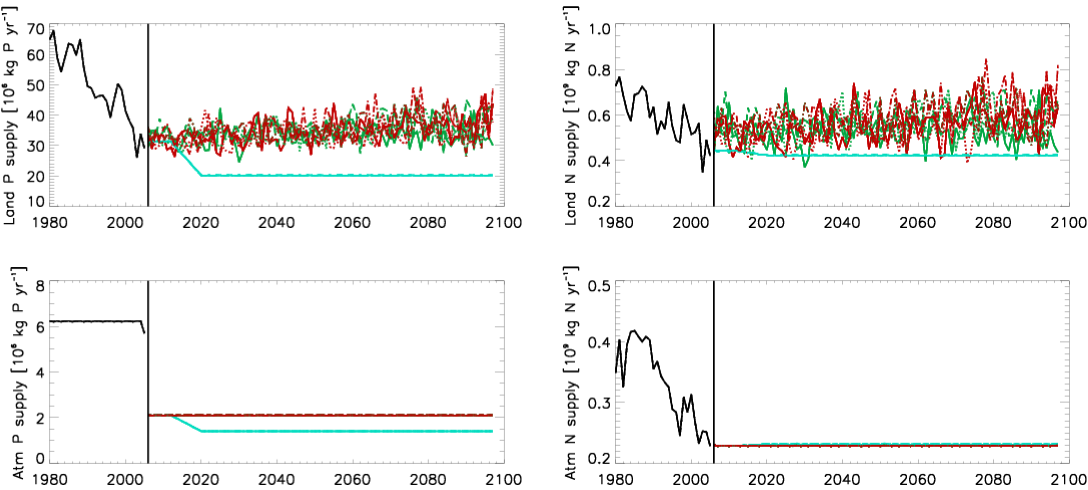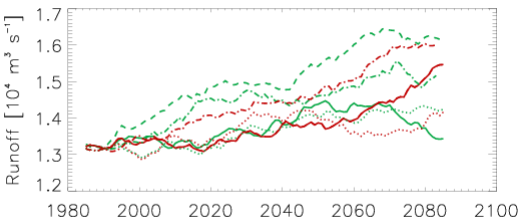
Figure 3. Projections of river discharge and nutrient inputs from land and atmosphere into the entire Baltic Sea according to the BalticAPP and CLIMSEA scenario simulations. Upper panel: Low-pass filtered runoff data (in $m^3 s^{-1}$) using a cut-off period of 30 years in four regionalised Earth System models (ESMs; illustrated by different line types) under RCP4.5 (green) and RCP8.5 (red) scenarios. Lower panels: Bioavailable phosphorus (in 10^6 kg P year^{-1}, left panels) and nitrogen inputs (in 10^9 kg N year^{-1}, right panels) from land (upper panels) and the atmosphere (lower panels) under RCP4.5, BSAP (blue), RCP4.5, REF (green), RCP8.5, BSAP (orange) and RCP8.5, REF (red) scenarios. Nutrient inputs during the historical period are depicted in black. The nutrient input scenario WORST of the BalticAPP scenario simulations (Saraiva et al., 2019a; their Fig. 4) is not displayed, neither are the ECOSUPPORT nutrient input scenarios (Gustafsson et al., 2011; their Fig. 3.1). (Source: Meier et al., 2021)

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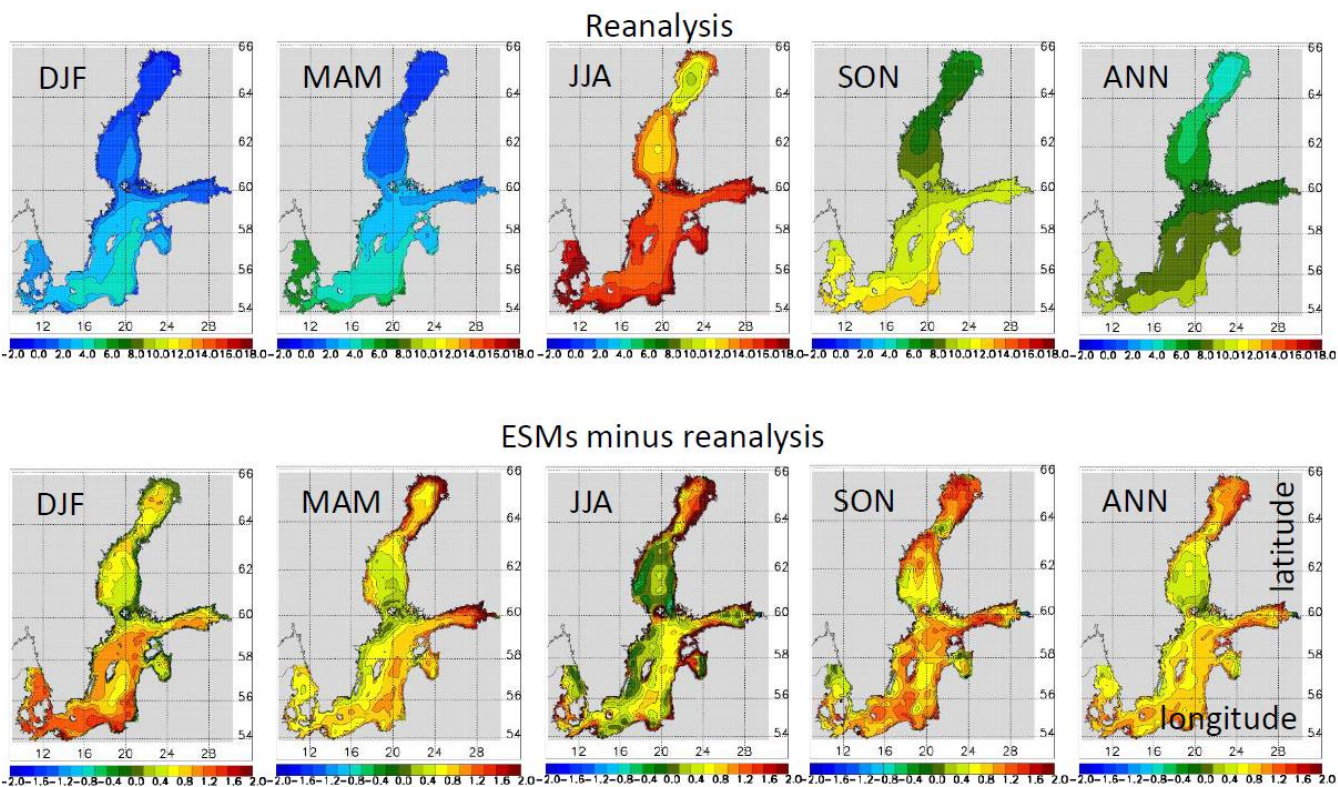
Figure 4: Upper panels: Annual and seasonal mean sea surface temperature (SST, in °C) in a reanalysis of data from 1970 to 1999 (Liu et al., 2017). Lower panels: Difference between the climatologies of the ensemble mean of the regionalised ESMs used in BalticAPP (Saraiva et al., 2019a) and CLIMSEA (Meier et al., 2021) during the historical period (1976-2005) and those of the reanalysis data. From the left to right: winter (December–February, DJF), spring (March–May, MAM), summer (June–August, JJA), autumn (September–November, SON) and annual (ANN) mean SSTs or SST differences.

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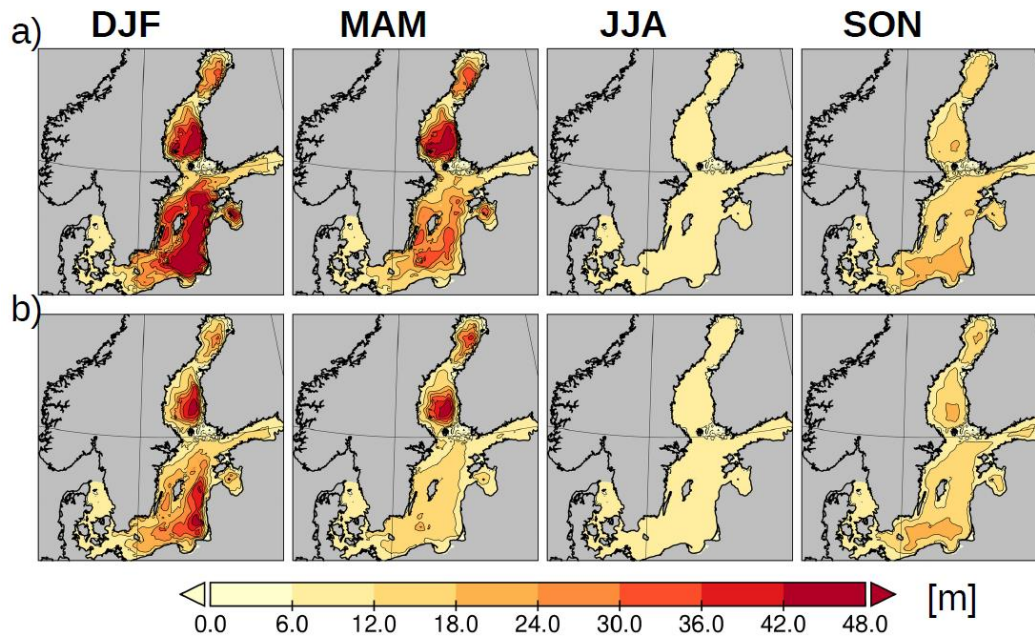
Figure 5: Mixed-layer thickness calculated according to the criterion following de Boyer Montégut et al. (2004). a) Reanalysis data (Liu et al., 2017). b) Ensemble mean over the four models (Saraiva et al., 2019a). Shown are the averages during 1976–1999.
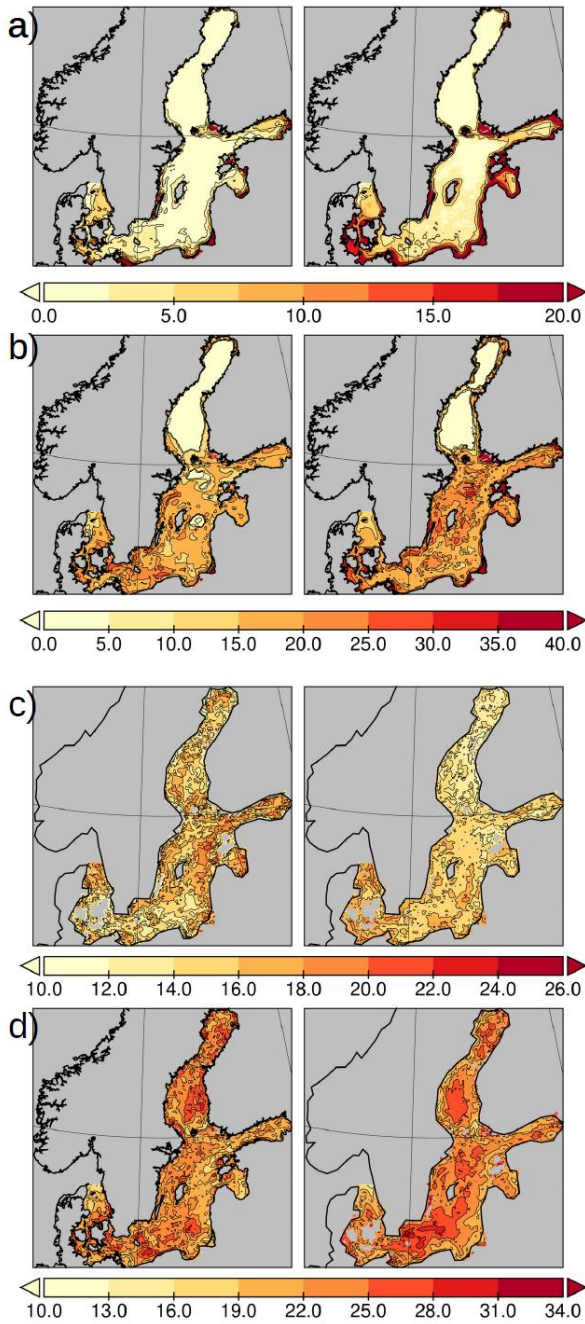
Figure 6: a) Number of >10-day periods in which the SST exceeds 20°C. b) Average duration of the periods displayed in a). c) Number of 10-day periods in which the SST exceeds the 95th percentile. d) Average duration of the periods displayed in c). Left column: reanalysis data (Liu et al., 2017). Right column: ensemble mean of the scenario simulations driven by four ESMs (Saraiva et al., 2019a). The analysis period is 1976–1999. Note the different colour scales used in c) and d).

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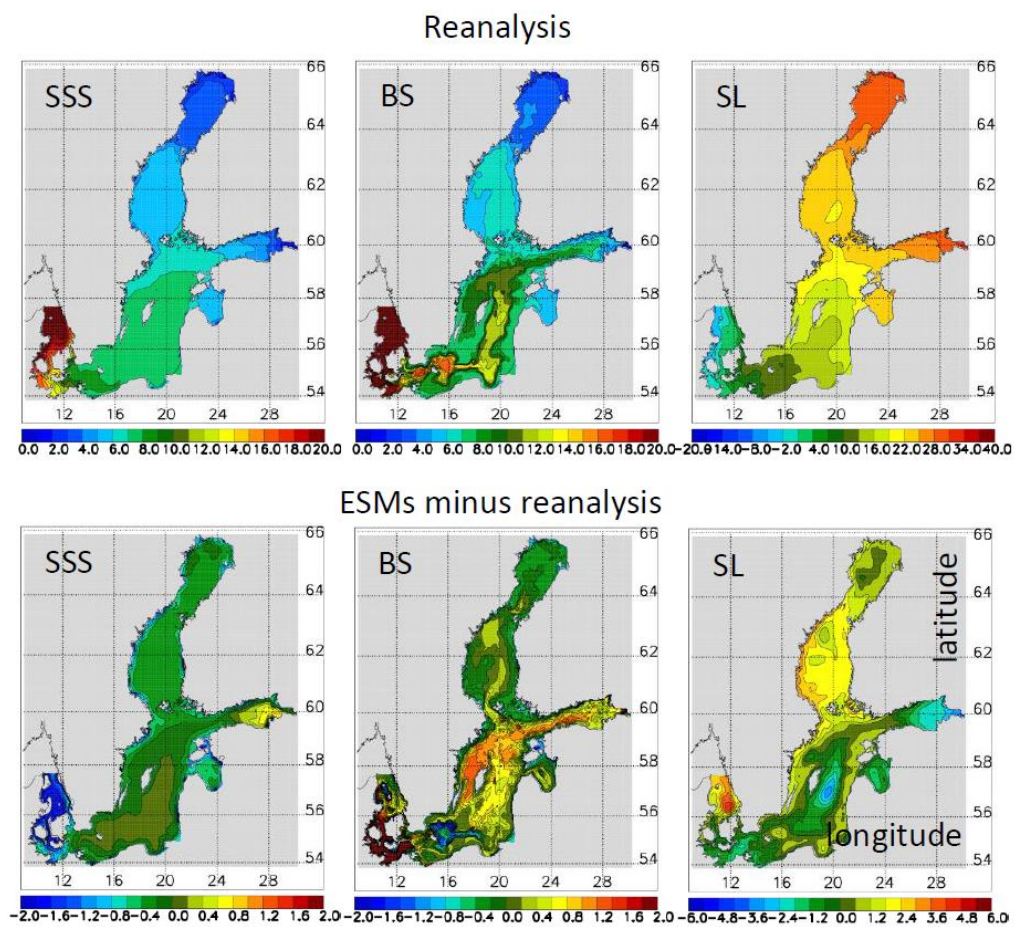
Figure 7: Upper panels: Annual mean sea surface salinity (SSS) and bottom salinity (BS) (in g kg^{-1}) and the winter (December–February) mean sea level (SL; in cm) in the reanalysis data of 1971–1999 (Liu et al., 2017; from left to right). Note that the model results for sea level are given in the Nordic height system 1960 (NH60) by Ekman and Mäkinen (1996). Lower panels: Difference between the climatologies of the ensemble mean of the regionalised ESMs used in BalticAPP (Saraiva et al., 2019a) during the historical period (1976–2005) and those of the reanalysis data.

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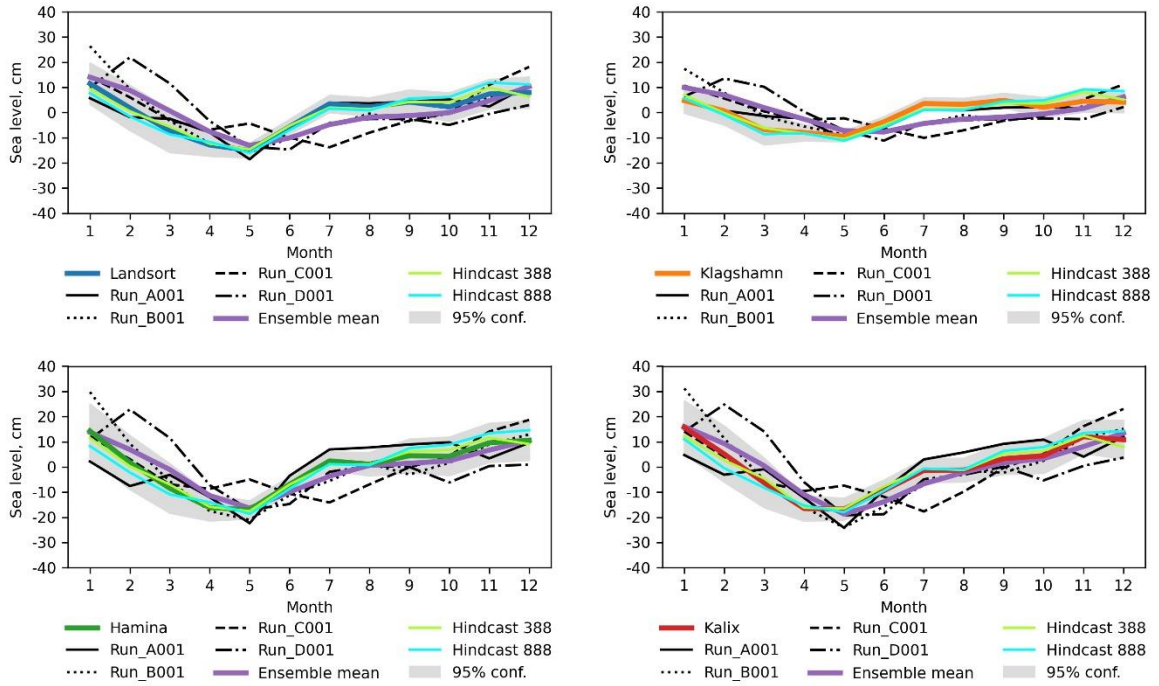
Figure 8: Monthly mean sea level according to a hindcast (driven by a regionalised reanalysis of atmospheric surface fields, i.e. RCA4 driven by ERA-40; Hindcast 388), reanalysis with the data assimilation of Liu et al. (2017) (Hindcast 888) and four climate simulations following Saraiva et al. (2019a) (Run_A001, …, Run_D001), the ensemble mean and observations for the historical period 1976–2005 at the sea level stations Klagshamn, Landsort, Hamina and Kalix (for the locations, see Figure 1). The 95% confidence interval of the observations is shown as a grey shaded area.

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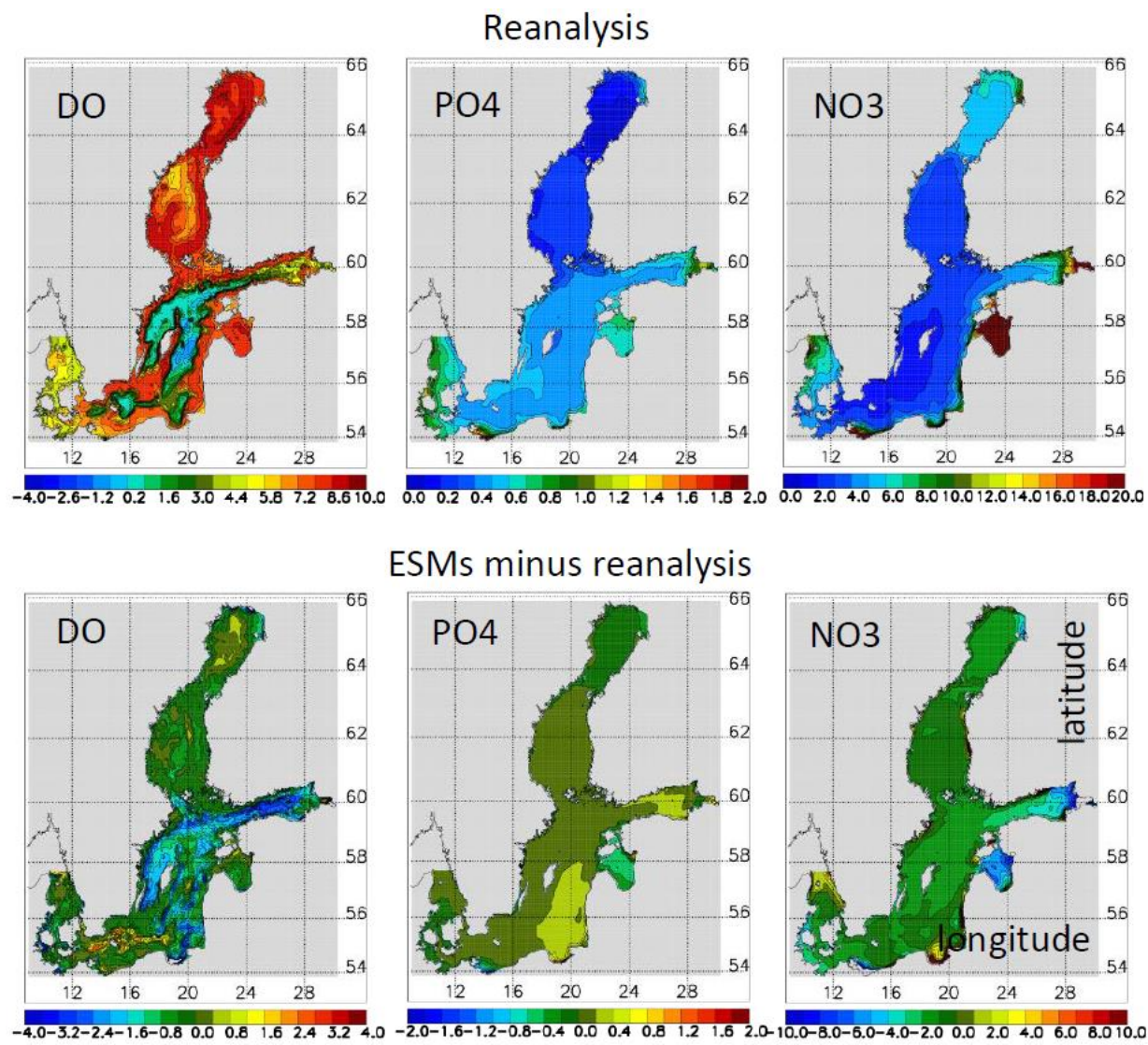
Figure 9: Upper panels: Summer (June–August) mean bottom dissolved oxygen (DO) concentrations (in mL L^{-1}), winter (December–February) mean surface phosphate (PO4) concentrations (in mmol P m^{-3}) and winter (December–February) mean surface nitrate (NO3) concentrations (in mmol N m^{-3}) in the reanalysis data of 1976–1999 (Liu et al., 2017). Negative oxygen concentration equivalents denote hydrogen sulphide concentrations, with 1 mL H$_2$S L^{-1} = –2 mL O$_2$ L^{-1}. Nutrient concentrations are vertically averaged for the upper 10 m. Lower panels: Difference between the climatologies of the ensemble mean of the ESMs (Saraiva et al., 2019a) and those of the reanalysis data of the historical period (1976–2005).

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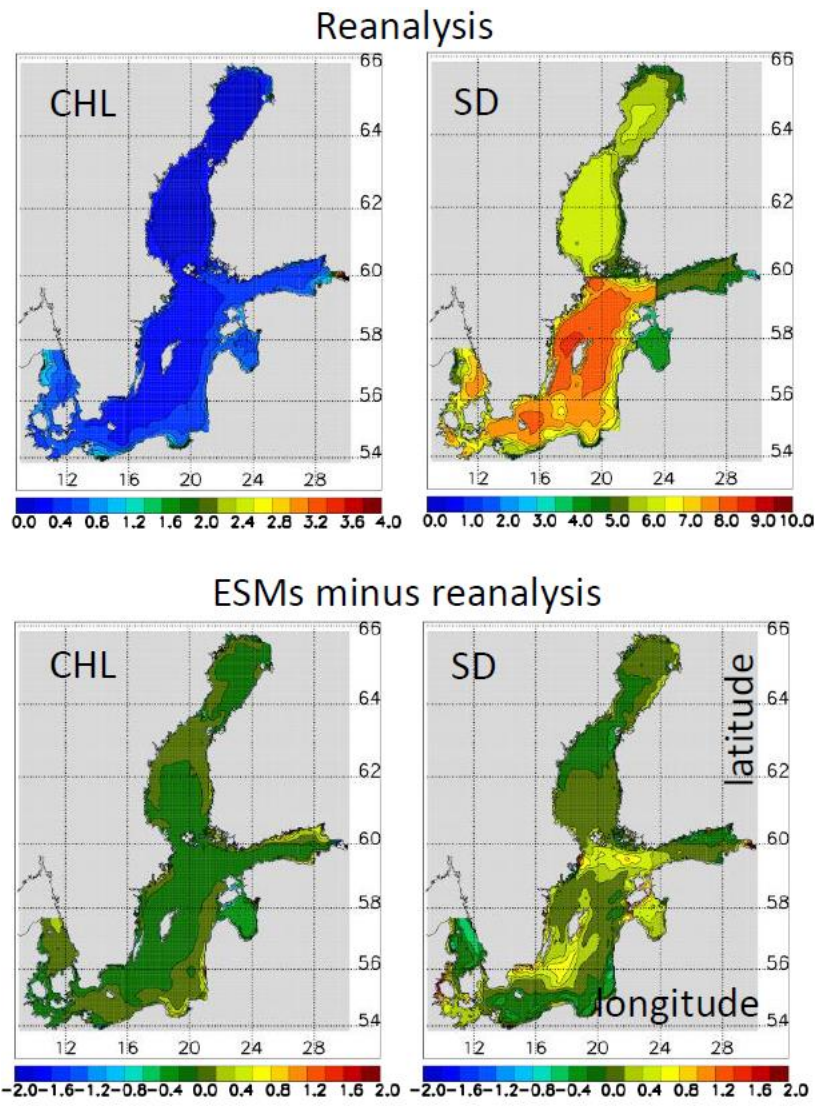
Figure 10: Upper panels: Annual mean phytoplankton concentrations (CHL; in mg Chl m^{-3}) and annual mean Secchi depth (SD; in cm) of the reanalysis data for 1976–1999 (Liu et al., 2017). Phytoplankton concentrations are vertically averaged for the upper 10 m. Since in the calculation of the Secchi depth as background only one value for the concentration of yellow substances per sub-basin is available, artificial borders between sub-basins become visible. Lower panels: Difference between the climatologies of the ensemble mean of the ESMs (Saraiva et al., 2019a) and those of the reanalysis data for the historical period (1976–2005).
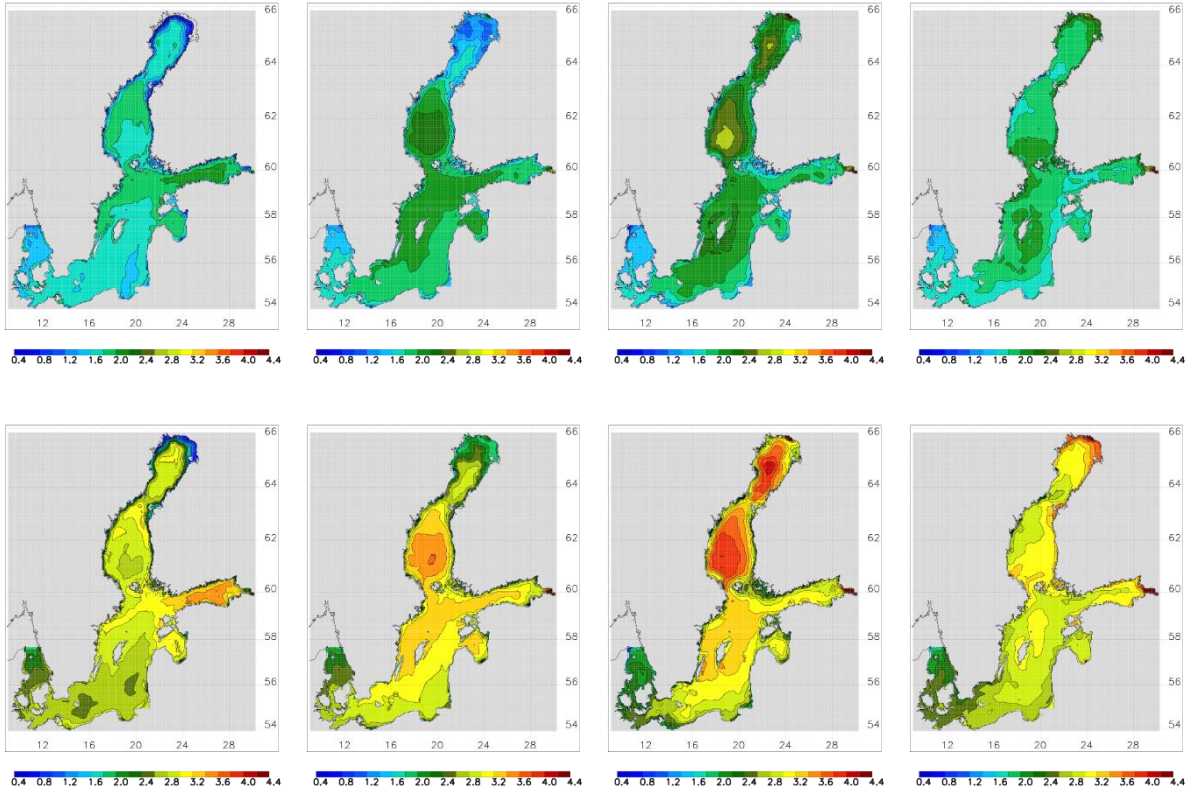
Figure 11. Changes in seasonal mean SST as simulated by the CLIMSEA ensemble (Meier et al., 2021). From left to right, mean SST changes (in °C) in winter (December, January and February; DJF), spring (March, April and May; MAM), summer (June, July and August; JJA) and autumn (September, October and November; SON) between 1976–2005 and 2069–2098 under RCP4.5 (upper panels) and RCP8.5 (lower panels).

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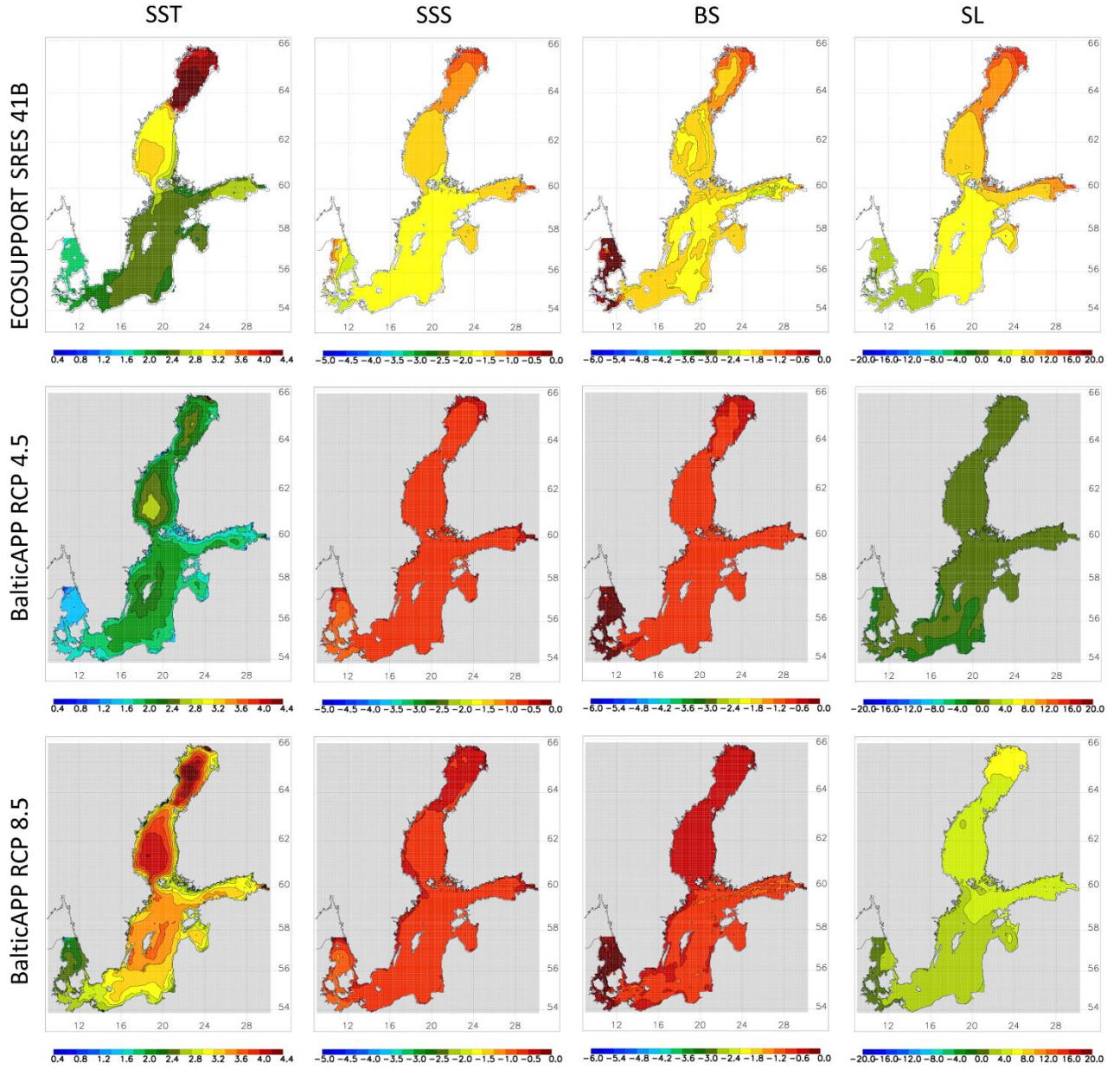
Figure 12: From left to right, changes in the mean SST (°C) in summer (June–August), the annual mean sea surface salinity (SSS; g kg⁻¹), annual mean bottom salinity (BS; g kg⁻¹), and winter (December–February) mean sea level (SL; cm) between 1978–2007 and 2069–2098. From top to bottom, the results of the ensembles ECOSUPPORT (white background, Meier et al., 2011b), BalticAPP RCP4.5 (grey background, Saraiva et al., 2019a) and BalticAPP RCP8.5 (grey background, Saraiva et al., 2019a).
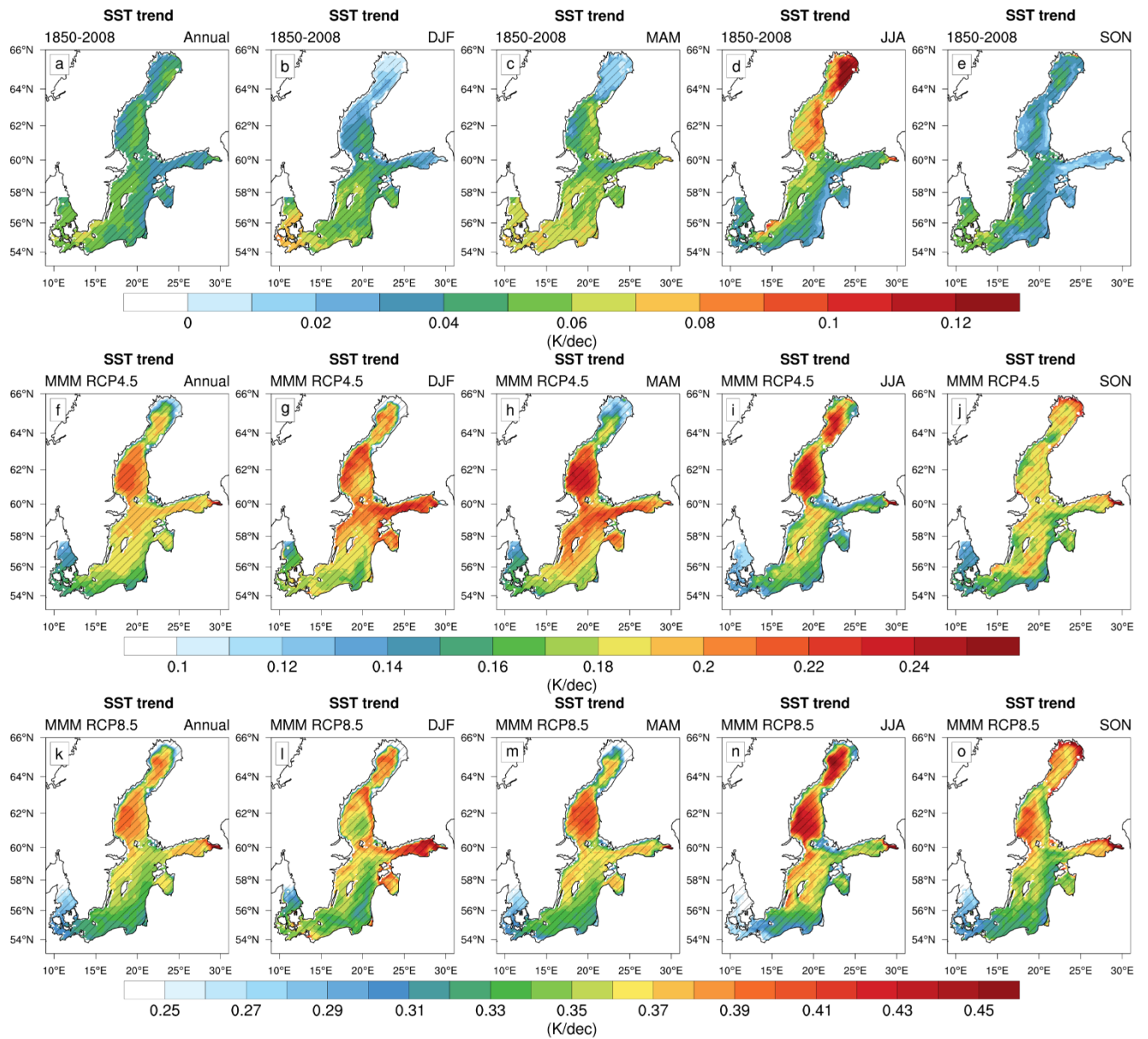
Figure 13: Multi-model mean (MMM) of annual (a, f, k) and seasonal (b–e, g–j, l–o) SST trends (in °C decade^{-1}) computed for the period 1850–2008 (top), 2006–2099 in RCP4.5 (middle) and RCP8.5 (bottom) scenario. Hatched areas represent the regions where the trend is statistically significant ($p < 0.05$, Mann-Kendall test). Data for historical reconstructions and projections are from Meier et al. (2019d) and Saraiva et al. (2019a) respectively.

Relative SST trend

MMM   RCP4.5

Relative SST trend

MMM   RCP8.5

-0.08 -0.06 -0.04 -0.02   0   0.02 0.04 0.06 0.08

(K/dec)
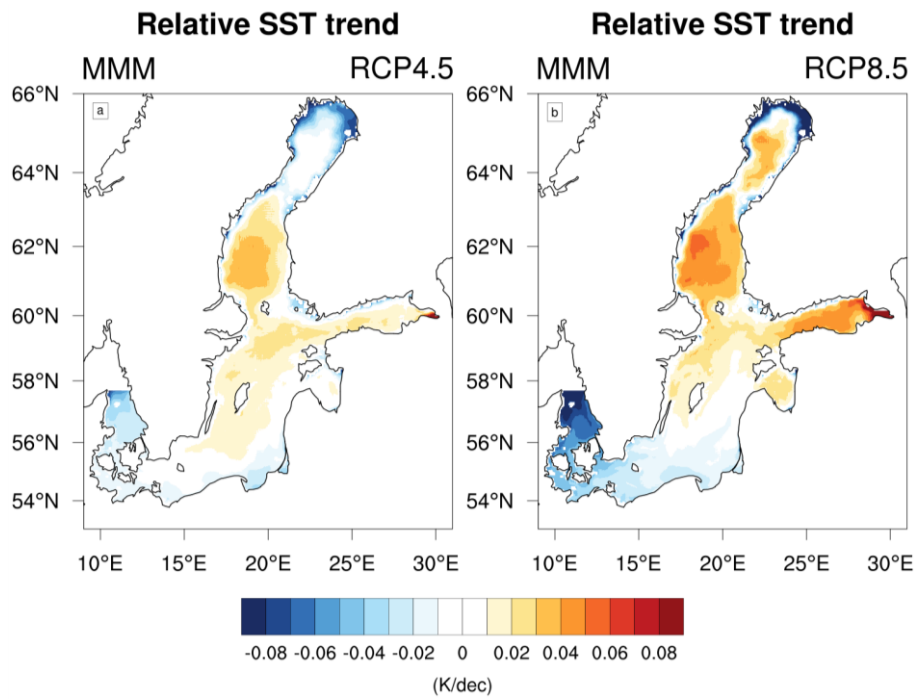
Figure 14: Multi-model mean (MMM) of the annual SST trends relative to the spatial average (in °C decade^{-1}) for a) RCP4.5 and b) RCP8.5 scenario simulations. (Data source: Saraiva et al., 2019a)
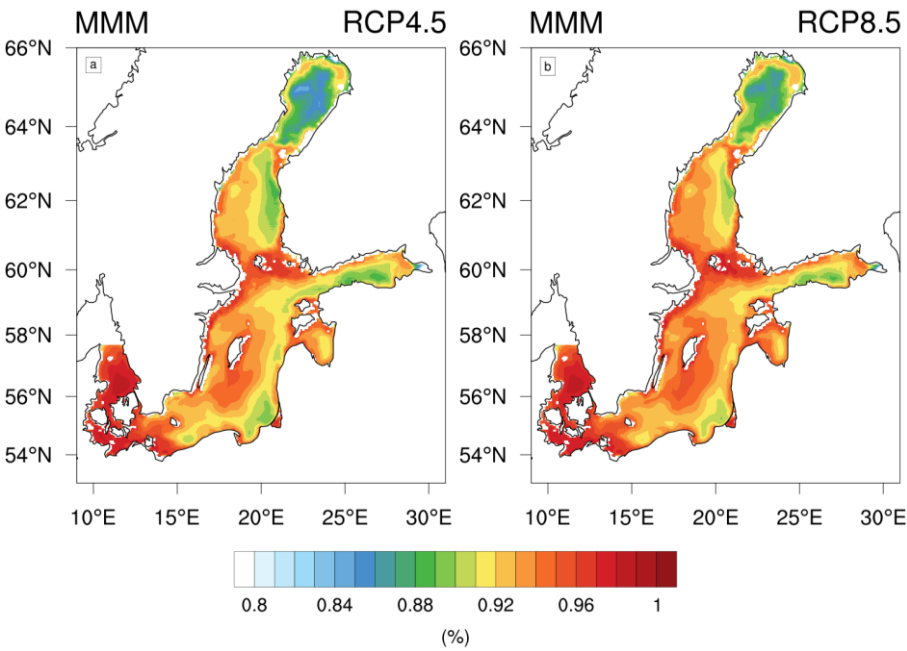
Figure 15: Multi-model mean (MMM) explained variance (in percent) between the monthly mean SST and the forcing air temperature over the period 2006–2099 in the a) RCP4.5 and b) RCP8.5 scenario simulations. (Data source: Saraiva et al., 2019a)
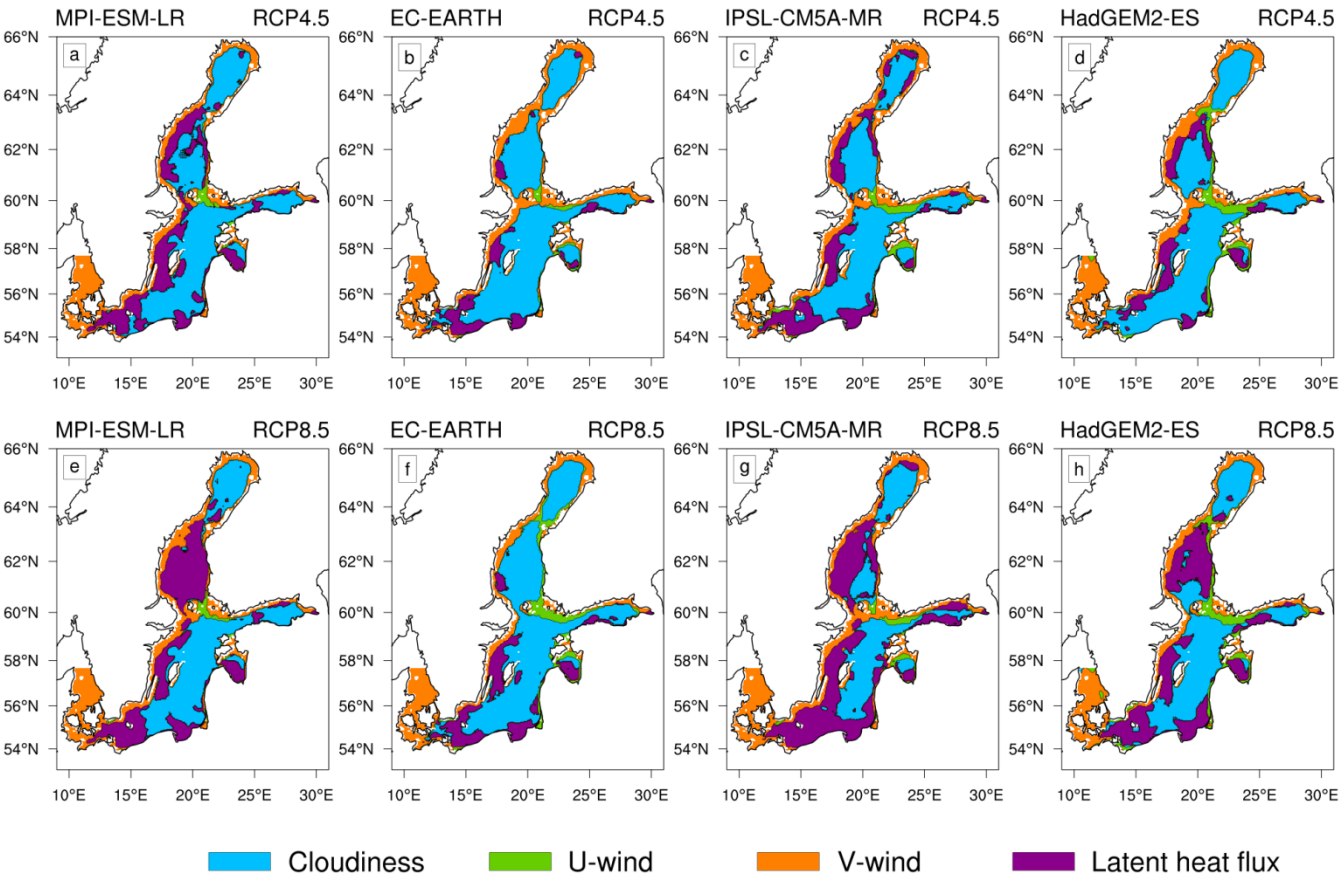
Figure 16: Results of the cross-correlation analysis of the detrended SST (monthly mean) with the wind components, latent heat flux and cloudiness. Maps of the atmospheric drivers with the highest cross-correlations in the RCP4.5 (top) and RCP8.5 (bottom) scenarios for various GCM forcings (Saraiva et al., 2019a). From left to right: MPI-ESM-LR, EC-EARTH, IPSL-CM5A-MR, HadGEM2-ES.
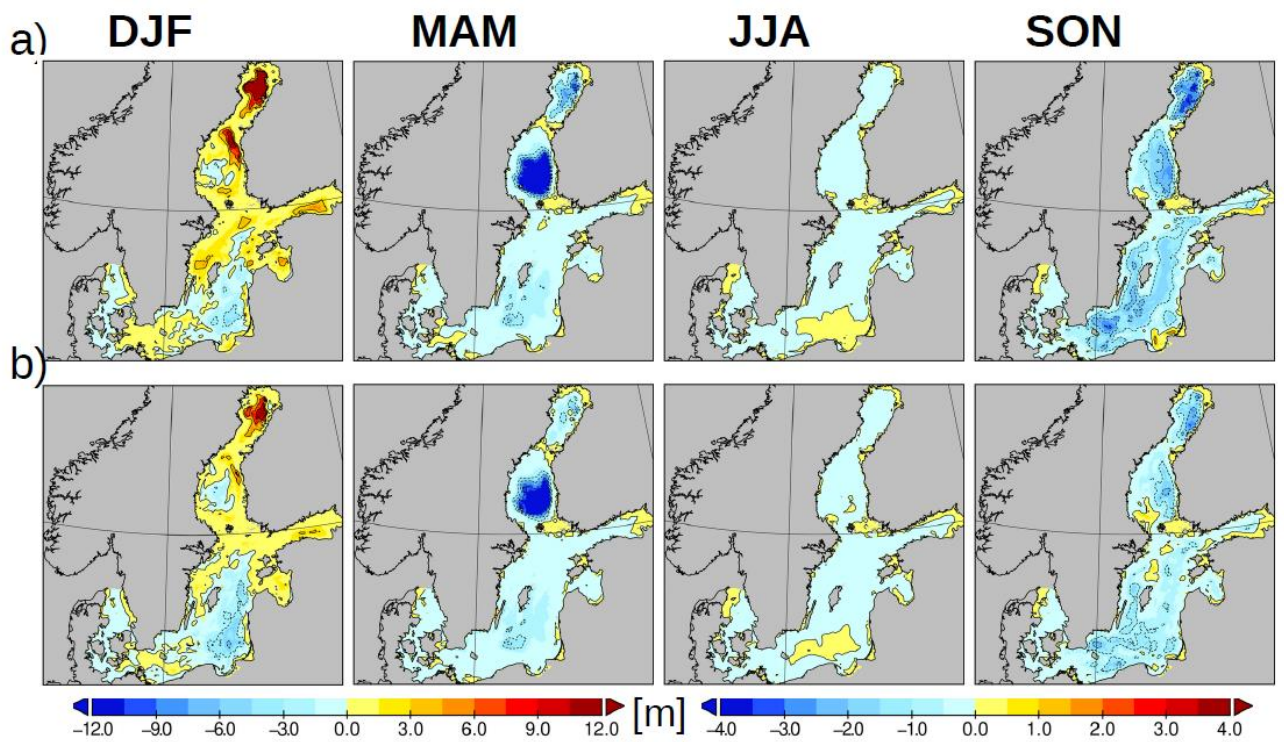
Figure 17. Mixed-layer depth calculated according to the criterion of de Boyer Montégut et al. (2004). Shown are the ensemble average changes of four different ESMs between 1976–2005 and 2069–2098 with the mean sea level rises a) 0.90 m (RCP8.5) and b) 0.54 m (RCP4.5). (Data source: Meier et al., 2021)
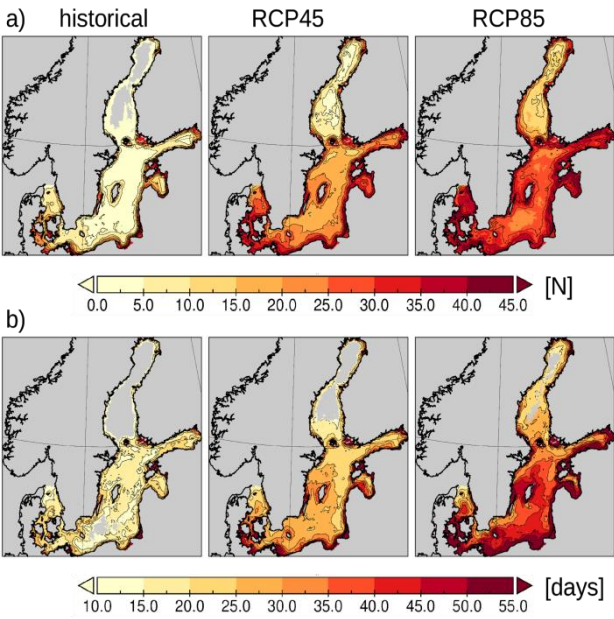
Figure 18. a) Number of heat waves (defined as the number of periods ≥ 10 days in which the water temperature is ≥ 20°C) for historical (1976–2005), and future (2069–2098) climates. b) Average duration of the heat waves. Note that no temperature bias adjustment was done prior to the analysis. Shown are the ensemble averages of four different ESMs with the mean sea level rises a) 0.54 m (RCP4.5) and b) 0.90 m (RCP8.5). (Data source: Meier et al., 2021)
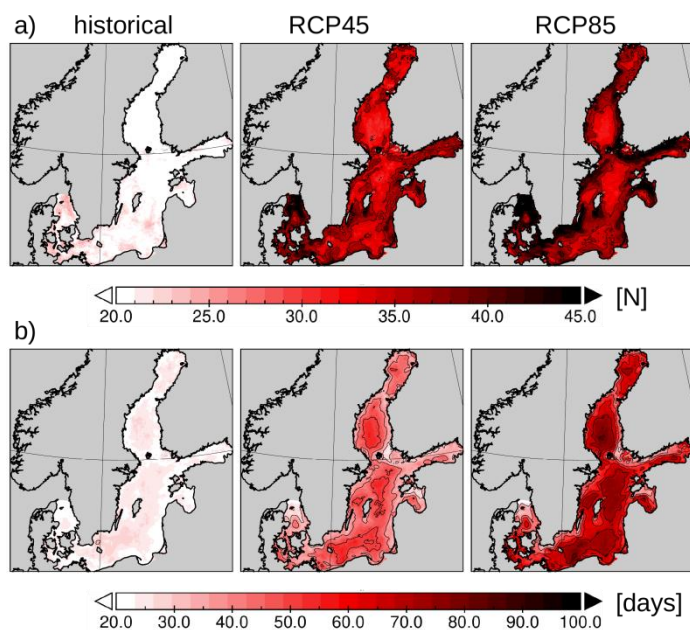
Figure 19. Same as in Figure 18 but for heat waves defined as periods of ≥10 days in which the water temperature exceeded the 95th percentile of the historical reference temperature. (Data source: Meier et al., 2021)
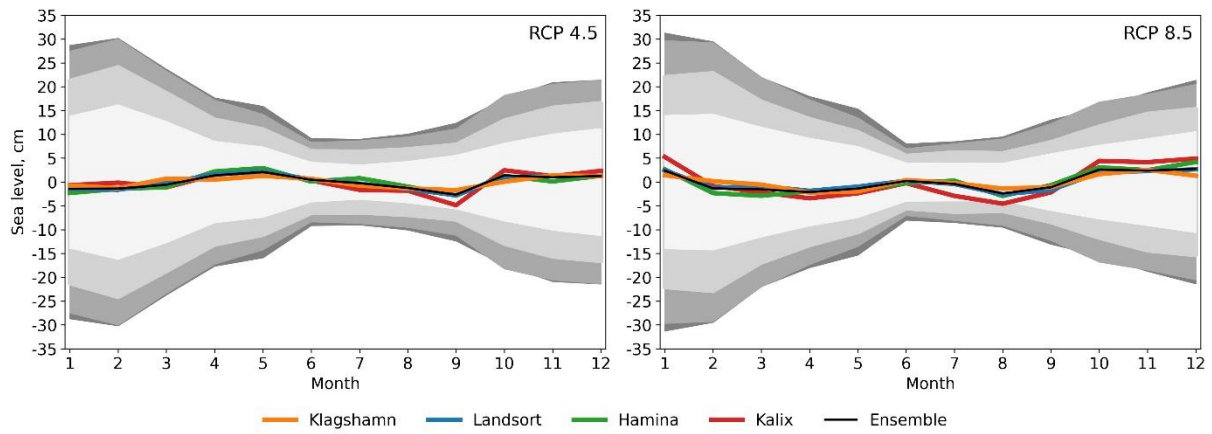
Figure 20: Monthly mean sea level changes between 1976–2005 and 2069–2098 at Klagshamn, Landsort, Hamina and Kalix (for the locations, see Figure 1) for RCP4.5 (left panel) and RCP8.5 (right panel). Shown are the changes relative to the mean sea level rise a) 0.54 m (RCP4.5) and b) 0.90 m (RCP8.5). Shaded areas in white to dark grey denote +/- two standard deviations at Klagshamn to Kalix respectively. The chosen model approach does not indicate any non-linear effects for scenarios with a larger rise in sea level. (Data source: Meier et al., 2021)

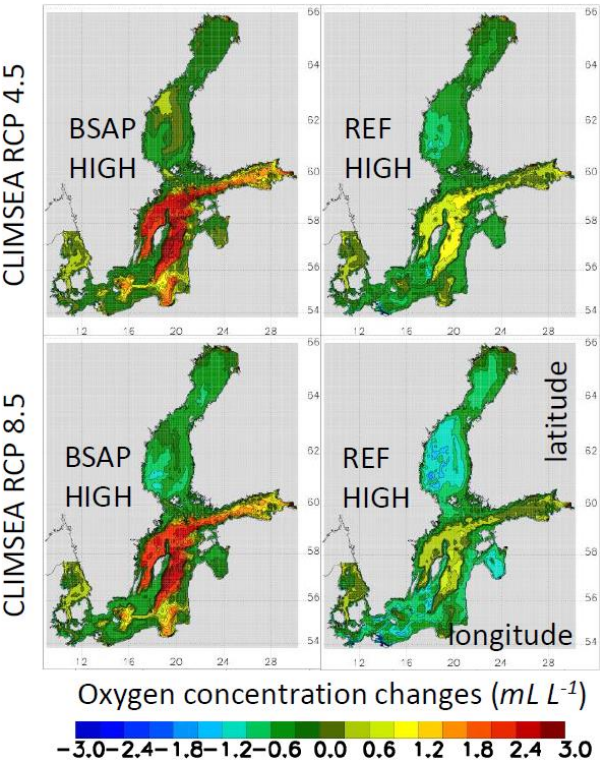1049   **(a)**
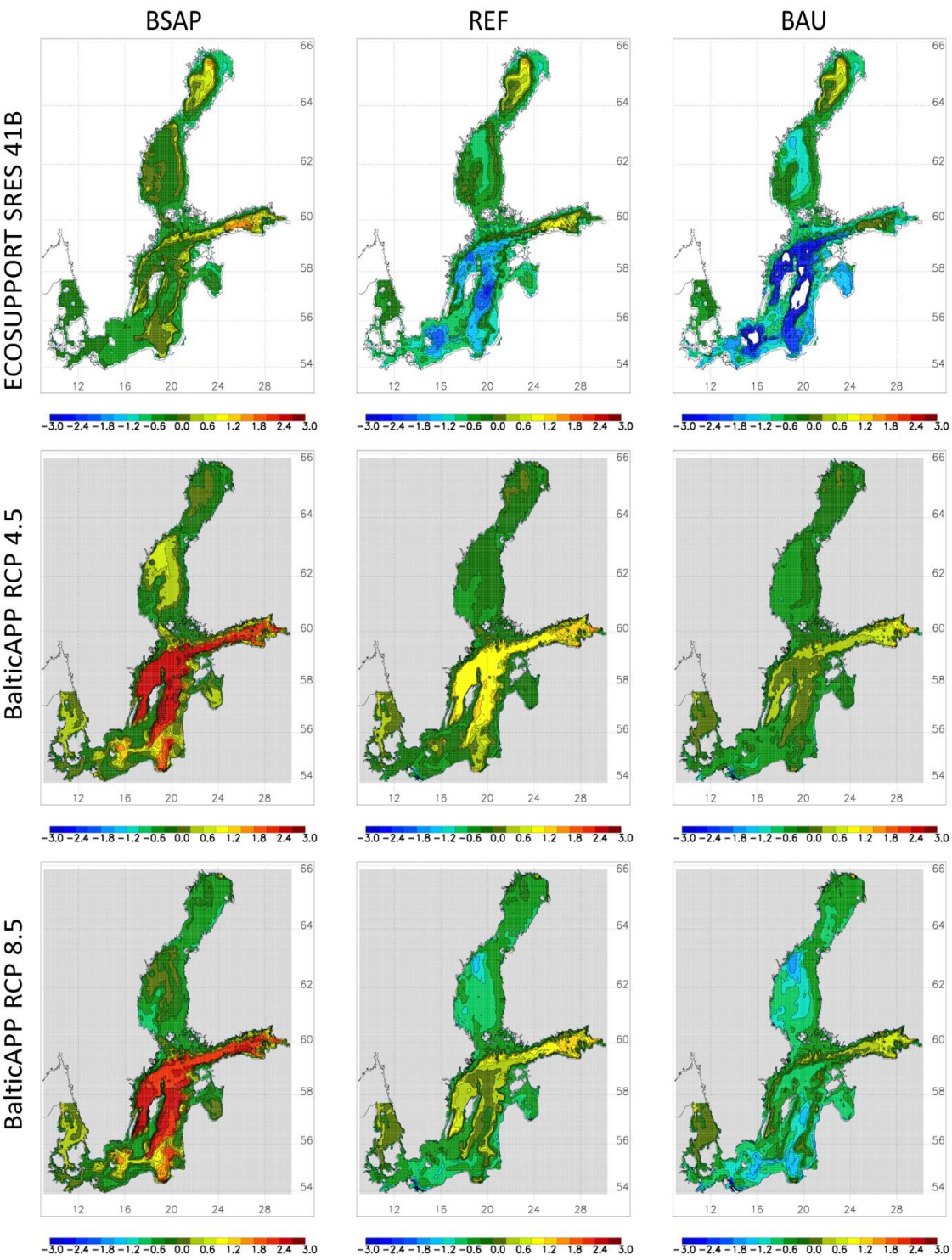
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(b)

Oxygen concentration changes (*mL L⁻¹*)

−3.0 −2.4 −1.8 −1.2 −0.6 0.0 0.6 1.2 1.8 2.4 3.0

Figure 21: a) Ensemble mean changes in the bottom dissolved oxygen concentration (mL L^{-1}) in summer (June–August) between 1978–2007 and 2069–2098. From left to right, the results of the nutrient input scenarios of the Baltic Sea Action Plan (BSAP), Reference (REF) and Business-As-Usual (BAU). From top to bottom, the results of the ensembles ECOSUPPORT (white background; Meier et al., 2011b), BalticAPP RCP4.5 (grey background; Saraiva et al., 2019a) and BalticAPP RCP8.5 (grey background; Saraiva et al., 2019a). b) As in panel a) but for CLIMSEA RCP4.5 (upper panels) and CLIMSEA RCP8.5 (lower panels) under a high sea level rise scenario, i.e. 1.26 m (RCP4.5) and 2.34 m (RCP8.5). Left and right columns show the BSAP and REF scenarios respectively. (Source: Meier et al., 2021)

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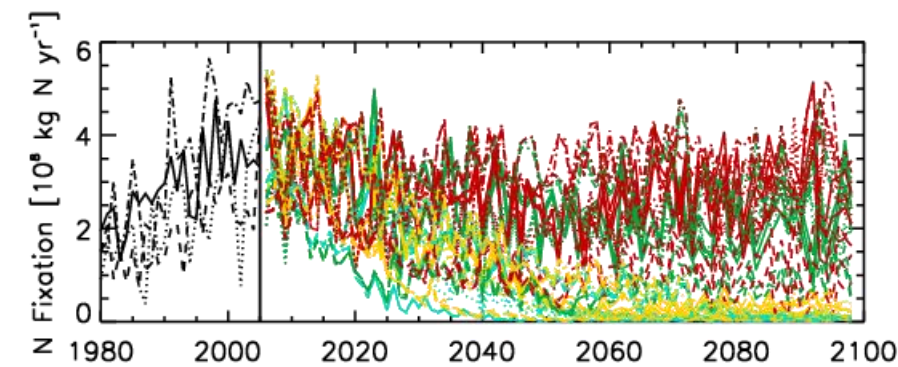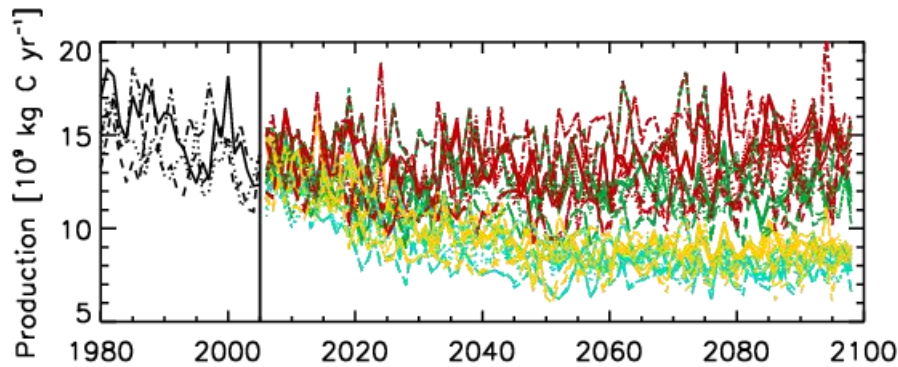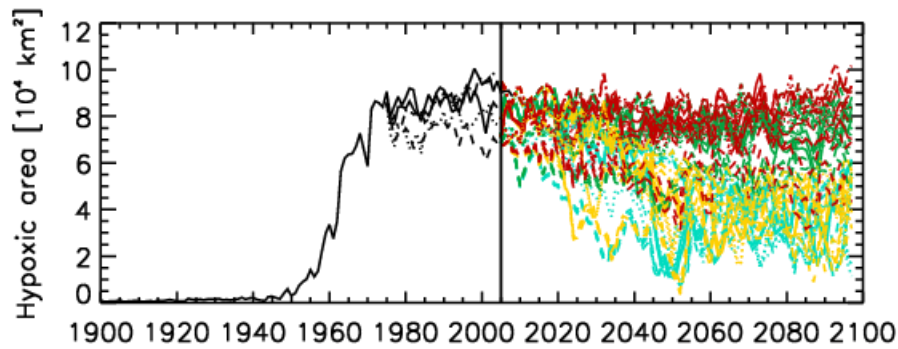
1073 **Figure 22:** From top to bottom: hypoxic area (in km^2), volume-averaged primary production (in kg C year^{-1}) and

1074 volume-averaged nitrogen fixation (in kg N year^{-1}) for the entire Baltic Sea, including the Kattegat (see Figure 1)

1075 in the historical (until 2005, black lines) and scenario (after 2005, coloured lines) simulations driven by four

1076 regionalised ESMs (illustrated by different line types) under RCP4.5, BSAP (blue), RCP4.5, REF (green), RCP8.5,

1077 BSAP (orange) and RCP8.5, REF (red) scenarios. A spin-up simulation since 1850 was performed, as illustrated

1078 by the evolution of hypoxia. (Source: Meier et al., 2021)

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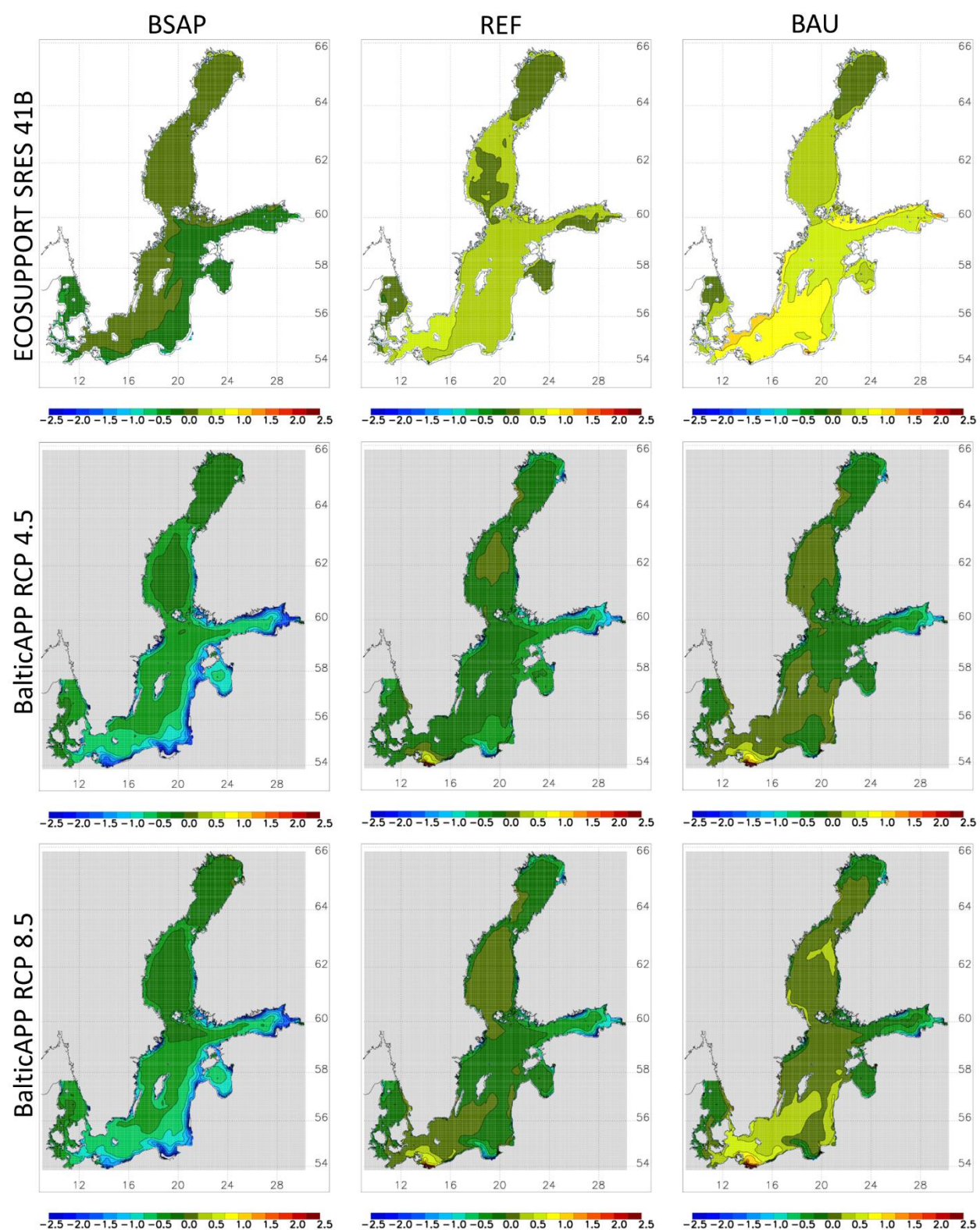
Figure 23: As in Figure 21a but for annual mean surface phytoplankton concentration changes (mg Chl m^{-3}).
Concentrations are vertically averaged for the upper 10 m. (Source: Meier et al., 2011b; Saraiva et al., 2019a)

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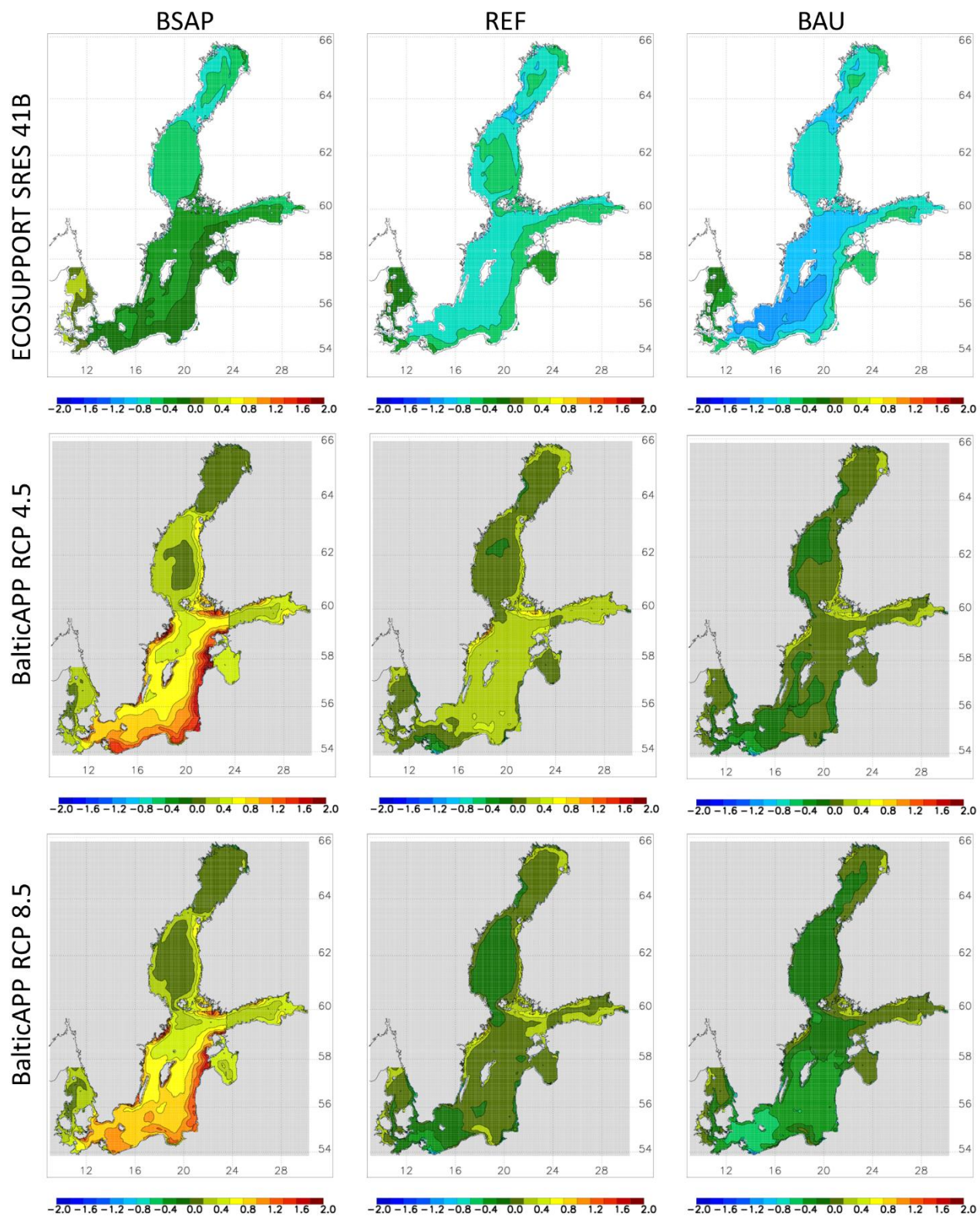
Figure 24. As in Figure 21a but for changes in the annual mean Secchi depth (m). (Source: Meier et al., 2011b; Saraiva et al., 2019a)

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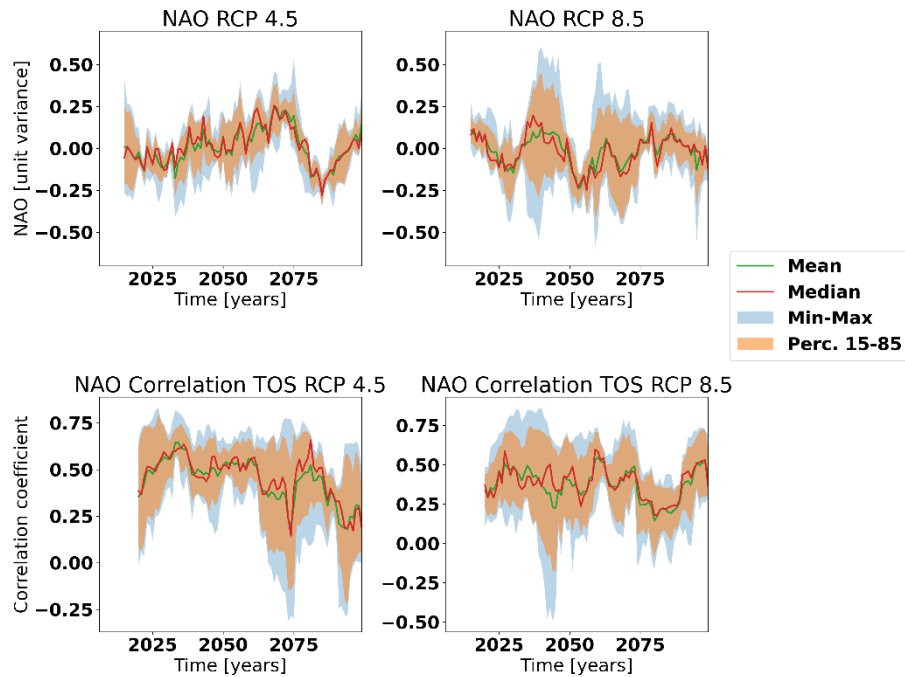
1090 **Figure 25.** Ensemble 10-year running mean North Atlantic Oscillation (NAO) index (upper panels) and 10-year

1091 running correlation between the NAO and area averaged SST in the Baltic Sea (lower panels) under RCP4.5 (left

1092 panels) and RCP8.5 (right panels) scenarios. Depicted are winter (December–February) mean, median, minimum,

1093 maximum and 15th and 85th percentiles. (Data source: Meier et al., 2021)

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Tables

1096 **Table 1.** Selected ensembles of the scenario simulations for the Baltic Sea carried out in international projects (AR

1097 = IPCC Assessment Report, GCM = General Circulation Model, RCSM = Regional Climate System Model,

1098 RCAO = Rossby Centre Atmosphere Ocean model, RCA4 = Rossby Centre Atmosphere model Version 4, NEMO

1099 = Nucleus for European Modelling of the Ocean, REMO = Regional Model, MPIOM = Max Planck Institute

1100 Ocean Model, HAMSOM = Hamburg Shelf Ocean Model)

Project	Swedish Regional Climate Modelling Program	Advanced modelling tool for scenarios of the Baltic Sea ECOsystem to SUPPORT decision making	Holocene saline water inflow changes into the Baltic Sea, ecosystem responses and future scenarios	Building predictive capability regarding the Baltic Sea organic/inorganic carbon and oxygen systems	Wellbeing from the Baltic Sea - applications combining natural science and economics	Impacts of Climate Change on Waterways and Navigation	Regionally downscaled climate projections for the Baltic and North Seas
Acronym	SWECLIM	ECOSUPPORT	INFLOW	Baltic-C	BalticAPP	KLIWAS	CLIMSEA
Duration	1997-2003	2009-2011	2009-2011	2009-2011	2015-2017	2009-2013	2018-2020
Project summaries	Rummukainen et al., 2004	Meier et al., 2014	Kotilainen et al., 2014	Omstedt et al., 2014	Saraiva et al., 2019a	Bülow et al., 2014	Meier et al., 2021
GCMs	AR3	AR4	AR4	AR4	AR5	AR4/AR5	AR5
RCSM	RCAO	RCAO	RCAO	RCA	RCA4-NEMO	REMO-MPIOM, REMO-HAMSOM, RCA4-NEMO	RCA4-NEMO
Horizontal resolution atmosphere/ ocean	50 km/10.8 km	25 km/3.6 km	25 km/3.6 km and 50 km/3.6 km for paleoclimate	25 km /horizontally integrated	25 km /3.6 km	varying	25 km/3.6 km
Period(s)	1961–1990 and 2071–2100	1961-2099	1961-2099 and 950-1800 AD	1960-2100	1976-2100, improved initial conditions	1961-2099	1976-2100
Ocean model	One physical Baltic Sea model	Three physical-biogeochemical Baltic Sea models	See ECOSUPPORT	One physical-biogeochemical Baltic Sea model including the carbon cycle	One physical-biogeochemical Baltic Sea model	Two physical regional models with focus on the Baltic Sea and North Sea regions and one physical-biogeochemical ocean model	One physical-biogeochemical Baltic Sea model
References	Döscher and Meier, 2004; Meier et al., 2004a; Meier et al., 2004b	Meier et al., 2011b; Meier et al., 2012c; Neumann et al., 2012	See ECOSUPPORT, Schimanke and Meier, 2016	Omstedt et al., 2012	Saraiva et al., 2019a, b; Meier et al., 2019b	Bülow et al., 2014; Gröger et al., 2019; Dieterich et al., 2019	Gröger et al., 2021b; Meier et al., 2021

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1102 **Table 2.** Salinity projections assessed by the BACC Author Team (2008), BACC II Author Team (2015) and
1103 BEAR (this study). Salinity changes depend on the changes in the wind field (in particular, the west wind
1104 component), river discharge and sea level rise (SLR).

	West wind	River discharge (%)	SLR	Salinity
BACC (2008)	Large increase	-8 to +26	0	0 to −3.7 g kg^{-1}
BACC II (2015)	Small increase	+15 to +22	0	−1 to −2 g kg^{-1}
BEAR (this study)	No significant change	+2 to +22	Medium SLR +0.54 to +0.90 m	No robust change, with a considerable spread

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Table 3. List of scenario simulations of three ensembles. From left to right, the columns show the Earth System
Model (ESM), the Regional Climate System Model (RCSM), the Baltic Sea Ecosystem Model, the greenhouse
gas (GHG) emission or concentration scenario, the nutrient input scenario, the sea level rise (SLR) scenario and
the simulation period, including historical and scenario periods. The four nutrient input scenarios were: Baltic Sea
Action Plan (BSAP), Reference (REF), Business-As-Usual (BAU) and Worst Case (WORST). For the three SLR
scenarios in the CLIMSEA ensemble, the mean sea level changes at the end of the century are given in meters.

ECOSUPPORT (28 scenario simulations, Meier et al., 2011b)						
ESM	RCSM	Baltic Sea Model	GHG scenario	Nutrient input scenario	SLR scenario	Period
HadCM3	RCAO	BALTSEM	A1B	BSAP/REF/BAU	0	1961–2099
ECHAM5/MPI-OM-r1	RCAO	BALTSEM	A1B	BSAP/REF/BAU	0	1961–2099
ECHAM5/MPI-OM-r3	RCAO	BALTSEM	A1B	BSAP/REF/BAU	0	1961–2099
ECHAM5/MPI-OM-r1	RCAO	BALTSEM	A2	BSAP/REF/BAU	0	1961–2099
HadCM3	RCAO	MOM-ERGOM	A1B	BSAP/REF	0	1961–2099
ECHAM5/MPI-OM-r1	RCAO	MOM-ERGOM	A1B	BSAP/REF	0	1961–2099
HadCM3	RCAO	RCO-SCOBI	A1B	BSAP/REF/BAU	0	1961–2099
ECHAM5/MPI-OM-r1	RCAO	RCO-SCOBI	A1B	BSAP/REF/BAU	0	1961–2099
ECHAM5/MPI-OM-r3	RCAO	RCO-SCOBI	A1B	BSAP/REF/BAU	0	1961–2099
ECHAM5/MPI-OM-r1	RCAO	RCO-SCOBI	A2	BSAP/REF/BAU	0	1961–2099
BalticAPP (21 scenario simulations, Saraiva et al., 2019a)						
MPI-ESM-LR	RCA4-NEMO	RCO-SCOBI	RCP4.5	BSAP/REF/WORST	0	1976–2099
MPI-ESM-LR	RCA4-NEMO	RCO-SCOBI	RCP8.5	BSAP/REF/WORST	0	1976–2099
EC–EARTH	RCA4-NEMO	RCO-SCOBI	RCP4.5	BSAP/REF/WORST	0	1976–2099
EC-EARTH	RCA4-NEMO	RCO-SCOBI	RCP8.5	BSAP/REF/WORST	0	1976–2099
IPSL-CM5A-MR	RCA4-NEMO	RCO-SCOBI	RCP4.5	BSAP/REF/WORST	0	1976–2099
HadGEM2-ES	RCA4-NEMO	RCO-SCOBI	RCP4.5	BSAP/REF/WORST	0	1976–2098
HadGEM2-ES	RCA4-NEMO	RCO-SCOBI	RCP8.5	BSAP/REF/WORST	0	1976–2098
CLIMSEA (48 scenario simulations, Meier et al., 2021)						
MPI-ESM-LR	RCA4-NEMO	RCO-SCOBI	RCP4.5	BSAP/REF	0/0.54/1.26	1976–2099

MPI-ESM-LR	RCA4-NEMO	RCO-SCOBI	RCP8.5	BSAP/REF	0/0.90/2.34	1976–2099
EC-EARTH	RCA4-NEMO	RCO-SCOBI	RCP4.5	BSAP/REF	0/0.54/1.26	1976–2099
EC-EARTH	RCA4-NEMO	RCO-SCOBI	RCP8.5	BSAP/REF	0/0.90/2.34	1976–2099
IPSL-CM5A-MR	RCA4-NEMO	RCO-SCOBI	RCP4.5	BSAP/REF	0/0.54/1.26	1976–2099
IPSL-CM5A-MR	RCA4-NEMO	RCO-SCOBI	RCP8.5	BSAP/REF	0/0.90/2.34	1976–2099
HadGEM2-ES	RCA4-NEMO	RCO-SCOBI	RCP4.5	BSAP/REF	0/0.54/1.26	1976–2098
HadGEM2-ES	RCA4-NEMO	RCO-SCOBI	RCP8.5	BSAP/REF	0/0.90/2.34	1976–2098

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1116 **Table 4:** Summary of the characteristics of the ECOSUPPORT, BalticAPP and CLIMSEA scenario simulations
1117 discussed in this study. For further details, the reader is referred to Tables 1 and 3. Acronyms are defined in
1118 Table 5.

Acronym	Atmospheric forcing	GHG emission or concentration scenario	Hydrological forcing	Nutrient input scenario	Special features
ECOSUPPORT	Regionalised CMIP3 data, two GCMs	SRES A1B, A2	STAT	BSAP, REF, BAU	Three Baltic Sea models
BalticAPP	Regionalised CMIP5 data, four ESMs	RCP4.5, 8.5	E-HYPE	Revised scenarios BSAP, REF, WORST	Four ESMs
CLIMSEA	Regionalised CMIP5 data, four ESMs	RCP4.5, 8.5	E-HYPE	Revised scenarios BSAP, REF	SLR is considered

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1122   **Table 5.** List of acronyms (in alphabetical order), their definitions and references

Acronym	Definition	Comment	Reference
AMO	Atlantic Multidecadal Oscillation	Mode of climate variability	Knight et al. (2005)
BACC	Assessment of climate change for the Baltic Sea basin	Regional climate change assessment	BACC Author Team (2008), BACC II Author Team (2015)
BalticAPP	Well-being from the Baltic Sea: applications combining natural science and economics	Climate modelling project for the Baltic Sea	Saraiva et al. (2019a)
BEAR	Baltic Earth assessment reports	Regional climate change assessment	https://baltic.earth
BSAP	Baltic Sea Action Plan	Nutrient load abatement strategy for the Baltic Sea	HELCOM (2013b)
BSAP, REF, BAU, WORST	Baltic Sea Action Plan, Reference, Business-As-Usual and WORST	Nutrient load scenarios	Meier et al. (2012a), Saraiva et al. (2019b)
CLC	Cyanobacteria Life Cycle Model	Advanced biogeochemical model including a cyanobacteria life cycle	Hense and Beckmann (2006, 2010)
CLIMSEA	Regionally downscaled climate projections for the Baltic and North Seas	Climate modelling project for the Baltic Sea	Meier et al. (2021)
CMIP	Coupled Model Intercomparison Project of the World Climate Research Programme	In this study, GCM/ESM results from CMIP3 and CMIP5 were assessed	https://www.wcrp-climate.org/wgcm-cmip
EC-EARTH	European Countries Earth System Model	ESM, CMIP5	https://www.knmi.nl/home
ECHAM5-MPI-OM	Max Planck Institute Global Climate Model	GCM, CMIP3	Roeckner et al. (2006), Jungclaus et al. (2006)
ECOSUPPORT	Advanced modelling tool for scenarios of the Baltic Sea ECOsystem to SUPPORT decision making	Climate modelling project for the Baltic Sea	Meier et al. (2014)
E-HYPE	Hydrological Predictions For The Environment applied for Europe	Process-based multi-basin model for the land surface	https://hypeweb.smhi.se/, Arheimer et al. (2012), Hundecha et al. (2016), Donnelly et al. (2013), Donnelly et al. (2017)

ERA-40	40-year reanalysis of the European Centre for Medium Range Weather Forecast	Reanalysis data used, e.g. as atmospheric forcing for ocean models	Uppala et al. (2005)
ESM	Earth System Model	Model applied for global climate simulations including the carbon cycle	Heavens et al. (2013)
EURO-CORDEX	Coordinated Downscaling Experiment: European Domain	High-resolution climate change projections for European impact research	Jacob et al. (2014), https://euro-cordex.net/
GCM	General Circulation Model	Model applied for global climate simulations	Meehl et al. (2004)
GHG	Greenhouse gas	Emission or concentration scenarios	Nakićenović et al. (2000), Mosset al. (2010), van Vuuren et al. (2011)
HadCM3	Hadley Centre Global Climate Model	GCM, CMIP3	Gordon et al. (2000)
HadGEM2-ES	Hadley Centre Global Environment Model version 2: Earth System	ESM, CMIP5	http://www.metoffice.gov.uk
HELCOM	Helsinki Commission	Consists of the Baltic Sea countries and the European Union	https://helcom.fi
IOW	Leibniz Institute for Baltic Sea Research Warnemünde	German research institute	http://io-warnemuende.de
IPCC	Intergovernmental Panel of Climate Change	Generated assessment reports (AR) of past and future changes in 1990, 1995, 2001, 2008, 2013 *inter alia* based upon CMIP results	http://www.ipcc.ch
IPSL-CM5A-MR	Institut Pierre Simon Laplace Climate Model: Medium Resolution	ESM, CMIP5	http://cmc.ipsl.fr/
MPI-ESM-LR	Max Planck Institute Earth System Model: Low Resolution	ESM, CMIP5	https://www.mpimet.mpg.de
NAO	North Atlantic Oscillation	Mode of climate variability	Hurrel (1995)
NOSCCA	North Sea Region Climate Change Assessment	Regional climate change assessment	Quante and Colijn (2016)
RCA3	Rossby Centre Atmosphere Model version 3	Regional climate model	Samuelsson et al. (2011)

RCA4-NEMO	Rossby Centre Atmosphere model Version 4: Nucleus for European Modelling of the Ocean	Coupled atmosphere-ocean model applied to the Baltic Sea and North Sea	Dieterich et al. (2013), Wang et al. (2015), Kupiainen et al. (2014), Madec (2016)
RCAO	Rossby Centre Atmosphere Ocean Model	Regional climate model	Döscher et al. (2002)
RCM	Regional Climate Model	Regional atmosphere or coupled atmosphere-ocean model applied to the dynamical downscaling of a changing climate	Giorgi (1990), Rummukainen (2010, 2016), Rummukainen et al. (2015), Feser et al. (2011), Rockel (2015), Schrum (2017)
RCO-SCOBI, BALTSEM, MOM-ERGOM	Model abbreviations	Coupled physical-biogeochemical models for the Baltic Sea	Meier et al. (2018a), their Tables 1 and 2 and references therein
RCP	Representative Concentration Pathway	Greenhouse gas concentration scenario	Moss et al. (2010), van Vuuren et al. (2011)
RCSM	Regional Climate System Model	Regional coupled atmosphere–sea ice–ocean–wave–land surface–atmospheric chemistry–marine ecosystem model	Giorgi and Gao (2018)
SRES	Special Report on Emission Scenarios	Described greenhouse gas emission scenarios, e.g. A1B, A2	Nakićenović et al. (2000)
STAT	Hydrological model	Statistical model for river runoff calculated from precipitation and evaporation over land	Meier et al. (2012a)

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1125   **Table 6.** Projected ensemble mean changes in total (land and atmosphere) bioavailable annual phosphorus (ΔP)
1126   and nitrogen (ΔN) inputs (in ktons) into the Baltic Sea between historical (1980–2005) and future (2072–2097)
1127   climates under the scenarios REF and BSAP. (Source: Meier et al., 2018a; their Fig. 3)

Nutrient input scenario	REF		BSAP	
Nutrient input changes	ΔP	ΔN	ΔP	ΔN
ECOSUPPORT BALTSEM	+2	−17	−15	−208
ECOSUPPORT MOM-ERGOM	+1	−15	−8	−180
ECOSUPPORT RCO-SCOBI	+4	+72	−11	−230
BalticAPP/CLIMSEA RCO-SCOBI (RCP4.5)	−18	−129	−34	−269

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1131 **Table 7.** Ensemble mean changes in sea surface temperature (SST; in °C) in the ECOSUPPORT, BalticAPP

1132 RCP4.5, BalticAPP RCP8.5, CLIMSEA RCP4.5 and CLIMSEA RCP8.5 scenario simulations averaged for the

1133 Baltic Sea including the Kattegat (data sources: Meier et al., 2011b; 2021; Saraiva et al., 2019a). (DJF = December,

1134 January, February, MAM = March, April, May, JJA = June, July, August, SON = September, October, November)

Δ SST	Winter (DJF)	Spring (MAM)	Summer (JJA)	Autumn (SON)	Annual mean
ECOSPPORT SRES A1B	+2.5	+2.8	+2.8	+2.5	+2.6
BalticAPP RCP4.5	+1.7	+1.9	+2.0	+1.8	+1.8
BalticAPP RCP8.5	+2.9	+3.2	+3.3	+3.0	+3.1
CLIMSEA RCP4.5	+1.7	+1.9	+2.0	+1.9	+1.9
CLIMSEA RCP8.5	+2.8	+3.0	+3.0	+2.9	+2.9

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1137   **Table 8.** Ensemble mean changes in annual mean sea surface salinity (SSS; in g kg^{-1}), annual mean bottom salinity

1138   (BS; in g kg^{-1}) and winter mean sea level (SL) relative to the global mean SL (in cm) in ECOSUPPORT, BalticAPP

1139   RCP4.5, BalticAPP RCP8.5, CLIMSEA RCP4.5 and CLIMSEA RCP8.5 scenario simulations averaged for the

1140   Baltic Sea including the Kattegat. For CLIMSEA, both the ensemble mean and the high SL scenarios are listed.

1141   In ECOSUPPORT and BalticAPP/CLIMSEA, the changes between 1978–2007 and 2069–2098 and between

1142   1976–2005 and 2069–2098 were calculated respectively. (Data sources: Meier et al., 2011b; 2021; Saraiva et al.,

1143   2019a)

Annual/winter changes	ECOSUPPORT A1B/A2	BalticAPP RCP4.5	BalticAPP RCP8.5	CLIMSEA RCP4.5 mean	CLIMSEA RCP4.5 high	CLIMSEA RCP8.5 mean	CLIMSEA RCP8.5 high
Δ SSS	−1.5	−0.7	−0.6	−0.3	+0.2	−0.2	+0.6
Δ BS	−1.6	−0.6	−0.6	−0.0	+0.6	−0.0	+1.1
Δ SL	+5.5	+0.4	+3.7	+0.2	+0.1	+3.4	+3.2

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1146 **Table 9**. Salinity changes averaged for the Baltic Sea in 1988–2007 relative to 1850 as a function of sea level rise
1147 (SLR). In the reference simulation, the mean salinity is 7.42 g kg^{-1}. In method 1, the increase in the water level
1148 was added to the first vertical grid box of the RCO-SCOBI model, while in method 2 the increase in the water
1149 level was evenly divided between the first and second grid boxes.

SLR (in m)	−0.24	+0.5	+1.0	+1.5 (method 1)	+1.5 (method 2)
Salinity (in g kg^{-1})	-0.35	+0.71	+1.41	+2.10	+2.03

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1151 **Table 10.** As in Table 8, but showing the ensemble mean changes in the summer mean bottom oxygen
1152 concentration (in mL L^{-1}) in ECOSUPPORT, BalticAPP RCP4.5, BalticAPP RCP8.5, CLIMSEA RCP4.5 and
1153 CLIMSEA RCP8.5 scenario simulations averaged for the Baltic Sea including the Kattegat. The projected changes
1154 depend on the nutrient input scenario: Baltic Sea Action Plan (BSAP), Reference (REF), Business-As-Usual
1155 (BAU) or Worst Case (WORST). (Data sources: Meier et al., 2011b; 2021; Saraiva et al., 2019a)

Summer changes	ECOSUPPORT A1B/A2	BalticAPP RCP4.5	BalticAPP RCP8.5	CLIMSEA RCP4.5 mean	CLIMSEA RCP4.5 high	CLIMSEA RCP8.5 mean	CLIMSEA RCP8.5 high
BSAP	−0.1	+0.6	+0.5	+0.6	+0.5	+0.4	+0.3
REF	−0.6	+0.1	−0.2	+0.0	−0.1	−0.2	−0.4
BAU	−1.1	-	-	-	-	-	-
WORST	-	−0.1	−0.5	-	-	-	-

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Table 11. As in Table 8, but showing the ensemble mean changes in the annual Secchi depth (in m) in the ECOSUPPORT, BalticAPP RCP4.5, BalticAPP RCP8.5, CLIMSEA RCP4.5 and CLIMSEA RCP8.5 scenario simulations averaged for the Baltic Sea including the Kattegat. The projected changes depend on the nutrient input scenario: BSAP, REF, BAU or WORST. (Data sources: Meier et al., 2011b; 2021; Saraiva et al., 2019a)

Annual changes	ECOSUPPORT A1B/A2	BalticAPP RCP4.5	BalticAPP RCP8.5	CLIMSEA RCP4.5 mean	CLIMSEA RCP4.5 high	CLIMSEA RCP8.5 mean	CLIMSEA RCP8.5 high
BSAP	−0.3	+0.6	+0.6	+0.6	+0.6	+0.6	+0.6
REF	−0.6	+0.2	0.1	+0.2	+0.2	+0.1	+0.1
BAU	−0.8	-	-	-	-	-	-
WORST	-	0.1	−0.1	-	-	-	-

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