

~~Evolution of the climate in the next million years: A r~~Reduced-complexity model for glacial cycles and impact of anthropogenic CO₂ emissions on future glacial cycles

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Abstract. We propose a reduced-complexity process-based model for the long-term evolution of the global ice volume, atmospheric CO₂ concentration and global mean temperature. The model only external forcings are the orbital forcing and anthropogenic CO₂ cumulative emissions. The model consists of a system of three coupled non-linear differential equations, representing physical mechanisms relevant for the evolution of the ~~C~~climate – Ice Sheets – Carbon cycle ~~S~~system in timescales longer than thousands of years. ~~The m~~Model parameters are calibrated using paleoclimate reconstructions and the results of two Earth system models of intermediate complexity. For a range of parameters values, the model is successful in reproducing the glacial-interglacial cycles of the last 800 kyr, in good agreement with the timing and amplitude of paleorecord fluctuations, with the best correlation between modelled and paleo global ice volume of 0.86. Using different model realisations, we produce ~~an~~a probabilistic assessment of forecast possible trajectories of the evolution of the Earth system over for the next 1 million years under natural and several fossil-fuel CO₂ release scenarios. In the natural scenario, the model assigns high probability of occurrence of long interglacials in the periods between present and 120 kyr after present, and between 400 kyr and 500 kyr after present. The next glacial inception is most likely to occur ~ 50 kyr after present with full glacial conditions developing ~ 90 kyr after present. The model shows that even already achieved cumulative CO₂ anthropogenic emissions (500 PgC) are capable of affecting the climate evolution for up to half million years, indicating that the beginning of the next glaciation is highly unlikely in the next 120 kyr. High cumulative anthropogenic CO₂ emissions (3000 PgC or higher), which could potentially be achieved in the next two to three centuries if humanity does not curb the usage of fossil-fuels, will most likely provoke Northern Hemisphere landmass ice-free conditions throughout the next half million years, postponing the natural occurrence of the next glacial inception to 600 kyr after present or later.

1 Introduction

The long-term evolution of the Earth system is not only of theoretical interest but a matter of utter practical relevance at present, when its assessment is required in the decision-making process for guaranteeing the safe permanent storage of radioactive waste. The safe disposal of such materials constitutes one of the urgent environmental and societal challenges that mankind faces in the 21st century. Disposal in deep geologically-stable rock formations is the only currently viable option to ensure long-term safety (IAEA Report 2007, Kim et al., 2011). While low- and intermediate-level nuclear wastes require at least 100 kyr years of safe storage, for high-level nuclear waste and spent nuclear fuel safety for 1 Myr should be envisaged. The

evaluation of geological disposal systems in response to future environmental changes, driven by a combination of natural and anthropogenic forcings is, therefore, mandatory.

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Numerous paleoclimate records show that during the last 2.7 Myr the natural evolution of the Earth system of the last 2.7 Myr is-has been characterized by persistent alternations of glacial and interglacial episodes. The last 800 kyr were dominated by strongly asymmetric glacial cycles with average periodicity of about 100 kyr, a rather slow glacial build-up and much more rapid deglaciations (Hays et al. 1976; Lisiecki and Raymo, 2005). Antarctic ice core records ~~also~~ show that atmospheric ~~Carbon~~carbon dioxide (CO₂) concentration fluctuated nearly synchronously with the global ice volume, and that CO₂ concentration during glacial times was up to 100 ppm lower than during preindustrial time (Petit et al., 1999, Lüthi et al., 2008).

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The astronomical theory of glacial cycles (Milankovitch, 1941) postulates that growth and shrinkage of ice sheets is primarily controlled by changes in Earth's orbital parameters (eccentricity, obliquity and precession) through their effect on the amount of solar radiation received at high latitudes of the Northern Hemisphere (NH) during boreal summer. Different aspects of Milankovitch theory have been tested and ~~confirmed~~supported with climate and Earth system models of different degree of complexity (e.g. Pollard, 1983; Berger et al., 1999; Ganopolski and Calov, 2011; Abe-Ouchi et al., 2013; Tabor and Poulsen, 2016; Willeit et al., 2019).

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For the future million years, however, not only natural processes are of relevance. Different types of models have been used also to study long term future evolution of the Earth system in response to anthropogenic influence, primarily, fossil fuel combustion. It It has been shown that anthropogenic CO₂ emissions can affect future climate variability ~~even~~ at very long timescales (Archer et al., 1997; Lenton and Britton, 2006; Ridgwell and Hargreaves, 2007), ~~and it has been demonstrated that~~ ~~e~~Even the current already achieved cumulative anthropogenic emission of the order of 500 GtC PgC, –when assuming no negative ~~Carbon~~carbon emission in the future, can postpone the onset of the next glaciation by at least 50 kyr while an emission of 5000 PgGtC could completely change climate variability over the next 0.5 Myr (Archer and Ganopolski, 2005; Ganopolski et al., 2016).

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The modelling of the climate evolution at 1Myr timescales is not “accessible” for complex Earth system models and results remains problematic even for most of the intermediate-complexity models due to prohibiting computational cost. Moreover, while Earth orbital parameters are accurately known for the next millions of years (Laskar et al., 2004), large uncertainties associated with the total amount of future fossil fuel combustion, requires a computationally efficient modelling approach which permits to perform numerous long-term simulations to asses seimpact of these uncertainties.

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Most of simple (conceptual) models of glacial cycles (e.g. Imbrie and Imbrie 1980; Paillard 1998; Tziperman 2006; Crucifix 2012) simulate only the response of global ice volume to orbital forcing changes and thus are not suitable for an assessment of long-term consequences of anthropogenic CO₂ emission. Some simple models do describe also CO₂ concentration (for example: Saltzman, 1987; Saltzman and Verbitsky, 1993) but these models are designed accounting only for description of the natural evolution of carbon cycle. Thus, to our knowledge, there is no computationally inexpensive modelling tool which can simulate the future evolution of the global ice volume and CO₂ concentration in response to anthropogenic perturbation.

To provide a set of possible scenarios of the future coupled climate – ice sheets – carbon cycle evolution we develop a reduced-complexity model based on paleoclimate data and results of physically-based Earth system models. This approach constitutes a significant further development of the method proposed in Archer and Ganopolski (2005) and Lord et al. (2019). In the former, Paillard’s conceptual model of glacial cycles was combined with the results of future CO₂ simulations performed with a simple marine carbon cycle. In the latter, Paillard’s model was combined with simulations of the future evolution of anthropogenic CO₂ anomaly modelled as in Lord et al. (2016) based on results of the Earth system model of intermediate complexity cGENIE (Lenton et al., 2006).

In this manuscript, a more advanced semi-empirical model is calibrated using paleoclimate records from the last 800 kyr and the results of the CLIMBER-2 EMIC (Petoukhov et al., 2000; Ganopolski et al., 2001), while the future evolution of anthropogenic CO₂ anomaly is modelled following Lord et al. (2019, 2016).

Different types of models have been used also to study long-term future evolution of the Earth system in response to anthropogenic influence, primarily, fossil fuel combustion. It has been shown that anthropogenic CO₂ emissions can affect future climate variability even at very long timescales (Archer et al., 1997; Lenton and Britton, 2006; Ridgwell and Hargreaves, 2007) and it has been demonstrated that even the current cumulative anthropogenic emission of 500 GtC, when assuming no negative Carbon emission in the future, can postpone the onset of the next glaciation by at least 50 kyr while an emission of 5000 GtC could completely change climate variability over the next 0.5 Myr (Archer and Ganopolski, 2005; Ganopolski et al., 2016).

Aforementioned results not only demonstrate the long-term impact of the human activity but are also of practical importance for another type of anthropogenic activity: the permanent storage of the radioactive waste. Indeed, low- and intermediate-level nuclear wastes must be safely stored for at least 100 kyr years, while for high-level nuclear wastes and spent nuclear fuel, safety for 1 Myr should be envisaged. The safe disposal of such materials constitutes one of the urgent environmental and societal challenges that mankind faces in the 21st century. Disposal in deep geologically stable rock formations is the only currently viable alternative to ensure long-term safety (IAEA Report 2007, Kim et al., 2011). The evaluation of geological disposal systems in response to future environmental changes, driven by a combination of natural (orbital variations) and

anthropogenic (fossil fuel emissions) forcings (e.g. Lord et al., 2016) is, therefore, mandatory. However, these timescales are not “accessible” by complex Earth system models and result problematic even for most of the intermediate complexity models. Moreover, while Earth orbital parameters are accurately known for the next millions of years (Laskar et al., 2004), large uncertainties associated with the total amount of fossil fuel combustion as well as the sensitivity of the Earth system to anthropogenic perturbations at multimillion years timescale, requires a computationally efficient modelling approach which permits the generation of quantitative probabilistic assessments. To this end we develop a reduced complexity process based model of the coupled climate—ice sheets—Carbon cycle evolution, whose only external forcings are insolation and cumulative anthropogenic CO₂ emissions. By formulation, this model is similar to the so-called “conceptual models of glacial cycles” (e.g. Imbrie and Imbrie 1980; Paillard 1998; Tziperman 2006; Crucifix 2012). Some of these conceptual models simulate only global ice volume evolution while others also simulate variations in natural (i.e. past) CO₂ concentration (e.g. Saltzman 1987; Saltzman and Verbitsky 1993; Paillard and Parrenin 2004). To account for the anthropogenic influence on the future glacial cycles, Archer and Ganopolski (2005) used Paillard’s conceptual model in conjunction with the results of future CO₂ simulations performed with a box model of marine Carbon cycle. Here, for the same purpose, we use a parameterization developed in Lord et al. (2016) based on simulations performed with the Earth system model of intermediate complexity eGENIE (Lenton et al., 2006).

We organize ~~The manuscript is organized~~ as follows. In Sec. 2 we introduce the model and datasets. In Sec. 3 Results are we present the model calibration in sections 3 and 4. ~~Section 3 is devoted to the assessment of model performance.~~ In Sec. 4 we present ~~a the results of the simulations for the next 1 Myr, probabilistic forecast for the future 1 Myr, under considering~~ scenarios with and without anthropogenic influence. In Sec. 5 we analyse the sensitivity of the results to several aspects of the model design and parameter selection strategy. Finally, in Sec. ~~6~~5 we summarize the results and present the conclusions.

2 Model and datasets

2.1 Modelling approach

~~To develop a probabilistic prediction of future climate w~~We develop a process-based model for the coupled evolution of global ice volume (v), atmospheric CO₂ concentration (CO₂) and global mean surface air temperature (T). For the past, the model’s only external forcing is orbital forcing ~~represented by the maximum summer insolation at 65°N.~~ For the future, we accounted additionally for the impact of the cumulative anthropogenic CO₂ emissions on atmospheric CO₂ concentration. The model consists of a system of three coupled, non-linear differential equations, describing Earth system evolution at the timescales from thousands to millions of years. While the model is designed for simulations of future glacial cycles, global temperature is a useful diagnostic which can become necessary in other potential applications. ~~Using different sets of model parameters, we generate a probabilistic forecast of the evolution of the system over the next 1 million years under natural and anthropogenic CO₂ release scenarios.~~

135 It has been shown (Crucifix, 2012) that very different conceptual models of glacial cycles, based on completely different
assumptions, can simulate past glacial cycles with similar and sufficiently high skill. Thus, paleodata alone are not sufficient
to select the “right” model. Our approach to the construction of such model is principally different from previous studies
because the model is designed based on the results of physically based Earth system climate models. In particular, based on
the CLIMBER-2 model simulations (1) we define orbital forcing as the maximum summer insolation at 65°N; (2) it is required
that for any orbital forcing there is at least one stable equilibria and for a certain range of orbital forcings two equilibria (Calov
and Ganopolski, 2005); (3) the transition from interglacial ($v=0$) to glacial state occurs when the insolation drops below a
certain threshold value which linearly depends on the logarithm of CO₂ concentration (Ganopolski et al., 2016). We only
selected those model realization which have this relationship consistent with the analysis presented in Ganopolski et al. (2016)
(see discussion below).

145 For model calibration, we use a hybrid approach, based on a combination of paleodata and CLIMBER-2 model simulations.
We use paleoclimate reconstructions of the late Quaternary (last 800 kyr; see below). This period was selected for two reasons:
(1) so far this is the only period of time for which accurate reconstructions of atmospheric CO₂ concentration are available; (2)
this period dominated by the long glacial cycles which are expected to continue in the future, at least, for some time in the
absence of significant anthropogenic influence (see discussion below). However, during that period the CO₂ atmospheric
concentration was most of time lower than preindustrial levels and, always below the present one. Therefore, those records
alone are insufficient to derive a model suitable for a modelling future climate evolution under anthropogenic influence. Given
the lack of accurate CO₂ reconstructions from periods with CO₂ atmospheric concentration higher than preindustrial level, we
calibrate the model behaviour in such a greenhouse world using results of the physically-based CLIMBER-2 model. The
selected model realizations are used to simulate the next million years under natural and several anthropogenic scenarios.
Future CO₂ concentration scenarios were computed using results of the studies performed with the cGENIE EMIC. Since there
are no available simulations of CO₂ concentration under the combined effect of anthropogenic CO₂ emissions and future ice
sheet growth (growth that can only happen after the CO₂ anomalies decrease considerably) we assume that natural and
anthropogenic CO₂ anomalies can be simply added up.

160 The model contains 12 parameters (see Table 1). Several of them were fixed using prior information, while others were treated
as “tunable” model parameters (see Sec. 2.5). For model calibration and testing, we use paleoclimate reconstructions for the
last 800 kyr (see below). This period was selected because it is dominated by the long glacial cycles which are expected to
continue in the future, at least, for some time (see discussion below) in the absence of significant anthropogenic influence.
Using empirical data and by varying model parameters, we generate a large subset of model realizations of which we select a
subset with the highest correlations to paleodata. This subset has been then additionally reduced by requiring consistency with
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the results of Ganopolski et al. (2016). This reduced subset of model realizations has been finally applied to modelling of the next million years for the natural (no anthropogenic CO₂ emission) and several anthropogenic scenarios.

2.2 Equation for global ice volume

170 The first model equation describes the temporal evolution of global ice volume v expressed in non-dimensional units (with a value of 1 ~~representative of corresponding to the reconstructed~~ Last Glacial Maximum (LGM, 21 kyr before present) ice volume ~~levels, according to empirical data~~). Since it is set that at present $v=0$, this variable in fact represents the anomaly of global ice volume relative to the preindustrial state. As the principal natural forcing for global ice volume is summer NH insolation, ~~it is natural to interpret it as the~~ this equation is only valid for ~~only~~ NH ice volume. To account for the contribution of the Southern Hemisphere (SH), we use the same approach as in Ganopolski and Calov (2011) and assume that the ice volume anomaly in the SH closely follows the NH anomaly and makes up a constant, relatively small fraction of the global ice volume anomaly. This approach, obviously, is not applicable for a possible future Antarctic and Greenland ice sheets melting under high CO₂ concentrations. This is why we do not consider future sea level rise above the preindustrial level and it is required that $v \geq 0$ at any time (see below).

180 Although the constraint of $v \geq 0$ for the future is a strong one, we expect the approach not to be a problem even under the most pessimistic anthropogenic scenario we consider in this manuscript (cumulative emission of 3000 PgC). On one hand, a scenario leading to the complete melting of the Greenland ice sheet would change the global ice volume by a small value compared to the glacial-interglacial variability and, thus, could be neglected. On the other hand, a complete Antarctic deglaciation would
185 definitely affect the global climate carbon cycle and could not be ignored. However, according to Garbe et al. (2020), a significant Antarctic mass loss occurs only if global temperature anomaly stays above 8°C for a very long time and, as a consequence, even under the 3000 PgC emission scenario the contribution of the Antarctic ice sheet to changes in global ice volume could be neglected, at least, to a first approximation (see below for the estimated temperature change under different anthropogenic scenarios).

190 The mass-conservation equation states that:

$$\frac{dv}{dt} = (\text{Accumulation} - \text{Ablation}) \quad (1)$$

195 “Accumulation” represents here the global ice accumulation (in mass per time units) and “Ablation” represents all mass losses from ice sheets including surface and basal melt, and icebergs calving. We assume that the total accumulation is proportional

to the total ice sheet area (Ganopolski et al., 2010). In turn, ice sheet area is closely related to the ice volume. It is often assumed that $A = v^\gamma$, where γ is about 0.8. Here for simplicity we assume $\gamma=1$, and thus:

$$200 \quad \text{Accumulation} = b v \quad (2)$$

where b is the proportionality constant.

On the other hand, the ablation is controlled by several processes of which the energy balance at the surface of the ice sheets is the most important one. Total ablation depends on the size of the ice sheets (and in the NH, especially, on the position of their southern margin), summer insolation (orbital forcing) and CO_2 atmospheric concentration (CH_4 and N_2O also play a role but we assume their radiative forcing to be proportional to the radiative forcing of CO_2 as in Willeit et al., 2019). The effect of CO_2 on the mass balance of ice sheet is introduced via the logarithm of its concentration because radiative forcing of CO_2 is proportional to the logarithm of concentration. As the metric for orbital forcing, ~~as in Ganopolski (2021)~~, we use the maximum summer insolation at 65°N computed using Laskar et al. (2004). ~~Similar to Ganopolski et al. (2021), To reproduce a rapid deglaciation process~~ we introduced ~~an additional term depending on proportional to dv/dt representing the “domino effect” during glacial termination.~~ To ensure existence of at least one equilibrium solution (solution with $dv/dt = 0$) for any combination of orbital and CO_2 forcing (Calov and Ganopolski, 2005; Abe-Ouchi et al. 2013) a negative term proportional to the ice volume in power large than one is required ~~(Calov and Ganopolski, 2005; Abe Ouchi et al. 2013).~~ ~~While Ganopolski (2021) used term v^2 , here~~ to this end we use the term $v^{3/2}$ since it gives a better model performance. Thus, the total ablation is represented as:

$$215 \quad \begin{aligned} \text{Ablation} &= b_{01}v - b_2v^{3/2} - b_3(f - \bar{f}) - b_4\log(\text{CO}_2) + b_5\frac{dv}{dt}M_v - b_{06} \\ \text{Ablation} &= b_{01}v + b_2v^{3/2} + b_3(f - \bar{f}) + b_4\log(\text{CO}_2) + b_5\frac{dv}{dt}M_v \end{aligned} \quad (3)$$

220 where $b_{01}, b_2, \dots, b_{06}$ are constants, \bar{f} is the average orbital forcing, and M_v is a memory term that reflects the history of the ice volume during the last τ yr, as long as dv/dt is negative. At any time t , $M_v(t)$ is defined as:

$$M_v(t) = \delta \frac{\int_{t-\tau}^t v(x) dx}{\tau} \quad (4)$$

$$225 \quad \delta = 1 \quad \text{if} \quad \frac{dv}{dt} < 0; \quad \text{and} \quad \delta = 0 \quad \text{otherwise} \quad (5)$$

Finally, ~~setting $b_1=b_{01}$ and $b_6=b_{06}/(1+b_5M_v)$~~ the mass-balance Eq. (1) ~~can be is~~ re-written as:

$$\frac{dv}{dt} = \frac{b_1 v - b_2 v^{3/2} - b_3 (f - \bar{f}) - b_4 \log(CO_2)}{1 - b_5 M_v} + b_6 \quad (6)$$

Where b_1, b_2, b_3, b_4, b_5 and b_6 are additional model parameters.

235 2.3 Equation for atmospheric CO₂ concentration

It is generally recognized that CO₂ (together with other well-mixed greenhouse gases such as CH₄ and N₂O) represents an important amplifier of glacial cycles and that the anthropogenic CO₂ emission will affect the climate during long periods of time (Archer and Brovkin, 2008) and, in particular, the timing and magnitude of future glacial cycles (Archer and Ganopolski, 2005). Although the mechanisms of glacial-interglacial atmospheric CO₂ variability are still not fully understood, significant progress in modelling the global ~~Carbon~~carbon cycle operation during glacial cycles has been achieved in recent times (Brovkin et al., 2012; Menviel et al., 2012; Willeit et al., 2019). Since CO₂ and ice volume are highly correlated (during the past 800 kyr the coefficient of correlation between atmospheric CO₂ concentration and ice volume is -0.71), it is not surprising that simple conceptual models of glacial cycles can describe the ice volume evolution without an explicit treatment of CO₂ (e.g. Paillard 1998; Crucifix 2013; Ganopolski 2021). However, this close relationship between CO₂ and global ice volume is not valid for the Anthropocene and, thus, the modelling of the future climate evolution requires an explicit treatment of atmospheric CO₂ concentration which is the second equation of our model. The response of the global ~~Carbon~~carbon cycle to external perturbations involves a broad range of timescales: from very short (annual) to geological. For the purpose of this study we treat natural and anthropogenic anomalies of CO₂ separately but consider them additive. Namely, we assume that on the relevant timescales (10³-10⁵ yrs), the natural component of CO₂ concentration is in equilibrium with external conditions and ~~can be~~is expressed through a linear combination of global temperature and global ice volume.

The most direct effect of temperature on CO₂ is through changes in the CO₂ solubility in ocean water (Zeebe and Wolf-Gladrow, 2001). Temperature, directly and through sea ice, also affects ocean circulation (Watson et al., 2015), ventilation rate of the deep water (Kobayashi et al., 2015), changes in relative volume of different water masses (Brovkin et al., 2007) and metabolic rates of living organisms (Eppley, 1972; Laws et al., 2000).

The direct effect of ice volume on CO₂ changes is less straightforward (~~please~~note that the strong effect of ice sheets on atmospheric temperature is already accounted for in the temperature term). An increase in ice volume leads to a decrease of the ocean volume and, as a consequence, increased ocean salinity. As CO₂ solubility in the ocean decreases with salinity this would provoke to an increase of atmospheric CO₂ concentration, counteracting the temperature effect (Zeebe and Wolf-

Gladrow, 2001). Higher global ice volume would also mean a smaller area of the globe is covered by forests and, therefore, a diminished carbon storage in the terrestrial reservoir, driving an increase in atmospheric CO₂ (Prentice et al., 2011). Increased glacial supply of iron-rich dust may help suppress the iron limitation in some areas (in particular in the south Atlantic Ocean) leading to increased biological productivity and thus lower atmospheric CO₂ levels (Martin, 1990; Watson et al., 2000).

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In addition, based on results of Ganopolski and Brovkin (2017) we parameterized CO₂ “overshoots” during glacial terminations by an additional term proportional to dv/dt . Note that this additional term is only applied when $dv/dt < 0$. The justification for such parameterization is the results seen in many models (Gottschalk et al., 2019) that a shutdown of the Atlantic Meridional Overturning Circulation (AMOC) causes an atmospheric CO₂ rise of about 10-20 ppm. A shutdown of the AMOC is caused by a large meltwater flux which is, in turn, controlled by changes in ice volume (i.e. dv/dt).

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To describe the anthropogenic component of CO₂, we assume that anthropogenic CO₂ emission is a relatively short (order of 10² yrs) pulse followed by zero emissions. In this case, the temporal dynamics of the anthropogenic component of CO₂ anomaly depends only on the cumulative ~~Carbon~~carbon emission. To describe the long anthropogenic tail of CO₂ (*Anth*-CO₂ term in the Eq. (7) below) we use the analytical parameterization defined in Lord et al. (2016) based on the results of experiments with the Earth system model of intermediate complexity cGENIE (Lenton et al., 2006). In addition, we assume that natural and anthropogenic CO₂ anomalies can be simply summed up and that at the preindustrial time the global ~~Carbon~~carbon cycle was in equilibrium. This is, obviously, a very strong assumption since even a rather small imbalance in the global ~~Carbon~~carbon cycle which is impossible to detect at the millennial timescales can result in a very large “drift” of the Earth system from its preindustrial state at the million years timescale. However, due to the absence of any practical alternative, we proceeded with this assumption.

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After these considerations, the equation that governs the CO₂ evolution has the following shape:

$$285 \quad CO_2 = c_1 dT + c_2 v + c_3 \min\left(\frac{dv}{dt}, 0\right) + c_4 + Anth_{CO_2} \quad (7)$$

Where $c_1 \dots c_4$ are adjustable model parameters. The first three terms on the right-hand-side of equation (7) represent the already discussed effects of global temperature (T), ice volume and AMOC weakening or shutdown. The fourth term (c_4) is simply a constant and the last term represents ~~the effect of anthropogenic CO₂ emissions. external forcing due to anthropogenic processes (and we are, thus, assuming that anthropogenic emissions occur on top of the natural Carbon cycle as a linear addition). Please, refer~~ to Sec. 4.1 for a detailed explanation on the assumptions and shape of the function $Anth_{CO_2}(t)$.

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2.4 Equation for global mean surface temperature

The last equation in our model describes global mean surface air temperature anomaly (ΔT ; relative to preindustrial values) as a linear combination of two terms: first, the direct effect function of global ice volume (the larger the ice volume, the higher the planetary albedo and the lower the global temperature) and, second, a term representing the radiative forcing of CO_2 , which is proportional to the logarithm of CO_2 concentration:

$$\Delta T = d_1 v + d_2 \log\left(\frac{\text{CO}_2}{278}\right), \quad (8)$$

where d_1 and d_2 are adjustable model parameters.

2.5 Model constraints, bounds and parameter selection strategy

The model of past and future glacial cycles is represented by the system of three equations (6), (7), (8) which contains three variables (v , CO_2 and ΔT), orbital forcing $f(t)$, anthropogenic CO_2 anomaly $\text{Anth}_{\text{CO}_2}(t)$ and twelve parameters (b_i $i=1:6$; c_j $j=1:4$, d_k $k=1:2$). The orbital forcing $f(t)$ depends only on is determined by astronomical parameters (eccentricity, precession and obliquity) which are accurately known for the entire period of interest (Laskar et al., 2004). The function $\text{Anth}_{\text{CO}_2}(t)$ is assumed to be null until $t = 0$ kyr, afterwards the function depends only on time and cumulative CO_2 emission as detailed in Sec. 4.1.

Model parameters are selected-calibrated to yield solutions that are in a good (see discussion below) agreement both with the paleoclimatic information over the last 800 kyr (see Sec. 2.6 for details on the paleoclimate reconstructions used) and CLIMBER-2 results. To determine some of model parameters and thus reduce the number of “free parameters” we use three conditions based on paleoclimatic data and current estimations of equilibrium climatic sensitivity.

We impose on the model a series of constraints based on paleorecords from the last 800 kyr: 1. Reproduction of present-day interglacial state; 2. Reproduction of LGM conditions; 3. Mid-Brühnes transition (MBT); 4. Compliance with CO_2 minimum level.

The first constraint, based on the observational fact that the present state is an interglacial one, The first constraint is related to preindustrial conditions. We require that:

$$\text{-if } v=0 \text{ and } \Delta T = 0, \text{ then } \text{CO}_2 = 278 \text{ ppm.} \quad (9)$$

This condition fixes the value for c_4 at 278.

As the confidence on empirical data for glacial conditions is higher at LGM than at any other glacial episode, the accurate reproduction of LGM conditions is ~~the only constraint that we will require from the model~~ important. At LGM, the empirical ice volume in non-dimensional units was 1, CO₂ concentration 194 ppm and global cooling around 5°C (Schneider et al., 2006; Annan et al., 2013; Tierney et al., 2020). We then require from the model that:

$$\text{if } v = 1 \ \& \ CO_2 = 194 \text{ ppm} \rightarrow \Delta T = -5^\circ C \quad (109)$$

It is known that glacial cycles prior to the mid-Brünhes Transition (MBT) about 400 kyr ago differ in some respects from those after MBT. In particular, pre-MBT interglacials were characterised by lower CO₂ and higher benthic d¹⁸O values than the post-MBT episodes. The later implies a large interglacial, probably due to remaining continental ice sheets in the NH. The mechanisms accounting for these differences are still not understood (Tzedakis et al., 2009; Berger et al., 2016) but they are unlikely related to differences in orbital forcing before and after MBT. This is why it is not expected that a simple model forced by orbital variations alone can simulate such behaviour. Thus, we prescribe that before the MBT the minimum ice volume must be 0.05 in normalized units:

$$v(t) = \max(v(t), 0) \text{ for all } t \quad (11)$$

$$v(t) = \max(v(t), 0.05) \text{ if } t < -400\text{kyr} \quad (12)$$

The last condition based on empirical data prevents CO₂ to drop to levels significantly lower than the minimum atmospheric CO₂ concentration registered in the last 800 kyr by paleorecords (172 ppm) at any given moment in time:

$$CO_2(t) = \max(CO_2(t), 150) \quad (13)$$

We also ~~The last imposed~~ impose a criteria condition ~~is~~ related to the current estimates for the equilibrium climatic sensitivity (ECS), i.e. global temperature response to a doubling in atmospheric CO₂ concentration from preindustrial levels. We select ~~a~~ equilibrium climate sensitivity (ECS) equal to 3.9°C, which coincides with the multi-model mean in the Coupled Model Intercomparison Project 6 (CMIP6; Zelinka et al., 2020):

$$\text{if } CO_2 = 2 \times 278 \text{ ppm} \ \& \ v = 0 \rightarrow \Delta T = 3.9^\circ C. \quad (140)$$

Equations (10) and (141) determine the values for the coefficients d_1 and d_2 at -3 and 5.56, respectively.

355 Recent studies (e.g. Nijsee et al., 2020) suggest that some of CMIP6 models have unrealistically high ECS and the IPCC AR6
report proposed a best estimate for ECS of 3°C. However, on the long time scales considered in this paper, the classical ECS
which accounts only for fast climate feedback, but not for the feedbacks related to vegetation, ice sheets and atmospheric
composition (apart from CO₂) likely underestimates the Earth system response to radiative perturbation. However, Nonetheless,
360 to assess uncertainties in future climate simulations associated with uncertainties in ECS, we performed an additional set of
simulation with ECS=3°C (see Sec. 5).

~~With these additional constraints, the initial set of twelve independent parameters shrinks to a total of nine: $P = \{b_i, c_j\}$ $i = 1-6, j = 1-3$.~~

365 ~~Furthermore, to prevent the selection of parameters that will lead to unrealistic values for the selected variables we also apply two limitations:~~

~~The first limitation is applied to the simulated ice volume v . First, we assume that for the recent interstadials and any future time, v cannot be negative. Second, it is known that glacial cycles prior to the mid-Brünhes Transition (MBT) about 400 kyr ago differ in some respects from those after MBT. In particular, pre-MBT interglacials were characterised by lower CO₂ and higher benthic d¹⁸O values than the post-MBT episodes. The latter implies a large interglacial, probably due to remaining continental ice sheets in the NH. The mechanisms accounting for these differences are still not understood (Tzedakis et al., 2009; Berger et al., 2016) but they are unlikely related to differences in orbital forcing before and after MBT. This is why it is not expected that a simple model forced by orbital variations alone can simulate such behaviour. This is why we prescribe that~~
375 ~~before the MBT the minimum ice volume must be 0.05 in normalized units:~~

~~$$v(t) = \max(v(t), 0.05) \text{ if } t < 400 \text{ kyr} \quad (11)$$~~

~~$$v(t) = \max(v(t), 0) \text{ if } t \geq 400 \text{ kyr} \quad (12)$$~~

380 ~~The second limitation, based on empirical data, prevents CO₂ to drop to levels significantly lower than the minimum atmospheric CO₂ concentration registered in the last 800 kyr by paleorecords (172 ppm) at any given moment in time:~~

~~$$CO_2(t) = \max(CO_2(t), 150) \quad (13)$$~~

385 In addition, we ~~applied a constraint which is not from the past but from the future. S~~account for several modelling studies (Berger and Loutre, 2002; Cochelin et al., 2006; Ganopolski et al., 2016) indicating that the conditions for the new glacial inception will not be met in the near future even in the absence of anthropogenic influence on climate. To assure that our model

satisfies this requirement, we require for all valid model versions the average global ice volume not to exceed 0.025 (in normalized units) in the time period between 0 and 20 kyr:

$$\frac{\sum_{t=0}^{20 \text{ kyr}} v(t)}{21} < 0.025 \quad (154)$$

For the Base experimental setting (other settings will be discussed in Section 5), the model is defined by equations 6-8 together with the conditions expressed in equations 9-15. A table of acronyms (Table 1) is provided as a quick reference.

Table 1: Table of acronyms

	Acronym	Units
<u>Global ice volume</u>	<u>v</u>	<u>Non-dimensional (varies between 0 and 1)</u>
<u>Atmospheric CO₂ concentration</u>	<u>CO₂</u>	<u>ppm</u>
<u>Global surface temperature anomaly (with respect to preindustrial levels)</u>	<u>T</u>	<u>°C</u>
<u>Orbital forcing: maximum summer insolation at 65°N</u>	<u>f</u>	<u>W m⁻²</u>
<u>Time</u>	<u>t</u>	<u>kyr</u>
<u>Memory term (see Equation (4))</u>	<u>M_v</u>	<u>Non-dimensional</u>
<u>Anthropogenic CO₂ concentration</u>	<u>Anth_{CO2}</u>	<u>ppm</u>
<u>Last Glacial Maximum</u>	<u>LGM</u>	<u>-</u>
<u>Mid-Brühnes transition</u>	<u>MBT</u>	<u>-</u>
<u>Cumulative anthropogenic CO₂ emissions</u>	<u>E</u>	<u>ppm</u>
<u>Equilibrium climate sensitivity (see Equation (14))</u>	<u>ECS</u>	<u>°C</u>
<u>Critical insolation – CO₂ relationship parameters (see Equation (18))</u>	<u>K, R</u>	<u>W m⁻²</u>
<u>Set of model parameters</u>	<u>P: b₁,...,b₆, c₁,..., c₄, d₁, d₂</u>	<u>see Table 2</u>
<u>Parameters for evolution of Anth_{CO2}, derived in Lord et al. (2016)</u>	<u>α_i, β_{ij}, γ_i, δ_{ij} i=1,...,5; j=1,...,3</u>	<u>see Lord et al. (2016)</u>
<u>Pearson correlation between modelled and paleo variable x (v, CO₂ or T) considering the time period [t_i,t_f]</u>	<u>C_x(P,t_i,t_f)</u>	<u>-</u>
<u>Root Mean Squared Error</u>	<u>RMSE</u>	<u>-</u>
<u>Set of solutions consistent with paleodata</u>	<u>Paleovalid</u>	<u>-</u>
<u>Subset of Paleovalid consistent with CLIMBER-2 results</u>	<u>Accepted</u>	<u>-</u>

Finally, we approach the task of the selection of set of parameters P as a non linear optimisation problem with equality and inequality constraints. We wish to find P to maximize the optimization target function C_p :

$$C_p(P, t_i, t_f) = \text{corr}(v_{\text{model}(P)}(t_i, t_f), v_{\text{paleo}}(t_i, t_f)) \quad (15)$$

where: $v_{\text{model}(P)}(t_i, t_f)$ denotes the modelled ice volume time series in the period $[t_i, t_f]$ (i.e. solution obtained for v when using the parameter set P in the system (6-8)); $v_{\text{paleo}}(t_i, t_f)$ denotes the paleo ice volume time series in the period $[t_i, t_f]$; $\text{corr}(x, y)$ denotes the linear Person correlation between x and y . The time interval for the optimisation is set to $[t_i, t_f] = [-800 \text{ kyr}, 0 \text{ kyr}]$. See Appendix A for a discussion of the dependence of model performance on the choice of this time interval. To select parameters that will optimise correlation at the same time as providing magnitudes in accordance to empirical estimations, an equality constraint is enforced: the maximum ice volume must be equal to 1 within a tolerance of 0.15 (in nondimensional units). Finally, the inequality constraint is given by Eq. (14).

The time step for the calculation of the time series is 1 kyr. We use the interior point optimisation algorithm under the Matlab environment. Approximately 1000 optimisation routines were performed, each starting from a different combination of the set of nine adjustable parameters (optimisation seeds).

Finally, the ensemble of valid solutions, *Valid*, is defined as the set:

$$\text{Valid} = \{P / C_p(P, -800 \text{ kyr}, 0 \text{ kyr}) \geq 0.7\} \quad (16)$$

The correlation level of 0.7, although arbitrary, guarantees a good fit to the paleo climatic ice volume record. For the selection of solutions, no conditions are imposed on the goodness of fit between modelled and paleo CO_2 or temperature. For each P , we calculate v , CO_2 and ΔT obtained using P in the system of equations (6), (7), (8) and denote them: $v_{\text{model}(P)}$, $\text{CO}_{2\text{model}(P)}$, $dT_{\text{model}(P)}$. This ensemble provides the basis to generate the probabilistic forecast in the remainder of the paper.

2.6 Empirical datasets

Paleo reconstructions covering the period $[-800 \text{ kyr}, 0 \text{ kyr}]$ are used as part of the learning and/or validation sets for the model. Reconstructed global ice volume is derived from sea level stack (Spratt and Lisiecki, 2016). Past CO_2 atmospheric concentration levels are derived from ice cores records (Lüthi et al., 2008). For global mean surface temperature anomalies (with respect to preindustrial conditions) we use two reconstructions: 1. Friedrich et al. (2016), which covers the period $[-784$

430 kyr, -10 kyr] and 2. Snyder (2016), covering [-800 kyr, -1 kyr]. All datasets are transformed into time-series with a 1 kyr time-step.

3 Model performance calibration

435 In this section we describe the choice of model parameters using paleoclimate data and CLIMBER-2 modelling results. First, we select a range of model parameters which enable us to simulate the past evolution of global ice volume, CO₂ and global temperature with a reasonable accuracy (see Sec. 3.2). We call this set of model realisation *Paleovalid*. We then apply an additional constraint based on the critical insolation – CO₂ relationship for glacial inceptions, which is of fundamental importance for the prediction of future glacial cycles in the response to anthropogenic CO₂ emissions. Since we find that
440 paleodata do not provide additional constraints on this relationship (see Sec. 3.3), we consider the estimation of the critical insolation – CO₂ curve derived from CLIMBER-2 in Ganopolski et al. (2016) and take forward only those solutions in line with it. All model realisations which satisfy this criterion are named *Accepted* and we use them for future scenarios simulations.

3.1 Empirical datasets

445 Paleo reconstructions covering the period [-800 kyr, 0 kyr] are used as part of the learning and/or validation sets for the model. Reconstructed global ice volume is derived from sea level stack (Spratt and Lisiecki, 2016). Past CO₂ atmospheric concentration levels are derived from ice cores records (Lüthi et al., 2008). For global mean surface temperature anomalies (with respect to preindustrial conditions) we use two reconstructions: 1. Friedrich et al. (2016), which covers the period [-784 kyr, -10 kyr] and 2. Snyder (2016), covering [-800 kyr, -1 kyr]. All datasets are transformed into time-series with a 1 kyr time-
450 step.

Obviously, the paleoclimate reconstructions are not perfect, and have associated uncertainties. While information for global ice volume and CO₂ atmospheric concentration are based directly on paleoclimate records and are considered quite accurate, global temperature reconstructions are based on a limited number of local temperature records, mostly from ocean sites, and
455 have large uncertainties (as easily noticeable by the two paleorecords in Fig 1c).

3.13.2 Comparison with Calibration using -paleoclimate reconstructions

For each set of parameters P , we calculate v , CO_2 and T using the system of equations 6-8 and denote them:

460 $v_{model(P)}, CO_{2model(P)}, dT_{model(P)}$

We approach the task of the selection of set of parameters P as a non-linear optimisation problem with constraints. We wish to find P to maximize the optimization target function C_v :

465 $C_v(P, t_i, t_f) = corr(v_{model(P)}(t_i, t_f), v_{paleo}(t_i, t_f))$ _____ (16)

where: $v_{model(P)}(t_i, t_f)$ denotes the modelled ice volume time-series in the period $[t_i, t_f]$; $v_{paleo}(t_i, t_f)$ denotes the paleo ice volume time-series in the period $[t_i, t_f]$; $corr(x, y)$ denotes the linear Pearson correlation between x and y . The time interval for the optimisation is set to $[t_i, t_f] = [-800 \text{ kyr}, 0 \text{ kyr}]$. See Appendix A for a discussion of the dependence of model performance on the choice of this time interval. To select parameters that will optimise correlation at the same time as providing magnitudes in accordance to empirical estimations, an equality constraint is enforced: the maximum ice volume must be equal to 1 within a tolerance of 0.15 (in nondimensional units). Eq. (14) is enforced as inequality constraint.

470

The time-step for the calculation of the time-series is 1 kyr. We use the interior-point optimisation algorithm under the Matlab environment. Approximately 1000 optimisation routines were performed, each starting from a different combination of the set of nine adjustable parameters (optimisation seeds).

475

Finally, the ensemble of valid solutions, *Paleovalid*, is defined as:

480 $Paleovalid = \{P / C_v(P, -800 \text{ kyr}, 0 \text{ kyr}) \geq 0.7\}$ _____ (17)

The correlation level of 0.7, although arbitrary, guarantees a good fit to the paleo climatic ice volume record. For the selection of solutions, no conditions are imposed on the goodness of fit between modelled and paleo CO_2 or temperature.

485 Naturally, there are many possible choices for model calibration procedures. In particular, it is possible to select a different optimisation target function. We selected to optimise the correlation between modelled and paleo ice volume because our main objective is to simulate future glacials and, thus, the reproduction of these cycles is extremely relevant. Livadiotis and McComas (2013) present the maximization of the correlation fitting method. Those authors show that the method is mathematically well defined under certain conditions and that it should be preferred over the classical least squares fitting in situations in which the data sets exhibit variations that need to be described, such as is the case for glacial cycles. The sensitivity of our results to a different optimisation target is evaluated in Sec. 5. While we opt for a frequentist perspective (estimates of

490

unknown parameters are obtained without assigning them probabilities) parameter estimation can also be approached through a Bayesian point of view (e.g. Crucifix and Rougier, 2009).

495 Figure 1 displays the ensemble probabilistic distribution of solutions in ~~Valid-Paleovalid~~ for v , CO_2 and ΔT as well as the corresponding paleoclimatic ~~analogues data~~ for the period [-800 kyr, 0 kyr]. The ranges of the different model parameters across the ~~PaleovValid~~ set are displayed in Table ~~24~~.

500 By definition, all the solutions derived from the ~~PaleovValid~~ ensemble closely follow the paleo ice volume curve (Fig. 1a), with the ensemble-mean correlation between modelled and paleo ice volume in [-800 kyr, 0 kyr] equal to 0.76. While most of the solutions derived from ~~PaleovValid~~ succeed in capturing the timing and magnitude of the major glaciations, there is a tendency for them to overestimate the ice volume in MIS 18 and MIS 14 (-560 kyr and -520 kyr, respectively)-. As model parameters were chosen to maximise ice volume correlation only, it is expected that the performance in terms of CO_2 and global temperature is inferior than for global ice volume. The ensemble-mean correlation between modelled and paleo CO_2 is 0.5 and ~~some approximately 50% of the~~ solutions display an amplitude range significantly larger than the observed one, reaching the imposed lower limit of 150 ppm (Fig. 1b). ~~Note however the better magnitude agreement when comparing simulated CO_2 with equivalent CO_2 concentration (as computed in Ganopolski et al., 2010) and which accounts for the additional radiative effect of N_2O and CH_4 . The imperfect CO_2 simulations represents however no major pitfall for a skill full ice volume reproduction, as shown by the 0.76 ensemble-mean correlation with paleodata. Interestingly, when using accurately~~

510 ~~known CO_2 time-series from ice cores, this ensemble-mean correlation for global ice volume only is slightly increased to 0.85.~~

For global mean surface temperature (Fig. 1c), the ensemble-mean correlation between modelled and paleo record is 0.56 and the solutions amplitude range within the paleo estimation limits.

515 **Table ~~24~~: Parameter ranges across solutions in the ~~PaleovValid~~ and *Accepted* sets as well as for the Best Solution.**

M model Parameter	Units	PaleovValid set	<i>Accepted</i> set	Best Solution
b₁	s^{-1}	0.075 – 0.27	0.12 – 0.26	0.22
b₂	$\text{m}^{-1/2} \text{s}^{-1}$	-0.49 – -0.15	-0.33 – -0.15	-0.29
b₃	$\text{m s}^{-1} \text{W}^{-1}$	-9×10^{-4} – -3×10^{-4}	-9×10^{-4} – -5×10^{-4}	-8×10^{-4}
b₄	$\text{m}^3 \text{s}^{-1}$	-0.62 – -0.02	-0.12 – -0.02	-0.095
b₅	<u>Non-dimensional</u>	-1 – -0.04	-0.76 – -0.13	-0.18
b₆	$\text{m}^3 \text{s}^{-1}$	0.1 – 3.49	0.1 – 0.66	0.53
c₁	ppm C^{-1}	10.6 – 18.84	12.37 – 18.84	17.28

c₂	<u>ppm m⁻³</u>	-35.1 – -20.0	-35.1 – -22.87	-31.95
c₃	<u>ppm m⁻³ s</u>	-120.1 – -119.9	-120.03 – -119.9	-120.0
c₄ (derived from constraints)	<u>ppm</u>	278	278	278
d₁ (derived from constraints)	<u>C m⁻³</u>	-3	-3	-3
d₂ (derived from constraints)	<u>C</u>	5.56	5.56	5.56

Table 32: Summary of ~~Valid~~Paleovalid and Accepted ensembles characteristics.

	Valid <u>Paleovalid</u>	Accepted
Correlation between modelled and paleo ice volume	≥0.7	≥0.7
K (W m⁻²)	Unrestricted	≥ -150
Number of elements	353	29

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While the model demonstrates a good performance when tested against the same data that was used for its training, its ability to produce predictions (situation in which previously unseen data is used as model input) can only be evaluated by the use of independent training and validation samples. A robust estimation of the model's predictive skill, using disjoint training and validation sets, ~~can be is~~ found in Appendix A. In general, we conclude that the model has a satisfactory ability also when used in predictive mode and, thus, we confidently venture to utilize it as a tool for the ~~forecast~~assessment of possible future scenarios. of the next 1 Myr climatic evolution.

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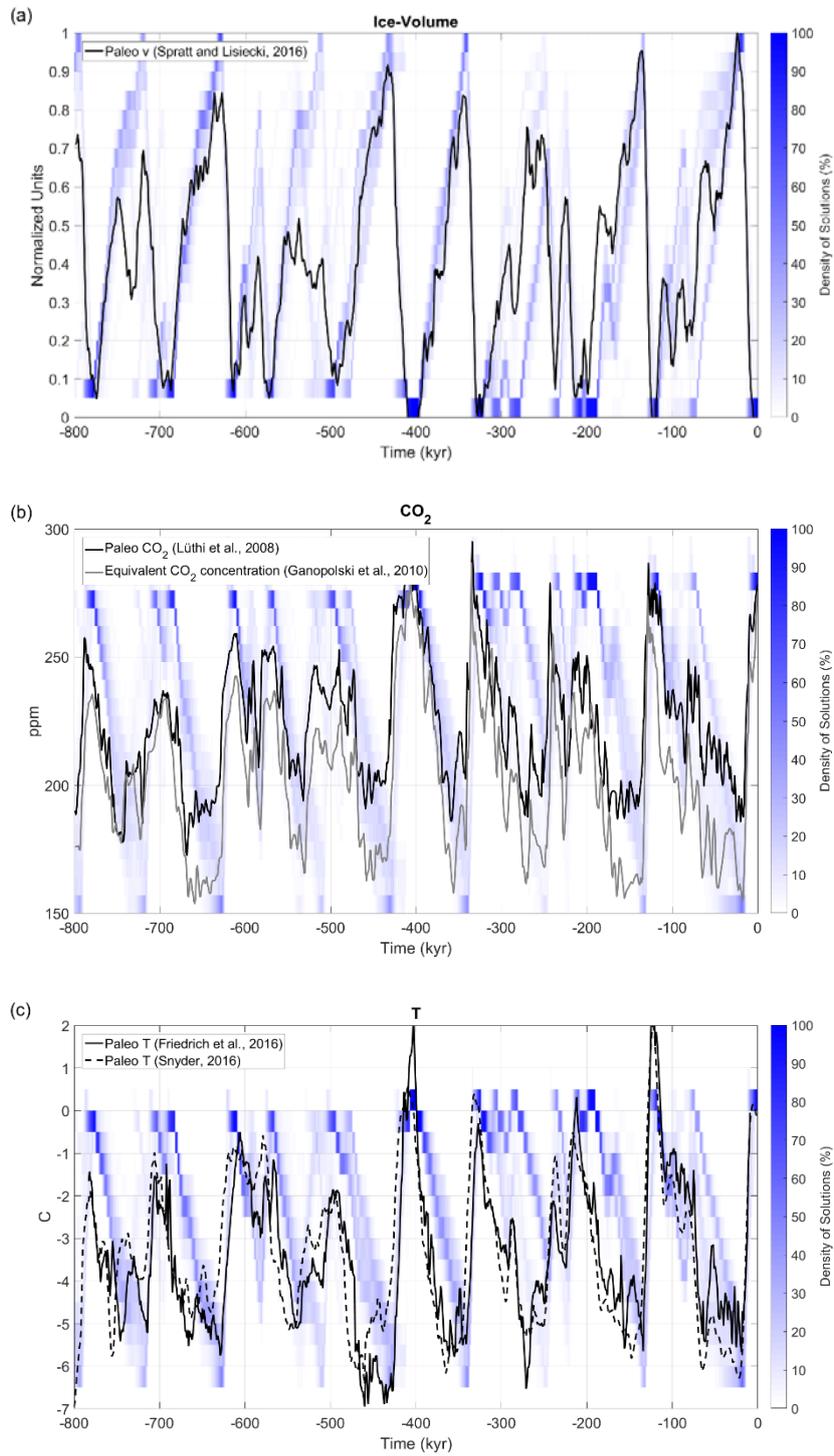


Figure 1: ~~Probabilistic distribution of *PaleoValid* solutions and paleoclimatic records in the period [-800 kyr, 0 kyr]:~~ a. Global landmass ice volume (normalized to unity); reconstruction (black line) from Spratt and Lisiecki (2016). b. Atmospheric CO₂ concentration (ppm); paleorecord from Lüthi et al. (2016) and ~~effective-equivalent CO₂ concentration from CLIMBER-2 simulation~~ Ganopolski et al. (2010) (grey line). c. Global mean surface temperature anomaly (C); paleorecords from Friedrich et al. (2016) (solid black) and Snyder (2016) (dashed). ~~For model simulations~~ Blue shaded areas represents ~~the probability density of the *Paleovalid* solutions~~ per unit of phase space- ~~(Units are: 1kyr for time, 0.05 for ice, 7ppm for CO₂ to 7 ppm and 0.5°C for temperature).~~

3.23.3 Glacial inception: Critical insolation – CO₂ relationship

Using an ensemble of CLIMBER-2 model realizations consistent with paleo-climatic constraints, Ganopolski et al. (2016) found that the critical threshold for summer insolation at 65° N to trigger glacial inception can be described by the following curve:

$$f_{critical} = K \log(CO_2/280) + R \quad (187)$$

where $K = -77 \text{ W m}^{-2}$ and $R = 466 \text{ W m}^{-2}$.

The parameter K is a measure of the sensitivity of the critical orbital forcing to atmospheric CO₂ concentration. The higher K is, the more important CO₂ is, i.e. for the same CO₂ emission scenario, the effect of anthropogenic perturbation on glacial cycles will last longer.

Following Ganopolski et al. (2016) approach, we calculate the corresponding coefficient K for each of the ensemble members in the *PaleoValid* set (Fig. 2). For the estimation of K we refer to Eq. (6), assuming equilibrium and interglacial conditions (i.e. $dv/dt=0$, $v=0$, $M_v=0$) it is derived that the critical insolation threshold:

$$f_{cr} = -\frac{b_4}{b_3} \log(CO_2) - \frac{b_6}{b_3} + \bar{f} \quad (19)$$

It follows then that K is estimated as $-b_4/b_3$.

Our results indicate that with the model derived in this study the possible values of the coefficient K range between -1279 and -31 W m^{-2} , with a median of -393 W m^{-2} . This highlights that, even though all the solutions derived from the *Valid-Paleovalid* set have a high level of agreement with the paleoclimatic records along last 800 kyr, the relationship between $f_{critical}$ and $\log(CO_2)$ is not well constrained by paleoclimate data, i.e model realisations with completely different values of K can produce similarly good agreement with the paleoclimate reconstructions. However, for the future simulations different K values lead to very different impact of anthropogenic CO₂ emissions on the Earth system evolution ~~(see Fig. 3 and the discussion below)~~. This is

why ~~we propose carrying forward only those~~for performing future simulations we select only those *Paleovalid* solutions ~~in~~ *Valid* which are consistent with CLIMBER-2 simulations. However, since ~~as in~~Ganopolski et al. (2016) presents only single values for K parameter and did not perform any uncertainty analysis of K, we applied a rather “soft” constraint on its value, namely, it should ~~i.e.~~ satisfy the condition $-K \geq -150 \text{ W m}^{-2}$. ~~With this choice we tried to account for the fact that this value can be strongly model dependent. This~~Even such “soft” constraints leads to a dramatic reduction of the remaining model ~~realizations—sub ensemble~~, which we call *Accepted*, and which consists only of 29 elements~~members~~. A summary of the characteristics of the *Valid-Paleovalid* and *Accepted* ensembles ~~can be is~~ found in Table 32. The probabilistic distribution of solutions for v , CO_2 and ΔT derived from *Accepted* is displayed in Fig. 4 and the ranges of the different model parameters across *Accepted* ~~can be is~~ found in Table 24.

In the next section where we present results of future simulations we use only *Accepted* model versions. However, to demonstrate the importance of additional constraint on the value of K, we perform future simulations also with all *Paleovalid* versions. Fig. 3 illustrates these results through the timing of the first (after the present) full glacial conditions as a function of K. Not surprisingly, for the natural scenario, there is no clear dependence on K and the first full glacial conditions are predicted 100 or 200 kyr from present. The situation is very different for the scenarios with anthropogenic emissions. For example, for low (500 PgC) emissions, the model versions with K value satisfying the criteria for the *Accepted* model version, predict the next full glacial conditions around 200 kyr from present. However, most of the model versions with higher K predict the next full glacial conditions 700 kyr and even 900 kyr from present. The situation is similar for other anthropogenic scenarios. This analysis clearly shows the importance of applying a tight constraint on K value to produce realistic future scenarios.

In addition, we define the Best Solution P_{Best} as:

$$P_{\text{Best}} = \text{Argmax}(C_v(P, -800 \text{ kyr}, 0 \text{ kyr}), P \in \text{Accepted}) \quad (2018)$$

P_{Best} values ~~can be find~~are shown in Table 42. The time-series $v_{\text{model}(P_{\text{Best}})}$, $\text{CO}_{2\text{model}(P_{\text{Best}})}$, $\Delta T_{\text{model}(P_{\text{Best}})}$ are shown in Fig. 4. The model version named Best Solution is rather successful in reproducing the observed climatic variability in the three considered variables for the whole [-800 kyr, 0 kyr] period. The model skill in reproducing both the temporal variability and amplitudes of the time-series fluctuations is remarkable. In particular, the performance for ice volume is excellent with a correlation between model and paleodata of 0.86 (Fig. 4a). The model is able to correctly capture the timing of all the glaciations, being the LGM the highest volume event. The ice volume model and paleorecord discrepancy is largest for MIS 18 ([-750 kyr, -710 kyr]). The frequency spectra of $v_{\text{model}(P_{\text{Best}})}$ and the empirical estimation are also in good agreement (Fig. 5): there is a clear dominant peak at 100 kyr, with additional power at 41 and 23 kyr. For atmospheric CO_2 concentration, the correlation between modelled and recorded series is 0.62 (Fig. 4b). While the model succeeds in reproducing glacial-

interglacial variability, it overestimates its magnitude: around 130 ppm instead of 80 ppm. It is, in particular, evident that during interglacials the CO₂ restores to its equilibrium value of 278 ppm, which according to the paleorecords is not accurate before the MBT. Finally, the correlation between modelled and paleo global mean surface temperature anomalies is 0.54 (Fig. 4c). The model slightly underestimates the variability of this variable, ranging between 0°C and -5°C, while the observational information suggests a larger range between 2°C and -7°C. The high positive temperature anomalies during some previous interglacials, however, are questionable.

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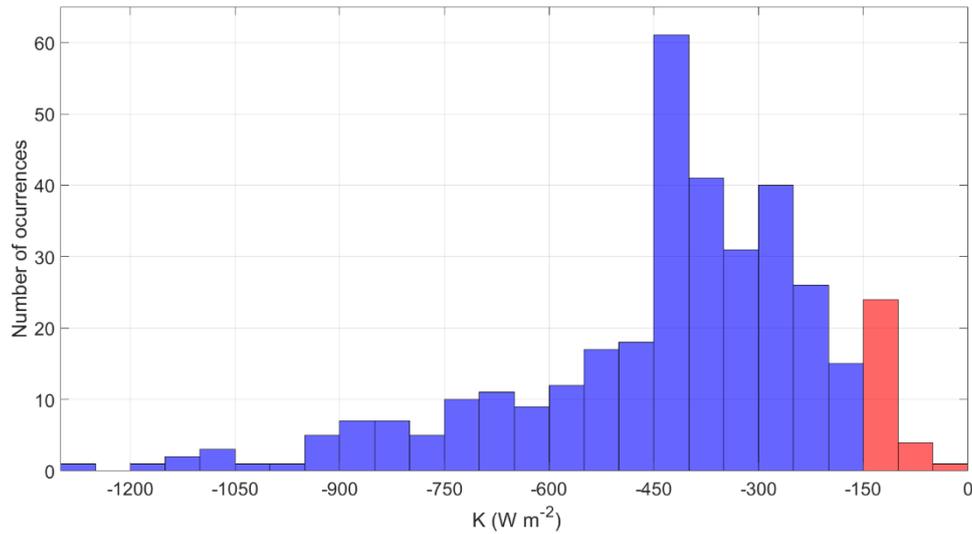
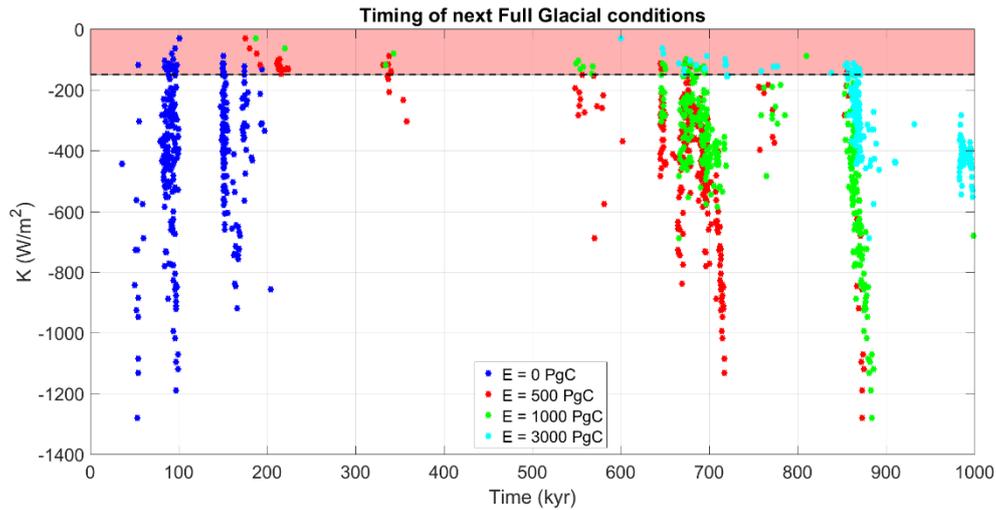


Figure 2: Histogram for K (see Eq. (17)) across the different members of the *PaleovValid* ensemble of solutions. Red bars correspond to the *Accepted* solutions.



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Figure 3: Dependence of timing of the next full glacial conditions ($v=0.5$) on parameter K , for all solutions in *PaleovValid* and under the different emission scenarios considered. The black dashed line indicates the value of $K = -150 \text{ W m}^{-2}$, which is the limit for the solutions in the red shaded area indicates the solutions in the *Accepted* set.

615 4 Next Simulations for the next 1 Myr

In this section we present results of simulations produce probabilistic of the evolution of the Earth system in the next 1 Myr under forecasts of the evolution of the Earth system in the next 1 Myr, considering natural and several anthropogenic emissions scenarios. For the probabilistic forecast we evaluate the future evolution of the solutions derived from using the *Accepted models* ensemble.

620 4.1 Description of CO₂ emission scenarios

Future scenarios are generated considering different temporal evolution paths for the excess atmospheric CO₂ concentration, which we assume depends only on the cumulative amount of fossil-fuel combustion, $\text{Anth}_{\text{CO}_2}(t)$ (see Eq. (7)). For the natural scenario, $\text{Anth}_{\text{CO}_2}(t)$ is assumed set to be zero at all times and, as a consequence, the climate system follows its natural evolution forced only by changes in orbital forcing.

625 For the fossil-fuel emission scenarios, we consider instantaneous releases of CO₂ at t = 0, of different magnitudes. If the
duration of the emission pulse release is rather short (centuries) and is followed by zero CO₂ emission, the future long-term
evolution of CO₂ depends primarily on the cumulative emission magnitude while the rate of release is only of secondary
relevance (Eby et al., 2009). For the period 1750-2017, estimations place the fossil-fuel cumulative emissions in Carboncarbon
equivalent at 660 +/- 95 PgC (Le Quéré et al., 2018). Projections for future cumulative emissions range between ~700 PgC
630 and ~3000 PgC, being the former an estimate of fossil-fuel reserves exploitable with today's technical and economic
constraints and the latter an estimation considering reserves that might become exploitable in the future (McGlade and Ekins,
2015). Taking into account the already achieved and future cumulative emissions estimations, we generate three different
scenarios: low, intermediate and high emissions, corresponding to instantaneous pulse releases of magnitude 500 PgC, 1000
PgC or 3000 PgC, respectively. In each scenario, after the pulse release, Anth_{CO₂}(t) follows a multi-exponential decay function
635 as proposed by Lord et al. (2016):

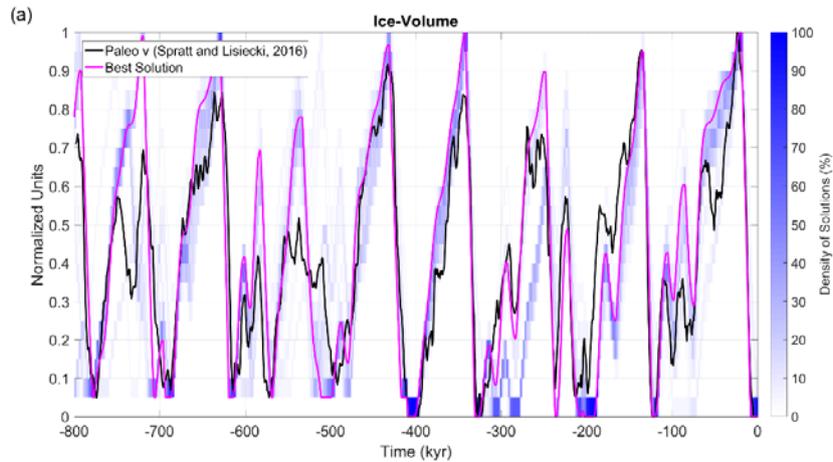
$$Anth_{CO_2}(t) = 0.469 * E \sum_{i=1}^5 A_i e^{-t/\tau_i} \quad (21+9)$$

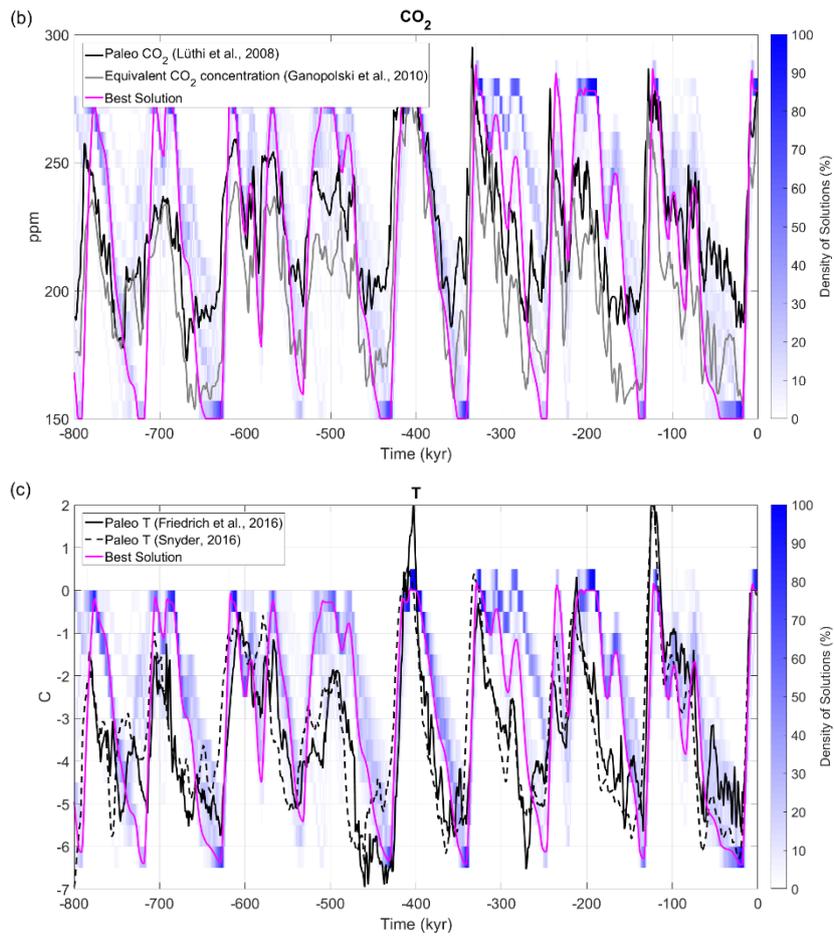
$$A_i(E) = \alpha_i + \beta_{i1}E + \beta_{i2}E^2 + \beta_{i3}E^3 \quad (22\theta)$$

$$\tau_i(E) = \gamma_i + \delta_{i1}E + \delta_{i2}E^2 + \delta_{i3}E^3 \quad (23+)$$

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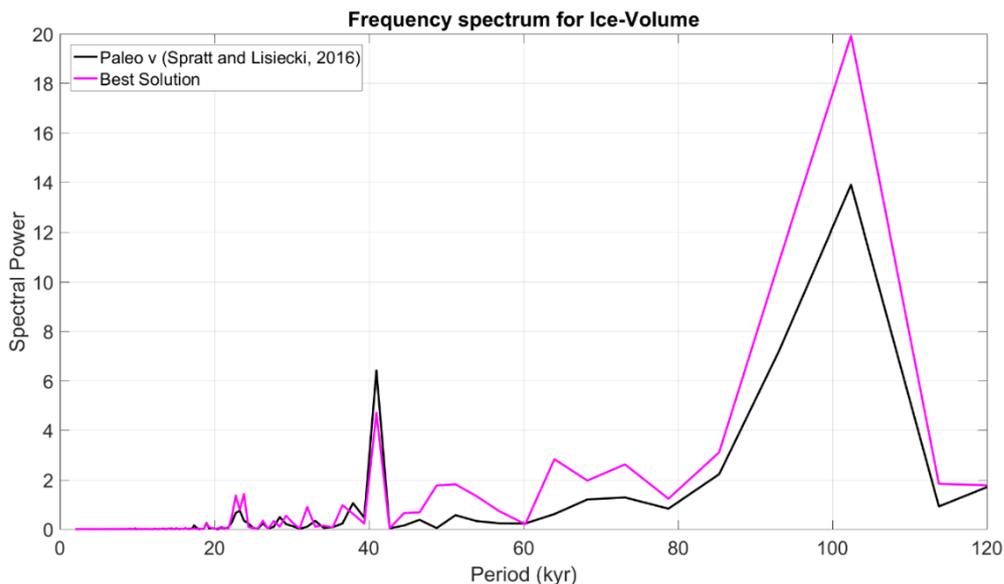
where E (PgC) represents the cumulative fossil-fuel emission magnitude, the constant 0.469 transforms PgC to ppm of CO₂
and Anth_{CO₂}(t) is expressed in units of ppm.





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Figure 4: Same as Fig.1 but for solutions in the *Accepted* ensemble. In addition, magenta lines show the evolution of the Best Solution.



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Figure 5: Frequency spectrum for simulated (Best Solution, magenta) and reconstructed ice volume (Spratt and Lisiecki, 2016; black).

655 Parameter values ($\alpha_i, \beta_{ij}, \gamma_i, \delta_{ij}$ $i=1, \dots, 5; j=1, \dots, 3$) are derived in Lord et al. (2016) in order to produce a good fit to a series of 1 Myr pulse-type of experiments performed with the *c*GENIE EMIC (Lenton et al., 2006). Lord et al. (2016) consider total anthropogenic CO₂ emissions ranging between 0 PgC to 20.000 PgC, where the latter value is justified assuming future techno-economic advances could make additional non-conventional resources such as methane clathrates available for extraction. The use of results from long-term simulations with the *c*GENIE model represents a substantial improvement compared to Archer and Ganopolski (2005) who used CO₂ scenarios obtained with a simple marine carbon cycle model which did not explicitly account for a climate – weathering feedback.

660

The function Anth_{CO₂}(t) from present until 1 million years into the future, under the low, intermediate and high emissions scenarios is shown in Fig. 6. 100 kyr after the (low, intermediate or high) emissions pulse release ~ 5% of the initial CO₂ anomaly is still present in the atmosphere. For the low emissions scenario, only 500 kyr after the pulse emission ~~the~~ the remaining anthropogenic CO₂ concentration anomalies drop below 1% of the original perturbation, marking the return of the system to its natural evolution state. For the medium and high emission scenarios, the return to natural conditions takes even longer.

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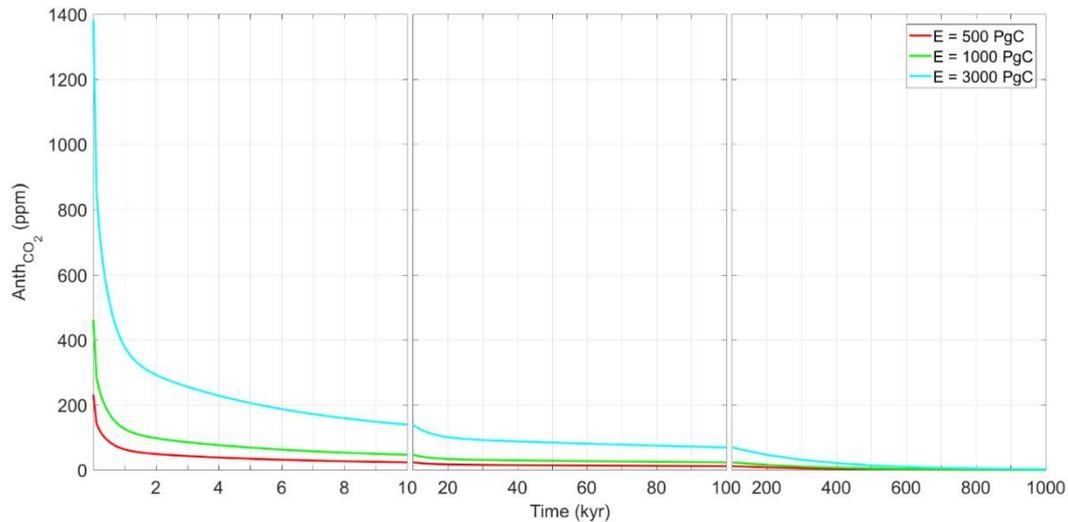


Figure 6: $\text{Anth}_{\text{CO}_2}(t)$ expressed in ppm, following Lord et al. (2016) for the three different emission scenarios considered.

675 **4.2 Possible future natural Earth system trajectories ~~Forecast under natural scenario~~**

Figure 7a displays future global ice volume evolution for the Best Solution and the probabilities of the trajectories of individual model realizations. It shows both periods of high agreement and periods of high discrepancy between solutions, highlighting the co-existence of more certain (deterministic) and less certain aspects of the Earth system response to orbital forcing. Most of the solutions agree that the planet will remain in a long interglacial state for the next 50 kyr, during which no large ice-growth is expected (however, there is one solution in which glacial inception starts at 20 kyr and full glacial conditions are reached by 36 kyr). This behaviour is predefined by one of the model constraints, namely, that v should be equal zero at present and by a low-weak orbital forcing during the next 520 kyr. With high level of confidence, a second long interglacial is expected to occur between 450 and 500 kyr, which is another prolonged period of low-weak orbital forcing related to 400 kyr periodicity of eccentricity. Also with high certainty, short (~ 10 kyr) interglacials are expected to occur in 110, 230, 250, 270, 370, 620, 820 and 900 kyr from present.

In the conceptual model we define that full glacial conditions occur when the ice volume (v) reaches 0.5 in normalized units. If full glacial conditions are reached at a time T , then the corresponding glacial inception is defined as the first time before T in which $v > 0$:

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$$\text{Glacial inception} = \sup\{t / t < T \ \& \ v(t) > 0\}$$
 (24)

695 ~~Following those definitions, R~~egarding the timing of the next full glacial conditions (~~defined here as the first moment after present in which ice volume reaches 0.5 in normalized units~~) the solutions tend to cluster into two different possibilities: predicted to occur in ~90 kyr or in ~150 kyr (see also Fig. 3). For those cases, the ~~ice growth begins~~glacial inception occurs at ~50 kyr or ~110 kyr, respectively. There is also some level of certainty in the timing of occurrence of major glaciations (defined here as periods with ice volume higher than 0.8 in normalized units): in ~100 kyr, ~210 kyr, ~410 kyr, ~450 kyr, ~680 kyr, ~790 kyr, ~870 kyr and 950 kyr from present. The ~~forecast~~possible future scenario indicates that only two of these major glaciations (the ones in ~210 kyr and ~870 kyr after present) have high probability of reaching LGM magnitudes (~1 in 700 normalized units).

If we consider the ~~forecast~~possible future scenario produced by the trajectory of the Best Solution (Fig. 8) as the most likely path, under natural circumstances: 1. The next glacial inception is expected to occur ~ 50 kyr after present (with full glacial conditions reached ~90 kyr after present); 2. In the next 1 Myr two long interglacials should take place; 3. Nine major 705 glaciations are expected, all of them with maximum magnitudes around 10-15% lower than LGM level. The natural glacial cycles of the future are strongly dominated by 100 kyr cyclicity as were the past ones.

4.3 ~~Forecast~~Possible future trajectories under different anthropogenic CO₂ emissions scenarios

The evolution of the ice volume even in the low-emissions anthropogenic scenario (E = 500 PgC, which is lower than already emitted) is ~~forecast~~possible future scenarioed to be significantly different from the natural one in the next 200 kyr, indicating 710 a long-lasting impact of fossil-fuel CO₂ releases (Fig. 7b). The uncertainty levels for the next 150 kyr and from 400 to 500 kyr are greatly reduced, with most of the solutions agreeing on ice-free NH conditions (~~excepting Greenland~~). Under this scenario, only after half a million years is the Earth system able to return to an evolution similar to the one expected under natural conditions. Under this setting, the next full glacial conditions are not expected to be able to occur before 180 kyr from present, implying a delay of at least 90 kyr from the natural expected evolution which is consistent with Archer and Ganopolski (2005) 715 and Ganopolski et al. (2016). In particular, the ice volume evolution of the Best Solution under this scenario differs significantly from the natural one during two periods: [0 kyr, 200 kyr] and [400 kyr, 500 kyr]. In both time spans glaciations fail to be triggered for this scenario (Figs. 7b and 8a). In [200 kyr, 350 kyr] the low-emissions Best Solution trajectory is similar to the natural one, except for a lag of ~10 kyr. CO₂ and global mean surface temperature anomalies (Figs. 8b and 8c) pathways are also restored to natural conditions only after 500 kyr from present.

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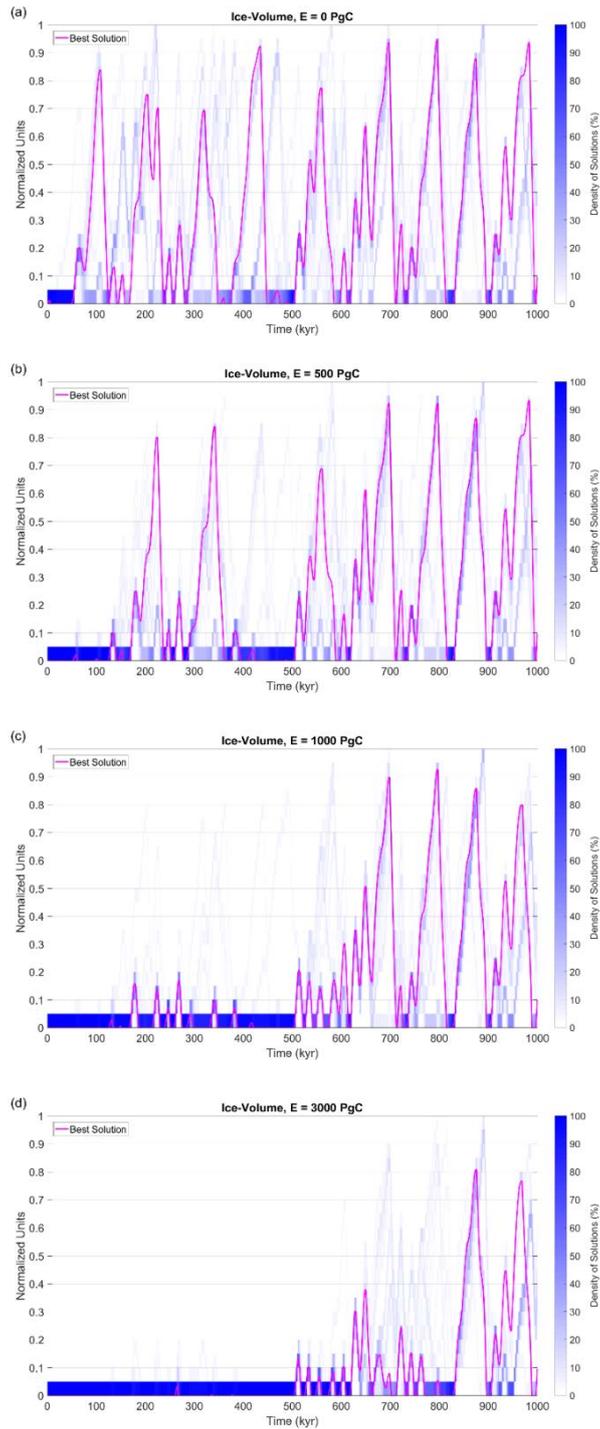


Figure 7: Probabilistic possible simulated ice volume for different future scenarios: (using the solutions in Accepted) for future NH landmass ice volume except Greenland (normalized to unity). a. Natural scenario. b. Low-emissions scenario (E = 500 PgC). c. Intermediate-emissions scenario (E = 1000 PgC) and d. High-emissions scenario (E = 3000 PgC). The magenta line depicts the

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behaviour of the Best Solution,- the blue shades represent the density of Accepted solutions per unit of phase space (Units are: 1kvr for time, 0.05 for ice, 7ppm for CO₂ and 0.5°C for temperature).

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For the intermediate-emissions case (E = 1000 PgC) the effect of anthropogenic CO₂ emission on the Earth system is even more notorious and long-lasting. The probabilistic ensemble of forecast possible future scenario indicates significantly different conditions from the natural ones in the whole next half million years, with very low probability of considerable ice growth during the totality of this time span (Fig. 7c). In most of the solutions the full glacial conditions occur for the first time either in ~ 550 kyr or between 600 and 700 kyr from present (Fig. 3). In particular, the Best Solution is part of the latter group, with the next full glacial condition taking place only at ~ 670 kyr after present. Additional insight into the relevance of the parameter K for the future behaviour under anthropogenic influence is presented in Figure 8: the larger the magnitude of K, the larger the delay in the next glacial inception.

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Finally, with a cumulative pulse-emission of 3000 PgC, the Earth system is expected, with high confidence, to experience almost ice-free conditions for the next 600 kyr (Fig. 7d). Under this scenario, the next full glacial conditions are not expected to occur before 670 kyr from present, implying a lag of almost 600 kyr from the natural scenario. For example, in the Best Solution, only two major glaciations are predicted, peaking in 880 kyr and with a magnitude 20% lower than LGM (Figs. 7d and 98a). In the 3000 PgC scenario, half a million years into the future, the CO₂ is expected to be 300 ppm (22 ppm higher than preindustrial levels) and the global mean surface temperature still 0.5°C higher than preindustrial level (Figs. 98b and 98c).

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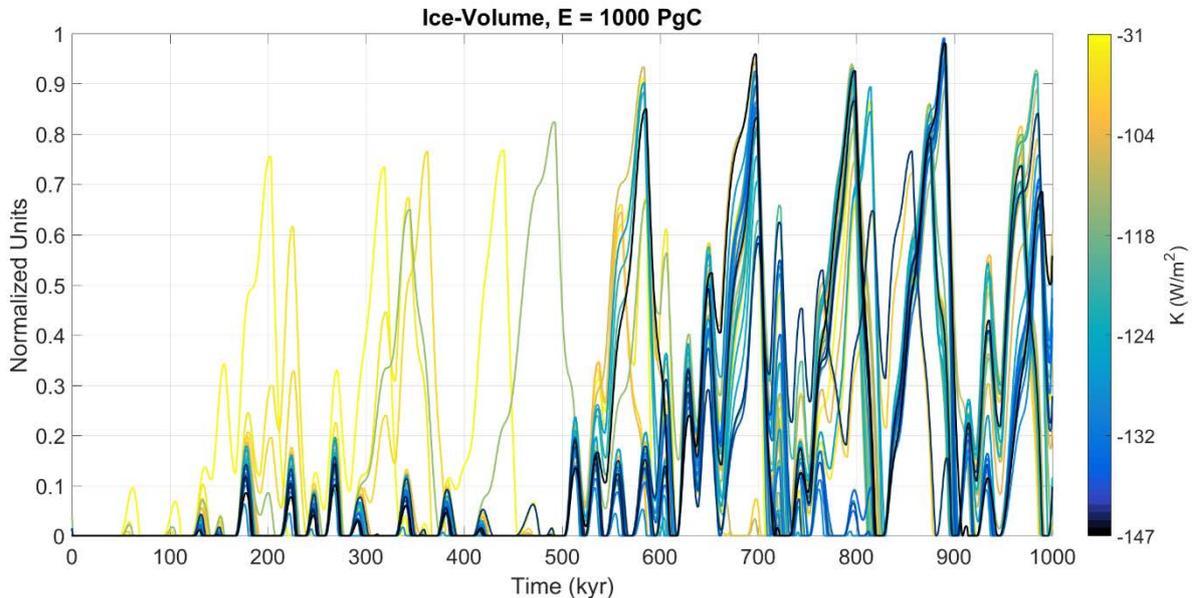
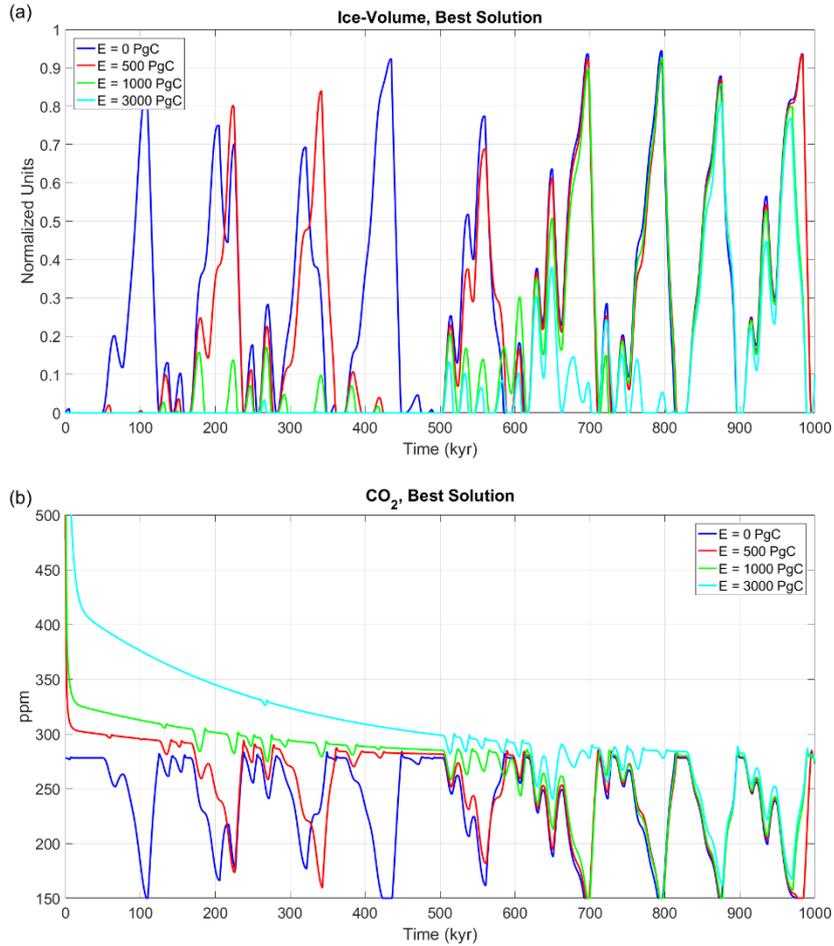


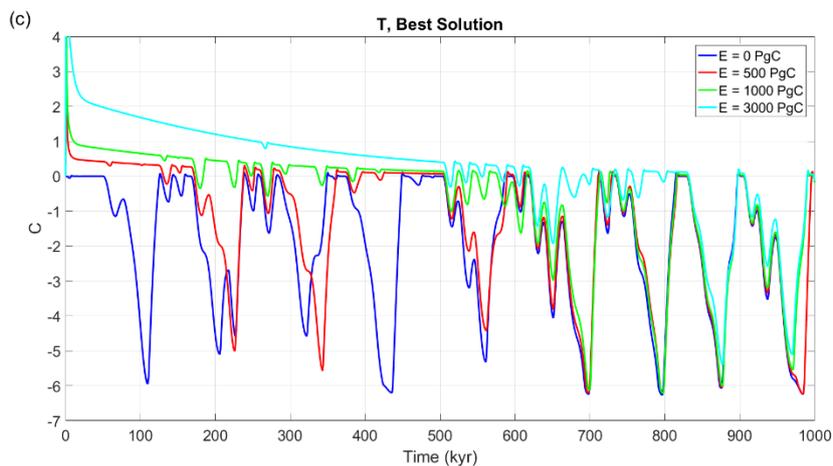
Fig. 8: Ice volume time-series for each solution in *Accepted* for the intermediate-emissions scenario ($E = 1000$ PgC). Each solution is coloured according to its corresponding K value.

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770 **Figure 98: Possible future scenario using the Best Solution under different emission scenarios: a. Ice volume (normalized to unity), b. CO₂ (ppm), c. Global mean surface temperature anomalies (C).**

5 Sensitivity of model results to the choice of constraints and parameters

775 In order to assess the robustness of the results to the model or parameter selection strategy, we performed a series of sensitivity experiments. Table 4 summarizes the settings for the Base experimental setup and five different sensitivity experiments. For each experiment, the optimisation process is carried out and a *Paleovalid* and *Accepted* sets of solutions are created (the definition of the *Paleovalid* might change from the standard one in equation 17 in some of the experiments, see Table 4). As a summary of results, Fig. 10 depicts the range of timings of the next full glacial conditions for each experiment and anthropogenic emissions scenario.

780 In the first sensitivity experiment (Sen_MBT) the ad-hoc imposition related to MBT (Eq. (12)) is lifted. Naturally, the correlation between modelled and paleo ice volume in Sen_MBT is slightly lower than in the base model formulation (the ensemble mean across *Accepted* drops from 0.79 to 0.72). The ensemble mean timing of next full glacial conditions under the natural, low and intermediate emissions scenarios for this setting does not substantially differ for the one in the Base experimental configuration. For the high emissions scenario, though, the timing for next full glacial conditions spans the full 200 kyr - 860 kyr period, with the ensemble mean being around 600 kyr in the future (~ 150 kyr earlier than in Base).

785 Second, we test the sensitivity of the results to the choice of ECS of the model. We report on the results ensuring an ECS of 3°C (instead of 3.9°C in Eq. (14)). The results of this experiment (Sen_ECS) indicate that the main conclusions are not overly
 790 sensitive to the change in the selected ECS for the natural, low and intermediate emissions scenarios. However, for the high

emissions scenario the timing of next full glacial conditions is confined to 520 kyr – 620 kyr, somewhat earlier than expected using the Base configuration.

In the third experiment, Sen_future, we modify Eq. (15) so that it does not include any conditions on the future. Without this constraint, the new *Paleovalid* ensemble contains three solutions (out of 400 ensemble members) for which glacial inception occurs at some point between present and 20 kyr into the future. However, those solutions do not fulfil the requirements for being included in *Accepted*. The main results remain largely unchanged with respect to the Base experiment.

Fourth and fifth, we test the sensitivity of the results to the selected optimisation metric. In the fourth experiment (Sen_Cv_C_CO2) we select parameters in order to jointly maximise correlations between modelled and paleo ice volume and CO₂. For that purpose we maximise C_v + C_{CO2}.

We find that for the Sen_Cv_C_CO2 (Base) experiment the *Accepted* ensemble mean correlation between modelled and paleo ice volume is 0.71 (0.79) and between modelled and paleo CO₂ is 0.55 (0.5), therefore showing similar performance in terms of correlations. The most notorious difference between this and the Base setting occurs in the natural scenario, for which the majority of the solutions agree next full glacial conditions will develop in 130 kyr from present.

In the fifth experiment, instead of optimising correlations, we select parameters to minimise the root mean squared error (RMSE) between modelled and paleo ice volume. Using a RSME instead of a correlation optimisation strategy leads to the strongest changes in the timing of the next full glacial conditions for the low and intermediate emissions scenarios: the ensemble mean occurs ~200 kyr and ~300 kyr earlier than in the Base configuration, respectively. While these differences in the ensemble mean behaviour are substantial, the spread of possible timings within the *Accepted* set is still comparable.

Table 4: Summary of settings for sensitivity experiments.

	Experiment Name					
	Base	Sen_ECS	Sen_MBT	Sen_future	Sen_Cv+CCO2	Sen_RMSEv
MBT condition (Equation 12)	Yes	Yes	No	Yes	Yes	Yes
ECS (in Equation 14)	3.9°C	3°C	3.9°C	3.9°C	3.9°C	3.9°C
Future condition (Equation 15)	Yes	Yes	Yes	No (modified to: v(t=0)<0.025)	Yes	Yes
Optimisation target	Maximise Cv	Maximise Cv	Maximise Cv	Maximise Cv	Maximise Cv + CCO2	Minimise RMSEv
Condition on P to be in Valid	Cv(P,-800kyr,0kyr)>=0.7	Cv(P,-800kyr,0kyr)>=0.7	Cv(P,-800kyr,0kyr)>=0.7	Cv(P,-800kyr,0kyr)>=0.7	Cv(P,-800kyr,0kyr)>=0.7	RMSE(P)<0.25 *

*0.25 corresponds to the RMSE of forecasting the average ice volume in the period [-800 kyr, 0kyr].

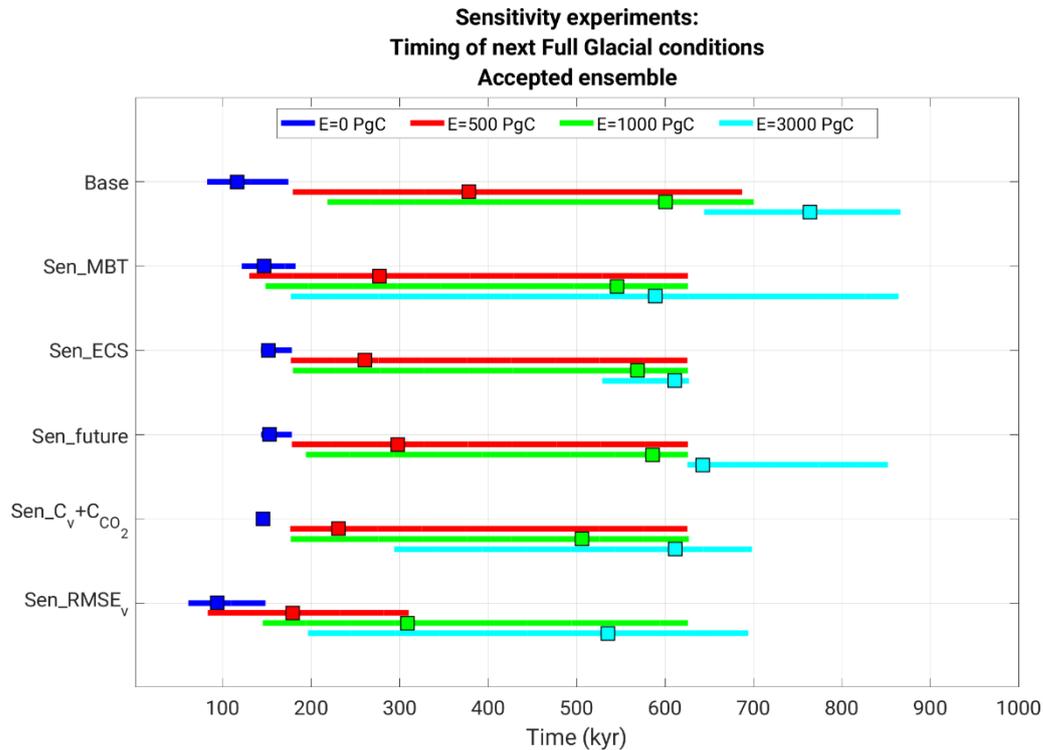


Fig. 10: Timing of next full glacial conditions (first time after present when $v=0.5$) for different model configurations and emission scenarios, considering the *Accepted* solutions.

For a certain configuration and emission scenario, line segments indicate the 5-95% percentiles of the timings while squares indicate the ensemble mean.

56 Summary and conclusions

825 We propose a reduced-complexity process-based model of the evolution of global ice volume, atmospheric CO₂ concentration and global mean temperature at multi-millennial and orbital timescales. The model only external forcings are the orbital forcing and (for the future) anthropogenic CO₂ cumulative emissions. The model consists of a system of three coupled non-linear differential equations representing physical mechanisms relevant for the evolution of the Climate-climate – Ice-ice Sheets-i
sheets – Carboncarbon cycle Ssystem-in-timescales longer than thousands of years. Model parameters are selected to guarantee
 830 achieve the best possible agreement with reconstructed ice volume variations during past 800 kyr and consistency with
CLIMBER-2 model results. Global temperature anomaly at the Last Glacial Maximum and the current best estimate for the equilibrium climate sensitivity have been used as additional constraints on model parameters. Furthermore, the solutions are

bounded to account for the difference between pre- and post- mid-Brühnes transition interglacial conditions and a minimum for CO₂ concentration is set to avoid unrealistic solutions.

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The model is successful in reproducing the natural glacial-interglacial cycles of the last 800 kyr, in agreement with paleorecords, replicating both the timing and amplitude of the major fluctuations, with a correlation between modelled and reconstructed global ice volume of up to 0.86. The model also shows a promising performance when evaluated against data not used for training, indicating a certain skill in predictive mode.

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Using different model realizations we generate an ensemble of possible future trajectories ~~of probabilistic forecast~~ ~~of for~~ the evolution of the Earth system over the next one million years under natural and three different anthropogenic CO₂ release emission scenarios.

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In the natural scenario, in which anthropogenic CO₂ emissions are assumed to be zero at all times, the model assigns high probability of occurrence of: 1. Long interglacials in two periods: ~~{0 kyr, 120 kyr}~~ from 0 till ca. 50 kyr after present and ~~{between 400 kyr and, 500 kyr after present}~~ and 2. The next glacial inception will most likely occur ~50 kyr after present with the development of full glacial conditions ~90 kyr after present. It is also clear, however, that even though there is a high level of agreement in the ~~solutions'~~ trajectories during the past 800 kyr, their paths tend to diverge for the future indicating that the past does not perfectly constraint the future evolution of the climate – ice sheets – ~~Carbon~~carbon cycle system.

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The selected model versions exhibit a large sensitivity to fossil-fuel CO₂ releases, with even already achieved emissions (500 PgC) causing a behaviour significantly different from the natural one over extremely long periods of time. The model predicts that a fossil-fuel CO₂ release of 500 PgC will make the beginning of the next Ice Age highly unlikely in the next 120 kyr and a delay of half a million years is a possibility that cannot be ruled out. Cumulative fossil-fuel CO₂ emissions of the order of 3000 PgC (which could be achieved in the next two to three centuries if humanity does not ~~constraint~~cease fossil-fuel usage; Meinshausen et al., 2011) will most likely lead to ice-free conditions over the NH ~~landmasses~~continents throughout the next half million years. Under this scenario, the probability for the next glacial inception to occur before 600 kyr is extremely low. Thus our study demonstrates that, in spite of large ~~uncertainties~~ in model parameters, there is a qualitative difference in the long-term Earth system trajectories (Steffen et al., 2018) between cumulative emissions of 500 and 3000 ~~GtC~~PgC.

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However, it is clear that large uncertainties pose serious challenges to our ability to model the deep future. First, ~~t~~The present study reveals ~~one serious challenge in modeling of the deep future: that~~ the results ~~show very strong~~are very sensitivity to the relationship between critical insolation threshold for glacial inception and CO₂ levels. In turn, this relationship is poorly constrained by the paleoclimatic data because during previous interglacials CO₂ was ~~close~~similar to or lower than the preindustrial CO₂ level. So far, this relationship has been analysed only with the CLIMBER-2 model (Calov and Ganopolski,

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2005; Ganopolski et al., 2016). This is why in this work, we assumed the uncertainty in this relationship of 100%. Reducing this uncertainty by performing experiments similar to those described in Ganopolski et al. (2016) but with more advanced Earth system models can help to reduce uncertainties in future projections. Second, the knowledge on the long-term (10^5 - 10^6 yrs) carbon cycle dynamics remains poor, primarily because of the lack of accurate empirical data. In particular, the results from Lord et al. (2016) used in this paper to simulate the decay of excess CO_2 in the atmosphere are largely theoretical and based on the assumption that the long-term carbon cycle is regulated solely through the silicate weathering omitting the role of other feedback processes (i.e. organic carbon burial in marine sediments, kerogen weathering).

We consider our approach potentially useful in providing projections of the possible future Earth system evolution. However, it has a number of limitations. In particular, one key assumption of our procedure is that ice volume is at all times equal or higher than the preindustrial level. While a possible future melting of Greenland will not substantially affect this hypothesis, a possible future melting of Antarctica definitely will. Therefore, we must emphasize that the methodology presented here is not applicable for anthropogenic scenarios in which a substantial melting of the Antarctic ice sheet could be expected (cumulative emissions higher than 3000 PgC).

Appendix A: Estimation of predictive skill

When a model is developed with a predictive goal (like in this study), it is important to have an assessment of its predictive skill. The best possible estimations of this skill are obtained by evaluating how well the model performs when presented with previously unseen data.

To follow the cross-validation assessment criterion (Stone, 1974), we divide the period [-800 kyr, 0 kyr] in halves: P1 = [-800 kyr, -400 kyr] and P2 = [-400 kyr, 0 kyr]. First, P1 is labelled training set and P2 validation set. Second, the model is trained using the information in P1: we run the Optimisation scheme to maximise the correlation between modelled and paleo ice volume during P1; we then accept all the models (parameter combinations) yielding correlations higher than 0.7. Third, all the accepted models are validated using the information in P2. We select three different performance metrics: the correlation between modelled and paleo ice volume, between modelled and paleo CO_2 or between modelled and global mean surface temperature anomaly, calculated during P2. Finally, we repeat the previous three steps but labelling P1 as validation and P2 as training sets, respectively. For each metric, the average that the models achieve on the two possible validation sets considered is the estimate of the model's predictive skill.

900 In Table A1 we summarize the results of this cross-validation assessment, showing the results for the three performance metrics. For reference, Table A1 also displays the values of the metrics when calculated over the training period (instead over the validation period). Results are the average of two possible cases: P1 as training and P2 as validation sets, or P2 as training and P1 as validation sets. As expected, when the same period is used both for model training and evaluation the results are in better agreement with paleoclimate data: the ice volume metric is 0.78, the CO₂-metric 0.52 and the temperature-metric 0.58.

905 These values drop to 0.49, 0.36 and 0.34 respectively, when the independent validation set is used to calculate the metrics. Even though the drop in performance is not negligible, the model skill is still substantial in particular for predicting ice volume.

Table A1: Estimation of Model Predictive Skill. Only solutions leading to a correlation between modelled and paleo estimation of ice volume higher than 0.7, calculated over the training period, are considered. Please see text for more details on the training and validation periods used.

Ensemble-Mean Correlations Between Modelled and Paleo-data	Calculated over the Learning period	Calculated over an independent validation period
Ice volume	0.78	0.49
CO ₂	0.52	0.36
Temperature	0.58	0.34

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Code/Data availability

Code and data used in this study can be obtained from: <https://doi.org/10.17605/OSF.IO/KB76G> (Talento, 2021).

Author contribution

920 AG conceived the original idea. ST performed all the calculations and produced figures. Both authors contributed to the analysis of results and writing of the manuscript.

Competing interests

The authors declare that they have no conflict of interest.

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Acknowledgments

Financial support for this study was provided by the Swiss National Cooperative for the Disposal of Radioactive Waste (NAGRA).

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