



1 **A continued role of Short-Lived Climate Forcers under the Shared Socioeconomic**
2 **Pathways**

3 *Marianne T. Lund^{1*}, Borgar Aamaas¹, Camilla W. Stjern¹, Zbigniew Klimont², Terje K.*
4 *Berntsen^{1,3}, Bjørn H. Samset¹*

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6 *1 CICERO, Center for International Climate Research, Oslo, Norway*

7 *2 International Institute for Applied Systems Analysis (IIASA), Laxenburg, Austria*

8 *3 Department of Geosciences, University of Oslo, Oslo, Norway*

9 **Corresponding author: m.t.lund@cicero.oslo.no*

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32 **Abstract**

33 Mitigation of non-CO₂ emissions plays a key role in meeting the Paris Agreement ambitions
34 and Sustainable Development Goals. Implementation of respective policies addressing these
35 targets mainly occur at sectoral and regional levels and designing efficient mitigation strategies
36 therefore relies on detailed knowledge about the mix of emissions from individual sources and
37 their subsequent climate impact. Here we present a comprehensive dataset of near- and long-
38 term global temperature responses to emissions of CO₂ and individual short-lived climate
39 forcers (SLCFs) from 7 sectors and 13 regions - for present-day emissions and their continued
40 evolution as projected under the Shared Socioeconomic Pathways. We demonstrate the key role
41 of CO₂ in driving both near- and long-term warming, and restate the importance of mitigating
42 methane emissions, from agriculture, waste management and energy productions, as the
43 primary strategy to further limit near-term warming. Due to high current emissions of cooling
44 SLCFs, policies targeting end-of-pipe energy sector emissions may result in net added warming
45 unless accompanied by simultaneous methane and/or CO₂ reductions. East Asia, North America
46 and Europe remain the largest contributors to total net warming until 2100, regardless of
47 scenario, while South Asia and Africa south of the Sahara overtakes Europe by the end of the
48 century in SSP3-7.0 and SSP5-8.5. We find that SLCFs will continue to play a role in many
49 regions, particularly those including low- to medium-income countries, under most of the SSPs
50 considered here. Our dataset is made available in an accessible format, aiming also at decision-
51 makers, to support further studies into the implications of policy implementation at the sectoral
52 and regional scales.

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69 **1 Introduction**

70 At the core of any strategy for sustained, long-term abatement of climate change are strong
71 reductions in emissions of CO₂ and other long-lived greenhouse gases (LLGHGs). However,
72 most anthropogenic activities emit a suite of additional species, with a range of climate impacts,
73 commonly termed short-lived climate forcers (SLCFs). While differing in characteristics and
74 contribution to temperature change, their common feature of a much shorter atmospheric
75 residence time compared to LLGHGs has resulted in significant discussion of the role of SLCFs
76 in climate mitigation strategies, in particular to limit near-term warming (e.g., Bowerman et al.,
77 2013; Pierrehumbert, 2014; Rogelj et al., 2015; Shindell et al., 2012; Shoemaker et al., 2013;
78 Stohl et al., 2015).

79 Many studies have placed particular emphasis on the subset of SLCFs with a warming impact
80 on climate, namely black carbon (BC), methane (CH₄) and tropospheric ozone (sometimes
81 collectively referred to as short-lived climate pollutants, or SLCPs) (e.g., AMAP, 2015; CCAC,
82 2019; UNEP, 2017). Assuming effective abatement of SLCPs, some studies estimate a
83 reduction in global temperature of 0.2-0.5°C increase by mid-century (e.g., Shindell et al.,
84 2012). Early studies brought particular attention to BC mitigation as a measure to limit near-
85 term (rate of) warming owing to the strong positive radiative forcing combined with short
86 atmospheric residence time of the aerosols (e.g., Ramanathan & Carmichael, 2008). More
87 recent work suggest that some of these early estimates may overestimate the effect of SLCP
88 mitigation (Rogelj et al., 2014; Smith & Mizrahi, 2013; Stohl et al., 2015; Takemura & Suzuki,
89 2019). While results from early studies brought some concern that the attractiveness of SLCP
90 mitigation could lead to delayed action on CO₂ emissions, most scientific studies emphasize
91 that SLCP measures should only be considered complimentary to early and stringent CO₂
92 mitigation for the achievement of long-term climate goals (e.g., Bowerman et al., 2013; Rogelj
93 et al., 2014).

94 SLCF mitigation may also give rise to potential trade-offs. As many species are commonly co-
95 emitted, any given mitigation measure or policy will affect a broad range of emitted
96 components. The combinations may, however, vary significantly between mitigation strategies
97 motivated by, and designed to address, different societal challenges. For instance, SLCFs are
98 inexorably linked to air quality (Anenberg et al., 2012; Lelieveld et al., 2015; Shindell et al.,
99 2012) and sustainable development (Haines et al., 2017; UNEP, 2019), in addition to their
100 climate impacts. The numerous environmental and societal co-benefits of SLFC reductions are
101 well recognized but may lead to adverse climatic consequences (Arneth et al., 2009). While
102 some SLCFs with a warming contribution to temperature change can, in part, be mitigated
103 individually (in particular methane), improving air quality requires consideration of all relevant
104 species, not just the warming BC particles. Removal of all present-day anthropogenic aerosols
105 may add as much as 0.5°C of additional global near-term warming according to recent work
106 (Aamaas et al., 2019; Hienola et al., 2018; Samset et al., 2018). Due to the co-emission, species
107 such as sulfur dioxide (SO₂) are also commonly affected by climate mitigation policies that
108 consider LLGHGs as a primary target. Hence, while it remains clear that deep reductions in
109 emissions of methane and BC play a key role in pathways for global emissions that limit global
110 warming to 1.5°C and 2°C warming (Harmsen et al., 2019; Rogelj et al., 2015; Rogelj et al.,



111 2018; Shindell & Smith, 2019; Xu & Ramanathan, 2017), any assessment of the mitigation
112 potential by SLCF reductions should encompass species such as sulfate, which have arguably
113 received considerably less attention so far.

114 A key characteristic of SLCFs is that the relative amount of SLCF emissions, as well as their
115 subsequent radiative forcing, can vary significantly between individual emission sources (Bond
116 et al., 2013; Lund et al., 2014b; Persad & Caldeira, 2018; Unger et al., 2010). Furthermore,
117 while previous scenarios for long-term evolution of SLCF emissions project a general, rapid
118 decline even in pathways with high climate forcing and GHG levels (Gidden et al., 2019; Rao
119 et al., 2017), the most recent generation scenarios, the Shared Socioeconomic Pathways (SSPs)
120 (O'Neill et al., 2014; Riahi et al., 2017) exhibit a much larger spatiotemporal heterogeneity in
121 projected future SLCF emissions. An up-to-date and detailed consideration of the emission
122 composition is therefore critical for the design of effective mitigation strategies and to provide
123 decision makers with a more integrated approach and guidance on how to best address linkages
124 between climate, sustainable development and air quality in policy processes (Melamed et al.,
125 2016). While studies comparing and quantifying the impacts of SLCFs and CO₂ exist, they
126 differ in selection of sectors and/or regions, methodology and emission inventory, making direct
127 comparison difficult (e.g., Harmsen et al., 2019; Kupiainen et al., 2019; Lund et al., 2014a;
128 Sand et al., 2015; Unger et al., 2010). Furthermore, studies often consider only the equilibrium
129 effect of present-day emissions, emission pulses or very simplified scenarios.

130 In the present work, we provide a comprehensive and updated investigation of the contribution
131 to near- and long-term global temperature impacts from individual SLCF and LLGHG
132 emissions. We first quantify the temperature response to an idealized pulse of present-day
133 emissions to demonstrate the methodology and temporal behavior of the various emitted
134 species, then calculate possible future evolutions of temperature impacts as they are projected
135 to develop under the pathways for future socioeconomic development, climate policy and air
136 pollution described by the SSPs. The temperature impact is calculated for seven economic
137 sectors and 13 source regions, accounting for best available knowledge and geographical
138 dependence of the forcing efficacy of different SLCFs, thereby providing a more
139 comprehensive overview than previous literature, focusing on both added benefits and trade-
140 offs offered by SLCF mitigation. By making our full data set openly available, we aim to
141 provide a toolkit for further studies of the implications of policy implementation at the sectoral
142 and regional level and demonstrate such use through calculations of the effect a set of idealized
143 sectoral policy packages.

144

145 **2 Methodology**

146 Using the concept of Absolute Global Temperature change Potential (AGTP) (Shine et al.,
147 2005), we calculate the global-mean temperature response over time to emissions of CO₂, CH₄,
148 ammonia (NH₃), BC, OC, SO₂, the ozone precursors nitrogen oxide (NO_x), carbon monoxide
149 (CO) and volatile organic compounds (VOCs) from the sectors and regions shown in Fig. 1.

150 The AGTP is an emission metric-based emulator of the climate response, and a well-established
151 method that enables us to quantify and compare global temperature impacts of a large number



152 of sources and scenarios in a transparent and, in terms of computer resources, cost-effective
153 manner. The approach is described in detail in the literature (Aamaas et al., 2013; Fuglestedt
154 et al., 2010; Shine et al., 2005); here we give a brief outline.

155 The ATGP gives the global-mean surface temperature response per kg emitted as a function of
156 time after an emission pulse, i.e., an instantaneous one-off emission. At time H after the
157 emission, the AGTP for species i is given (for each sector and region) by:

$$158 \quad AGTP_i(H) = \int_{t=0}^H F_i(t) IRF_T(H-t) dt$$

159 where F is the radiative efficiency and IRF is the impulse response function used to estimate
160 the temperature response to a given radiative forcing. See Aamaas et al. (2013) for further
161 details about AGTP calculations for individual species. For CO_2 and methane, we use the
162 global-mean F from the IPCC Fifth Assessment report (AR5) (Myhre et al., 2013), adjusted for
163 recent updates of the methane forcing (Etminan et al., 2016). For short-lived species (with the
164 exception of ammonia (NH_3), for which we also use the IPCC AR5 best estimate global forcing
165 value), we use values of F that depend on the location of the emission and calculate region-
166 specific AGTPs for BC, OC, SO_2 , and ozone precursors. These regional radiative efficiencies
167 (i.e., the global radiative forcing per unit of regional emissions) are derived from simulations
168 performed with the global chemistry transport model OsloCTM3 (Lund et al., 2018; Søvde et
169 al., 2012) for the second phase of the Hemispheric Transport of Air Pollution (HTAP2)
170 (Janssens-Maenhout et al., 2015) combined with radiative kernels. To account for the additional
171 negative RF resulting from aerosol-cloud interactions (or indirect aerosol effects), we scale the
172 regional AGTP of SO_2 by a factor of 2.1 based on the ratio of total global RF of sulfate to that
173 due to direct effects alone from the IPCC AR5 (Myhre et al., 2013). Due to lack of available
174 information, the same scaling factor is applied for all regions, recognizing that also the indirect
175 effect may vary with location of emission. We also account for the rapid adjustments of BC
176 which have been found to partly offset the positive direct radiative impact (Samset & Myhre,
177 2015; Stjern et al., 2017), by adjusting the AGTP of BC by -15% (based on Stjern et al. (2017),
178 in all regions except South Africa, where the rapid adjustments were positive in that study.
179 Radiative forcing of BC deposition on snow and ice is not included in our estimates.

180 Following the methodology established in the literature (e.g., Fuglestedt et al., 2010), we use
181 an IRF that is the sum to exponentials representing the short and long mode of the climate
182 system response to a perturbation:

$$183 \quad IRF(t) = \lambda \sum_{j=1}^J \frac{c_j}{d_j} \exp\left(-\frac{t}{d_j}\right)$$

184 Here, c_j and d_j are constants and timescales of the two modes, respectively, and λ is the
185 equilibrium climate sensitivity (ECS) (Table 1). Values of c_j , d_j and λ are derived from the
186 analytical solution of the two-layer energy balance model used by Geoffroy et al. (2013), which
187 yields an ECS of $0.885 \text{ K (Wm}^{-2}\text{)}^{-1}$. This is somewhat lower than the ECS of $1.06 \text{ K (Wm}^{-2}\text{)}^{-1}$)
188 inherent in the IRF from Boucher and Reddy (2008) which has been used in a number of
189 previous studies including the IPCC AR5. The timescales from Geoffroy et al. (2013) are also



190 somewhat shorter than the corresponding Boucher and Reddy (2008) numbers. Combined, this
191 results in lower AGTPs values in the present study than previous literature.

192 For each region and species, the AGTPs are then multiplied by present-day (year 2014)
193 emissions from the Community Emission Data System (CEDS) (Hoesly et al., 2018) to
194 calculate the temperature impact at a given time horizon H , $\Delta T(H)$. In this study, $H=10$ years
195 and $H=100$ years are selected to present near-term and long-term impacts, respectively. The
196 AGTP framework can readily be extended from pulse-based calculations since any scenario can
197 be viewed as a series of pulse emissions and analyzed through convolution (Aamaas et al.,
198 2013). The temperature response ΔT at time t for species i is (for each region and sector) given
199 by:

$$200 \quad \Delta T_i(t) = \int_0^t E_i(t') AGTP_i(t - t') dt'$$

201 Using this approach, we also calculate the global-mean temperature response to full time series
202 of historical (CEDS) and future (the nine gridded and harmonized SSPs (Gidden et al., 2019))
203 regional and sectoral emissions.

204 We establish a range in total net global-mean temperature response on 10- and 100-year time
205 scales due to uncertainties in radiative forcing by performing a Monte Carlo analysis. Each RF
206 mechanism is treated as a random variable, following a probability density function (PDF)
207 defined based on existing literature, and the distribution for the total RF is derived by summing
208 the individual PDFs, i.e., assuming that each RF mechanisms is independent. For the aerosols
209 and their precursors, we use the multi-model results from the AeroCom Phase II experiment
210 (Myhre et al., 2013a), while for CO_2 , NH_3 , and ozone precursors, we use the uncertainties from
211 the IPCC AR5 (Myhre et al., 2013b). For further details, see Aamaas et al. (2019) and Lund et
212 al. (2017). Our temperature responses are also influenced by uncertainties in emissions and
213 climate sensitivity. A comprehensive analysis of uncertainty in all three factors is challenging
214 due to lack of data, but the potential impact is discussed in Sect. 4.

215

216 **3 Results**

217 **3.1 Near- and long-term temperature response to current emissions**

218 We first discuss the global mean surface temperature response to one year of present-day (i.e.,
219 year 2014) emissions, for global total emissions and broken down by key contributing sectors
220 and geographical source regions as shown in Fig.2. While the 10- and 100-year time horizons
221 are commonly used to represent near- and long-term impacts, we recognize that other choices
222 may affect the relative importance, and even sign, of some of SLCFs like aerosols and NO_x , or
223 be more relevant for certain applications.

224 Globally, current emissions result in an approximate balance between cooling and warming
225 SLCFs in the near-term, with main warming contributions from BC and CH_4 and cooling from
226 SO_2 and NO_x (Fig.2a). As the impact of the SLCFs decays rapidly over the first few decades
227 after emission, the net long-term temperature impact is predominantly determined by CO_2 . Note



228 that in Fig.2b-c we only show this net temperature effect (grey circles) for the 100-year time
229 scale, not individual contributions.

230 As clearly seen in Fig. 2, CO₂ emissions also cause a notable contribution to near-term warming.
231 While known in the scientific community, this role of CO₂ as both a near- and long-term climate
232 forcer is not always fully acknowledged in the discussions of LLGHGs versus SLCFs. Figure
233 2 also readily shows that the mitigation potential inherent in the present SLCF emissions is
234 highly inhomogeneous, and that co-emitted species – including CO₂ – must be taken into
235 account in any targeted climate policy.

236 Differences in emission composition result in net near-term impacts on global temperature (i.e.,
237 10 years after emission) that vary significantly, in both magnitude and sign, between sectors
238 and regions. Of the global economic sectors, the largest net near-term warming is estimated for
239 the energy (ENE), agriculture (AGR), and waste management (WST) sectors (Fig. 2b). The
240 AGR and WST sectors are primarily a source of methane-induced near-term warming. The
241 energy sector (ENE) is also characterized by a significant warming due to methane (originating
242 from fossil fuel mining and distribution) but also by a considerable cooling from high emissions
243 of SO₂. Our results hence reinforce the importance of methane as a driver of near-term warming
244 but show that benefits from reductions may be offset if accompanied by simultaneous
245 reductions in SO₂ in some cases. A particular feature of the energy sector, however, is that a
246 significant portion of methane mitigation from oil and gas as well as coal mining (production
247 and distribution) can be done independently from other energy-related (combustion) emissions.
248 An explicit distinction between production and combustion emissions was not available in the
249 gridded CEDS inventory, but, as illustrated in the following section, mitigation strategies
250 targeting one category or the other can result in distinctly different temperature outcomes. On
251 the global level, emissions from industry (IND) cause a small net cooling impact despite a
252 considerable warming from CO₂ emissions. The near- and long-term temperature impacts from
253 the aviation sector were recently quantified in a separate study (Lund et al., 2017).

254 Current SO₂ emissions are also the primary contributor to near-term cooling in all source
255 regions (Fig.2c), with smaller contribution from NO_x. The largest absolute contribution to net
256 near-term warming is caused by emissions in East Asia (EAS) and North America (NAM),
257 followed by South East Asia (SEA) and South Africa (SAF). However, the relative
258 contributions from individual species vary. In EAS and NAM, as well as Europe (EUR), the
259 impact of current emissions of cooling and warming SLCFs approximately balance in the near-
260 term and these regions cause comparable net warming impacts on 10- and 100-year time scales,
261 as seen by comparing the white and grey circles in Fig. 2c. These balancing characteristics do
262 not imply that SLCF emission reductions measures should not be implemented, but that the net
263 benefits on global temperature may be lower than expected if mitigation policies simultaneously
264 affect both cooling and warming SLFCs.

265 In SEA, SAF and South and Central America (SAM and MCA) emissions of methane and BC
266 are presently high while CO₂ emissions are low compared to other regions. This results in a net
267 warming impact after 10 years that is substantially higher than that of CO₂ alone. Combined
268 with low cooling contributions, this suggests that there is a higher potential for mitigation by
269 targeting only SLCF emissions in these regions. As in the global case, the higher potential stems



270 primarily from methane from the agriculture and waste management sectors, with additional
271 potential in the energy sector especially in MCA (see “Data Availability” for sectoral data
272 within each region). In SAF, emissions of BC from the residential and transport sectors also
273 play an important role. In contrast, emissions from the industry sector in most regions cause a
274 net negative impact on global temperature change. The energy sector is characterized by
275 competing cooling and warming SLCFs, leaving CO₂ as the primary driver of net near-term
276 warming when considering the sector as a whole, i.e., without accounting for production and
277 combustion sub-categories as discussed above.

278

279 **3.2 Temperature response to idealized policy cases and further applications**

280 The results above suggest that strategies for emission reductions clearly can play out very
281 differently in terms of net impact on global temperature across source region and sector. To
282 illustrate the importance of considering co-emissions and how our dataset may be used further,
283 we now calculate the effect on global temperature in the near- and long-term of emission
284 changes following example policies in three global sectors (ENE, AGR and SHP). The policies
285 are assumed to be motivated by either *i*) air quality improvements (policy 1, P1), *ii*) methane
286 reductions (as part of the SDG agenda or climate mitigation) (P2) or *iii*) CO₂ reductions/climate
287 targets (P3), each resulting in a different package of emission reductions (Table 2). The global
288 temperature effect resulting from elimination of these emissions after 10 and 100 years is shown
289 in Fig.3, for each individual policy and the combination of all three.

290 The energy sector can be sub-divided into fossil fuel production/distribution and combustion
291 categories. An air quality-driven policy (P1), implementing end-of-pipe measures, would
292 strongly reduce SO₂ and NO_x emissions but would not affect the methane contribution. As
293 shown by the top bar in Fig.3, the subsequent near-term temperature impact would be a
294 warming due to removal of cooling aerosols, adding to the warming of methane for the sector
295 as a whole. A significant fraction of methane emissions, originating from the production stage,
296 could be mitigated separately from most other SLCFs, with added benefits in terms of reduced
297 CO₂ and/or BC (P2, P3), resulting in a notable reduction in both the near- and long-term impact
298 of the sector. Similarly, policies for the agriculture sector can be designed to target different
299 sources addressing either primarily nitrogen losses (bringing air quality benefits but unmasking
300 nitrate cooling) or focusing on methane sources. However, unsurprisingly only policies with
301 strong methane reductions (here, P2) would give a significant change in the temperature impact
302 of the sector. The net impact of the shipping sector (SHP) is a cooling in the near-term, which
303 has been shown in several previous studies (e.g., Berntsen & Fuglestedt, 2008; Fuglestedt et
304 al., 2009). Policies that reduce shipping emissions of SO₂ and NO_x (P1) hence result in an
305 added near-term warming, also when simultaneous elimination of the sector’s CO₂ emissions
306 occur (P2). A hypothetical CO₂-only policy (P3) gives a net cooling on both time scales but
307 would fail to address the environmentally detrimental impacts of the sector pollution emissions.

308 This example is simplified but meant to illustrate the applicability of our dataset and how it
309 allows for detailed analyses without further use of complex models. Furthermore, while we here
310 calculate the temperature impact following a pulse of emissions, i.e., assuming that the policies



311 instantaneously affect the sectoral emission composition, our pulse-based emission metrics can
312 easily be used to study changes over time to any emission or policy scenario through
313 convolution (Aamaas et al., 2013), allowing for a broad potential for further use of our data (see
314 Sect. 2). In the next section, we use precisely this method to quantify the impact of temporally
315 evolving emissions according to the most recent set of scenarios.

316

317 **3.3 Contributions from SLCFs and CO₂ to global temperature change under the SSPs**

318 While knowledge of the present-day emission composition and net temperature impact over
319 time is essential to support mitigation design and implementation, real-world emissions will
320 evolve following a combination of socioeconomic developments, technological advancement
321 and policy adoption. Next, we investigate plausible pathways for the future impact of SLCFs
322 and CO₂ by quantifying the global temperature change over the period 1900-2100 to regional
323 and sectoral emissions following the SSPs. In the following paragraphs, we focus on four of the
324 nine SSPs (SSP1-1.9, SSP2-4.5, SSP3-7.0 and SSP5-8.5) that span the range of future emission
325 evolutions. See “Data Availability” for results from remaining scenarios. Figure 4 shows the
326 evolution of temperature response under the SSPs for our source regions, with corresponding
327 results for the global economic sectors given in Fig. S1.

328 Our emissions regions not only have large differences in terms of present-day emissions, but
329 also in past evolution. This historical contribution, which was not captured in the analysis of
330 the first half of the paper, brings NAM and EUR as the two largest contributors to the present-
331 day warming (Fig. 4a) due to their much higher past CO₂ emissions, in line with previous
332 literature (Höhne et al., 2011; Skeie et al., 2017). While presently being the largest emission
333 source, EAS only surpasses EUR and NAM in net temperature impact between 2020 and 2030
334 when the cumulative effect of CO₂ is accounted for. In SSP1-1.9, where emissions of CO₂
335 decline strongly during the first half of the century in all regions, the net temperature response
336 levels off or starts to decline in the second half of the century. We note that negative CO₂
337 emissions are not included in these calculations. In the remaining scenarios, the net temperature
338 impact increases over the century for all regions. EAS remains the largest contributor, whereas
339 in SSP5-8.5 SAS overtakes NAM as the second most important region by 2100 and SAF
340 reaches the same order of magnitude as EUR.

341 Globally, the net temperature response following emissions from the ENE sector becomes
342 larger than that due to AGR and RES in the early 2000s under this emission evolution (Fig.
343 S1a), and ENE remains the largest individual sector until 2100 in all scenarios. The global mean
344 temperature impact of IND switches from a net cooling to a net warming in the late 20th century
345 as the warming due to CO₂ accumulates and overwhelms the cooling from SO₂.

346 While the contribution from CO₂ to the net warming becomes dominant by 2100 for most
347 regions and sectors under all SSP scenarios, the relative importance of SLCFs and CO₂ continue
348 to be highly variable across emission source over time, in particular under SSP3-7.0 and SSP5-
349 8.5. This can be seen in Fig.4b, where we break down the future net temperature response in
350 2030, 2050 and 2100 into individual contributions from methane, CO₂, BC and the sum of SO₂



351 and NO_x. Here we show a selection of the source regions that differ notably in composition and
352 temporal trend. See Fig. S2 for remaining regions and Fig.S1b for breakdown by global sector.

353 The SSPs differ in both climate forcing targets and stringency of air pollution control, as well
354 as underlying socioeconomic development. SSP1-1.9 is characterized by low societal
355 challenges to mitigation and adaptation, and strong climate and air quality policies, resulting in
356 rapidly declining emissions of both SLCFs and CO₂. However, even for strong air pollution
357 there is a differentiation between high-, medium- and low-income countries, with a substantial
358 time lag in the latter two (Rao et al., 2017). For example, emissions of SO₂ in SAS and SAF
359 decline less than in other regions, subsequently maintaining a significant cooling contribution
360 to the temperature change. In the intermediate scenario, SSP2-4.5, there is a reduction in
361 emissions, but this is delayed and slower compared to SSP1-1.9. In SSP3-7.0, the world follows
362 a path with more inequality and conflict, where only weak air pollution control is implemented
363 and the end-of-century climate forcing, and hence CO₂ emissions, is higher. Subsequently,
364 emission trends and SLCF contributions display more regional heterogeneity. There is a
365 particularly strong projected increase in methane emission in South Asia, Africa and South
366 America in this scenario. While EAS remains the region with the largest warming impact by
367 2100 in all scenarios, the contributions to warming from methane and BC in SAF and SAS
368 surpasses those of EAS in 2100 in both SSP3-7.0 and SSP5-8.5. As CO₂ emissions increase,
369 the net temperature response to emissions in SAS increases from close to zero to a significant
370 warming. SSP5-8.5 is characterized by high challenges to mitigation and high climate forcing
371 in 2100, but strong air pollution control since high use of fossil fuels would otherwise result in
372 unbearable air pollution levels. Combined, this leads to increasing temperature impact due to
373 increasing CO₂ emissions, but lower SLCF impacts than in SSP3-7.0, but with a non-negligible
374 contribution from methane for several regions. Hence, in medium- and low-income regions,
375 SLCFs, and in particular methane, are projected to play a continued important role for future
376 temperature change. Or put another way, the potential for climate mitigation highlighted in
377 Fig.2 is only realized in SSP1-1.9.

378 Clearly, and as expected, the largest difference in SLCF contributions to temperature response
379 is between SSP1-1.9 and SSP3-7.0. To see where the largest additional climatic benefit can be
380 gained from moving from an SSP3-7.0 world to one in line with SSP1-1.9, we show the
381 difference in temperature between these two scenarios in 2030, 2050 and 2100 in Fig.5. Results
382 are shown by region and sector, for all combinations where the temperature difference is greater
383 than $\pm 0.01^{\circ}\text{C}$. Our results emphasize the importance, for both near- and long-term climate
384 change, of the strong sources of methane; agriculture, energy and waste management, especially
385 in Africa, South Asia and South America. Fig. 5. also shows how the strong SLCF mitigation in
386 SSP1-1.9, relative to SSP3-7.0, results in a net warming contribution to climate for some region-
387 sector combinations, such as industry in East and South Asia. As shown by the panel on the
388 right-hand side of Fig. 5, for most sector/region combinations, around 10% of the avoided (or
389 added) warming from strong mitigation would be realized already by 2030, and around 40-50%
390 by 2050.

391

392 **4 Discussion**



393 In terms of avoided global warming, there is much to be gained by moving from a global
394 emission pathway following SSP3-7.0 to one following SSP1-1.9, including contributions from
395 reductions of SLCFs, as discussed above. Such rapid reductions of air pollutants' emissions are
396 technically possible drawing on experience in both developed and developing countries (Crippa
397 et al., 2016; Kanaya et al., 2019; Klimont et al., 2017) but would require simultaneous
398 strengthening of institutions to enforce the laws. The focus of such policies would differ
399 between OECD countries and the developing world. Further measures in the OECD would
400 primarily focus on reducing emissions from residential heating, non-road transportation, and
401 agriculture while assuring enforcement of legislation in power and industry. For methane,
402 reducing venting and increasing utilization of associated petroleum gas in oil and gas
403 exploration, increased use of biogas from waste, as well as addressing agriculture emissions
404 should be a priority, and the technical potential for considerable reductions until 2050 exists
405 (Höglund-Isaksson et al., 2020). A recent study suggests that anthropogenic fossil methane
406 emissions may be significantly underestimated (Hmiel et al., 2020), and as such, reductions
407 may be even more critical. The rapidly industrializing and developing countries would need to
408 further strengthen legislation for the power, industry, transport sectors, introduce new laws to
409 improve waste management, reduce emissions from agriculture, and provide wide access to
410 clean fuels securing cooking and heating needs. Several of these policies would secure
411 achievements of SDG goals (Rafaj et al., 2018). For methane, similar suite of measures is
412 needed as for the developed world although waste management requires larger transformation
413 and there is additional significant potential to reduce emissions from coal mining sector.
414 Specific measures for improving air quality while contributing to climate change mitigation
415 have recently been assessed for South East Asia (UNEP, 2019) and Latin America (UNEP,
416 2018). While previous decades have seen a southeastward shift in air pollution emissions, from
417 high income regions at northern latitudes to East and South Asia, recent trends and the SSPs
418 suggest that a second shift may be underway, where, as shown above, contributions from SLCFs
419 to temperature change increase in the Middle East and Africa. An increasing carbonization in
420 Africa south of the Sahara, primarily due to the increasing use of oil in the transport sector, has
421 already been observed (Steckel et al., 2019), highlighting the need for further focus on this
422 region.

423

424 SSP3-7.0 and SSP1-1.9 not only differ in the stringency of the assumed air pollution control,
425 but also in socioeconomic development and end-of-century climate forcing. To isolate the role
426 of air pollution policies in the transition to a low warming pathway, a companion scenario to
427 SSP3-7.0 has been developed, the SSP3-lowNTCF (Gidden et al., 2019). Here, the
428 socioeconomic narrative is the same, but emission factors for the short-lived species are
429 assumed to be in line with those in SSP1-1.9. The result is similar global CO₂ emission but up
430 to 60% reductions in global SLCF emissions in SSP3-lowNTCF relative to SSP3-7.0. Using
431 the SSP3-lowNTCF emissions as input, we find that this in turn leads to a net temperature
432 response to total global emissions in 2100 that is 13% lower in SSP3-LowNTCF than in SSP3-
433 7.0 (an absolute difference of 0.5°C, from 3.7°C to 3.2°C in our calculations). For comparison
434 the net temperature response is 71% (or 2.6°C) lower in SSP1-1.9 compared to SSP3-7.0.

435



436 The potential for near-term mitigation by targeting BC emissions in the transport and residential
437 sectors has been highlighted earlier (e.g., UNEP, 2011). We also find notable BC contributions
438 from the residential sector in some regions, mainly South Asia and Africa, but estimate quite
439 low BC effects from the transport sector. This has three main reasons. Firstly, since earlier
440 studies (done about 10 years ago) there have been significant changes in legislation, and new
441 diesel trucks and cars are (in several regions) equipped with particulate filters removing
442 effectively BC. By now these vehicles represent a significant part of the fleet in many regions
443 and the trend is expected to continue. Secondly, by accounting for the rapid adjustments
444 associated with BC (Stjern et al., 2017), which reduces the net warming climate impact of the
445 aerosols, we estimate a lower temperature response than earlier literature. Finally, we account
446 for cooling from nitrate aerosols from emissions of NO_x, for which the transport sector is a
447 significant source, even in regions where stricter vehicle emission standards (e.g., Euro 5) have
448 been adopted.

449 The AGTP is a well-established framework that has been applied in several studies of
450 attribution of temperature impacts to emission sources (e.g., Lund et al., 2017; Sand et al.,
451 2015; Stohl et al., 2015). It allows us to investigate the effects of individual species, sources
452 and scenarios, which would be confounded by the low signal-to-noise ratio in fully coupled
453 models, but also introduces caveats. Importantly, the AGTP metric is linear and does not
454 include saturation effects as emissions and atmospheric concentrations increase. We emphasize
455 that the absolute magnitude of temperature changes should therefore be interpreted with care,
456 as this method is primarily designed to study relative importance and relationships between
457 individual emissions and sources. Our analysis reflects the best estimate input data, but results
458 have considerable uncertainty. As shown in Fig. 2a, we estimate a 1 standard deviation range
459 in the total net temperature response on the 10-year time horizon of $\pm 0.01^{\circ}\text{C}$, about 38% of the
460 net temperature response of 0.03°C (the range is considerably lower on the 100-year time scale
461 as the RF of SLCFs is much more uncertain than that of CO₂). This excludes uncertainties in
462 emissions and climate sensitivity. Uncertainties in emission inventories are difficult to quantify,
463 but generally considered lowest for CO₂ and SO₂ emissions, and high for carbonaceous aerosols
464 (Hoesly et al., 2018). Moreover, recent studies point to emission trends that are not accurately
465 represented in the global inventory, such as SO₂ and NO_x in China (Zheng et al., 2018) and
466 fossil fuel CH₄ emissions (Hmiel et al., 2020). However, due to high spatiotemporal variability
467 and lack of consistent data, a comprehensive uncertainty analysis at the regional and sectoral
468 level is challenging. The impulse response function (IRF) used in the present analysis yields an
469 equilibrium climate sensitivity (ECS) of $0.885\text{ K (Wm}^{-2}\text{)}^{-1}$, which is in the upper range reported
470 by Bindoff et al. (2013), but lower than many recent estimates (Forster et al., 2019). While the
471 former has a spatiotemporal dependence, changes in the ECS mostly act to scale estimates for
472 all sectors and regions but is less important for their ranking. Furthermore, our analysis is
473 limited to temperature change as a measure of climate impacts. SLCFs, and in particular
474 aerosols, also play a key role in shaping local and regional hydrology and dynamics. Comparing
475 the SSP3-7.0 and SSP3-lowNTCF scenarios, Allen et al. (2020) recently found a significant
476 precipitation increase due to removal of aerosols, with the strongest moistening trends over
477 Asia. An increase in the Asian summer monsoon precipitation in scenarios with strong air
478 pollution reductions was also recently found by Wilcox et al. (2020). Hence, further studies



479 using coupled models are needed to fully capture the effects of the SLCFs under SSPs on local
480 climate and environment.

481

482 **5 Conclusions**

483 Complimentary mitigation of CO₂ and other LLGHG with SLCFs is of key importance for
484 achieving the climate ambitions of the Paris Agreement and meeting the Sustainable
485 Development Goals. Here we show that there is significant potential for mitigation of near- and
486 long-term temperature change, but also possible trade-offs, inherent in the present-day
487 emissions from the major source regions and economic sectors. In terms of contributions from
488 SLCFs, we reinforce the importance of the major emitters of methane, in particular agriculture
489 and waste management, but also energy production, for reducing near-term warming. In
490 contrast to the existing potential, we find that SLCFs are projected to continue to play an
491 important role for global temperature change over the 21st century under most of the Shared
492 Socioeconomic Pathway (SSP) scenarios. Several of the SSPs project a particularly strong
493 increase in emissions in Africa south of the Sahara. In addition to the focus on South and East
494 Asia as the major current sources of SLCFs, enabling technological development and legislation
495 implementation on the African continent may be of key importance for a transition from high
496 air pollution SSP3-7.0 pathway towards one in line with SSP1-1.9, which in turn would add
497 reduction in global warming already over the next couple of decades. The large spatiotemporal
498 heterogeneity in emissions trends and subsequent temperature responses underlines the need to
499 go beyond global emission scenarios. By assessing the global temperature response to
500 emissions from 13 regions, 7 sectors and 4 scenarios we provide a more comprehensive dataset
501 than, to our knowledge, currently exists, enabling further analysis of mitigation strategies and
502 economic analyses at a detailed level.

503

504

505 **Data availability**

506 All output data is publicly available via Figshare
507 (<https://doi.org/10.6084/m9.figshare.11386455>)

508

509 **Author contributions**

510 Lund led the study, prepared the input data and wrote the paper. Aamaas performed the emission
511 metric and uncertainty calculations. Stjern and Samset produced the graphics. Klimont and
512 Berntsen contributed to the design of the analysis. All authors contributed to the manuscript
513 preparation.

514

515 **Competing interests**

516 The authors declare that they have no competing interests.



517

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521

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768 **Tables:**

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770 *Table 1: Constants of the Geoffroy et al. (2013) IRF.*

	Mode 1	Mode 2
c_j	0.587	0.413
d_j (years)	4.1	249

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773 *Table 2: Summary of policies and species reduced.*

Sector	Policy 1 (P1)	Policy 2 (P2)	Policy 3 (P3)
ENE	End-pipe measures	Reducing losses in oil and gas and flaring	Capture methane in coal mining
	SO ₂ , NO _x	CH ₄ , BC, CO ₂	CH ₄ , CO ₂
AGR	Nitrogen use efficiency	Meat reduction	Increase in biogas use
	NH ₃ , NO _x	CH ₄ , NH ₃ , NO _x	NH ₃ , CO ₂
SHP	Scrubbers and particulate filters	Slow-steaming	CO ₂ -only policy
	SO ₂ , NO _x , BC	CO ₂ , SO ₂ , NO _x , BC	CO ₂

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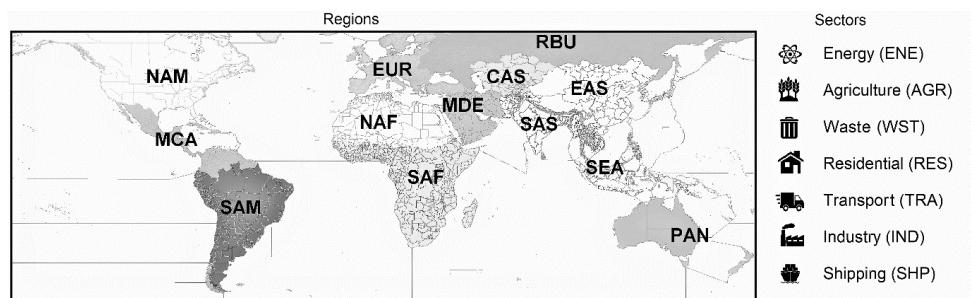
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786 **Figures:**

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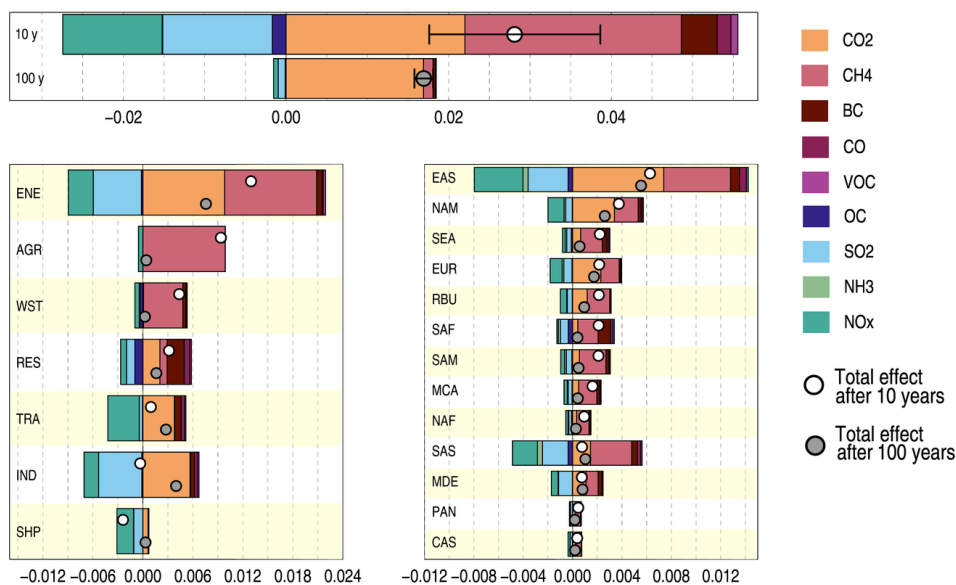


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789 *Figure 1: Emission source regions and sectors used in the analysis.*

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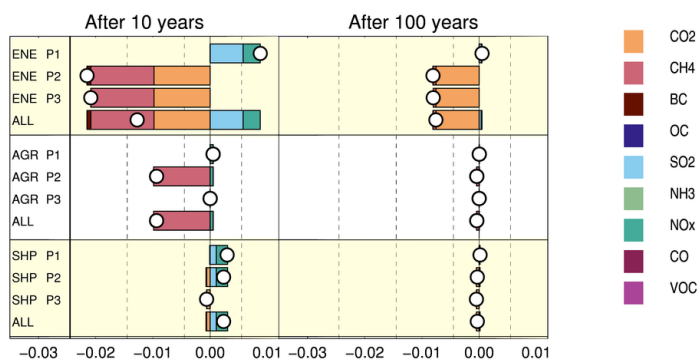


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793 *Figure 2: Global-mean surface temperature impact 10 and 100 years after one year of present-day (i.e., year 2014) emissions of SLCFs and CO₂ for: a) global total emissions, b) emissions from seven major economic sectors, and c) total (i.e., sum of all sectors) emissions in 13 sources regions. Panels b and c are sorted by total net effect on the 10-year timescale (white circle). Error bars (± 1 standard deviation) in the top panel represent the range in total net temperature impact due to uncertainties in radiative forcing.*

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802 *Figure 3: Global-mean surface temperature impact after 10 and 100 years resulting from*
 803 *instantaneous reductions of different sets of SLCFs and CO₂ emissions under three different*
 804 *policies, as well as for these three combined. White circles indicate the net impact.*

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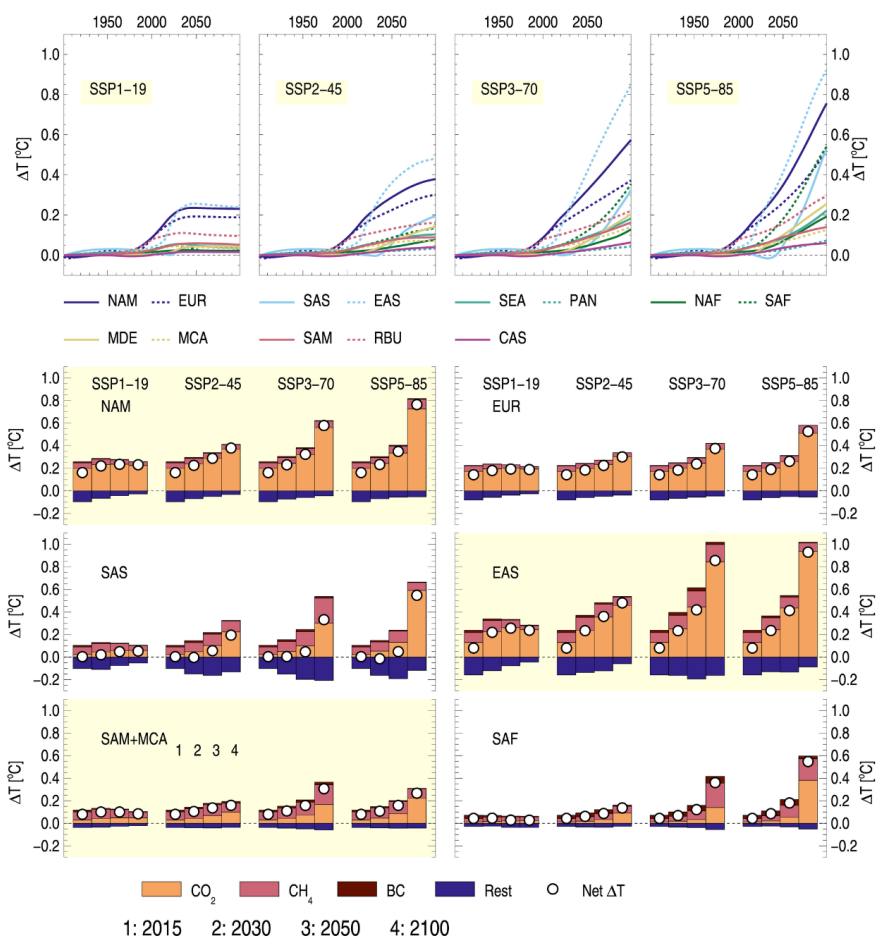
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815 *Figure 4: Global mean temperature response to historical emissions and future SSP pathways:*
 816 *a) Net (i.e., sum over all species and sectors) response over the period 1900 to 2100 for each*
 817 *region and scenario and b) net response in 2015, 2030, 2050 and 2100 to emissions in six*
 818 *regions broken down by contributions from CO₂, BC, methane and the sum of SO₂, OC, NH₃*
 819 *and ozone precursors (i.e., “Rest”).*

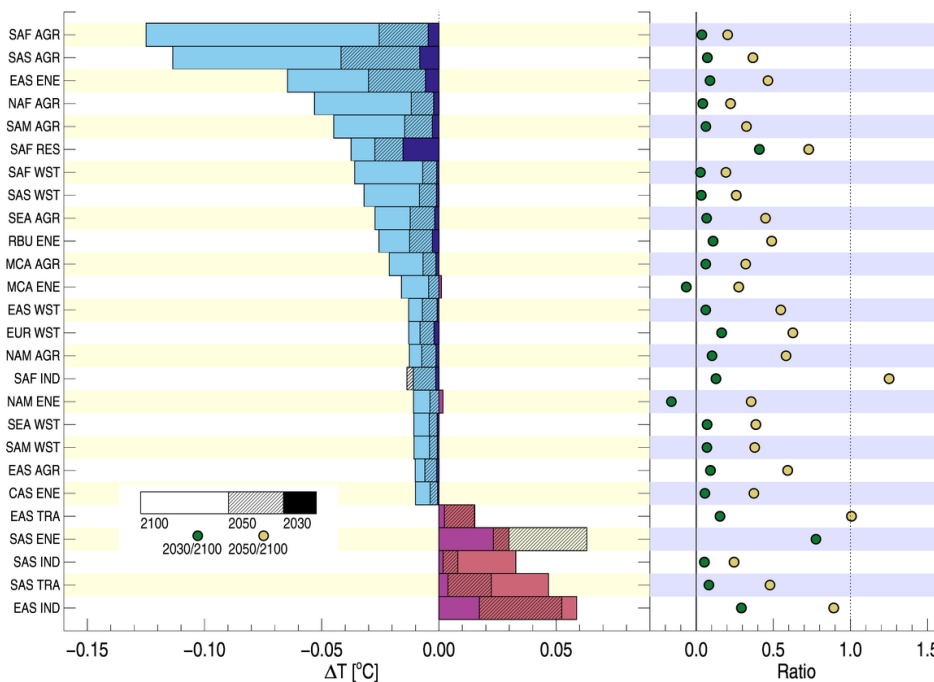
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826 *Figure 5: Difference in net SLCF (i.e., sum of all components except CO₂) temperature*
 827 *response between SSP1-1.9 and SSP3-7.0 in 2030, 2050 and 2100 by region and sector. Only*
 828 *combinations of sectors and regions where the differences in global temperature response is*
 829 *larger than ±0.01 °C are shown. For each of these combinations, the panel on the right shows*
 830 *the ratio between the temperature response difference in 2030 and 2100 and between 2050 and*
 831 *2100.*

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