

1 **A continued role of Short-Lived Climate Forcers under the Shared Socioeconomic**
2 **Pathways**

3 *Marianne T. Lund^{1*}, Borgar Aamaas¹, Camilla W. Stjern¹, Zbigniew Klimont², Terje K.*
4 *Berntsen^{1,3}, Bjørn H. Samset¹*

5

6 *1 CICERO, Center for International Climate Research, Oslo, Norway*

7 *2 International Institute for Applied Systems Analysis (IIASA), Laxenburg, Austria*

8 *3 Department of Geosciences, University of Oslo, Oslo, Norway*

9 **Corresponding author: m.t.lund@cicero.oslo.no*

10

11

12

13

14

15

16

17

18

19

20

21

22

23

24

25

26

27

28

29

30

31

32 **Abstract**

33 Mitigation of non-CO₂ emissions plays a key role in meeting the Paris Agreement ambitions
34 and Sustainable Development Goals. Implementation of respective policies addressing these
35 targets mainly occur at sectoral and regional levels and designing efficient mitigation strategies
36 therefore relies on detailed knowledge about the mix of emissions from individual sources and
37 their subsequent climate impact. Here we present a comprehensive dataset of near- and long-
38 term global temperature responses to emissions of CO₂ and individual short-lived climate
39 forcers (SLCFs) from 7 sectors and 13 regions - for present-day emissions and their continued
40 evolution as projected under the Shared Socioeconomic Pathways. We demonstrate the key role
41 of CO₂ in driving both near- and long-term warming, and highlight the importance of mitigating
42 methane emissions, from agriculture, waste management and energy productions, as the
43 primary strategy to further limit near-term warming. Due to high current emissions of cooling
44 SLCFs, policies targeting end-of-pipe energy sector emissions may result in net added warming
45 unless accompanied by simultaneous methane and/or CO₂ reductions. We find that SLCFs are
46 projected to play a continued role in many regions, particularly those including low- to medium-
47 income countries, under most of the SSPs considered here. East Asia, North America and
48 Europe remain the largest contributors to total net warming until 2100, regardless of scenario,
49 while South Asia and Africa south of the Sahara overtake Europe by the end of the century in
50 SSP3-7.0 and SSP5-8.5. Our dataset is made available in an accessible format, aiming also at
51 decision-makers, to support further assessments of the implications of policy implementation
52 at the sectoral and regional scales.

53

54

55

56

57

58

59

60

61

62

63

64

65

66

67

68 **1 Introduction**

69 At the core of any strategy for sustained, long-term abatement of climate change are strong
70 reductions in emissions of CO₂ and other long-lived greenhouse gases (LLGHGs). However,
71 most anthropogenic activities emit a suite of additional species, with a range of climate impacts,
72 commonly termed short-lived climate forcers (SLCFs). While differing in characteristics and
73 contribution to temperature change, their common feature of a much shorter atmospheric
74 residence time compared to LLGHGs has resulted in significant discussion of the role of SLCF
75 mitigation in strategies to reduce climate change, in particular to limit near-term warming
76 (e.g., Bowerman et al., 2013; Pierrehumbert, 2014; Rogelj et al., 2015; Shindell et al., 2012;
77 Shoemaker et al., 2013; Stohl et al., 2015).

78 Many assessments have placed particular emphasis on the subset of SLCFs with a warming
79 impact on climate, namely black carbon (BC), methane (CH₄) and tropospheric ozone
80 (sometimes collectively referred to as short-lived climate pollutants, or SLCPs) (e.g., AMAP,
81 2015; CCAC, 2019; UNEP, 2017). Assuming effective abatement of SLCPs, some studies
82 estimate a reduction in global temperature increase of 0.2-0.5°C by mid-century (e.g., Shindell
83 et al., 2012). More recent work suggest that some of these early estimates may overestimate
84 the effect of SLCP mitigation (Rogelj et al., 2014; Smith & Mizrahi, 2013; Stohl et al., 2015;
85 Takemura & Suzuki, 2019). While results from early studies brought some concern that the
86 attractiveness of SLCP mitigation could lead to delayed action on CO₂ emissions, most
87 scientific studies emphasize that SLCP measures should only be considered complementary to
88 early and stringent CO₂ mitigation for the achievement of long-term climate goals (Ramanathan
89 & Carmichael, 2008; Rogelj et al., 2014).

90 SLCF mitigation may also give rise to potential trade-offs. Due to co-emission, any given
91 mitigation measure or policy can affect a broad range of species. The combinations may,
92 however, vary significantly between sources and mitigation strategies motivated by, and
93 designed to address, different societal challenges. For instance, many SLCFs are tightly linked
94 to air quality (Anenberg et al., 2012; Lelieveld et al., 2015; Shindell et al., 2012) and sustainable
95 development (Haines et al., 2017; UNEP, 2019), in addition to their climate impacts. The
96 numerous environmental and societal co-benefits of SLCF reductions are well recognized but
97 may lead to adverse climatic consequences (Arneth et al., 2009). While some SLCFs with a
98 warming contribution to temperature change can, in part, be mitigated individually (in particular
99 methane), improving air quality requires consideration of all relevant species. Removal of all
100 present-day anthropogenic aerosols may add as much as 0.5°C of additional global near-term
101 warming according to recent work (Hienola et al., 2018; Samset et al., 2018; Aamaas et al.,
102 2019). Due to co-emission, species such as sulfur dioxide (SO₂) are also commonly affected by
103 measures to reduce climate warming even if these have LLGHGs as the primary target. Hence,
104 while it remains clear that deep reductions in emissions of methane and BC play a key role in
105 pathways for global emissions that limit global warming to 1.5°C and 2°C warming (Harmsen
106 et al., 2019; Rogelj et al., 2015; Rogelj et al., 2018; Shindell & Smith, 2019; Xu & Ramanathan,
107 2017), co-emitted species such as sulfate need to be carefully considered.

108 A key characteristic of SLCFs is that the composition of emissions, as well as their subsequent
109 radiative forcing, can vary significantly between individual emission sources (Bond et al., 2013;

110 Lund et al., 2014b; Persad & Caldeira, 2018; Unger et al., 2010). While previous scenarios for
111 long-term evolution of aerosols and ozone precursor emissions projected a general, rapid
112 decline even in pathways with high climate forcing and GHG levels (Gidden et al., 2019; Rao
113 et al., 2017), the most recent generation scenarios, the Shared Socioeconomic Pathways (SSPs)
114 (O'Neill et al., 2014; Riahi et al., 2017), exhibit a much larger spatiotemporal heterogeneity in
115 projections of these emissions. Additionally, the SSPs provide a framework for combining
116 future climate scenarios (Representative Concentration Pathways – RCPs) with socioeconomic
117 development, and hence more detailed information about plausible future evolutions of society
118 and natural systems. Up-to-date and detailed knowledge of the climate impact of individual
119 emission sources is critical for the design of effective mitigation strategies and to provide
120 decision makers with more integrated guidance on how to best address linkages between
121 climate, sustainable development and air quality in policy processes (Melamed et al., 2016).
122 While studies comparing and quantifying the impacts of SLCFs and CO₂ exist, they differ in
123 selection of sectors and/or regions, methodology and emission inventory, making direct
124 comparison difficult (e.g., Harmsen et al., 2019; Kupiainen et al., 2019; Lund et al., 2014a;
125 Sand et al., 2015; Unger et al., 2010). Furthermore, studies often consider only the equilibrium
126 effect of present-day emissions, emission pulses or very simplified scenarios.

127 In the present work, we provide a comprehensive and updated investigation of the contribution
128 to near- and long-term global temperature impacts from individual SLCF and LLGHG
129 emissions. We first quantify the temperature response to an idealized pulse of present-day
130 emissions to demonstrate the methodology and temporal behavior of the various emitted
131 species, focusing on both added benefits and trade-offs offered by SLCF mitigation. Then we
132 calculate the future evolutions of temperature impacts as they are projected to develop under
133 the pathways for future socioeconomic development, climate policy and air pollution described
134 by the SSP-RCP scenarios. The temperature impact is calculated for seven economic sectors
135 and 13 source regions, accounting for best available knowledge and geographical dependence
136 of the forcing efficacy of different SLCFs, thereby providing a more detailed breakdown than
137 previous literature. By making our full data set openly available, we aim to provide a toolkit for
138 further studies of the implications of policy implementation at the sectoral and regional level,
139 demonstrating the potential for such applications for a set of idealized sectoral emission
140 reduction packages.

141

142 **2 Methodology**

143 Using the concept of Absolute Global Temperature change Potential (AGTP) (Shine et al.,
144 2005), we calculate the global-mean temperature response over time to emissions of CO₂, CH₄,
145 ammonia (NH₃), BC, OC, SO₂, the ozone precursors nitrogen oxide (NO_x), carbon monoxide
146 (CO) and volatile organic compounds (VOCs) from 7 sectors and 13 regions (Fig. 1).

147 2.1 Calculations of global and regional AGTPs

148 The AGTP is an emission metric-based emulator of the climate response, and a well-established
149 method that enables us to quantify and compare global temperature impacts of a large number
150 of sources and scenarios in a transparent and, in terms of computer resources, cost-effective

151 manner. The approach is described in detail in the literature (Fuglestedt et al., 2010; Shine et
152 al., 2005; Aamaas et al., 2013); here we give a brief outline.

153 The ATGP gives the global-mean surface temperature response per kg species emitted as a
154 function of time after an emission pulse, i.e., an instantaneous one-off emission. At time H after
155 the emission, the AGTP for species i is given (for each sector and region) by:

$$156 \quad AGTP_i(H) = \int_{t=0}^H F_i(t) IRF_T(H - t) dt \quad (1)$$

157 where F_i is the radiative efficiency. Emissions of SLCFs can have both direct and indirect
158 radiative effects. For BC, OC and SO₂ we account for the direct, semi-direct and indirect RF as
159 described below. AGTPs for NO_x, CO and VOC includes the forcing due to tropospheric ozone
160 production and (for NO_x) nitrate aerosol formation, as well as the longer-term effect on
161 methane lifetime and methane-induced ozone loss. The AGTP for methane includes the direct
162 forcing, as well as the effect of OH-induced changes in its lifetime and effects on tropospheric
163 ozone and stratospheric water vapor. See Aamaas et al. (2013) for details and analytical
164 expressions for the AGTP of individual species.

165
166 For CO₂ and methane, we calculate the global-mean F for year 2014 global concentrations (i.e.,
167 the year that is considered present-day in our emissions data – see below) using the equations
168 from Etminan et al. (2016). Compared to the approach used on the IPCC Fifth Assessment
169 report (AR5) (Myhre et al., 2013), this increases the radiative efficiency of methane by 14%.
170 For NH₃, we use the IPCC AR5 best estimate for global mean radiative efficiency for all
171 regions. For the remaining short-lived species, we derive values of F_i that depend on the
172 location of the emission and calculate region-specific AGTPs for BC, OC, SO₂, and the ozone
173 precursors. The regional radiative efficiencies (i.e., the global radiative forcing per unit of
174 regional emissions) for BC, OC, sulfate, nitrate and ozone (in response to NO_x, CO and VOC)
175 are derived using radiative kernels (Samset & Myhre, 2011) and atmospheric concentrations
176 from simulations performed with the global chemistry transport model OsloCTM3 (Søvde et
177 al., 2012) for the second phase of the Hemispheric Transport of Air Pollution (HTAP2)
178 (Janssens-Maenhout et al., 2015). Details about the chemistry and aerosol parameterizations
179 and properties can be found in Lund et al. (2018). In addition to their direct radiative effects,
180 aerosols also affect the energy balance through modifications of clouds and atmospheric heating
181 rates (indirect and semi-direct effects). To account for the additional negative RF resulting from
182 aerosol-cloud interactions, we scale the AGTP of SO₂ by a factor of 2.1 based on the ratio of
183 total global RF of sulfate to that due to direct effects alone from the IPCC AR5 (Myhre et al.,
184 2013). Due to lack of available information about geographical dependence of the radiative
185 efficiency, the same scaling factor is applied for all regions, recognizing that this is a
186 simplification as the indirect effect also likely varies with location of emission. We also account
187 for the semi-direct effect of BC (i.e., the rapid adjustments of the atmosphere to the local heating
188 (Smith et al., 2018)). Here we use the multi-model ratio between semi-direct and direct BC RF
189 from Stjern et al. (2017) and calculate an average adjustment factor for the rapid adjustments
190 of -15%. This is then applied to the AGTP of BC for all regions, except South Africa where
191 Stjern et al. (2017) found a small positive forcing from rapid adjustments. Radiative forcing of
192 BC deposition on snow and ice is not included in our estimates.

193

194 IRF_T in Eq.1 is the impulse response function used to estimate the temperature response to a
195 given radiative forcing:

$$196 \quad IRF_T(t) = \lambda \sum_{j=1}^J \frac{c_j}{d_j} \exp\left(-\frac{t}{d_j}\right) \quad (2)$$

197 where c_j and d_j are constants and timescales of the fast and slow model of the climate system
198 response, respectively, and λ is the equilibrium climate sensitivity (ECS). An IRF is also used
199 to represent the atmospheric decay of CO_2 . Several different IRFs exist in the literature. Here
200 we use the IRF_T from Geoffroy et al. (2013) (G13) and the IRF_{CO_2} from Joos et al. (2013).
201 Values of c_j , d_j and λ derived from the analytical solution of the two-layer energy balance model
202 used by G13 are given in Table 1. Compared to the IRF_T from Boucher and Reddy (2008)
203 (B&R08) used in the bulk of previous metrics studies including IPCC AR5, G13 has shorter
204 timescales and yields a lower ECS ($0.885 \text{ K (Wm}^{-2}\text{)}^{-1}$) compared to $1.06 \text{ K (Wm}^{-2}\text{)}^{-1}$) from
205 B&R08. To place our values in the context of previous literature and explore sensitivities to the
206 choice of IRFs, we perform additional calculations using different combinations of IRF_T and
207 IRF_{CO_2} – see section Sect. 1 of the Supplementary Information (SI).

208 Finally, we consistently account for the climate-carbon feedback (CCf) in the AGTPs. The
209 IRF_{CO_2} , derived from complex models, implicitly includes the CCf. However, this is not the
210 case for other components. This inconsistency was first highlighted in Myhre et al. (2013),
211 where a first attempt to include the CCf was made for halocarbons based on an earlier study by
212 Collins et al. (2013). This method has since been refined. Here we use the framework developed
213 by Gasser et al. (2017) where a separate IRF for the CCf was derived using the simple Earth
214 system model OSCARv2.2. This IRF is used to calculate a $\Delta AGTP_i(H)$ which is then added to
215 the $AGTP_i(H)$ without CCf. The difference between this method and the approach taken by
216 Myhre et al. (2013) is discussed in Gasser et al. (2017). We also perform a sensitivity test to
217 quantify the impact on our estimated temperature responses of excluding the CCf – see Sect.
218 4.1. Furthermore, as different methods to account for the CCf exist in the literature, we provide
219 both sets of AGTPs for further use (see “Data Availability”).

220

221 2.2 Emission data and temperature response calculations

222 As described above, we investigate the role and global temperature impacts of SLCF and CO_2
223 from two different perspectives. First, the AGTPs at two given time horizons H (here 10 and
224 100 years) are multiplied by year 2014 emissions from the Community Emission Data System
225 (CEDS) (Hoesly et al., 2018) for each species, sector and region. The result is the near- and
226 long-term global temperature response, $\Delta T_i(H)$, to present-day regional and sectoral emissions.

227 Next, we quantify the temperature response to temporally evolving emissions from 1900 to
228 2100. The AGTP framework can readily be extended from pulse-based calculations since any
229 scenario can be viewed as a series of pulse emissions and analyzed through convolution
230 (Aamaas et al., 2013). The temperature response ΔT at time t for species i is (for each region
231 and sector) given by:

232 $\Delta T_i(t) = \int_0^t E_i(t')AGTP_i(t - t')dt'$

233 Importantly, the AGTPs are linear in that they do not account for the potential changes in
234 radiative efficiency with changing background pollution levels – see Sect. 4 for further
235 discussion.

236 Historical emissions are from the CEDS database, while future emissions follow the SSP-RCP
237 scenarios. Gridded and harmonized emissions are available via ESFG from the Integrated
238 Assessment Modeling Community (IAMC) for nine SSP-RCP combinations that form the core
239 of the Coupled Model Intercomparison Project Phase 6 (CMIP6) experiments (Gidden et al.,
240 2019): SSP1-1.9, SSP1-2.6, SSP2-4.5, SSP3-7.0, SSP3-7.0 lowNTCF, SSP4-3.4, SSP4-6.0,
241 SSP5-3.4, and SSP5-8.5. The gridded SSP-RCP data product, including the methodology for
242 country and sector level emission mapping, is documented by Feng et al. (2020). We extract
243 regional emission scenarios using the geographical definitions and spatial mask from HTAP2
244 (Janssens-Maenhout et al., 2015). Furthermore, we consider the energy (ENE), agriculture
245 (AGR), waste (WST), residential (RES), industry plus solvents (IND), transport (TRA) and
246 shipping (SHP) sectors, as they are defined in the CEDS-SSP inventory (Feng et al., 2020;
247 Hoesly et al., 2018). Due to the large spread in historical estimates and lack of emissions
248 consistent with CEDS, we do not include CO₂ emissions due to land-use/land cover.
249 Additionally, agricultural waste burning is excluded as these are more difficult to mitigate and
250 estimates of future CO₂ emissions are not available.

251

252 2.3 Uncertainties

253 We establish a range in total net global-mean temperature response on 10- and 100-year time
254 scales due to uncertainties in radiative forcing by performing a Monte Carlo analysis. Each RF
255 mechanism is treated as a random variable, following a probability density function (PDF)
256 defined based on existing literature, and the distribution for the total RF is derived by summing
257 the individual PDFs, i.e., assuming that each RF mechanisms is independent. For the aerosols
258 and their precursors, we use the multi-model results from the AeroCom Phase II experiment
259 (Myhre et al., 2013a), while for CO₂ NH₃, and ozone precursors, we use the uncertainties from
260 the IPCC AR5 (Myhre et al., 2013b). For further details, see Aamaas et al. (2019) and Lund et
261 al. (2017). Our temperature responses are also influenced by uncertainties in emissions and
262 climate sensitivity. A comprehensive analysis of uncertainty in all three factors is challenging
263 due to lack of data, but the potential impact is discussed in Sect. 4.

264

265 3 Results

266 3.1 Near- and long-term temperature response to current emissions

267 We first discuss the global mean surface temperature response to one year of present-day (i.e.,
268 year 2014) emissions, for global total emissions and broken down by key contributing sectors
269 and geographical source regions as shown in Fig.2. While we here select 10- and 100-year time
270 horizons to represent near- and long-term impacts, we recognize that other choices may affect

271 the relative importance, and even sign, of the temperature response from some of the SLCFs or
272 be more relevant for certain applications. For this reason, we also provide the full time series
273 of our AGTPs (see Data Availability).

274 Globally, current emissions result in an approximate balance between cooling and warming
275 SLCFs in the near-term, with main warming contributions from BC and CH₄ and cooling from
276 SO₂ and NO_x (Fig.2a). The total net effect after 10 years is therefore only slightly larger than
277 that due to CO₂ alone. As the impact of the SLCFs decays over years to decades upon emission,
278 the total net temperature impact after 100 years is predominantly determined by CO₂. As clearly
279 seen in Fig. 2a, CO₂ emissions also cause a notable contribution to near-term warming. While
280 both of these features are well known in the scientific community, the role of CO₂ as driver also
281 of near- term warming is not always fully acknowledged in the discussions of LLGHGs versus
282 SLCFs.

283 Differences in the mix of emissions result in net impacts on global temperature that vary
284 significantly, in both magnitude and sign, between sectors and regions. Of the economic sectors,
285 energy (ENE), agriculture (AGR), and waste management (WST) give the largest net near-term
286 warming (i.e., after 10 years) (Fig. 2b). For AGR and WST, this is a result of strong methane-
287 induced warming. The energy sector (ENE) is also characterized by a significant warming due
288 to methane (originating from fossil fuel mining and distribution), as well as CO₂, but also by a
289 considerable cooling from high emissions of SO₂. Our results hence reinforce the importance
290 of methane as a driver of near-term warming but show that the net effect on global temperature
291 of SLCF mitigation may be small in the case of the energy sector if simultaneous reductions in
292 SO₂ take place. A particular feature of the energy sector, however, is that a significant portion
293 of methane mitigation from oil and gas (production and distribution) can be done independently
294 from other energy-related (combustion) emissions. An explicit distinction between production
295 and combustion emissions was not available in the gridded CEDS inventory, but, as illustrated
296 in Sect. 3.2, mitigation strategies targeting one category or the other can result in distinctly
297 different temperature outcomes. Global emissions from industry (IND) and shipping (SHP)
298 cause a net cooling impact despite a considerable warming from CO₂ emissions. In the long
299 term, the net impact of AGR and WST is small, while energy is the largest individual
300 contributor to warming due to its high CO₂ emissions (note that N₂O is not included in the
301 present analysis as emissions are not included in the gridded CEDS and SSP database, but
302 would add a small contribution to the long-term impact of AGR). The second largest driver of
303 long-term temperature change is IND, demonstrating the importance of non-CO₂ emissions for
304 shaping relative weight over different time frames. Aviation is not included here, but was
305 recently evaluated by Lund et al. (2017).

306 The largest regional contribution to net near-term warming is caused by emissions in East Asia
307 (EAS) and North America (NAM), followed by South East Asia (SEA) and South Africa (SAF)
308 (Fig.2c). However, the relative contributions from individual species vary. In EAS and NAM,
309 as well as Europe (EUR), the impact of current emissions of cooling and warming SLCFs
310 approximately balance in the near-term and these regions cause comparable net warming
311 impacts on 10- and 100-year time scales, as seen by comparing the white and grey circles in
312 Fig. 2c. These balancing characteristics do not imply that SLCF emissions should not be

313 reduced, but that the net benefits on global temperature may be low if mitigation measures that
314 simultaneously affect both cooling and warming SLFCs are implemented, in turn placing added
315 focus on the need to reduce CO₂ in order to mitigate warming in both the near- and long-term.
316 In SEA, SAF and South and Central America (SAM and MCA) methane and BC emissions are
317 presently high while emissions of CO₂ and cooling aerosols are low compared to other regions,
318 resulting in a net warming impact after 10 years that is substantially higher than that of CO₂
319 alone. This, in turn, suggest that using SLCF emission reduction to limit near-term warming
320 would be more effective here than in many other regions. Such detailed characteristics at the
321 emission source level are needed for the design of effective mitigation strategies.

322 Breaking the temperature impacts further down into economic sectors within each region (not
323 shown), we find that the results largely mirror the relative role of species and sectors on the
324 global level in Fig. 2b. The warming contributions in South America and Africa, and hence
325 higher potential for net temperature reductions, stem primarily from the agriculture, waste
326 management, and energy production sectors. In SAF, mitigation of BC emissions from the
327 residential and transport sectors also play an important role. In most regions, emissions from
328 IND cause a net negative impact on global temperature change, while in the ENE sector,
329 impacts of cooling and warming SLFCs compete and warming from CO₂ is a key driver of both
330 near- and long-term warming.

331 Overall, the potential for global temperature reductions inherent in the present SLCF emissions
332 is highly inhomogeneous, and co-emitted species – including CO₂ – must be taken into account
333 in any targeted climate policy for reduction of near-term warming. We emphasize that
334 mitigation of SLFCs, while important, need to be sustained and complimentary to strong cuts
335 in CO₂ for long-term reduction in global warming.

336

337 **3.2 Temperature response to example mitigation measures**

338 The results above suggest that strategies for emission reductions clearly can play out very
339 differently in terms of net impact on global temperature across source region and sector. To
340 further illustrate the importance of considering co-emissions and demonstrate the applicability
341 of our dataset, we calculate the effect on global temperature in the near- and long-term
342 following simplified examples of emission reduction packages in three of the global sectors
343 (ENE, AGR and SHP). The measures are broadly assumed to be motivated by either *i*) air
344 quality improvements (package 1, P1), *ii*) methane reductions (as part of the SDG agenda or
345 climate mitigation) (P2) or *iii*) CO₂ reductions/climate targets (P3). Table 2 shows the set of
346 species reduced in each case, with the percentage reduction given in parentheses. We note that
347 these reductions are based on expert judgement given underlying assumptions, e.g., for the
348 reduction in shipping speed, and are associated with uncertainties. Furthermore, they are
349 assumed to occur instantaneously. However, as the primary purpose here is illustrative, the
350 examples are kept idealized and should be interpreted as such.

351 The global temperature effect resulting from elimination of emissions in each package on 10-
352 and 100-year time horizons is shown in Fig.3. The energy sector can be sub-divided into fossil
353 fuel production/distribution and combustion categories. An air quality-driven set of measures

354 (P1), e.g., end-of-pipe measures such as scrubbers, filters and catalysts, could therefore be
355 implemented that would strongly reduce SO₂ and NO_x emissions but not noticeably affect the
356 key methane or CO₂ contribution. Such measures are well understood, i.e., their efficiencies,
357 costs, and technical implementation has been well documented and real-life application is
358 already widespread but there is still large potential, especially in fast-growing economies. As
359 shown by the top bar on the left in Fig.3, the subsequent near-term temperature impact would
360 be a warming contribution due to removal of cooling aerosols, adding to the already large net
361 warming impact of the sector (Fig. 2b). As seen from the right-hand side of Fig. 3, the long-
362 term effect would also be minor, leaving the dominating CO₂ warming. A significant fraction
363 of methane emissions, originating from the production and distribution of fossil fuels, could be
364 mitigated separately from several other SLCFs, for instance by addressing venting and leaks
365 from oil, gas and coal exploration, and upstream and downstream gas flaring. Respective
366 measures would include capture/recovery and use of gas, as well as reduced and improved
367 flaring, with added benefits in terms of reduced CO₂ and BC (P2). This results in a notable
368 reduction in the near-term impact of the sector. Finally, P3 shows the impact of a dedicated
369 climate strategy, here illustrated by the difference between a middle-of-the-road and a below-
370 two-degrees scenario (in 2050, obtained from the GAINS model (Klimont et al., 2017)), where
371 more substantial CO₂ mitigation also result in larger reduction of the sector's long-term
372 temperature impact than in P2.

373 Due to the dominating contribution from methane to the temperature impact of the agriculture
374 sector, measures that primarily target other emissions, such as improving nitrogen use
375 efficiency (P1), unsurprisingly bring low net climate benefits unless accompanied by
376 simultaneous measures for methane reductions (P2). Examples of the latter is promoting dietary
377 changes, leading to lower meat consumption and consequently lower livestock numbers.
378 Reducing NH₃ and NO_x (P1) could, however, bring important local air quality benefits, and our
379 results suggest that these would come with relatively small trade-offs from unmasking of
380 aerosol cooling, at least in terms of global mean temperature on this time scale. Only small
381 additional benefits (at a global scale) were estimated for the increased use of biogas (P3) based
382 in utilization of livestock manures. The net impact of the shipping sector (SHP) is a cooling in
383 the near-term, as shown in several previous studies (e.g., Berntsen & Fuglestvedt, 2008;
384 Fuglestvedt et al., 2009). Measures that eliminate shipping emissions of SO₂ (low sulfur fuels,
385 scrubbers) and NO_x (selective catalytic reduction) hence result in an added near-term warming,
386 also when simultaneous elimination of the sector's CO₂ emissions occur (P2, P3).

387 This example is simplified and illustrative, and we calculate pulse-based temperature impacts
388 following instantaneous emission reductions. However, since our pulse-based emission metrics
389 can easily be used to study changes over time to any emission or policy scenario through
390 convolution (Aamaas et al., 2013), our dataset has broad applicability. In the next section, we
391 use precisely this method to quantify the impact of temporally evolving emissions according to
392 the most recent scenarios.

393 **3.3 Temperature response to SLCFs and CO₂ under the SSP-RCP scenarios**

394 While knowledge of the present-day emission composition and net temperature impact over
395 time is essential to support mitigation design and implementation, real-world emissions will

396 evolve following a combination of socioeconomic developments, technological advancement
397 and policy adoption. Next, we investigate plausible pathways for the future impact of SLCFs
398 and CO₂ by quantifying the global temperature change over the period 1900-2100 to regional
399 and sectoral emissions following the SSP-RCP scenarios. In the following paragraphs, we show
400 results from four of the nine SSP-RCP scenarios used in the present analysis (SSP1-1.9, SSP2-
401 4.5, SSP3-7.0 and SSP5-8.5). These span the range of future emission evolutions, but we
402 recognize that the realism of SSP5-8.5 is debated in the literature due to its very high emissions
403 (e.g., Ritchie & Dowlatabadi, 2017).

404 Figure 4 shows the evolution of temperature response under the SSP-RCPs for our source
405 regions, with corresponding results for the global economic sectors given in Fig. S3. Our
406 emissions regions not only have large differences in terms of present-day emissions, but also
407 of past evolution. This historical contribution, which was not captured in the analysis of the
408 first half of the paper, brings NAM and EUR as the two largest contributors to the present-day
409 warming (Fig. 4a) due to their much higher past CO₂ emissions, in line with previous literature
410 (Höhne et al., 2011; Skeie et al., 2017). While presently being the largest emission source, EAS
411 only surpasses EUR and NAM, in terms of contribution to temperature change, between 2020
412 and 2030 when the cumulative effect of CO₂ is accounted for. In SSP1-1.9, where emissions of
413 CO₂ decline strongly during the first half of the century in all regions, the net temperature
414 response levels off or starts to decline in the second half of the century. In the remaining
415 scenarios, the net temperature impact increases over the century for all regions. EAS remains
416 the largest contributor, whereas in SSP5-8.5 SAS overtakes NAM as the second most important
417 region by 2100 and SAF reaches the same order of magnitude as EUR. This shows a projected
418 shift in emissions and increasing importance of the developing world. We note that since our
419 primary focus here is on quantifying the contributions to, and potential for further reduction of,
420 near- and long-term temperature impacts, we do not include negative CO₂ emissions which is
421 already a mitigation measure. Furthermore, the gridded SSP-RCP inventory only provide
422 negative CO₂ as a separate category without information for mapping these emissions to
423 economic sectors. We do, however, include the negative CO₂ category in our inventory of
424 regional scenarios for further analyses beyond our study (see “Data Availability”).

425 In our calculations, the net temperature response to emissions from the global energy (ENE)
426 sector becomes larger than that due to AGR and RES in the early 2000s (Fig. S1a), after which
427 ENE remains the largest individual sector until 2100 in all scenarios. The relative importance
428 of AGR and ENE historically is yet another example of how including SLCFs can change
429 relevance over different time frames, as also demonstrated by Reisinger and Clark (2018) for
430 non-CO₂ livestock emissions. In our results, both the warming due to CH₄ from AGR and the
431 contributions from cooling emissions from ENE act to shape the relative role of the two sectors
432 over time. The global mean temperature impact of IND switches from a net cooling to a net
433 warming in the late 20th century as the warming due to CO₂ accumulates and overwhelms the
434 cooling from SO₂.

435 While the contribution from CO₂ to the net warming becomes dominant by 2100 for most
436 regions and sectors in all scenarios, the relative importance of SLCFs and CO₂ continue to be
437 highly variable across emission source over time, in particular under SSP3-7.0 and SSP5-8.5.

438 This can be seen in Fig.4b, where we break down the future net temperature response in 2030,
439 2050 and 2100 into individual contributions from methane, CO₂, BC and the sum of SO₂ and
440 NO_x. Here we show a selection of the source regions that differ notably in composition and
441 temporal trend. See Fig. S4 for remaining regions.

442 The SSP-RCPs differ in both climate forcing targets and stringency of air pollution control, as
443 well as underlying socioeconomic development. SSP1-1.9 is characterized by low societal
444 challenges to mitigation and adaptation, and strong climate and air quality policies, resulting in
445 rapidly declining emissions of both SLCFs and CO₂. However, even for strong air pollution
446 there is a differentiation between high-, medium- and low-income countries, with a substantial
447 time lag in the latter two (Rao et al., 2017). For example, emissions of SO₂ in SAS and SAF
448 decline less than in other regions, subsequently maintaining a significant cooling contribution
449 to the temperature impact. In the intermediate scenario, SSP2-4.5, there is a reduction in
450 emissions, but this is delayed and slower compared to SSP1-1.9. In SSP3-7.0, the world follows
451 a path with more inequality and conflict, where only weak air pollution control is implemented
452 and the end-of-century climate forcing, and hence CO₂ emissions, is higher. Subsequently,
453 emission trends and SLCF contributions display more regional heterogeneity. There is a
454 particularly strong projected increase in methane emission in South Asia, Africa and South
455 America in this scenario. While previous decades have seen a southeastward shift in air
456 pollution emissions, from high income regions at northern latitudes to East and South Asia,
457 these findings suggest that a second shift may be underway, towards low- and middle-income
458 countries in the developing world. Further studies are needed to improve the knowledge about
459 the resulting climate and environmental consequences, as well as how to strengthen the
460 mitigation options, in these regions. While EAS remains the region with the largest warming
461 impact by 2100 in all scenarios, the contributions to warming from methane and BC in SAF
462 and SAS surpasses those of EAS in 2100 in both SSP3-7.0 and SSP5-8.5. The net temperature
463 response to emissions in SAS increases from close to zero to a significant warming as CO₂
464 emissions increase. SSP5-8.5 is characterized by high challenges to mitigation and high climate
465 forcing in 2100, but still assumes strong air pollution control since the high use of fossil fuels
466 would otherwise result in unbearable air pollution levels. Combined, this leads to increasing
467 temperature impact due to increasing CO₂ emissions, but lower SLCF impacts than in SSP3-
468 7.0, but with a non-negligible contribution from methane for several regions. Hence, in
469 medium- and low-income regions, SLCFs, and in particular methane, are projected to play a
470 continued important role for future temperature change.

471 Clearly, and as expected, the largest difference in SLCF contributions to future temperature
472 response is between SSP1-1.9 and SSP3-7.0. To see where the largest additional climatic
473 benefit can be gained from mitigating SLCF emissions in line with SSP1-1.9, relative to SSP3-
474 7.0, we show the difference in temperature between these two scenarios in 2030, 2050 and 2100
475 in Fig.5. Results are shown by region and sector, for all combinations where the temperature
476 difference is greater than $\pm 0.01^{\circ}\text{C}$. For comparison, the CMIP6 mean difference in projected
477 surface temperature between SSP3-7.0 and SSP1-2.6 (which is close to SSP1-1.9 in emissions)
478 is around 0.5°C in 2050 and 2°C in 2100 when accounting for all global emissions (Tokarska
479 et al., 2020). As seen from Fig. 4 and Fig. S3, CO₂ is the key driver of this long-term temperature
480 difference between the scenarios for most sectors and regions. However, as seen in Fig.5, there

481 are also important SLCF contributions, most notably from the large sources of methane;
482 agriculture, energy and waste management. Furthermore, 9 of the 12 top contributions are from
483 regions in Africa, South Asia or South and Central America, again demonstrating the
484 importance of the development in low- and middle-income countries for future levels of SLCFs.
485 Fig.5. also shows how the strong SLCF mitigation in SSP1-1.9, relative to SSP3-7.0, can result
486 in a net warming contribution to climate for some region-sector combinations, as exemplified
487 by the industry sector in East and South Asia. As shown by the panel on the right-hand side of
488 Fig. 5, for most sector/region combinations, around 10% of the avoided (or added) warming
489 from strong mitigation would be realized already by 2030, and around 40-50% by 2050.

490

491 **4 Discussion**

492 In terms of avoided global warming, there is much to be gained by moving from a global
493 emission pathway following SSP3-7.0 to one following SSP1-1.9, including contributions from
494 reductions of SLCFs, as discussed above. While a comprehensive assessment of policy and
495 technological interventions required to translate this potential to actual emission cuts is beyond
496 the scope of the present study, we outline key general features and discuss specific examples in
497 the case of methane, in the following paragraphs.

498

499 Available literature suggest that rapid reductions of air pollutants' emissions are technically
500 possible drawing on experience in both developed and developing countries (Crippa et al.,
501 2016; Kanaya et al., 2020; Klimont et al., 2017) but would require simultaneous strengthening
502 of institutions to enforce the laws. The focus of policies would differ between OECD countries
503 and the developing world. As demonstrated by our findings, further measures in the OECD
504 would primarily focus on reducing emissions from residential heating, non-road transportation,
505 and agriculture while assuring enforcement of legislation in power and industry sectors. The
506 rapidly industrializing and developing countries would need to further strengthen legislation for
507 the power, industry, transport sectors, implement improved measures to reduce waste
508 management emissions, reduce emissions from agriculture, and provide wide access to clean
509 fuels securing cooking and heating needs. Several of these policies would contribute positively
510 to the SDGs (Rafaj et al., 2018). For methane, the non-CO₂ component found here to be most
511 important for future warming, reducing venting and increasing utilization of associated
512 petroleum gas in oil and gas exploration and increased use of biogas from waste should be a
513 priority, and the technical potential for considerable reductions until 2050 exists (Höglund-
514 Isaksson et al., 2020). Integrated response options that can deliver significant mitigation also
515 exist for the agriculture sector, including increased productivity of land used for food
516 production and improved livestock management (Smith et al., 2019). A similar suite of methane
517 measures is needed as for the developed and developing world, although waste management
518 requires larger transformation and there is additional significant potential to reduce emissions
519 from coal mining sector in the latter. A recent study suggests that anthropogenic fossil methane
520 emissions may be significantly underestimated (Hmiel et al., 2020), and as such, reductions
521 may be even more critical. Specific measures for reducing aerosols and ozone precursors in
522 order to improve air quality while contributing to climate change mitigation have recently been

523 assessed for South East Asia (UNEP, 2019) and Latin America (UNEP, 2018). As shown in the
524 present analysis, contributions from SLCFs to temperature change are projected to increase
525 strongly in the Middle East and Africa in several scenarios. An increasing carbonization in
526 Africa south of the Sahara, primarily due to the increasing use of oil in the transport sector, has
527 already been observed (Steckel et al., 2019). This underlines the need for further focus on these
528 regions in future studies and assessments.

529
530 SSP3-7.0 and SSP1-1.9 not only differ in the stringency of the assumed air pollution control,
531 but also in socioeconomic development and end-of-century climate forcing. To isolate the role
532 of air pollution policies in the transition to a low warming pathway, a companion scenario to
533 SSP3-7.0 has been developed, the SSP3-lowNTCF (Gidden et al., 2019). Here, the
534 socioeconomic narrative is the same, but emission factors for the short-lived species are
535 assumed to be in line with those in SSP1-1.9. The result is similar global CO₂ emissions but up
536 to 60% reductions in global SLCF emissions in SSP3-lowNTCF relative to SSP3-7.0. Using
537 the SSP3-lowNTCF emissions as input, we find that this in turn leads to a net temperature
538 response to total global emissions in 2100 that is 13% lower in SSP3-LowNTCF than in SSP3-
539 7.0 (an absolute difference of 0.5°C, from 3.7°C to 3.2°C in our calculations). For comparison,
540 the net temperature response is 71% (or 2.6°C) lower in SSP1-1.9 than in SSP3-7.0 in our
541 calculations.

542
543 The potential for reducing near-term warming by targeting BC emissions in the transport and
544 residential sectors has been highlighted earlier (e.g., UNEP, 2011). We also find notable BC
545 warming contributions from the residential sector in some regions, mainly South Asia and
546 Africa, but estimate quite low BC effects from the transport sector. This has three main reasons.
547 Firstly, since earlier studies (done about 10 years ago) there have been significant changes in
548 legislation, and new diesel trucks and cars are (in several regions) equipped with particulate
549 filters effectively removing BC. By now these vehicles represent a significant part of the fleet
550 in many regions and the trend is expected to continue. Secondly, as described in Sect.2, we use
551 an AGTP for BC that is 15% lower than in previous studies using the same methodology. This
552 is done to account for the rapid adjustments associated with BC short-wave absorption (Stjern
553 et al., 2017), which has been found to reduce the effective RF in a range of global climate
554 models via changes in stability and cloud formation (Smith et al., 2018). For our study, this
555 factor applies to BC emissions from all sources and hence results in a reduced the net warming
556 impact. Finally, we account for cooling from nitrate aerosols from emissions of NO_x, for which
557 the transport sector is a significant source, even in regions where stricter vehicle emission
558 standards (e.g., Euro 5) have been adopted.

559

560 4.1 Caveats and uncertainties

561 The AGTP is a well-established framework that has been applied in several studies of
562 attribution of temperature impacts to emission sources and scenarios (e.g., Collins et al., 2013;
563 Lund et al., 2017; Sand et al., 2015; Stohl et al., 2015; Aamaas et al., 2019). Here we have also
564 consistently included the carbon-climate feedback in the AGTP for all species. This increases

565 the non-CO₂ AGTPs, however, less than initially suggested by Myhre et al. (2013) as discussed
566 by Gasser et al. (2017). Figure S5 shows the global mean net temperature response to total
567 emissions under 6 of the 9 SSP-RCPs, with and without the feedback. By the end of century,
568 there is a 5-9% difference depending on scenario.

569 A key strength of the AGTP framework is that allows us to investigate the effects of individual
570 species, sources and scenarios, which would be confounded by the low signal-to-noise ratio in
571 fully coupled models, in a transparent manner. However, there are also caveats. Importantly,
572 the AGTP metric is linear, while in reality the radiative efficiency can have non-linear
573 dependencies on the background atmospheric conditions. In this study, we account for one part
574 such non-linearities by using radiative efficiencies for the aerosols and ozone precursors that
575 vary with emission location to calculate region-specific AGTPs. The part of the non-linearities
576 caused by changing background levels of pollutants over time is, however, not included. For
577 the well-mixed greenhouse gases CO₂, CH₄ and N₂O, the radiative efficiency (RE) is reduced
578 with increasing atmospheric background concentrations. Previous literature suggests that the
579 sensitivity to emission scenario is small, and the relationship between emissions and
580 temperature response more linear, for CO₂ (Caldeira & Kasting, 1993). However, the same has
581 not been shown for methane (and N₂O – which is not considered here). We therefore perform
582 an additional sensitivity test where we calculate an AGTP(t) that is adjusted to the global
583 atmospheric concentrations over time (using the equation from Etminan et al. (2016) and global
584 concentrations for each SSP-RCP from the IIASA SSP database (IIASA, 2020; Riahi et al.,
585 2017)). Figure S5 shows the resulting temperature response, compared to the temperature
586 response calculated with and without the CCf. As expected, using a dynamically adjusted RE
587 results in a lower warming in the high emission scenarios and a slightly higher temperature
588 response under low emissions. In the case of extreme scenario SSP5-8.5, the effect is of the
589 same order of magnitude as that from adding the CCf, but of opposite sign. For aerosols and
590 ozone precursors, potential saturation effects involve complex, spatially heterogeneous
591 chemistry, cloud and climate interactions that require detailed chemistry-climate simulations to
592 be resolved, and even then, may not be fully captured due to e.g., the coarse resolution of current
593 models. We emphasize that the absolute magnitude of temperature changes quantified with the
594 AGTP framework should be interpreted with care, as this method is primarily designed to study
595 relative importance and relationships between individual emissions and sources.

596 Our analysis reflects best estimate input data to the extent possible, but results have
597 considerable uncertainty, in emissions, RF and climate sensitivity. As shown in Fig. 2a, we
598 estimate, due to uncertainty in RF alone, a 1 standard deviation range in the total net temperature
599 response on the 10-year time horizon of $\pm 0.01^\circ\text{C}$, about 38% of the net temperature response
600 of 0.03°C (the range is considerably lower on the 100-year time scale as the RF of SLCFs is
601 much more uncertain than that of CO₂). Uncertainties in emission inventories are difficult to
602 quantify, but are generally considered lowest for CO₂ and SO₂ emissions, and high for
603 carbonaceous aerosols (Hoesly et al., 2018). The level of uncertainty also differs across regions
604 and sectors, with emissions from nature related emissions (e.g., agriculture, landfills) more
605 uncertain than emissions in the fossil-fuel sector (Amann et al., 2013; Jonas et al., 2019).
606 Moreover, recent studies point to emission trends that are not accurately represented in the
607 global inventory, such as SO₂ and NO_x in China (Zheng et al., 2018) and fossil fuel CH₄

608 emissions (Hmiel et al., 2020). However, due to high spatiotemporal variability and lack of
609 consistent data, a comprehensive uncertainty analysis at the regional and sectoral level is
610 challenging. The equilibrium climate sensitivity (ECS) inherent in the climate response in IRF
611 used in the present analysis is $0.885 \text{ K (Wm}^{-2}\text{)}^{-1}$. This is in the upper range reported by Bindoff
612 et al. (2013), but lower than many recent estimates (Forster et al., 2019; Zelinka et al., 2020).
613 While emission uncertainties can have a strong spatiotemporal character, changes in the ECS
614 mostly act to scale estimates for all sectors and regions but is less important for their relative
615 ranking.

616 Our analysis is limited to temperature change as a measure of climate impacts. SLCFs, and in
617 particular aerosols, also play a key role in shaping local and regional hydrology and dynamics.
618 Comparing the SSP3-7.0 and SSP3-lowNTCF scenarios, Allen et al. (2020) recently found a
619 significant precipitation increase due to removal of aerosols, with the strongest moistening
620 trends over Asia. An increase in the Asian summer monsoon precipitation in scenarios with
621 strong air pollution reductions was also recently found by Wilcox et al. (2020). Hence, further
622 studies using coupled models are needed to fully capture the effects of the SLCFs under SSPs
623 on local climate and environment.

624

625 **5 Conclusions**

626 Complimentary mitigation of CO₂ and other LLGHG with SLCFs is of key importance for
627 achieving the ambitions of the Paris Agreement and meeting the Sustainable Development
628 Goals. Using the concept of Absolute Global Temperature change Potential (AGTP), an
629 emission metric-based emulator of the climate response, we here investigate the contribution
630 from emissions of SLCFs and CO₂ from 7 economic sectors in 13 source regions to global
631 temperature change. In addition to quantifying the near- and long-term temperature response to
632 present-day emissions, i.e., in line with the traditional emission metric studies, we evaluate the
633 role of individual SLCFs and CO₂ as projected by the most recent generation scenarios, the
634 Shared Socioeconomic Pathways (SSPs), with greater regional and sectoral detail than previous
635 literature. We account for the geographical dependence of the radiative forcing of SLCF
636 emissions, as well as the current understanding of global-scale indirect and semi-direct aerosol
637 forcing. A key update to our method relative to the bulk of comparable literature, is a treatment
638 of the carbon-climate feedback in the AGTPs of the SLCFs.

639

640 As is well established, CO₂ is the dominant driver of warming on longer time scales and any
641 strategy for limiting long-term temperature change critically depends on deep cuts in CO₂
642 emission. As shown by our results, CO₂ also give a significant contribution to near-term
643 warming. The potential for additional reductions in near-term temperature change from
644 reductions in present-day SLCF emissions is highly inhomogeneous across region and sector.
645 Key in all regions are the major emitters of methane, in particular agriculture and waste
646 management, but also energy production. In contrast, some sectors and regions, notably
647 industry, energy and transport in East and South Asia and the Middle East, have strong
648 contributions from cooling SLCFs resulting in a net negative near-term temperature impact or
649 an approximate balance between cooling and warming SLCFs. While this does not imply that

650 mitigation measures should not be implemented, understanding of the detailed characteristics
651 and relevance over time at the emission source level is needed for the design and assessment of
652 mitigation strategies.

653
654 The regional heterogeneity in SLCF emissions and subsequent contributions to global
655 temperature change continues under most of the nine SSP-RCP scenarios considered here.
656 While CO₂ becomes the dominant contributor to warming in all regions over time, SLCFs are
657 projected to continue to play an important role for global temperature change over the 21st
658 century in many regions. In particular, emissions of SLCFs in East and South Asia are projected
659 to remain high, at least until the mid-21st century. Moreover, there is a shift in emissions towards
660 low- and middle-income countries in the developing world. Notably, a strong increase in
661 emissions in Africa south of the Sahara is projected under most of the SSP-RCPs considered,
662 and is especially pronounced in SSP3-7.0 and SSP5-8.5. Hence, in addition to the focus on the
663 current major current sources of SLCFs, enabling technological and legislative development
664 on the African continent will likely be of key importance for a transition from high emission
665 pathways towards one in line with SSP1-1.9 and the ambitions of the Paris Agreement, which
666 in turn could give reductions in global warming already over the next couple of decades.
667 Technological advancement could bring benefits even if there is no dedicated climate policy
668 addressing SLCFs, simply by reduced emission factors, as demonstrated by the SSP3-lowNTPCF
669 scenario.

670
671 The large spatiotemporal heterogeneity in emissions trends and subsequent temperature
672 responses underlines the need to go beyond global emission scenarios. By quantifying the
673 global temperature response to emissions from 13 regions, 7 sectors and 9 scenarios in a
674 consistent and transparent framework, we provide a more comprehensive dataset than, to our
675 knowledge, currently exists. We note that the AGTP framework is primarily designed to study
676 relative importance and relationships between individual emissions and sources, and that the
677 absolute magnitude of temperature responses should be interpreted with care due to the linearity
678 of the AGTP. The uncertainties in emissions could also affect the regional and sectoral ranking
679 but are poorly known. However, by making our full dataset publicly available, we provide a
680 tool that enables further analysis and comparison of e.g., mitigation strategies at the sectoral
681 and regional level without the use of complex models.

682

683

684 **Data availability**

685 Regional and sectoral emission timeseries, AGTPs and temperature responses are publicly
686 available via Figshare under the DOI <https://doi.org/10.6084/m9.figshare.11386455> (Lund et
687 al., 2020). The full set of gridded SSP anthropogenic emission data are available from the
688 ESGF system (<https://esgf-node.llnl.gov/search/input4mips/>, Cinquini et al. (2014), last
689 access: December 2019). Code available upon request from Marianne T. Lund
690 (m.t.lund@cicero.oslo.no).

691

692 **Author contributions**

693 Lund led the study, prepared the input data and wrote the paper. Aamaas performed the emission
694 metric and uncertainty calculations. Stjern and Samset produced the graphics. Klimont and
695 Berntsen contributed to the design of the analysis. All authors contributed to the manuscript
696 preparation.

697

698 **Competing interests**

699 The authors declare that they have no competing interests.

700

701 **Acknowledgements**

702 The authors acknowledge funding from the Research Council Norway grant no. 248834
703 (QUISARC). We thank Glen Peters and Robbie Andrews (CICERO) for assistance with the
704 technical implementation of the carbon-climate feedback. We also thank the two anonymous
705 referees for their comments and suggestions.

706

707 **References**

708 Allen R. J., Turnock S., Nabat P., et al. Climate and air quality impacts due to mitigation of
709 non-methane near-term climate forcers, *Atmos. Chem. Phys.* 2020(20), 9641–9663,
710 <https://doi.org/10.5194/acp-20-9641-2020>, 2020.

711 Amann M., Klimont Z. & Wagner F. Regional and Global Emissions of Air Pollutants: Recent
712 Trends and Future Scenarios, *Annual Review of Environment and Resources.* 38(1), 31-55,
713 10.1146/annurev-environ-052912-173303, 2013.

714 AMAP AMAP Assessment 2015: Black carbon and ozone as Arctic climate forcers. Arctic
715 Monitoring and Assessment Programme (AMAP), Oslo, Norway, 2015.

716 Anenberg S. C., Schwartz J., Shindell D., et al. Global Air Quality and Health Co-benefits of
717 Mitigating Near-Term Climate Change through Methane and Black Carbon Emission Controls,
718 *Environmental Health Perspectives.* 120(6), 831-839, doi:10.1289/ehp.1104301, 2012.

719 Arneth A., Unger N., Kulmala M., et al. Clean the Air, Heat the Planet?, *Science.* 326(5953),
720 672-673, 10.1126/science.1181568, 2009.

721 Berntsen T. & Fuglestedt J. Global temperature responses to current emissions from the
722 transport sectors, *Proc Natl Acad Sci U S A.* 105(49), 19154-19159, 10.1073/pnas.0804844105, 2008.

723 Bindoff N. L., Stott P. A., AchutaRao K. M., et al. Detection and Attribution of Climate Change:
724 from Global and Regional, in: *Climate Change 2013: The Physical Science Basis, Contribution of*
725 *Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*,
726 edited by: Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia,
727 Y., Bex, V., and Midgley, P. M., Cambridge University Press, Cambridge, UK and New York, NY, USA.,
728 2013.

729 Bond T. C., Doherty S. J., Fahey D. W., et al. Bounding the role of black carbon in the climate
730 system: A scientific assessment, *Journal of Geophysical Research: Atmospheres.* 118(11), 5380-5552,
731 10.1002/jgrd.50171, 2013.

732 Boucher O. & Reddy M. S. Climate trade-off between black carbon and carbon dioxide
733 emissions, *Energy Policy.* 36(1), 193-200, <https://doi.org/10.1016/j.enpol.2007.08.039>, 2008.

734 Bowerman N. H. A., Frame D. J., Huntingford C., et al. The role of short-lived climate
735 pollutants in meeting temperature goals, *Nature Climate Change*. 3, 1021, 10.1038/nclimate2034,
736 2013.

737 Caldeira K. & Kasting J. F. Insensitivity of global warming potentials to carbon dioxide
738 emission scenarios, *Nature*. 366(6452), 251-253, 10.1038/366251a0, 1993.

739 CCAC Air pollution in Asia and Pacific: Science-based solution, 2019.

740 Cinquini L., Crichton D., Mattmann C., et al. The Earth System Grid Federation: An open
741 infrastructure for access to distributed geospatial data, *Future Generation Computer Systems*. 36,
742 400-417, <https://doi.org/10.1016/j.future.2013.07.002>, 2014.

743 Collins W. J., Fry M. M., Yu H., et al. Global and regional temperature-change potentials for
744 near-term climate forcers, *Atmos. Chem. Phys.* 13(5), 2471-2485, 10.5194/acp-13-2471-2013, 2013.

745 Crippa M., Janssens-Maenhout G., Dentener F., et al. Forty years of improvements in
746 European air quality: regional policy-industry interactions with global impacts, *Atmos. Chem. Phys.*
747 16(6), 3825-3841, 10.5194/acp-16-3825-2016, 2016.

748 Etminan M., Myhre G., Highwood E. J., et al. Radiative forcing of carbon dioxide, methane,
749 and nitrous oxide: A significant revision of the methane radiative forcing, *Geophysical Research*
750 *Letters*. 43(24), 12,614-612,623, 10.1002/2016gl071930, 2016.

751 Feng L., Smith S. J., Braun C., et al. The generation of gridded emissions data for CMIP6,
752 *Geosci. Model Dev.* 13(2), 461-482, 10.5194/gmd-13-461-2020, 2020.

753 Forster P. M., Maycock A. C., McKenna C. M., et al. Latest climate models confirm need for
754 urgent mitigation, *Nature Climate Change*, 10.1038/s41558-019-0660-0, 2019.

755 Fuglestad J., Berntsen T., Eyring V., et al. Shipping Emissions: From Cooling to Warming of
756 Climate—and Reducing Impacts on Health, *Environmental Science & Technology*. 43(24), 9057-9062,
757 10.1021/es901944r, 2009.

758 Fuglestad J. S., Shine K. P., Berntsen T., et al. Transport impacts on atmosphere and
759 climate: Metrics, *Atmospheric Environment*. 44(37), 4648-4677,
760 <https://doi.org/10.1016/j.atmosenv.2009.04.044>, 2010.

761 Gasser T., Peters G. P., Fuglestad J. S., et al. Accounting for the climate-carbon feedback in
762 emission metrics, *Earth Syst. Dynam.* 8(2), 235-253, 10.5194/esd-8-235-2017, 2017.

763 Geoffroy O., Saint-Martin D., Olivé D. J. L., et al. Transient Climate Response in a Two-Layer
764 Energy-Balance Model. Part I: Analytical Solution and Parameter Calibration Using CMIP5 AOGCM
765 Experiments, *Journal of Climate*. 26(6), 1841-1857, 10.1175/jcli-d-12-00195.1, 2013.

766 Gidden M. J., Riahi K., Smith S. J., et al. Global emissions pathways under different
767 socioeconomic scenarios for use in CMIP6: a dataset of harmonized emissions trajectories through
768 the end of the century, *Geosci. Model Dev.* 12(4), 1443-1475, 10.5194/gmd-12-1443-2019, 2019.

769 Haines A., Amann M., Borgford-Parnell N., et al. Short-lived climate pollutant mitigation and
770 the Sustainable Development Goals, *Nature Climate Change*. 7(12), 863-869, 10.1038/s41558-017-
771 0012-x, 2017.

772 Harmsen M., Fricko O., Hilaire J., et al. Taking some heat off the NDCs? The limited potential
773 of additional short-lived climate forcers' mitigation, *Climatic Change*, 10.1007/s10584-019-02436-3,
774 2019.

775 Hienola A., Partanen A.-I., Pietikäinen J.-P., et al. The impact of aerosol emissions on the
776 1.5 °C pathways, *Environmental Research Letters*. 13(4), 044011, 10.1088/1748-9326/aab1b2, 2018.

777 Hmiel B., Petrenko V. V., Dyonisius M. N., et al. Preindustrial 14CH4 indicates greater
778 anthropogenic fossil CH4 emissions, *Nature*. 578(7795), 409-412, 10.1038/s41586-020-1991-8, 2020.

779 Hoesly R. M., Smith S. J., Feng L., et al. Historical (1750–2014) anthropogenic emissions of
780 reactive gases and aerosols from the Community Emission Data System (CEDS), *Geosci. Model Dev.*
781 2018(11), 369-408, <https://doi.org/10.5194/gmd-11-369-2018>, 2018.

782 Höglund-Isaksson L., Gómez-Sanabria A., Klimont Z., et al. Technical potentials and costs for
783 reducing global anthropogenic methane emissions in the 2050 timeframe –results from the GAINS
784 model, *Environmental Research Communications*. 2(2), 025004, 10.1088/2515-7620/ab7457, 2020.

785 Høhne N., Blum H., Fuglestedt J., et al. Contributions of individual countries' emissions to
786 climate change and their uncertainty, *Climatic Change*. 106(3), 359-391, 10.1007/s10584-010-9930-6,
787 2011.

788 IIASA SSP Database (Shared Socioeconomic Pathways) - Version 2.0,
789 <https://tntcat.iiasa.ac.at/SspDb/dsd?Action=htmlpage&page=10>. Accessed: June 2020., 2020.

790 Janssens-Maenhout G., Crippa M., Guizzardi D., et al. HTAP_v2.2: a mosaic of regional and
791 global emission grid maps for 2008 and 2010 to study hemispheric transport of air pollution, *Atmos.*
792 *Chem. Phys.* 15(19), 11411-11432, 10.5194/acp-15-11411-2015, 2015.

793 Jonas M., Bun R., Nahorski Z., et al. Quantifying greenhouse gas emissions, *Mitigation and*
794 *Adaptation Strategies for Global Change*. 24(6), 839-852, 10.1007/s11027-019-09867-4, 2019.

795 Joos F., Roth R., Fuglestedt J. S., et al. Carbon dioxide and climate impulse response
796 functions for the computation of greenhouse gas metrics: a multi-model analysis, *Atmos. Chem. Phys.*
797 13(5), 2793-2825, 10.5194/acp-13-2793-2013, 2013.

798 Kanaya Y., Yamaji K., Miyakawa T., et al. Rapid reduction of black carbon emissions from
799 China: evidence from 2009-2019 observations on Fukue Island, Japan, *Atmos. Chem. Phys.* . 20, 6339-
800 6356, <https://doi.org/10.5194/acp-20-6339-2020>, 2020.

801 Klimont Z., Kupiainen K., Heyes C., et al. Global anthropogenic emissions of particulate
802 matter including black carbon, *Atmos. Chem. Phys.* 17(14), 8681-8723, 10.5194/acp-17-8681-2017,
803 2017.

804 Kupiainen K. J., Aamaas B., Savolahti M., et al. Climate impact of Finnish air pollutants and
805 greenhouse gases using multiple emission metrics, *Atmos. Chem. Phys.* 19(11), 7743-7757,
806 10.5194/acp-19-7743-2019, 2019.

807 Lelieveld J., Evans J. S., Fnais M., et al. The contribution of outdoor air pollution sources to
808 premature mortality on a global scale, *Nature*. 525, 367, 10.1038/nature15371, 2015.

809 Lund M. T., Berntsen T. K. & Fuglestedt J. S. Climate Impacts of Short-Lived Climate Forcers
810 versus CO₂ from Biodiesel: A Case of the EU on-Road Sector, *Environmental Science & Technology*.
811 48(24), 14445-14454, 10.1021/es505308g, 2014a.

812 Lund M. T., Berntsen T. K., Heyes C., et al. Global and regional climate impacts of black
813 carbon and co-emitted species from the on-road diesel sector, *Atmospheric Environment*. 98, 50-58,
814 <https://doi.org/10.1016/j.atmosenv.2014.08.033>, 2014b.

815 Lund M. T., Aamaas B., Berntsen T., et al. Emission metrics for quantifying regional climate
816 impacts of aviation, *Earth Syst. Dynam.* 8(3), 547-563, 10.5194/esd-8-547-2017, 2017.

817 Lund M. T., Myhre G., Haslerud A. S., et al. Concentrations and radiative forcing of
818 anthropogenic aerosols from 1750 to 2014 simulated with the Oslo CTM3 and CEDS emission
819 inventory, *Geosci. Model Dev.* 11(12), 4909-4931, 10.5194/gmd-11-4909-2018, 2018.

820 Lund M. T., Aamaas B., Stjern C. W., et al. Data collection - "A continued role of Short-Lived
821 Climate Forcers under the Shared Socioeconomic Pathways". figshare. Dataset. ,
822 <https://doi.org/10.6084/m9.figshare.11386455>, 2020.

823 Melamed M. L., Schmale J. & von Schneidemesser E. Sustainable policy—key considerations
824 for air quality and climate change, *Current Opinion in Environmental Sustainability*. 23, 85-91,
825 <https://doi.org/10.1016/j.cosust.2016.12.003>, 2016.

826 Myhre G., Shindell D., Brøn F.-M., et al. Anthropogenic and natural radiative forcing. In:
827 Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth
828 Assessment Report of the Intergovernmental Panel on Climate Change [Stocker, T.F., D., Qin, G.-K.
829 Plattner, M. Tignor, S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P.M. Midgley (eds).
830 Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA 2013.

831 O'Neill B. C., Kriegler E., Riahi K., et al. A new scenario framework for climate change
832 research: the concept of shared socioeconomic pathways, *Climatic Change*. 122(3), 387-400,
833 10.1007/s10584-013-0905-2, 2014.

834 Persad G. G. & Caldeira K. Divergent global-scale temperature effects from identical aerosols
835 emitted in different regions, *Nature Communications*. 9(1), 3289, 10.1038/s41467-018-05838-6,
836 2018.

837 Pierrehumbert R. T. Short-Lived Climate Pollution, *Annual Review of Earth and Planetary*
838 *Sciences*. 42(1), 341-379, 10.1146/annurev-earth-060313-054843, 2014.

839 Rafaj P., Kiesewetter G., Gül T., et al. Outlook for clean air in the context of sustainable
840 development goals, *Global Environmental Change*. 53, 1-11,
841 <https://doi.org/10.1016/j.gloenvcha.2018.08.008>, 2018.

842 Ramanathan V. & Carmichael G. Global and regional climate changes due to black carbon,
843 *Nature Geoscience*. 1(4), 221-227, 10.1038/ngeo156, 2008.

844 Rao S., Klimont Z., Smith S. J., et al. Future air pollution in the Shared Socio-economic
845 Pathways, *Global Environmental Change*. 42, 346-358,
846 <https://doi.org/10.1016/j.gloenvcha.2016.05.012>, 2017.

847 Reisinger A. & Clark H. How much do direct livestock emissions actually contribute to global
848 warming?, *Global Change Biology*. 24(4), 1749-1761, 10.1111/gcb.13975, 2018.

849 Riahi K., van Vuuren D. P., Kriegler E., et al. The Shared Socioeconomic Pathways and their
850 energy, land use, and greenhouse gas emissions implications: An overview, *Global Environmental*
851 *Change*. 42, 153-168, <https://doi.org/10.1016/j.gloenvcha.2016.05.009>, 2017.

852 Ritchie J. & Dowlatabadi H. Why do climate change scenarios return to coal?, *Energy*. 140,
853 1276-1291, <https://doi.org/10.1016/j.energy.2017.08.083>, 2017.

854 Rogelj J., Schaeffer M., Meinshausen M., et al. Disentangling the effects of CO₂ and short-lived climate forcer mitigation, *Proceedings of the National Academy of Sciences*. 111(46),
855 16325-16330, 10.1073/pnas.1415631111, 2014.

857 Rogelj J., Meinshausen M., Schaeffer M., et al. Impact of short-lived non-CO₂ mitigation on
858 carbon budgets for stabilizing global warming, *Environmental Research Letters*. 10(7), 075001,
859 10.1088/1748-9326/10/7/075001, 2015.

860 Rogelj J., Shindell D., Jiang K., et al. Mitigation Pathways Compatible with 1.5°C in the Context
861 of Sustainable Development. In: Global Warming of 1.5°C. An IPCC Special Report on the impacts of
862 global warming of 1.5°C above pre-industrial levels and related global greenhouse gas emission
863 pathways, in the context of strengthening the global response to the threat of climate change,
864 sustainable development, and efforts to eradicate poverty [Masson-Delmotte, V., P. Zhai, H.-O.
865 Pörtner, D. Roberts, J. Skea, P.R. Shukla, A. Pirani, W. Moufouma-Okia, C. Péan, R. Pidcock, S.
866 Connors, J.B.R. Matthews, Y. Chen, X. Zhou, M.I. Gomis, E. Lonnoy, T. Maycock, M. Tignor, and T.
867 Waterfield (eds.)]. In Press., 2018.

868 Samset B. H. & Myhre G. Vertical dependence of black carbon, sulphate and biomass burning
869 aerosol radiative forcing, *Geophysical Research Letters*. 38(24), 10.1029/2011gl049697, 2011.

870 Samset B. H., Sand M., Smith C. J., et al. Climate Impacts From a Removal of Anthropogenic
871 Aerosol Emissions, *Geophysical Research Letters*. 45(2), 1020-1029, doi:10.1002/2017GL076079,
872 2018.

873 Sand M., Berntsen T. K., von Salzen K., et al. Response of Arctic temperature to changes in
874 emissions of short-lived climate forcers, *Nature Climate Change*. 6, 286, 10.1038/nclimate2880
875 <https://www.nature.com/articles/nclimate2880#supplementary-information>, 2015.

876 Shindell D., Kuylenstierna J. C. I., Vignati E., et al. Simultaneously Mitigating Near-Term
877 Climate Change and Improving Human Health and Food Security, *Science*. 335(6065), 183-189,
878 10.1126/science.1210026, 2012.

879 Shindell D. & Smith C. J. Climate and air-quality benefits of a realistic phase-out of fossil fuels,
880 *Nature*. 573(7774), 408-411, 10.1038/s41586-019-1554-z, 2019.

881 Shine K. P., Fuglestvedt J. S., Hailemariam K., et al. Alternatives to the Global Warming
882 Potential for Comparing Climate Impacts of Emissions of Greenhouse Gases, *Climatic Change*. 68(3),
883 281-302, 10.1007/s10584-005-1146-9, 2005.

884 Shoemaker J. K., Schrag D. P., Molina M. J., et al. What Role for Short-Lived Climate Pollutants
885 in Mitigation Policy?, *Science*. 342(6164), 1323-1324, 10.1126/science.1240162, 2013.

886 Skeie R. B., Fuglestedt J., Berntsen T., et al. Perspective has a strong effect on the calculation
887 of historical contributions to global warming, *Environmental Research Letters*. 12(2), 024022,
888 10.1088/1748-9326/aa5b0a, 2017.

889 Smith C. J., Kramer R. J., Myhre G., et al. Understanding Rapid Adjustments to Diverse Forcing
890 Agents, *Geophysical Research Letters*. 45(21), 12,023-012,031, 10.1029/2018gl079826, 2018.

891 Smith P., Nkem J., Calvin K., et al. Interlinkages Between Desertification, Land Degradation,
892 Food Security and Greenhouse Gas Fluxes: Synergies, Trade-offs and Integrated Response Options.
893 In: Climate
894 Change and Land: an IPCC special report on climate change, desertification, land degradation,
895 sustainable land management, food security, and greenhouse gas fluxes in terrestrial ecosystems
896 [P.R. Shukla, J. Skea, E. Calvo Buendia, V. Masson-Delmotte, H.- O. Portner, D. C. Roberts, P. Zhai, R.
897 Slade, S. Connors, R. van Diemen, M. Ferrat, E. Haughey,
898 S. Luz, S. Neogi, M. Pathak, J. Petzold, J. Portugal Pereira, P. Vyas, E. Huntley, K. Kissick, M. Belkacemi,
899 J. Malley, (eds.)]. In press., 2019.

900 Smith S. J. & Mizrahi A. Near-term climate mitigation by short-lived forcers, *Proceedings of*
901 *the National Academy of Sciences*. 110(35), 14202-14206, 10.1073/pnas.1308470110, 2013.

902 Steckel J. C., Hilaire J., Jakob M., et al. Coal and carbonization in sub-Saharan Africa, *Nature*
903 *Climate Change*, 10.1038/s41558-019-0649-8, 2019.

904 Stjern C. W., Samset B. H., Myhre G., et al. Rapid Adjustments Cause Weak Surface
905 Temperature Response to Increased Black Carbon Concentrations, *Journal of Geophysical Research:*
906 *Atmospheres*. 122(21), 11,462-411,481, 10.1002/2017JD027326, 2017.

907 Stohl A., Aamaas B., Amann M., et al. Evaluating the climate and air quality impacts of short-
908 lived pollutants, *Atmos. Chem. Phys.* 15(18), 10529-10566, 10.5194/acp-15-10529-2015, 2015.

909 Søvde O. A., Prather M. J., Isaksen I. S. A., et al. The chemical transport model Oslo CTM3,
910 *Geosci. Model Dev.* 5(6), 1441-1469, 10.5194/gmd-5-1441-2012, 2012.

911 Takemura T. & Suzuki K. Weak global warming mitigation by reducing black carbon emissions,
912 *Scientific Reports*. 9(1), 4419, 10.1038/s41598-019-41181-6, 2019.

913 Tokarska K. B., Stolpe M. B., Sippel S., et al. Past warming trend constrains future warming in
914 CMIP6 models, *Science Advances*. 6(12), eaaz9549, 10.1126/sciadv.aaz9549, 2020.

915 UNEP Integrated Assessment of Black Carbon and Tropospheric Ozone, 2011.

916 UNEP The Emissions Gap Report 2017. United Nations Environment Programme (UNEP),
917 Nairobi, 2017.

918 UNEP Integrated Assessment of Short-lived Climate Pollutants in Latin America and the
919 Caribbean., 2018.

920 UNEP Air Pollution in Asia and the Pacific: Science-based Solutions. UNEP, Paris. , 2019.

921 Unger N., Bond T. C., Wang J. S., et al. Attribution of climate forcing to economic sectors,
922 *Proceedings of the National Academy of Sciences*. 107(8), 3382-3387, 10.1073/pnas.0906548107,
923 2010.

924 Wilcox L. J., Liu Z., Samset B. H., et al. Accelerated increases in global and Asian summer
925 monsoon precipitation from future aerosol reductions, *Atmos. Chem. Phys. Discuss. (accepted)*. 2020,
926 1-30, 10.5194/acp-2019-1188, 2020.

927 Xu Y. & Ramanathan V. Well below 2 °C: Mitigation strategies for avoiding dangerous to
928 catastrophic climate changes, *Proceedings of the National Academy of Sciences*. 114(39), 10315-
929 10323, 10.1073/pnas.1618481114, 2017.

930 Zelinka M. D., Myers T. A., McCoy D. T., et al. Causes of Higher Climate Sensitivity in CMIP6
931 Models, *Geophysical Research Letters*. 47(1), e2019GL085782, 10.1029/2019gl085782, 2020.

932 Zheng B., Tong D., Li M., et al. Trends in China's anthropogenic emissions since 2010 as the
933 consequence of clean air actions, *Atmos. Chem. Phys.* 18(19), 14095-14111, 10.5194/acp-18-14095-
934 2018, 2018.

935 Aamaas B., Peters G. P. & Fuglestedt J. S. Simple emission metrics for climate impacts, *Earth*
936 *Syst. Dynam.* 4(1), 145-170, 10.5194/esd-4-145-2013, 2013.

937 Aamaas B., Berntsen T. K. & Samset B. H. The regional temperature implications of strong air
938 quality measures, *Atmos. Chem. Phys.* 19(24), 15235-15245, 10.5194/acp-19-15235-2019, 2019.

939

940

941

942

943

944

945

946

947

948

949

950

951

952

953

954

955

956

957

958

959

960

961

962

963

964

965

966

967 **Tables:**

968

969 *Table 1: Constants of the Geoffroy et al. (2013) IRF.*

	Mode 1	Mode 2
c _j	0.587	0.413
d _j (years)	4.1	249

970

971

972 *Table 2: Summary of species considered in the idealized emission reduction packages, the*
 973 *percentage reduction assumed, and example polices. All percentages refer the total emissions*
 974 *of a given sector, not total anthropogenic.*

Sector	Package 1 (P1)	Package 2 (P2)	Package 3 (P3)
ENE ^{a)}	End-of-pipe measures	Reduced loss in fossil fuel production and distribution	Climate strategy
	SO ₂ (85%) NO _x (75%)	CH ₄ (75%), BC (85%) CO ₂ (3%) ^{b)}	CO ₂ (65%), CH ₄ (40%) SO ₂ (65%), NO _x (45%) BC (35%)
AGR	Nitrogen use efficiency and technical improvements	Meat reduction	Increase in biogas use
	NH ₃ (65%) NO _x (60%)	CH ₄ (35%) NH ₃ (75%) NO _x (75%)	CH ₄ (2%) NH ₃ (10%) CO ₂ (negligible)
SHP	Scrubbers and particulate filters	Slow-steaming ^{d)}	Strong increase in LNG capacity
	SO ₂ (95%) ^{c)} NO _x (75%) BC (85%)	CO ₂ (35%) SO ₂ , NO _x , (35%) BC (20%)	CO ₂ (5%) SO ₂ , (90%) NO _x , (55%) BC (30%)

975

a) Here stationary combustion in power and industry.

976

b) Through use of recovered CH₄ instead of coal as fuel in oil, gas and coal industry.

977

c) The reduction level is based on a year 2015 baseline with relatively high sulfur content for international shipping

978

d) Assuming about 20% reduction in speed

979

980

981

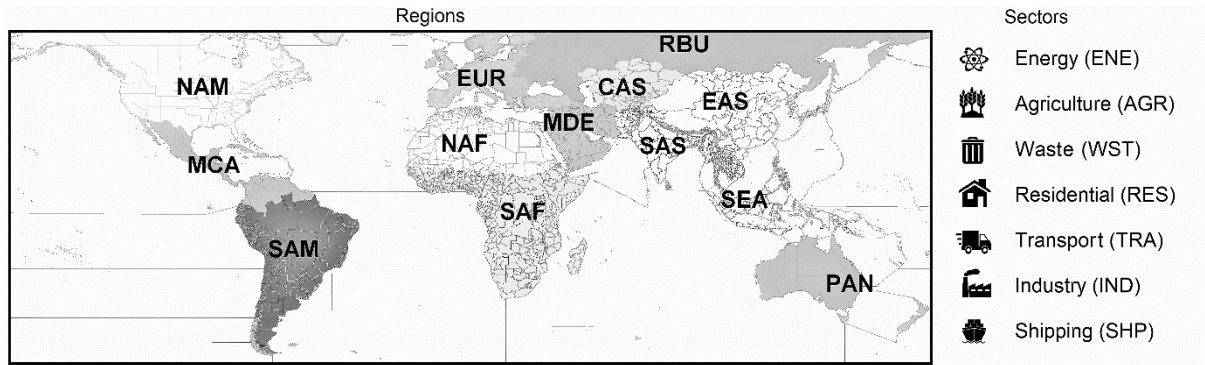
982

983

984

985 **Figures:**

986



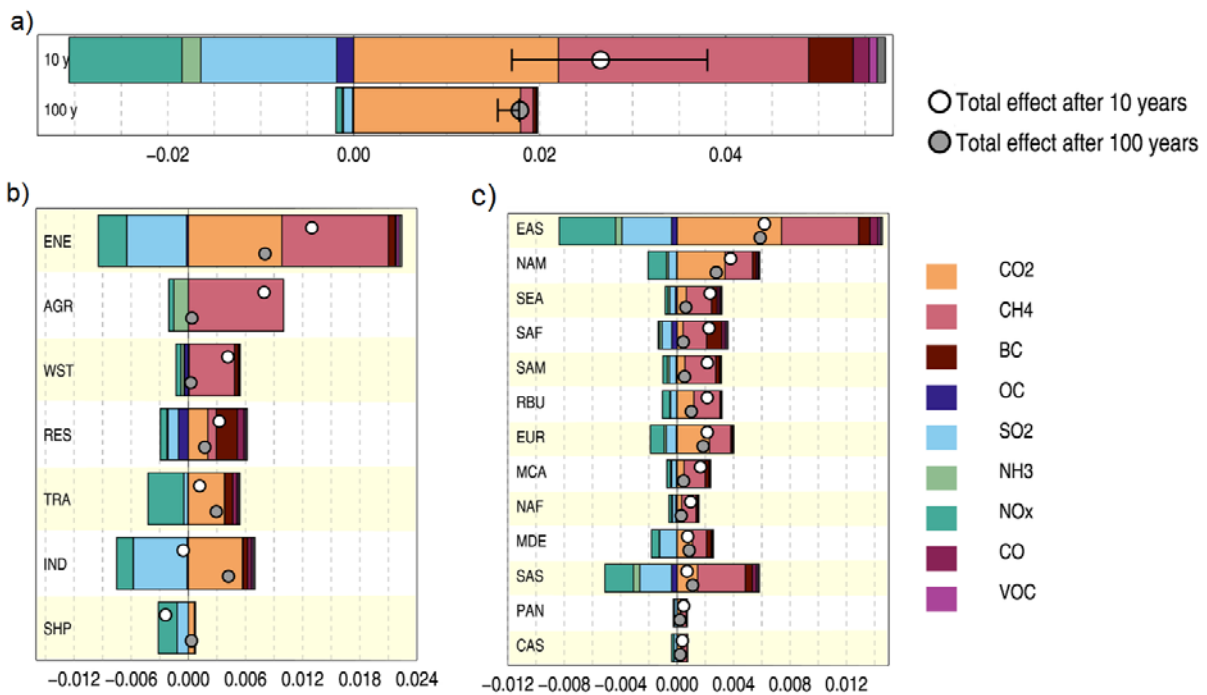
987

988 *Figure 1: Emission source regions and sectors used in the analysis.*

989

990

991

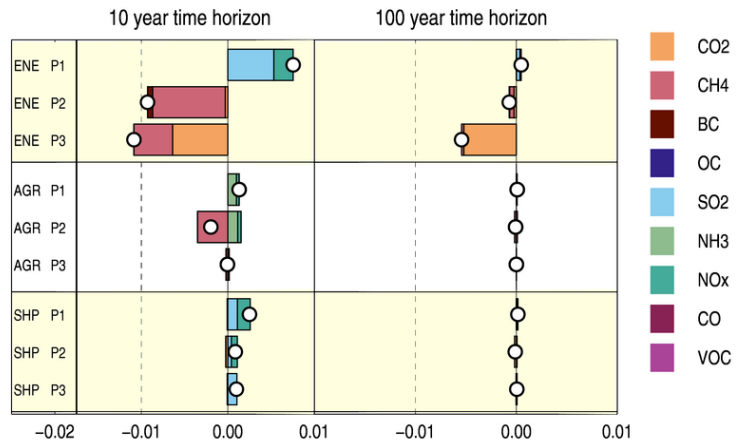


992

993 *Figure 2: Global-mean surface temperature impact 10 and 100 years after one year of present-*
994 *day (i.e., year 2014) emissions of SLCFs and CO₂ for: a) global total emissions, b) emissions*
995 *from seven major economic sectors, and c) total (i.e., sum of all sectors) emissions in 13 sources*
996 *regions. Panels b and c are sorted by total net effect on the 10-year timescale (white circle).*
997 *Error bars (± 1 standard deviation) in the top panel represent the range in total net temperature*
998 *impact due to uncertainties in radiative forcing.*

999

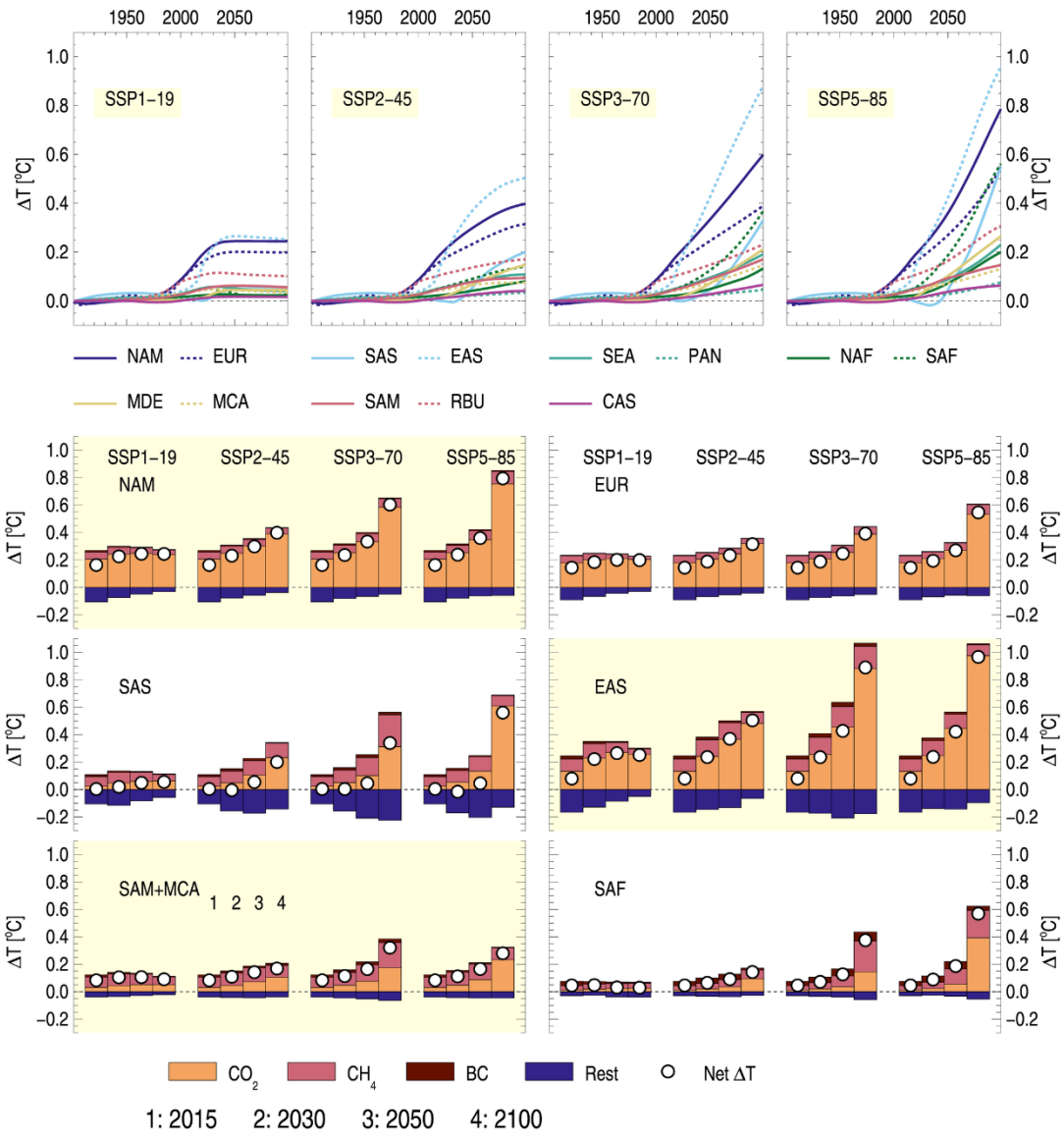
1000



1001

1002 *Figure 3: Global-mean surface temperature impact on 10 and 100 year time horizons resulting*
 1003 *from instantaneous reductions of different sets (listed in Table 2) of SLCFs and CO₂ emissions.*
 1004 *White circles indicate the net impact of these reductions.*

1005



1006

1007 *Figure 4: Global mean temperature response to historical emissions and future SSP pathways:*
 1008 *a) Net (i.e., sum over all species and sectors) response over the period 1900 to 2100 for each*
 1009 *region and scenario and b) net response in 2015, 2030, 2050 and 2100 to emissions in six*
 1010 *regions broken down by contributions from CO₂, BC, methane and the sum of SO₂, OC, NH₃*
 1011 *and ozone precursors (i.e., “Rest”).*

1012

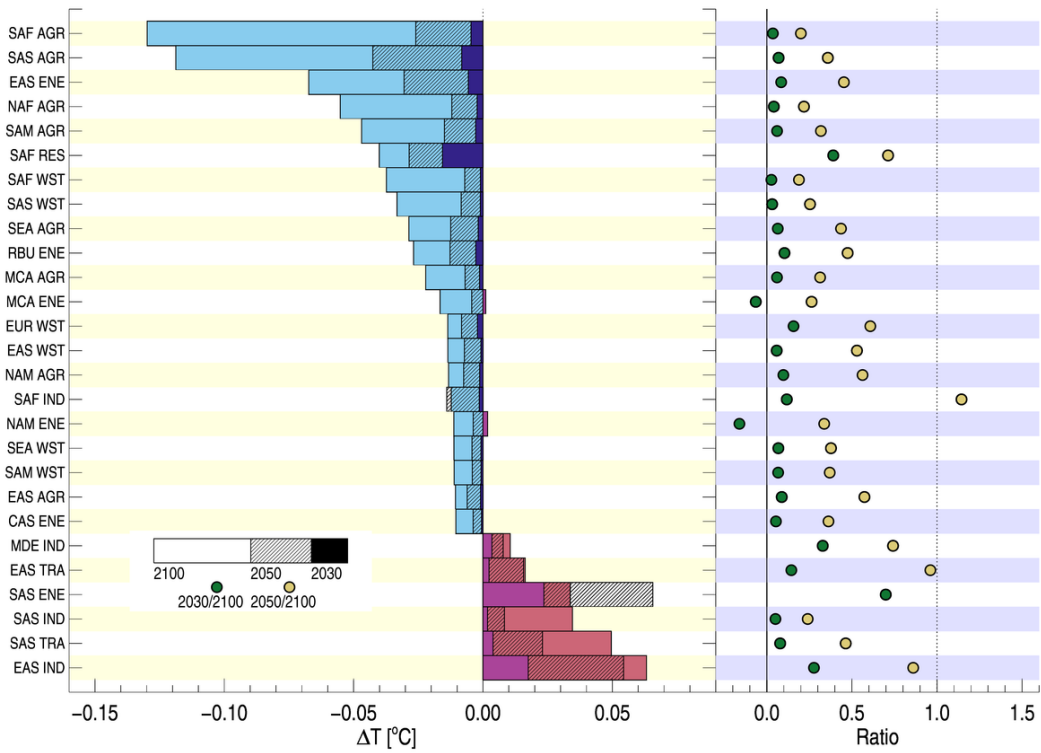
1013

1014

1015

1016

1017



1018

1019 *Figure 5: Difference in net SLCF (i.e., sum of all components except CO₂) temperature*
1020 *response between SSP1-1.9 and SSP3-7.0 in 2030, 2050 and 2100 by region and sector. Only*
1021 *combinations of sectors and regions where the differences in global temperature response is*
1022 *larger than ± 0.01 °C are shown. For each of these combinations, the panel on the right shows*
1023 *the ratio between the temperature response difference in 2030 and 2100 and between 2050 and*
1024 *2100.*

1025

1026

1027

1028

1029

1030

1031

1032

1033

1034

1035

1036