Daytime low-level clouds in West Africa – occurrence, associated drivers and shortwave radiation attenuation

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Abstract. This study focuses on daytime low-level clouds (LLC) that occur within the first 2 km of the atmosphere over West Africa (WA). These daytime LLCs play a major role in the earth's radiative balance, yet, their understanding is still relatively low in WA. We use the state-of-the-art ERA5 dataset to understand their occurrence and associated drivers as well as their impact on the incoming surface solar radiation in the two contrasting Guinean and Sahelian regions of WA. The diurnal cycle of the daytime occurrence of three LLC classes namely No LCC, LLC Class-1 (LLCs with lower fraction), and LLC Class-2 (LLCs with higher fraction) are first studied. The monthly evolutions of hourly and long-lasting LLC (for at least 6 consecutive hours) events are then analyzed as well as the synoptic-scale moisture flux associated with the long-lasting LLC events. Finally, the impact of LLC on the surface heat fluxes and the incoming solar irradiance is investigated. During the summer months in the Guinean region, LLC Class-1 occurrence is low while LLC Class-2 is frequent (occurrence frequency around 75 % in August). In the Sahel, LLC Class-1 is dominant in the summer months (occurrence frequency more than 80 % from June to October), however the peak occurrence frequency of Class-2 is also in the summer. In both regions, events with No LLC do not present any specific correlation with the time of the day. However, a diurnal evolution that appears to be strongly different from one region to the other is noted for the occurrence of LLC Class-2. LLC occurrence in both regions is associated with high moisture flux driven by strong southwesterly winds from the Gulf of Guinea and significant background moisture levels. LLC Class-2 in particular leads to a significant reduction in the upward transfer of energy and a net downward energy transfer caused by the release of large amounts of energy in the atmosphere during the cloud formation. In JAS, most of the LLC Class-2 events may likely be the low-level stratiform clouds that occur frequently over the Guinean region while they may be deep convective clouds in the Sahel. Additionally, LLC Class-2 causes high attenuation of the incoming solar radiation, especially during JAS where about 49 % and 44 % of the downwelling surface shortwave radiation is lost on average in Guinea and Sahel respectively.

1 Introduction

30 In West Africa (WA), the prediction of key features of the climate such as the West African Monsoon (WAM) is known to have large uncertainties (Christensen et al., 2013). Most of the uncertainties are essentially linked to the inability of climate models to efficiently represent convection and cloudiness, in particular low-level clouds (LLCs) (Hannak et al., 2017). Climate models struggle to properly reproduce the planetary boundary layer above the ocean and the continental surface as well as the land-ocean flux gradients that drive the triggering and the variability of LLC systems. These clouds and their temporal variation

35 strongly influence the reflectivity of the atmosphere. As a result, large uncertainties exist in the prediction of surface solar radiation in the region (Knippertz et al., 2011). In the light of the recent increased interest in solar energy projects in WA (World Bank, 2019) after the Paris Agreement in 2015, predictability of surface solar radiation (Armstrong and Hurley, 2010; Kosmopoulos et al., 2015) thus, remains a key issue for any feasibility study concerning solar energy during either dry or wet seasons in the region.

40 WA is characterized by a frequent occurrence of different cloud types. Near the Gulf of Guinea in the southern part of WA, LLCs are prevalent all year round (Schrage and Fink, 2012; van der Linden et al., 2015; Adler et al., 2017; Danso et al., 2019). Farther north in the Sahelian region, these clouds are frequently observed during the WAM season from July to September (Bouniol et al., 2012). These LLCs consist of stratified clouds, most of which are nocturnal low stratus clouds covering wide areas and persisting into the early afternoon (Schuster et al., 2013; Babić et al., 2019a) as well as shallow convective clouds.

45 LLCs composed of liquid water are well known to have a large impact on radiative transfer (Liou, 2002; Turner et al., 2007; Hill et al., 2018;). As a result, they drive the diurnal cycle of convection and the planetary boundary layer (Grabowski et al., 2006; Gounou et al., 2012). During the daytime, these clouds block a large portion of the incoming shortwave radiation reducing the irradiance at the earth's surface (Kniffka et al., 2019).

Several studies have been carried out in WA to understand LLCs, their variability, and the processes that aid in their formation,

50 maintenance, and dissipation (e.g., Schrage and Fink, 2012; Schuster et al., 2013; Kalthoff et al., 2018; Dione et al., 2019; Lohou et al., 2020; Pedruzo-Bagazgoitia et al., 2020). The majority of those studies were carried out within the framework of two major field campaigns; the African Monsoon Multidisciplinary Analysis (AMMA) (Redelsperger et al., 2006) in 2006 and the Dynamics–Aerosol–Chemistry–Cloud Interactions in West Africa (DACCIWA) (Knippertz et al., 2015) in 2016. Both of these field campaigns provided an unprecedented database of surface-based observations of different cloud properties. Other

55 studies on LLCs in the region have also been performed with satellite observations and model data (van der Linden et al., 2015; Adler et al., 2017; Hannak et al., 2017). These studies suggest that LLCs formation is linked to a number of processes including but not limited to cooling caused by horizontal cold air advection and the occurrence and strengthening of the nocturnal low-level jet (Schrage and Fink 2012; Adler et al. 2017; Babić et al. 2019b).

Most of these studies, however, focused on night-time LLCs in the southern part of WA (e.g., Schrage and Fink, 2012; Schuster

60 et al., 2013; Adler et al., 2017, 2019; Babić et al., 2019b). In the night, the nocturnal LLCs have no influence on the surface solar radiation due to the absence of sunlight. However, they persist long into the morning and early afternoon, thus, directly

influencing the amount of incoming solar irradiance and having a considerable impact on the energy balance of the earth surface as shown by Lohou et al., (2020). The conditions associated with these low clouds during the daytime in the region are less documented. As a result, our understanding of the impacts of these clouds on surface solar irradiance during the daytime is relatively limited. Moreover, most of the previous studies have been limited to the WAM season (e.g., Knippertz et al., 2011; Schuster et al., 2013; van der Linden et al., 2015; Dione et al., 2019). This is understandable as WA is cloudiest during the WAM season. However, it was shown that daytime LLCs exist in all seasons especially over southern WA (Danso et al., 2019). It is therefore important to understand the nature of such daytime LLCs over the whole region and not only during the WAM season. This is particularly challenging due to the lack of a long-term surface-based observational dataset in WA, thus, limiting our understanding of these cloud systems. Few studies which were done with simulations, reanalysis, and satellite data in the region (e.g., Knippertz et al., 2011; Schuster et al., 2013; van de Linden et al., 2015; Adler et al., 2017) have nevertheless shown results similar to those of ground observational studies (e.g., Adler et al., 2019; Babić et al., 2019b) to some extent.

In this general context, the main contribution of this study is to enhance our understanding of daytime LLCs in WA by focusing on daytime LLCs in both the dry and wet seasons with a state-of-the-art reanalysis dataset. Firstly, we will present the seasonal and diurnal distributions of the occurrence frequency of the daytime LLCs in two areas of WA with contrasting climate regimes (the Sahelian and the Guinean regions). We will then identify synoptic-scale conditions (specifically the moisture flux) associated with the presence of the daytime LLCs. Additionally, we will investigate the impact of LLC on surface heat fluxes and estimate the percentage of incoming surface shortwave (SW) radiation that is attenuated when these clouds are present.

This paper is organized as follows: Sect. 2 describes the study area, datasets, and methodology. Section 3 presents the diurnal and monthly distribution of occurrence of daytime LLCs while the synoptic and local conditions associated with the occurrence of long-lasting LLCs are discussed in Sect. 4. Section 5 presents the estimated attenuation of incoming shortwave radiation during the occurrence of LLCs. Finally, conclusions are drawn in Sect. 6.

2 Data and Methods

2.1 Data

In this study, we analyze the 5th generation of the European Centre for Medium-range Weather Forecasts (ECMWF) Re-analysis – ERA5[1] (Hersbach et al., 2020) to understand the occurrence of LLCs and their associated synoptic conditions and land surface fluxes and to discuss their impacts on incoming shortwave radiation in WA (Fig. 1). The reanalysis is available at a temporal resolution of an hour and has a native horizontal resolution of 0.25°. However, due to more emphasis on large-scale conditions in this study, the data was directly extracted on a regular 1° x 1° grid from ECMWF. A period of 10 years from 2006 to 2015 was analyzed for the study. For this period a total of 43800 data points consisting of only daytime hours from 0600 h to 1700 h, were thus considered.

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[1] Detailed documentation at https://confluence.ecmwf.int/display/CKB/ERA5+data+documentation

ERA5 provides different variables of cloud cover properties. These cloud variables are produced in CY41R2 of the ECMWF's Integrated Forecasting System (IFS) and documented in Part IV of the IFS Documentation[2]. The cloud cover variables consist of cloud fraction (CF) and cloud hydrometeors (liquid or ice water) and are available as 3D or 2D distributions. Here, we use the 3D distribution of cloud liquid and ice water content to show the atmospheric vertical profiles of clouds in selected areas over WA. The occurrence frequency of clouds may be defined by the hydrometeor content or the CF value. In this study, we determine the occurrence frequency based on the value of the 2D single-level CF, although, the liquid and ice water paths mostly determine the attenuation of incoming solar radiation in the atmosphere. The ERA5 CF depends on the subgrid-scale cloud scheme in the ECMWF IFS and can be non-zero when the cloud water content is zero (see Appendix A). The atmospheric conditions and radiative impacts are analyzed based on this occurrence In this study, we focus on low clouds defined by ECMWF in the ERA5 product as all clouds integrated from the surface to 2 km altitude in the atmosphere ($\approx$ 800hPa). This includes all clouds with bases below 2 km.

Other ERA5 variables are analyzed to show some of the atmospheric and surface conditions during the occurrence of LLCs in order to understand the possible interactions between the surface and the lower atmosphere. These include specific humidity, zonal and meridional wind components, and the surface sensible and latent heat fluxes. Furthermore, we use the incoming downwelling shortwave radiation variable in ERA5 to estimate the percentage of incoming solar radiation that is attenuated in the presence of LLCs. We first performed a detailed evaluation of the ERA5 cloud fraction and radiation data with surface and satellite-based observations for some locations in WA. This evaluation revealed a reasonably good performance by ERA5 in reproducing the observations (see supplementary material).

2.2 Methods

2.2.1 Identification of LCC occurrence

Our analysis is based on the occurrence of LLCs in two sub-regional windows over WA (rectangular boxes in Fig. 1) denoted here as Guinea (10° W to 0° E and 6° N to 9° N) and Sahel (10° W to 0° E and 12° N to 15° N). These areas are chosen primarily due to their contrasting climate regimes (Gbobaniyi et al., 2014). The Guinea region which includes countries like Ghana and Cote d'Ivoire is currently undergoing significant economic and infrastructural developments and thus the need for clean energy keeps increasing. Consequently, there are many ongoing solar energy projects in this region such as the 155 MW Nzema solar PV power project in Ghana (Kuwonu, 2016), which when completed will be the largest solar farm on the African continent. Côte d'Ivoire has planned to increase its share of variable renewable energy up to 11 % by the end of 2020 and up to 16 % by 2030. Thus, a 25 MW solar power plant is under construction in the Northern part of the country in Korhogo while two solar power plants projects with a capacity of 30 MW each have been approved and will be built in the localities of Tuoba and Laboa in the northwestern part of the country (MPEER, 2019). Additionally, there are several planned and existing solar

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[2] Available at https://www.ecmwf.int/en/elibrary/16648-part-iv-physical-processes

energy projects in the Sahel window, especially in Burkina Faso with the 33 MW-capacity Zagtouli solar power plant (Moner-Girona et al., 2017) which is currently the largest solar energy project in the Sahelian part of WA.

Occurrence or no-occurrence of LLC are first identified from the extracted data points based on the quantitative value of the CF in the two selected windows. This is used to show the diurnal and monthly distribution of the LLC occurrence. All timestamps with zero CF in the lower atmospheric level (up to 2 km) are designated as No LLC. Thus, hourly events of LLC occurrence are all timestamps with non-zero CF values including those with very small CF values. LLC events with very low CF values may show no or weak signals of the conditions that exist during their occurrence. Moreover, cloud radiative forcing is strongly dependent on the fractional coverage of clouds (Liu et al., 2011). Each LLC occurrence event is therefore classified as one of two classes depending on the CF value for that event. The definitions of the classes are provided in Table 1.

Though the focus of this study is on LLC occurrence and the definition of the different classes is based on the CF in the lower atmosphere only, the convoluted cloud climatology in WA with frequent multilayer clouds (Stein et al., 2011) may also mean that other clouds at higher altitudes may exist in addition to LLCs. Fig. 2 illustrates this multilayer nature of clouds in WA in connection with LLC occurrence, by showing the mean (of all hourly events) vertical profile of cloud water content of each class in the two selected regions. For LLC Class-1 and Class-2 in both windows, the trimodal vertical distribution of tropical convection (Johnson et al., 1999) indicates the presence of higher altitude clouds in addition to the LLCs. In the case of No LLC, high- and mid-level clouds could exist but with very low water content. This suggests and agrees with the satellite-based climatology of Stein et al., (2011) that LLCs frequently exist together with upper-level clouds in WA.

2.2.2 Identification of synoptic conditions during LLC occurrence

Recent studies (e.g., Adler et al., 2019; Babic et al., 2019b) have identified some physical processes relevant to the formation and maintenance of LLC in WA. They include but are not limited to the strengthening of the nocturnal low-level jet, horizontal cold air advection and background moisture level. Babic et al., (2019b) suggested the latter could play an important role in the formation of LLCs but this has not received as much attention as the other processes. Based on the different classes defined in Table 1, the regional atmospheric circulation and moisture flux were analyzed for the 10-year period of this study. For that, a composite approach was used to produce robust results based on a large number of events. We only take into account the occurrence of the different LLC classes that last for a sufficiently long time for their associated mean atmospheric and surface conditions to be significant at the regional scale. Therefore, an event here refers to the occurrence of a given LLC class that lasts for at least six consecutive hours during the daytime. Composite analysis of the moisture flux (Q_{flux}) and wind at 950 hPa was performed. Here, Q_{flux} is defined as:

$$Q_{flux} = \frac{1}{N}\sum_{i=1}^{N} q_i v_i \qquad (1)$$

where q is the specific humidity (in $g \cdot kg^{-1}$), v is the meridional wind component (in $m\,s^{-1}$) and N is the total number of events in a given LLC occurrence class. Since q is always positive, Q_{flux} can either be negative or positive depending on the

wind direction, with positive or negative values indicating northward and southward air movements respectively. Over the sea, positive values of Q_{flux} from the ocean toward the continent thus suggest that moisture is likely transported onto the land along with the wind.

Lohou et al., (2020) have shown that LLC occurrence also has a considerable impact on the surface heat fluxes in WA. Here, we examine the local surface sensible and latent heat fluxes in the selected windows to understand the daytime interaction of the land surface with the boundary layer and LLC occurrence. Due to the definition of LLCs used, the present study does not consider the impact of the atmospheric boundary layer (ABL) that may considerably modify the altitude of the low cloud. Indeed, ABL clouds have their bases higher than 2 km especially in the Sahel region during the dry season. In addition, some studies (Lohou et al., 2020; Pedruzo-Bagazgoitia et al., 2020; Zouzoua et al., 2020) have shown that some LLCs can be decoupled from the surface and are therefore not influenced by the surface heat fluxes. Therefore, having a higher cloud base threshold for LLCs may alter the analysis of the surface heat fluxes. The analysis was performed for two seasons representing two contrasting climate regimes associated with the southernmost position of the Inter-Tropical Convergence Zone (ITCZ) over Guinea (in December, January, February (DJF)) and the northernmost position of the ITCZ over Sahel (in July, August, September (JAS)).

2.2.3. Determination of cloud shortwave attenuation effects

To investigate the quantity of incoming downwelling shortwave radiation attenuated in the presence of LLCs (in $W\ m^{-2}$) at a given time t, the cloud shortwave radiative effect ($CRE_{SW\downarrow}$) was estimated by finding the difference between the downwelling shortwave radiation in all-sky conditions ($SW^{\downarrow}$) and the corresponding downwelling shortwave radiation in clear-sky conditions ($SW^{\downarrow}_{CS}$). The $SW^{\downarrow}_{CS}$ which is directly available in ERA5, is a theoretical value computed at a given time based on the same atmospheric conditions (temperature, humidity, aerosols, etc.) as the $SW^{\downarrow}$ but assuming there are no clouds. To eliminate the influence of the wide range of daytime hours on its mean value, we expressed $CRE_{SW\downarrow}$ as a percentage of the attenuated $SW^{\downarrow}_{CS}$. Since $SW^{\downarrow}$ is usually lower than $SW^{\downarrow}_{CS}$, the mean $CRE_{SW\downarrow}$ over all events in a given LLC class will be negative. To avoid presenting a negative mean $CRE_{SW\downarrow}$, the absolute of this value is used. Thus, the mean $CRE_{SW\downarrow}$ of all events for a given LLC class is given by:

$$CRE_{SW\downarrow} = \left| \frac{1}{N} \sum_{i=1}^{N} \frac{SW^{\downarrow}(i) - SW^{\downarrow}_{CS}(i)}{SW^{\downarrow}_{CS}(i)} \times 100 \right| \qquad (2)$$

3. Temporal distribution of occurrence of daytime LLCs

The monthly distribution of the occurrence frequency (hourly) of the three LLC classes is presented in Fig. 3 for the two selected regions (occurrences that last for at least six consecutive hours are in shown as black bars). In Guinea, events of No LLC are very rare (less than 200 in the 10 years data) and occur only in the DJF. On the other hand, LLCs occur in all months in this area. The occurrence of LLC events with lower fractional coverage (LLC Class-1) shows a clear seasonal cycle with

the highest number of events in the dry and transitional seasons and the lowest number of events during JJAS which corresponds to the core of the monsoon season (Sultan et al., 2003). Coincidentally, LLC events with higher fractional coverage

185   (LLC Class-2) occur most during the monsoon season. This period is associated with increased moisture influx from the Gulf of Guinea (Sultan and Janicot, 2003; Thorncroft et al., 2011) and cold air advection which enhances low clouds formation over the region during this period as shown by Adler et al., (2019) and Babic et al. (2019b). Interestingly, the peak of LLC Class-2 events occurs in August, the so-called "little dry season" in southern WA (Adejuwon and Odekunle, 2006; Chineke et al., 2010; Froidurot and Diedhiou, 2017). Whether this peak of LLC Class-2 events in August which coincides with reduced

190   rainfall in southern WA is due to an increased occurrence of non-rain-bearing low clouds (and a decrease in rain-bearing ones) or not is beyond the scope of this paper but needs to be investigated in future.

In Sahel, events of No LLC are rather frequent (unlike Guinea) and occurs in every month except in August when the monsoon is fully developed over the Sahelian region of WA (Sultan et al., 2003; Thorncroft et al., 2011). Similarly, LLC Class-1 events are frequent throughout the year but with the highest occurrences from July to October (with a slight decrease in August) and

195   lowest occurrences during the dry season. LLC Class-2 events are marked with a strong seasonal signature, occurring only from July to October with a distinct peak occurrence in August when the Inter-Tropical Convergence Zone (ITCZ) is at its northernmost latitude over WA (Sultan and Janicot, 2000; Nicholson, 2009, 2018). This period also coincides with the maximum rainfall over the Sahelian region (Nicholson, 2013) and suggests a strong link between the occurrence of LLC Class-2 events and the Sahelian rainfall. This is an indication that LLC Class-2 in the Sahel are likely related to rainfall events in the

200   region. Again, Mathon et al., (2002b) showed that cloud cover associated with a sub-population of Mesoscale Convective Systems (MCS) in the Sahel is modulated at the synoptic-scale during the African easterly wave activity, with an increase of the cloud cover in and ahead of the trough and maximum rainfall behind of the wave trough.

The daytime distribution of occurrence frequency of the three LLC classes is presented in Fig. 4 for the two selected areas. The distribution presents the total occurrence frequency of all seasons combined. In Guinea, the occurrence frequency of No

205   LLC events does not present a very well-marked diurnal cycle with a slightly higher number of events occurring at 1700 UTC, though these events are extremely rare (less than 25 for the 10 years). In contrast, both LCC Class-1 and Class-2 occurrences present well-defined diurnal cycles. The occurrence of LLC Class-1 events is low during the early morning (lowest at 0700 UTC) and increases progressively throughout the day to late in the afternoon. On the contrary, LLC Class-2 events are high early in the morning (highest at 0700 UTC) but decrease continuously over the course of the day to the late afternoon (1700

210   UTC). The early morning peak in the LLC Class-2 events could be as a result of the contribution of residuals from nocturnal low-level stratus clouds which persist long into the day (Kalthoff et al., 2018). These clouds usually have large fractional coverage during their occurrence in the night (van der Linden et al., 2015) but start to dissipate during the daytime (Kalthoff et al., 2018). The continuous decreasing and increasing in the event number of LLC Class-2 and LLC Class-1 respectively, from morning to late afternoon could, therefore, be linked to the gradual dissipation of those nocturnal low-level stratus clouds

215   during the day. Additionally, the early morning peak in the events of LLC Class-2 could also be partly linked to contributions from tropical oceanic low-level convection which is maximum during the early morning (Yang and Slingo, 2001). The

proximity of the Guinean region to the ocean means that cold air over the Gulf of Guinea could be advected inland which will, in turn, enhance LLC formation.

In the Sahel region, events of No LLC occur without a well-marked diurnal cycle but are slightly lower around midday. Similarly, LLC Class-1 events do not present a very distinct diurnal cycle. On the other hand, events of LLC Class-2 show a clear and well-marked diurnal cycle with a higher occurrence frequency in the late morning to midday (highest at 1100 UTC) and lower values during early mornings and late afternoon. This tendency seems to be strongly controlled by surface processes such as continental daytime solar heating (Mallet et al., 2009). Moreover, it could further be related to the occurrence of deep MCSs whose convective parts will be considered as low clouds if found within the lowest 2 km of the atmosphere. These MCSs present a fairly similar evolution (both seasonally (Fig. 3) and diurnally) as the LLC Class-2 in this region of WA (Vizy and Cook, 2018). Therefore, there is a higher likelihood of an MCS in the LLC Class-2 events than Class-1 and it is reasonable to assume that most of the LLC Class-2 events in the Sahel are the well documented deep MCSs (Figure A1 in Appendix A) that are responsible for around 90 % of total rainfall in the region (Mathon et al., 2002a; Goyens et al., 2012; Vizy and Cook, 2019).

4. Synoptic conditions related to the occurrence of daytime LLCs

The synoptic conditions associated with the occurrence of the different LLC classes are mainly presented for the JAS and DJF seasons. The total number of daytime hours in JAS and DJF during the 10 years are 11040 and 10800 respectively. Only occurrences that last for at least six consecutive hours are considered to analyze these conditions (i.e., here, an event refers to an occurrence that persists for at least six consecutive hours). Table 2 presents the number of the 'six-consecutive hours' events extracted for each class and for both seasons over the 10 years. During the JAS season, no occurrence of No LLC events was detected in both Guinea and Sahel window (i.e., daytime LLCs are always present in both regions during the JAS). LLC Class-1 events dominated the Sahel region during JAS (with 793 events) but on the contrary, LLC Class-2 events dominated the Guinea region (with 852 events). In DJF, No LLC events dominated the Sahel (with 877 events) and no LLC Class-2 occurrence has been identified in this area during this season. In Guinea, on the other hand, LLC Class-1 occurrences dominated the area during DJF (with 878 events). Composites of the different variables are thus analyzed and then discussed only based on the extracted events for each class and each season, as presented in Table 2.

4.1. Atmospheric circulation and moisture flux

The occurrence of daytime LLC events in the selected windows is related to large-scale dynamics of the atmosphere over the WA region. Fig. 5 and 6 present the composites of Q_{flux} and the atmospheric circulation at 950hPa for JAS and DJF respectively, and for the two studied regions. During JAS in both regions (Fig. 5), the occurrence of LLC Class-1 and Class-2 events is related to high moisture flux dominated by strong southwesterly winds (positive Q_{flux}) from the Guinean Gulf. The moisture flux is stronger during the occurrence of LLC Class-2 events than LLC Class-1 events. This is true for LLC occurrence

in both regions; however, the moisture flux is more intense for LLC occurrence in Sahel in terms of its inland penetration. There are no events of No LLC in both Sahel and Guinea during JAS. During DJF in both regions (Fig. 6), events with No LLC are characterized by dry northeasterly winds that inhibit LLC formation. The occurrence of LLC Class-1 events in Guinea and Sahel is characterized by southerly winds just at the southern coast but do not penetrate the area of the two windows. Nevertheless, both LLC Class-1 and Class-2 events are associated with significant specific humidity (true for both seasons) unlike No LLC events with very low specific humidity (see Figure B1 in Appendix B). This is in agreement with Babic et al., (2019b) who found that the most distinct difference between LLC and No LLC events is in the specific humidity, suggesting that the background moisture may play a crucial role in the formation and maintenance of low clouds in WA. Additionally, the LLC Class-1 events during DJF (in Sahel) are related to positive anomalies of specific humidity that seem to have been transported onto the continent from the North Atlantic (not shown). LLC Class-2 events in Guinea are associated with a much stronger moisture flux in the region (Fig. 6) as well as a positive anomaly of specific humidity within and around the window (not shown).

The moisture fluxes presented so far are those present when LLC events are identified. However, it is also interesting to investigate the moisture flux before and after the LLCs occurrence in order to analyse their contribution to the maintenance or triggering of these systems. To do this, we analyze the composite moisture flux (at 950 hPa) of the first hour just before the LLC event ($1h_{Before} LLC$) and the first hour just after the LLC disappears ($1h_{After} LLC$), and the difference between the two periods. This is illustrated for LLC Class-2 during the JAS season in both Guinea and Sahel (Fig 7). In both regions, the moisture flux during $1h_{Before} LLC$ is larger than $1h_{After} LLC$ and there is a significant difference between the two periods as shown in Fig. 7. This is also true for the composites of specific humidity only (not shown). During $1h_{Before} LLC$, the water vapour associated with the moisture flux is likely cooled by some cold air advection into the region which leads to saturation and consequently, LLC formation. Indeed, the vertical profile of the potential temperature in the lower atmosphere during the two periods shows a strong temperature decrease just before the LLC occurrence and an increase after its disappearance (Figure B2 in Appendix B). This result, which is in agreement with the finding of Babic et al., (2019b), may be an indication of the cooling caused by the cold air advection. Over the Guinea region, studies have shown that cooling by the horizontal advection of cold air from the Gulf of Guinea contributes significantly to the formation of LLC (e.g., Adler et al., 2019; Babic et al., 2019b). In the Sahel region, moist air advection may be crucial for LLC formation as suggested by Parker et al. (2005) and Monerie et al. (2020). However, before LLC formation in the Sahel region, there may also be some cooling due to meridional advection (Marsham et al., 2013) as well as cold air advection from the North Atlantic, which could explain the strong temperature reduction during $1h_{Before} LLC$ (Figure B2).

4.2. Surface heat fluxes

Fig. 8 and 9 respectively present the monthly mean and anomalies of the surface heat fluxes, computed over the two selected regions for the three LLC occurrence classes. For a particular hour *i*, on a given day and month during the 10 years, the heat

280 flux anomaly is the difference between the actual heat flux value and the 10-year mean. For example, the heat flux anomaly at 12 UTC on 1st January 2006 was determined by taking the difference between the actual heat flux at 12 UTC on 1st January 2006 (HF_i) and the 10-year mean of all heat flux values at 12 UTC on 1st January ($\overline{HF}$) i.e.,

$$HF_{anom_i} = HF_i - \overline{HF} \tag{3}$$

where, HF_{anom_i} is the heat flux anomaly. The mean monthly anomaly is then computed by finding the mean of hourly
285 anomalies in the month. The ECMWF convention for vertical fluxes including the sensible and latent heat fluxes, is positive downwards and negative upwards.

No LLC events occur almost throughout the entire year in the Sahel (except JAS) and only during January in the Guinean region. In both areas, the occurrence of these events leads to high negative sensible heating (Fig. 8) indicating an upward transfer of sensible heat. Indeed, during No LLC events, much of the incoming radiation heats the surface, making it warmer
290 than the air directly above it, leading to an upward sensible heat flux. Both regions are characterized by negative anomalies of sensible heat (Fig. 9) suggesting a much stronger than usual upward transfer of sensible heat. The mean latent heat fluxes during No LLC events are also negative but significantly lower than the sensible heat (Fig. 8). Thus, most of the energy from the surface is used to heat the atmosphere rather than evaporating moisture. This could be as a result of a low moisture gradient between the soil and the air during those events especially in the dry season when both soil and air moisture are extremely low
295 in the region. The lack of moisture in the region inhibits the formation of LLC during these events. In the Sahel region, there are positive anomalies of latent heat (from April to June), which indicates a downward transfer of energy from the atmosphere to the surface.

During LLC occurrence, a portion of the incoming SW radiation is blocked. Therefore, the warming of the surface is reduced during LLC occurrence in both regions. However, due to the higher cloud fraction during LLC Class-2 events, a larger quantity
300 of the downwelling shortwave radiation is blocked. Consequently, the associated surface warming is much lower during LLC Class-2 than the Class-1 events (lower negative sensible heat fluxes than Class-1 events in Fig. 8a). This result suggests that the LLC Class-2 events are most likely the low stratiform clouds which are known to cover large portions of the sky or the deep convective clouds which extends to the upper levels of the atmosphere, generating anvils in the process. With regards to the former, Lohou et al., (2020) have shown that these low stratiform clouds have significant impacts on the energy balance
305 of the earth's surface during the daytime in the Guinea region. The latter has also been shown to have large impacts on the incoming shortwave radiation, especially in the Sahel region during the monsoon season (Bouniol et al., 2012).

In terms of anomalies, the LLC Class-2 events present positive values (Fig. 9a) over both regions suggesting a net downward transfer of sensible heat. This could be explained by two main reasons. Firstly, a larger portion of the sky is blocked during these events leading to lower sensible heating at the surface. Additionally, large amounts of energy are released in the lower
310 atmosphere during the formation of these clouds, which in turn leads to a higher sensible heat flux in the atmosphere than when there are no LLCs or when the cloud fraction is low. On the other hand, LLC Class-1 in the Guinean region presents negative anomalies of sensible heat that suggest a net upward transfer of sensible heat during these events. As these LLCs do

not cover a large portion of the sky, a large amount of shortwave radiation still warms the surface. Moreover, the energy released in the atmosphere is not as much as for LLC Class-2 events. Thus, during these events, the atmosphere has a relatively lower sensible heat flux. It is rather surprising that the Sahel region presents positive anomalies (except in JAS) for sensible heat. As suggested by Bouniol et al., (2012), such a result should not be interpreted exclusively in terms of the radiative effects of clouds but rather the coupled interactions between fluctuations in temperature, water vapor, and clouds in this region. However, this is beyond the scope of the present study.

During LLC occurrence, the large mean negative latent heat values indicate more evaporation from the surface (Fig. 8b). Interestingly, the transfer of latent heat upwards is higher in Class 1 than in Class 2. Since only a small portion of the sky is covered during LLC Class-1 events, the surface still receives a significant amount of downwelling shortwave radiation. Therefore, the surface continues to be warmed better by the sun's radiation during LLC Class-1 events than Class-2 events. As a result, the evaporation of moisture at the surface is much higher during the occurrence of the former. There are negative anomalies of the surface latent heat (Fig. 9b) in both regions during LLC Class-1 events in all months (except in November and December in the Sahel). This signifies a more intense upward transfer of latent energy from evaporation during these events. On the other hand, LLC Class-2 events are associated with positive anomalies of latent heat. This is likely due to two main reasons. Firstly, the reduced incoming downwelling shortwave radiation decreases evaporation at the surface. Secondly, large amounts of latent heat of condensation are released in the atmosphere during the formation of the clouds. There is therefore a greater amount of latent energy in the atmosphere than at the surface and a net downward transfer of this energy.

5. Attenuation of incoming shortwave radiation during LLC occurrence

In Fig. 10 and 11, the regional distribution of cloud radiative effect on incoming downwelling shortwave radiation during JAS and DJF is presented respectively as the percentage of incoming shortwave radiation attenuated during the occurrence of LLCs in the two regions as computed with Eq. (3) ($CRE_{SW\downarrow}$). In this section, we will focus more on the mean $CRE_{SW\downarrow}$ in each of the regions. Firstly, there is a strong seasonal variability of $CRE_{SW\downarrow}$ over the entire WA region. It is obvious that shortwave attenuation is much higher during JAS than during DJF over the entire region as a result of increased cloud coverage during the monsoon season (Bouniol et al., 2012; Danso et al., 2019). This is true whether LLCs are present or not. During the occurrence of LLCs, the values of $CRE_{SW\downarrow}$ are much lower for LLC Class-1 than LLC Class-2 events, whatever the region and the season. This clearly shows the strong dependence of $CRE_{SW\downarrow}$ on the fractional coverage of clouds as previously discussed by (Liu et al., 2011).

In JAS, the strong impacts of LLCs on incoming shortwave radiation is seen from the high $CRE_{SW\downarrow}$ values during LLC occurrence. During the occurrence of LLC Class-1 events, the mean percentage attenuation of incoming shortwave radiation becomes 29.9 % and 17.5 % respectively in the Guinean and Sahelian regions. In the same way, 49.1 % and 44.2 % of incoming shortwave radiation are attenuated on average in Guinea and Sahel, respectively during events of LLC Class-2. Though these values are attenuated in the two regions under consideration, higher losses can be experienced in different areas over the region

345   during those events. For instance, during LLC Class-2 events in the Sahel, the areas around the southern coast and the Guinea highlands experience much higher $CRE_{SW\downarrow}$ values (up to 65 %). This could be due to the presence of more clouds in those areas which extend into the Sahel region.

During DJF, during events of No LLC in Guinea and Sahel, about 2.2 % and 2.3 % (Fig. 11) of the incoming shortwave radiation is lost on average respectively. Even though there is no LLC occurrence during these events, this attenuation value is due to the possible presence of other cloud types at the upper levels. However, similar to the JAS season, the occurrence of

350   LLCs leads to higher $CRE_{SW\downarrow}$ in both regions. In Guinea, 13.3 % of incoming shortwave radiation is lost on average during the occurrence of LLC Class-1 events while 32 % is lost during LLC Class-2 events. In Sahel, the mean attenuation during the occurrence of LLC Class-1 events is 16.3 %.

The large attenuation of incoming shortwave radiation during LLC occurrence (especially LLC Class-2) could be detrimental

355   to planned solar energy application projects if not properly planned. For instance, during the JAS season in the Guinea region (and southern part of WA in general), the dominance of LLC Class-2 events (Table 2 and Fig. 3) means that surface solar radiation will be significantly reduced for a large number of periods. Although there are equally high losses of incoming solar radiation in the Sahel during LLC Class-2 events in JAS, these events are much lower than in the Guinea region. Besides, the Sahel region has several persistent periods with no LLCs compared to the Guinea region in the other months of the year,

360   especially during DJF. However, other issues that are not discussed in this paper such as dust (e.g., Bonkaney et al., 2017) could severely hinder solar energy projects in the Sahel region. In all cases, the implementation of large-scale solar energy projects in both regions requires meticulous planning to ensure that the intense cloudiness over the region and the high variability of these clouds as shown by Danso et al., (2019) do not compromise the long-term aims of those developments.

6. Conclusion

365   In this paper, we made use of the state-of-the-art hourly reanalysis dataset from ECMWF – ERA5 from 2006 to 2015 to analyze the occurrence of daytime LLCs in WA. The analysis was performed for both the wet (JAS) and dry (DJF) seasons and focused on two climatically contrasting areas in the region (i.e., Guinea and Sahel). We have first identified events of LLC occurrence in both regions. We then analyzed some regional- and local-scale environmental conditions which occur during those events of LLC occurrence. This analysis mainly focused on the moisture flux before, during and after the LLC occurrence. We also

370   analyzed the impacts of the LLC occurrence of on the surface heat fluxes. Finally, the attenuation of incoming shortwave radiation during the occurrence of LLC in the region was estimated. To account for the influence of fractional coverage of LLC on the attenuation of incoming shortwave radiation, we classified the events of LLC occurrence into two groups: events with low fractional coverage (LLC Class-1) and events with high fractional coverage (LLC Class-2). The main findings of the study are summarized below:

375   -   The occurrence frequency of daytime LLCs is characterized by significant diurnal and seasonal variations: (i) Over Guinea, the occurrence of LLC Class-1 events is lower during the core of the monsoon season (JJAS) while LLC

Class-2 events are much higher during this period. During the day, LLC Class-2 events occur frequently during the early morning while LLC Class-1 is rather frequent during the late afternoon. (ii) Over Sahel, the occurrence of both LLC Class-1 and LLC Class-2 events is higher during the core of the monsoon season. Occurrence frequency of LLC Class-2 events is higher around midday while for LLC Class-1 events, it does not show a strong diurnal variation.

- During JAS, the occurrence of LLCs (both Class-1 and Class-2) in both Guinea and Sahel is associated with high moisture flux dominated by significant background moisture and strong southwesterly winds from the Gulf of Guinea. We also found that just before the occurrence of LLC, the temperature in the lower atmosphere is significantly reduced (Appendix B). This is due to cold air advection which leads to saturation as shown in previous studies (e.g., Adler et al., 2019 and Babic et al., 2020a).

- In the dry season, the moisture flux is not as intense as in JAS but the background moisture content is significant during occurrence of LLCs. The region is dominated by dry northeasterly winds during this period. Other processes such as the atmospheric waves and jets in the region (African Easterly wave, African Easterly Jet, etc.) may also contribute to the occurrence of LLCs but these are not analyzed in this study. These processes are, however, only relevant in the wet season.

- The occurrence of LLCs has significant impacts on surface heat fluxes. LLC Class-2 events lead to a significant reduction in the transfer of energy (both sensible and latent) upward and significant release of energy in the atmosphere which causes a net downward transfer of energy.

- Attenuation of incoming shortwave radiation is higher during JAS than DJF and depends on the fractional coverage of LLCs. In JAS, the occurrence of LLC Class-1 leads to mean attenuations of 29.9 % and 17.5 % respectively in Guinea and Sahel. LLC Class-2 leads to mean attenuations of 49.1 % and 44.2 % in Guinea and Sahel respectively. In DJF, mean attenuations of 13.3 % and 32 % are experienced in Guinea for LLC Class-1 and Class-2 events respectively. In Sahel, the occurrence of LLC Class-1 events leads to a 16.3 % attenuation of incoming shortwave radiation. It should however be noted that ERA5 slightly overestimates cloud fraction in the lower atmosphere over WA (see supplementary material).

By using the ERA5 cloud cover dataset in this study, we have provided important information on daytime LLCs in WA – a region notably known for its lack of surface observations. Previous studies have mostly focused on nighttime LLCs in the southern part of WA and only during the WAM season. Here, we expanded to the Sahel region and investigated the occurrence of LLCs during the dry season as well. The study helps to identify some conditions such as the background moisture level which is known to contribute to LLC formation in the region. Again, this study also helps to provide an idea on the impacts of these clouds on the surface energy budget especially on the attenuation of incoming shortwave radiation during LLC occurrence. The results on the seasonal and diurnal distributions of LLCs as well as on that on the attenuation of incoming shortwave radiation during LLC occurrence could be useful when making feasibility studies for large-scale solar energy projects.

Appendix A: Vertical profiles of cloud hydrometeor content during LLC occurrence

The occurrence frequency of clouds may be defined based on the value of the cloud water content (liquid and ice) or the cloud fraction (CF). In this study, we defined the occurrence frequency of low-level clouds (LLC) based on the value of CF. LLC occurrence was first classified as Class-1 ($0 < CF \le 0.5$) or Class-2 ($CF > 0.5$). In ERA5, CF can be larger than zero for zero liquid or ice water content. This is illustrated in Figure A1 which shows the vertical profiles of cloud water content during the occurrence of random LLC Class-1 and Class-2 events. It can be seen in some of the profiles (for Class-1) that the liquid and ice water content in the lowest 2km is null (but these events have $CF > 0$ in the lowest 2km in the profile).

Fig. A1 is also used to suggest the kind of cloud types that may be dominant in each class. In this case, we say that, in the Sahel region, there is a higher probability of getting a Mesoscale Convective System (MCS) in the LLC Class-2 than Class-1. Fig. A1 is the vertical profile of 12 random events of LLC Class-1 and Class-2 in the Sahel region. Based on the work of Storer and Van den Heever (2013), we designate an event as a 'MCS event' when there is a column containing at least 0.01 g·kg^{-1} cloud water (black dashed vertical line in Fig. A1) through a continuous depth of 8 km. From the twelve random profiles in Fig. A1, there are at least seven LLC Class-2 events (and just one for LLC Class-1) that satisfy this MCS definition. The number LLC Class-2 events are extremely low compared to LLC Class-1 (see Fig. 3), therefore, the likelihood that an event will be an MCS is relatively higher in the LLC Class-2 events.

Appendix B: Specific humidity associated with LLC occurrence and potential temperature before and after LLC formation

Figure B1 shows the composites of specific humidity (colours) and winds (vectors) during the occurrence of the different LLC occurrence classes in both Guinea and Sahel for the DJF season. During No LLC events, the moisture content in the region of the two windows is very low or zero. However, during LLC occurrence, the moisture content in the region is significant. This indicates that the background moisture plays an important role in the occurrence of these clouds. This result is confirmed in the recent study of Babic et al., (2019b) who found that in southern WA, the most distinct difference between LLC and No LLC events is in the specific humidity.

Figure B2 presents the composite vertical profile of the potential temperature for the hour before ($1h_{Before}\ LLC$) and after ($1h_{After}\ LLC$) the occurrence of LLC Class-2 in both regions for DJF and JAS. The profile is shown for the lowest 2 km of the atmosphere. There is no LLC Class-2 event in the Sahel during DJF, hence the empty plot. The profiles presented here suggest that before the LLC occurrence, there is a significant temperature decrease in the lower atmosphere. In the Guinea region, this is likely due to the well documented horizontal cold air advection from the Gulf of Guinea (Adler et al., 2019; Babic et al., 2019b) which leads to saturation of water vapour and consequent formation of LLCs. There could also be some cold air advection in the Sahel region, as suggested by Marsham et al., (2013). The temperature sharply increases after the disappearance of the LLC and this could be due to enhanced surface warming from incoming shortwave radiation (since attenuation of the incoming radiation is less in cloudless conditions).

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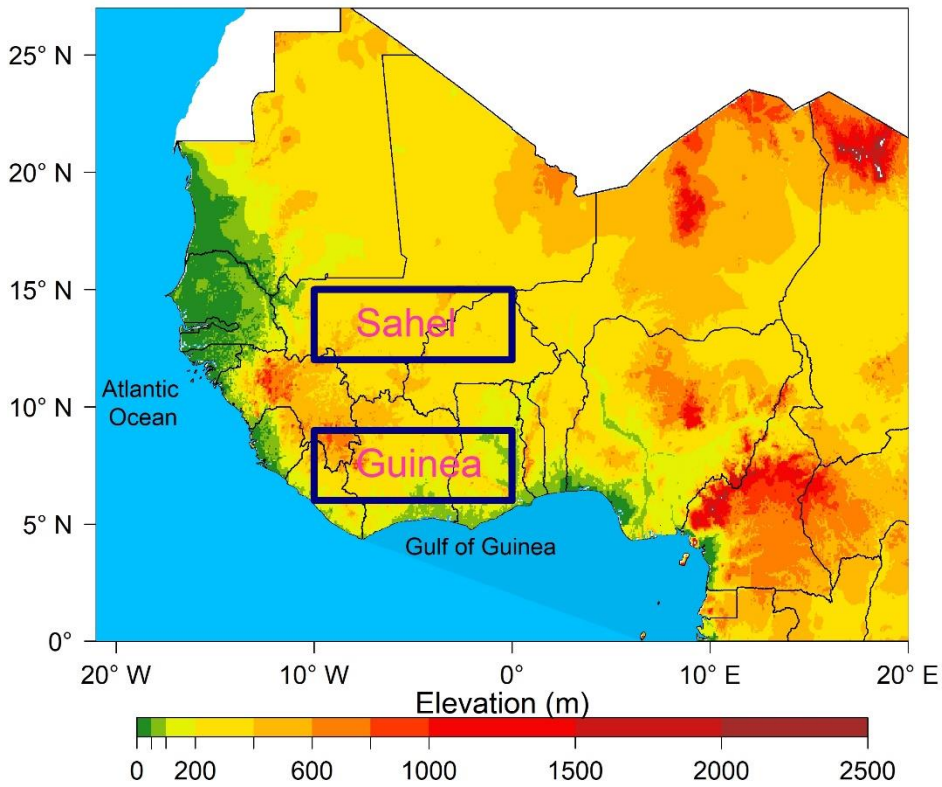
Figure 1: Map of West Africa showing the two selected windows (dark blue rectangles) used for identification of LLC events. The name denoted for each window is based on the climate zone of the window in the region.
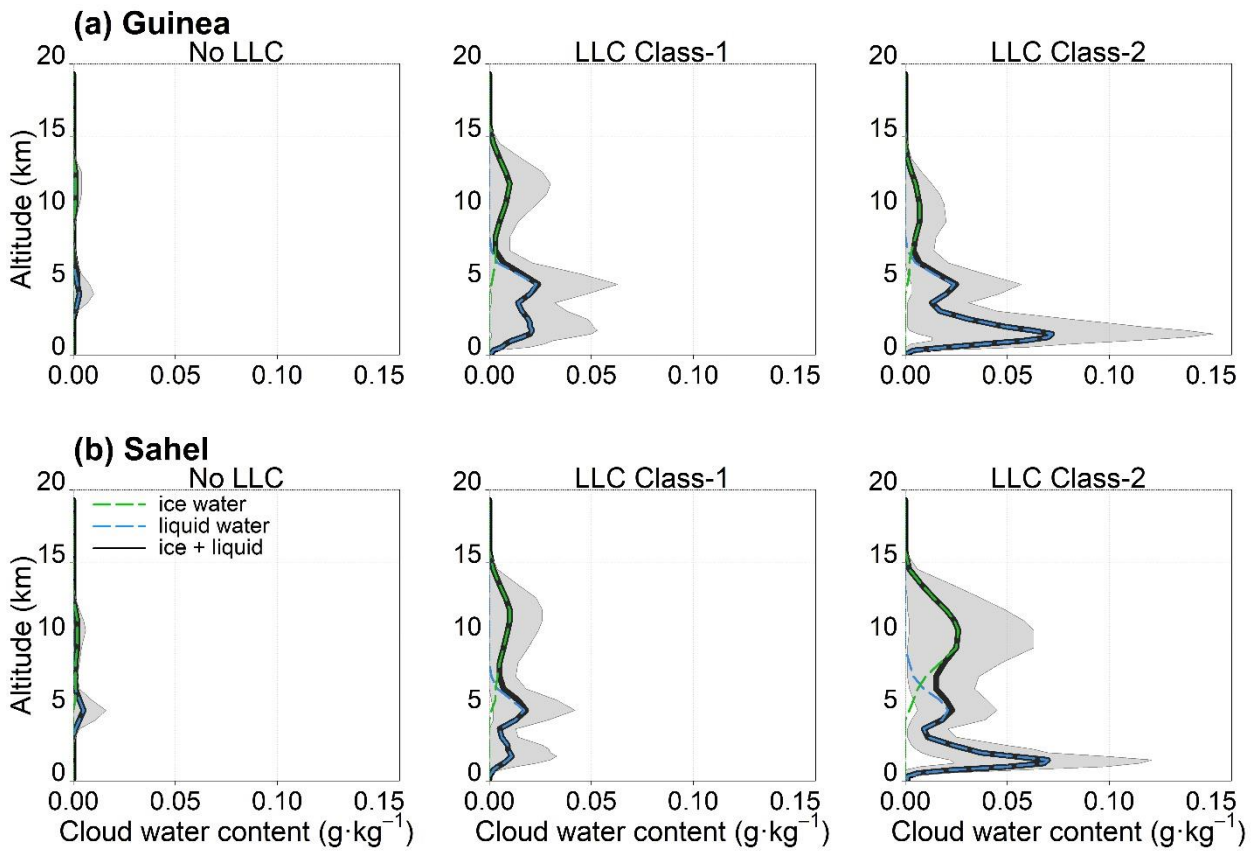
Figure 2: Mean vertical profiles of cloud water content (ice water (dotted green line), liquid water (dotted blue line), and total water content (black line)) during the occurrence of each LLC class in Guinea (top) and Sahel (below). Shaded areas refer to the 90th and 10th inter-percentile range.
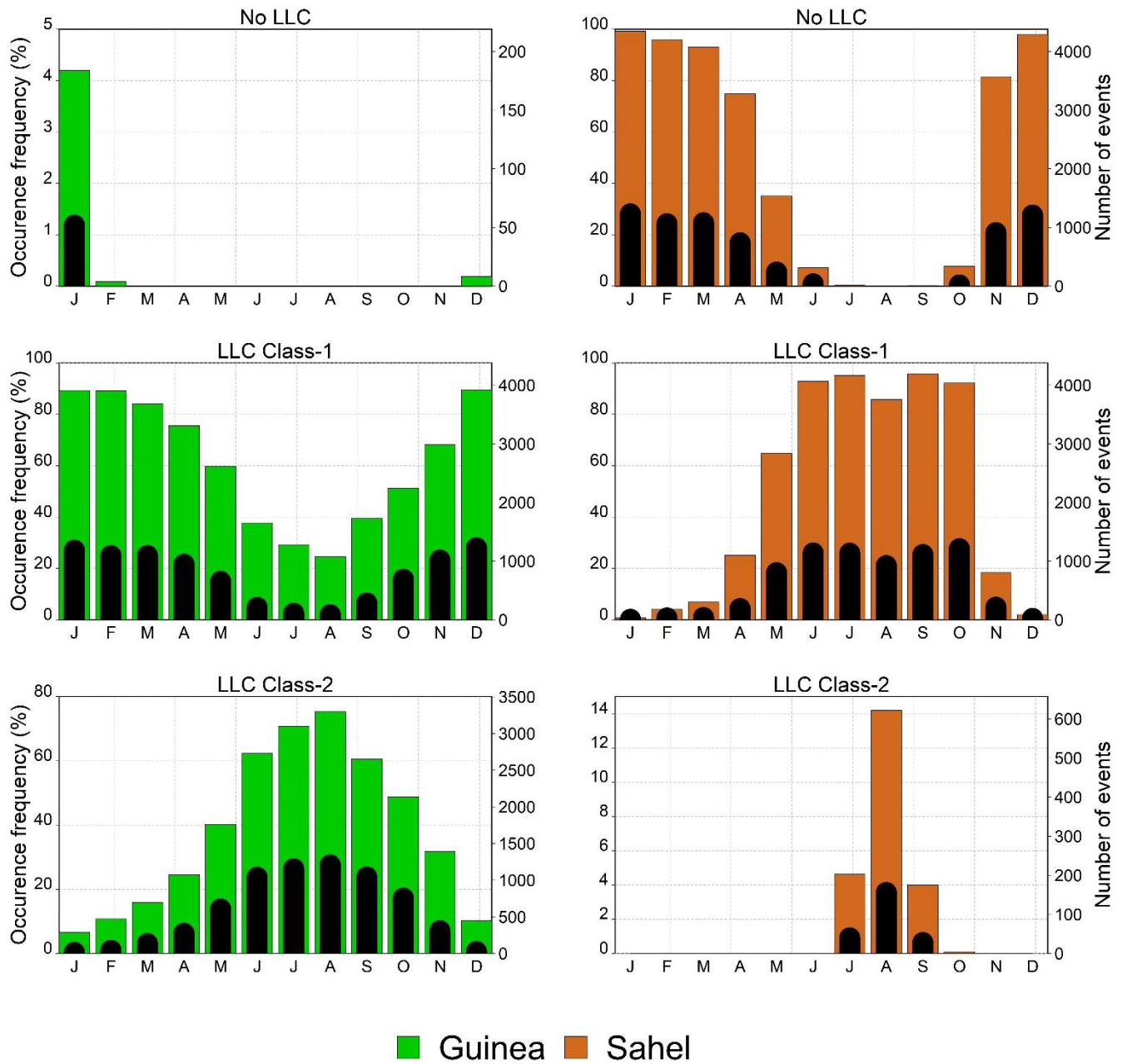
Figure 3: Monthly distribution of the occurrence frequency of hourly events for the three daytime LLC classes during the daytime in the Guinea (green bar) and Sahel (orange bar) regions defined in Fig. 1. The shorter black bars are the distribution of occurrences that last for at least six consecutive hours (the actual value has been multiplied by a factor or 4 to improve the legibility of the distribution).
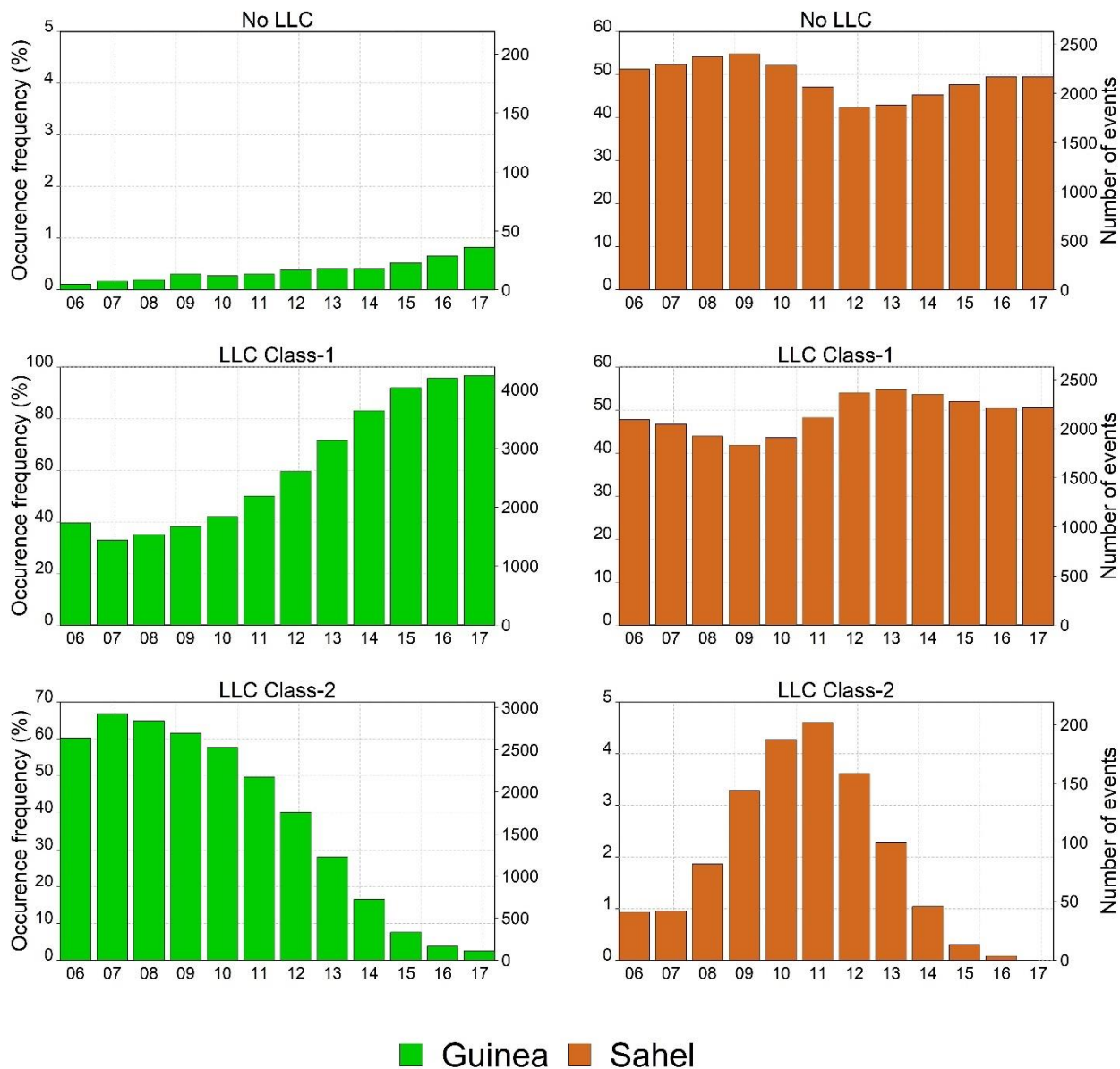
Figure 4: Diurnal distribution of occurrence frequency of the three daytime LLC classes during the daytime in the Guinea and Sahel regions as defined in Fig. 1. All seasons are included in the distribution shown.
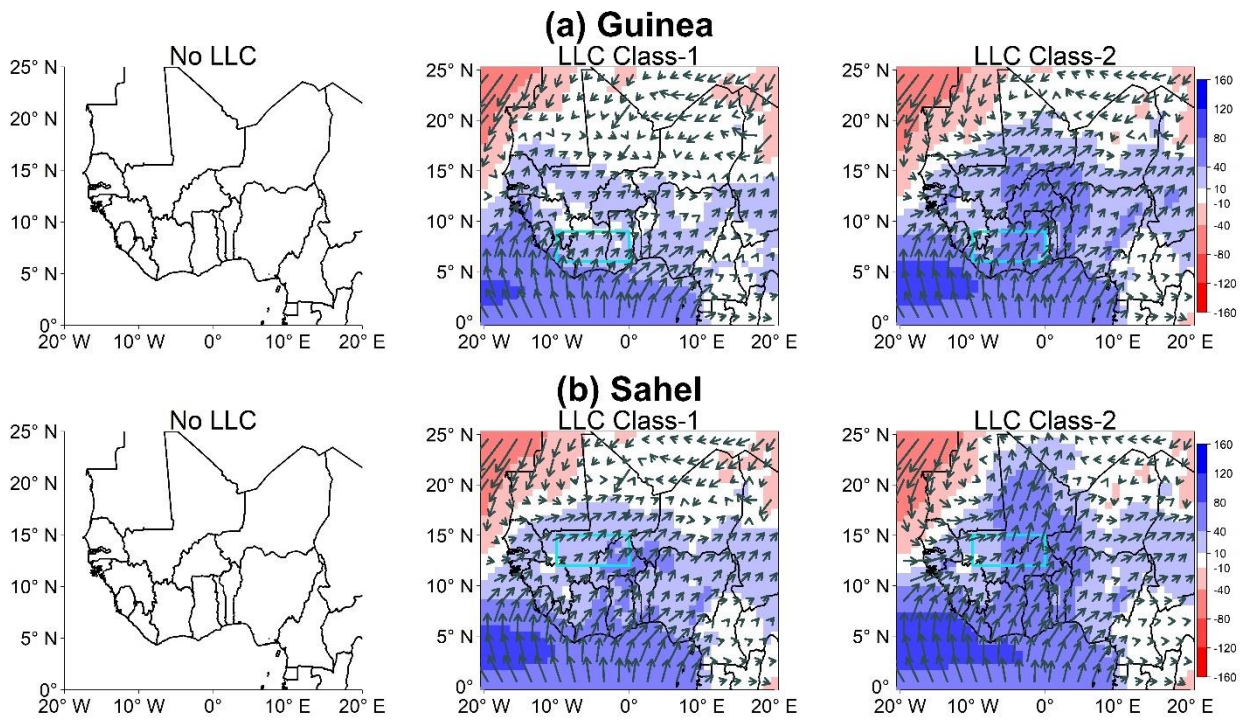
Figure 5: Composites of 950 hPa moisture flux (colours, g·kg⁻¹·ms⁻¹) and wind (vectors, m s⁻¹) over WA for the three LLC occurrence classes in the Guinea (top) and Sahel (below) windows during JAS from 2006 to 2015. Only occurrences that last for at least six consecutive hours are considered. Wind vectors are averaged over 2 grid points in both the x and y directions. No event of the No LLC class has been extracted in the Guinea and Sahel regions during JAS, hence the empty plots.

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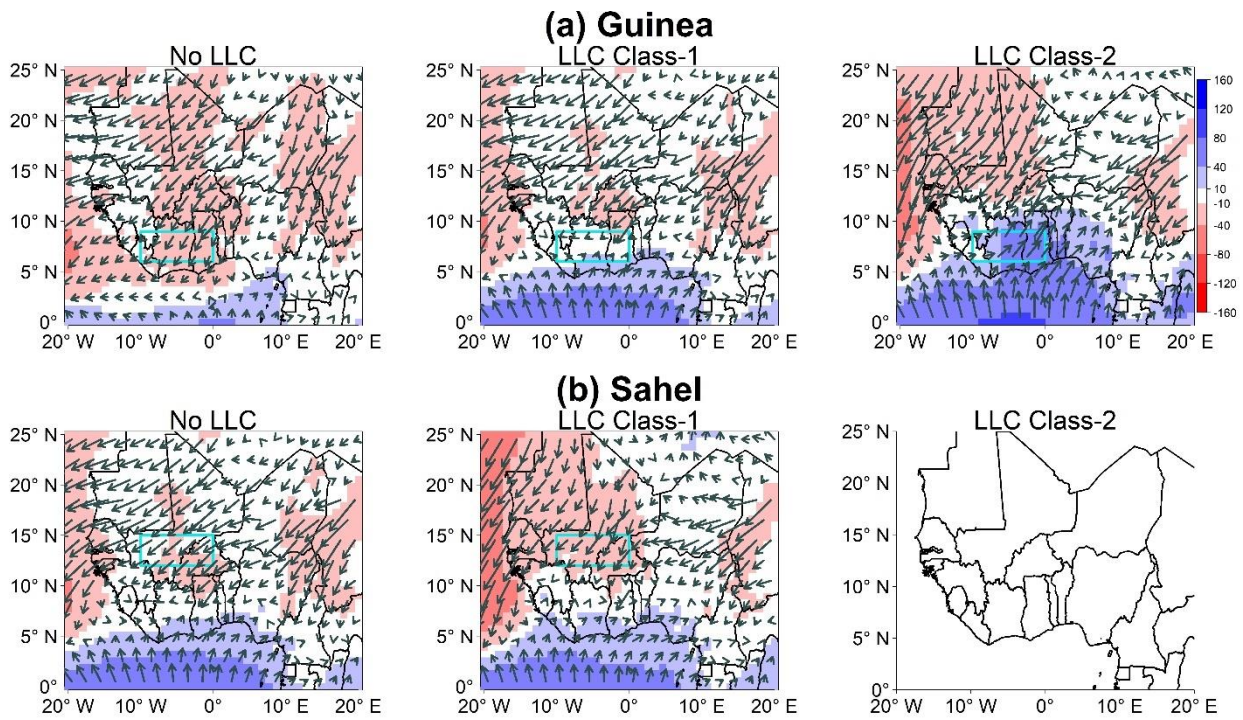


Figure 6: Same as Fig. 5 but for DJF. No event of LLC Class-2 has been extracted in the Sahel area during DJF, hence the empty plot.

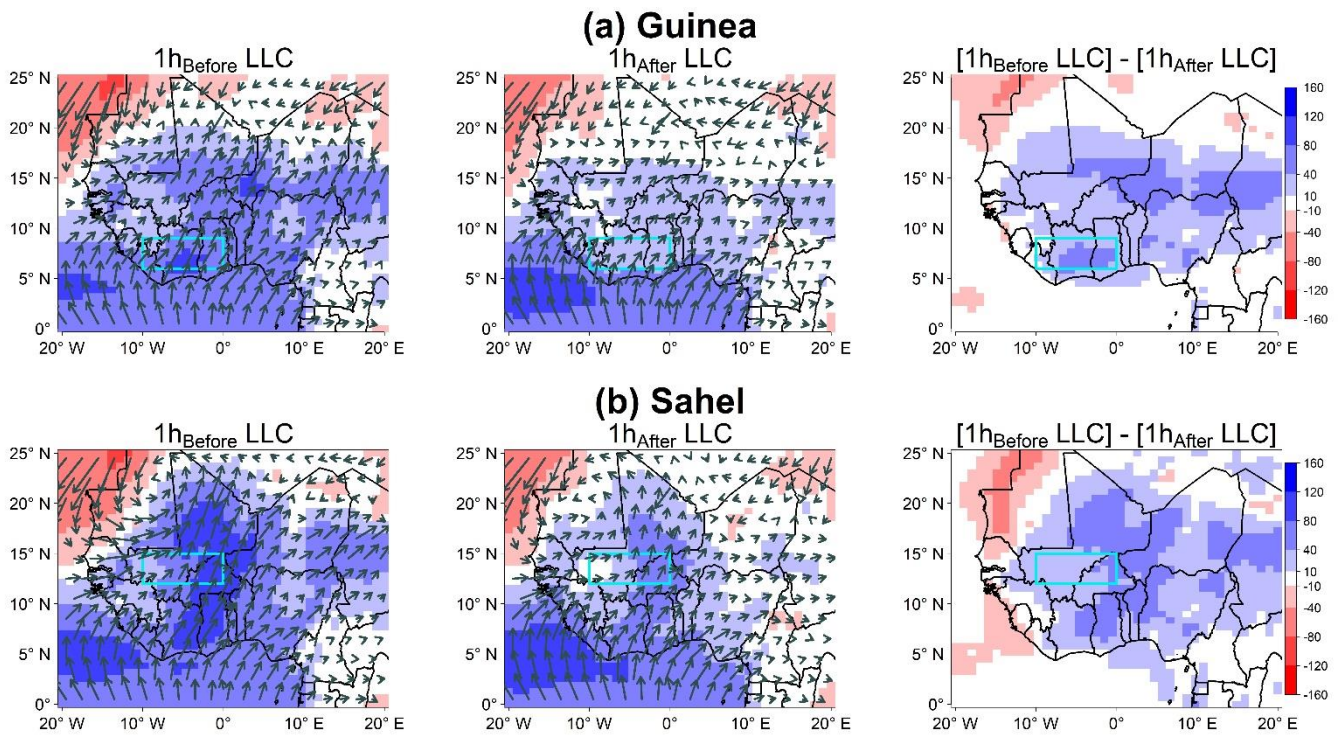



Figure 7: Composites of 950 hPa moisture flux (colours, g·kg^{-1}·ms^{-1}) and wind (vectors, m s^{-1}) over WA for the hour preceding LLC Class-2 occurrence ($\mathbf{1h_{Before}\ LLC}$) and after it disappears ($\mathbf{1h_{After}\ LLC}$) and the difference between the two periods during JAS from 2006 to 2015. Wind vectors are averaged over 2 grid points in both the x and y directions.

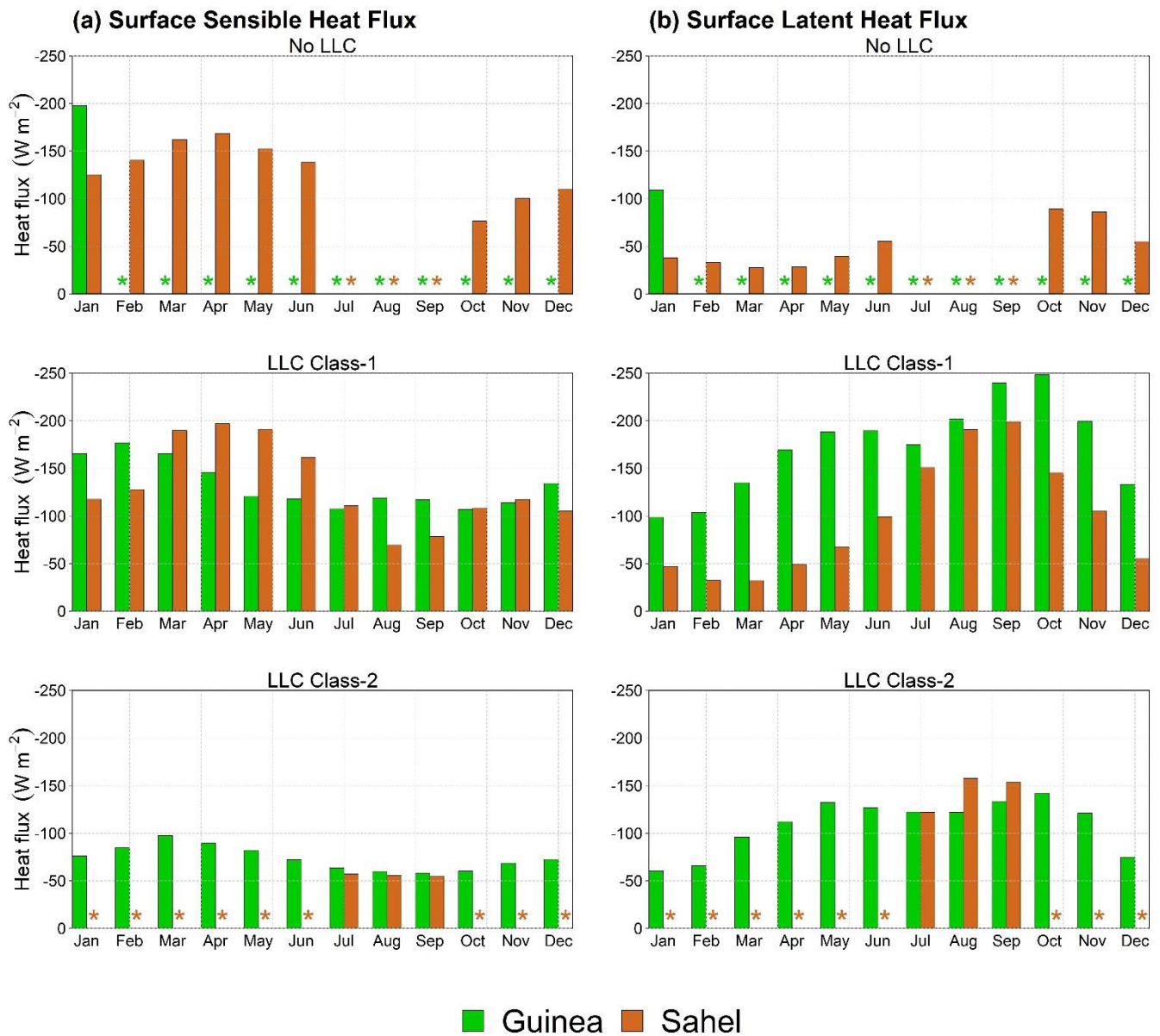



Figure 8: Monthly mean (a) surface sensible heat and (b) surface latent heat fluxes computed in the Guinea and Sahel regions during the occurrence of the three LLC classes from 2006 to 2015. The orange and green stars mean that there was no occurrence of the particular LLC class events in the region during that month, hence the composites were not computed. Only occurrences that last for at least six consecutive hours are considered.

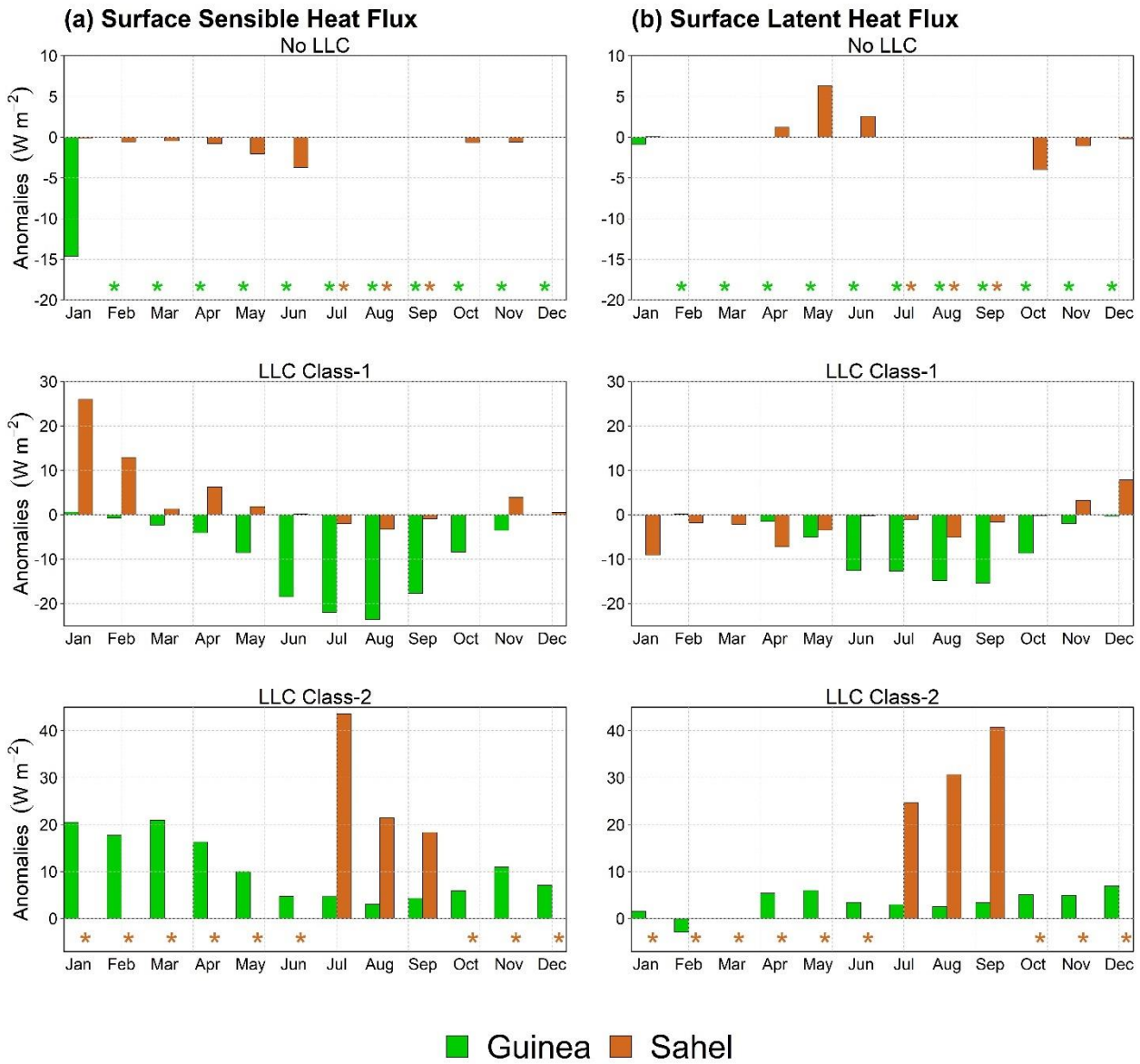
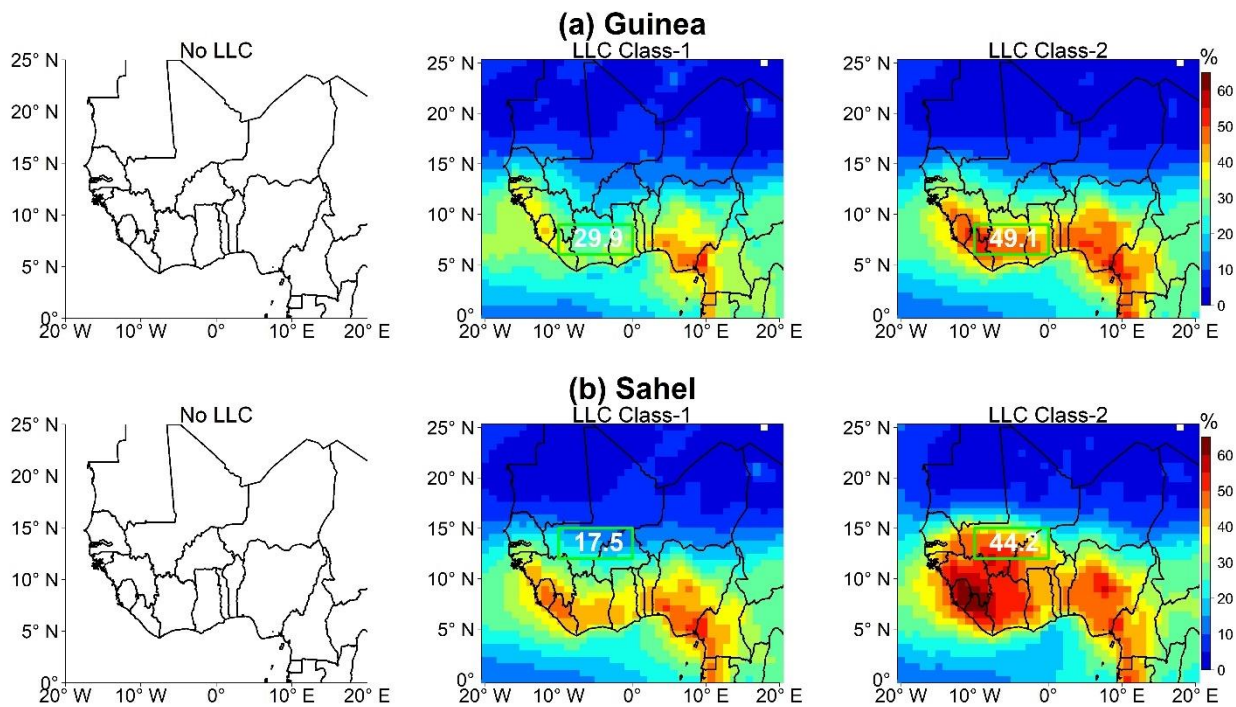



Figure 9: Monthly mean anomalies of (a) surface sensible heat and (b) surface latent heat fluxes computed in the Guinea and Sahel regions for the occurrence of the three LLC classes from 2006 to 2015. The orange and green stars mean that there was no occurrence of the particular LLC class events in the region during that month, hence the composites were not computed. Only occurrences that last for at least six consecutive hours are considered.

665 **Figure 10: Mean attenuation of incoming downwelling shortwave radiation (%) during JAS from 2006 to 2015 over WA for occurrence of the three LLC classes in the Guinea (top) and Sahel (below) regions. Only occurrences that last for at least six consecutive hours are considered. Numbers in the box refer to the mean percentage attenuation of incoming shortwave radiation in that specific area. No event of the No LLC class has been extracted in the Guinea and Sahel regions during JAS, hence the empty plots.**

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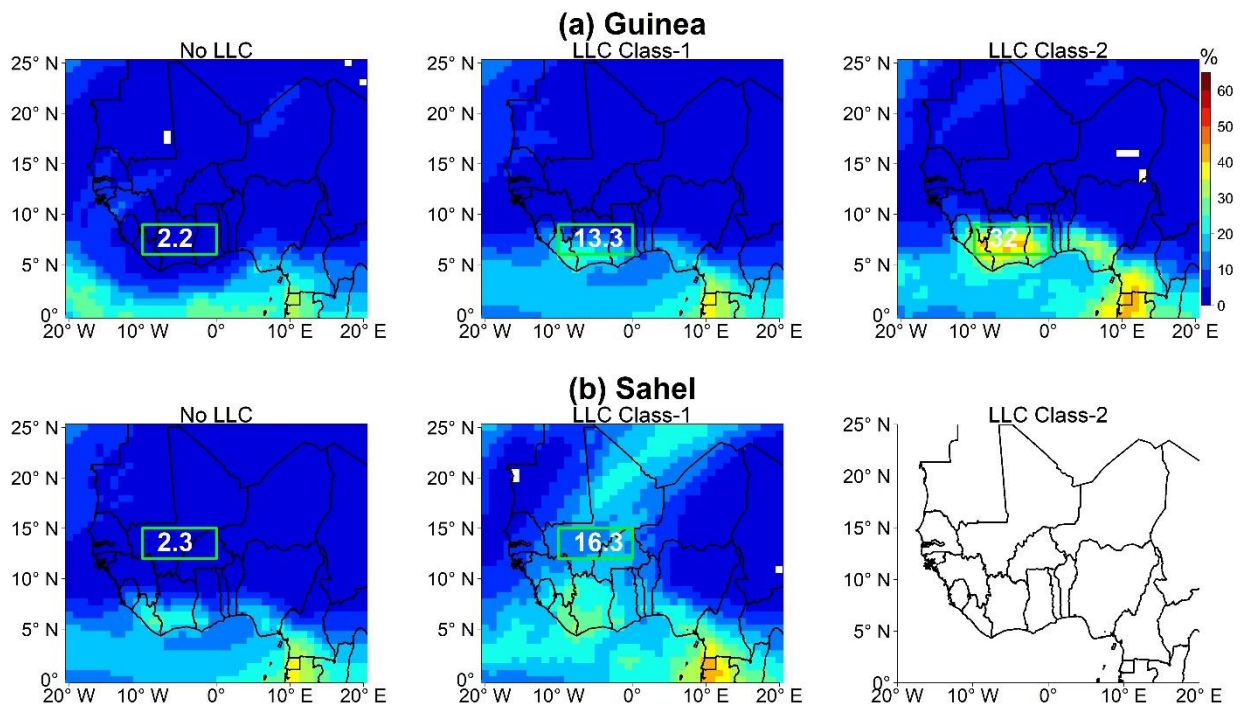


Figure 11: Same as Fig. 10 but for DJF. No event of LLC Class-2 has been extracted in the Sahel region during DJF, hence the empty plot.

675 **Table 1: Definitions of the three classes of LLC occurrence based on cloud fraction.**

Class of LLC occurrence	Description
No LLC	$CF = 0$
Class-1	$0 < CF \leq 0.5$
Class-2	$CF > 0.5$

Table 2: Number of events found for the occurrence of the three LLC classes in JAS and DJF. Composite analysis is performed based on these events for each class and season. For the study period, the total number of daytime hours are 11040 and 10800 in JAS and DJF respectively.

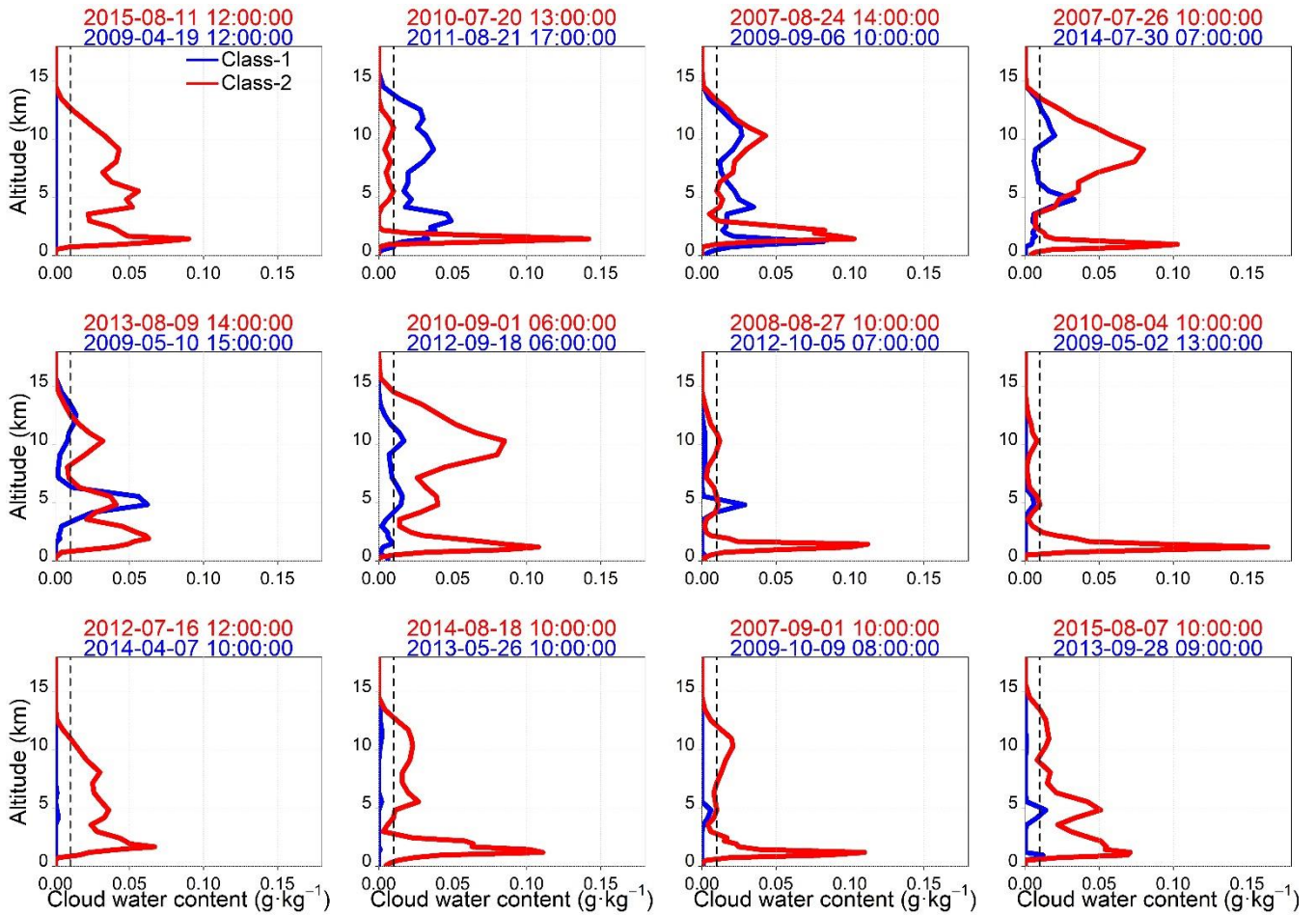
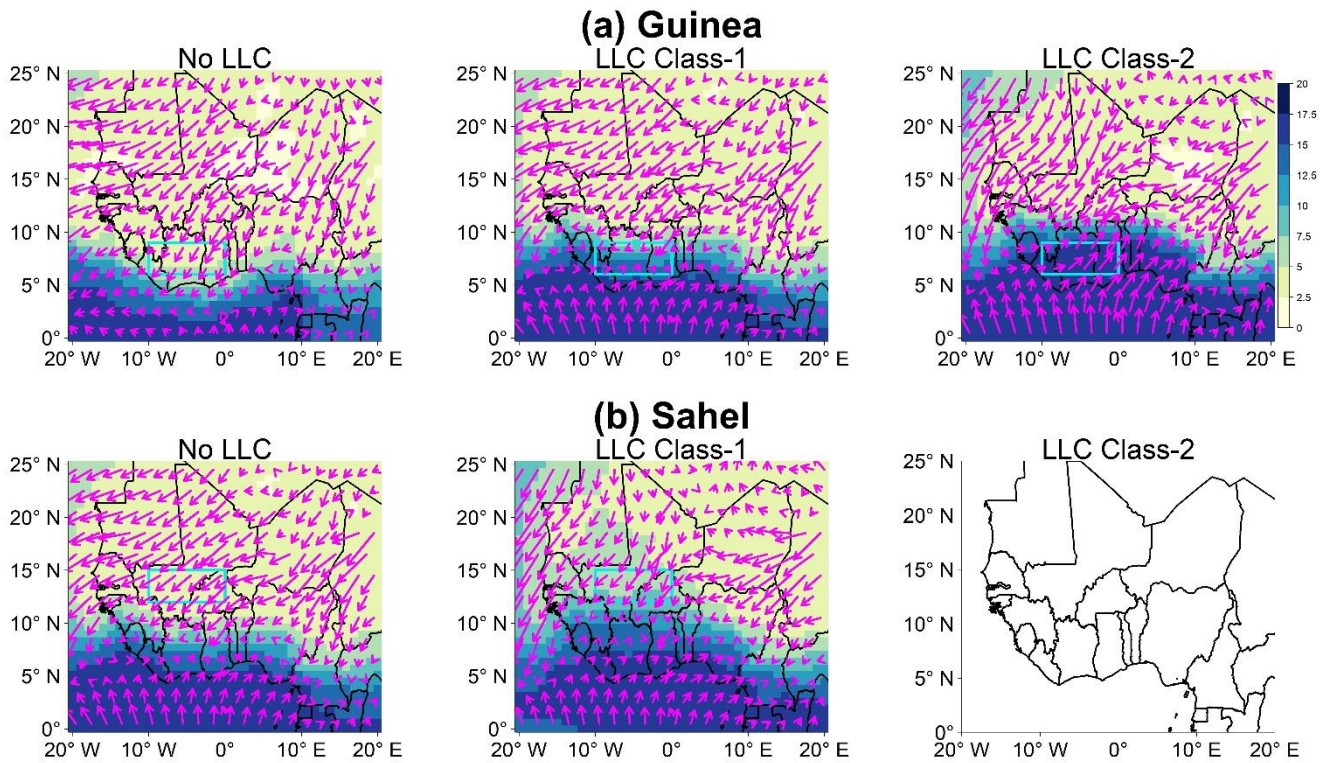
	JAS		DJF	
	Guinea	**Sahel**	**Guinea**	**Sahel**
Number of No LLC events	0	0	13	877
Number of LLC Class-1 events	122	793	878	16
Number of LLC Class-2 events	852	56	22	0
Total number of hours in season	11040		10800	

685 **Figure A1: Vertical profiles of cloud hydrometeor content (liquid and ice water) during the occurrence of LLC Class-1 and Class-2 events in the Sahel region. These events are randomly chosen. The timestamps at the top of each plot are the times for those LLC events. The black dashed vertical line shows the threshold for deep convective clouds as defined by Storer and Van den Heever (2013).**

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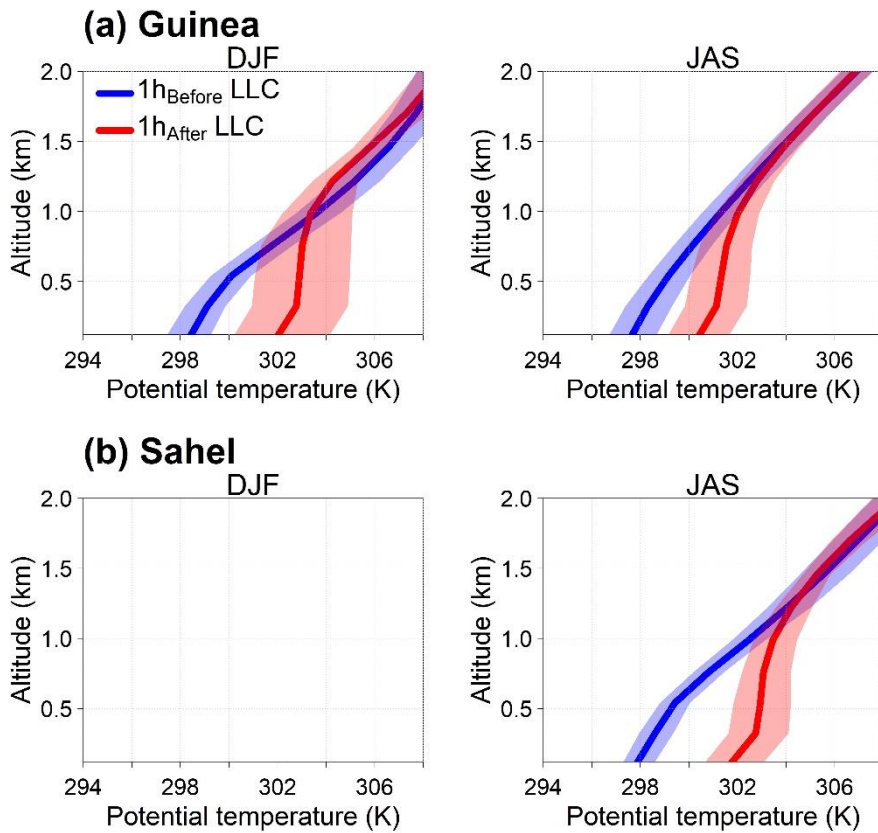
Figure B1: Composites of 950 hPa specific humidity (colours, g·kg⁻¹) and wind (vectors, m s⁻¹) over WA for the three LLC occurrence classes in the Guinea (top) and Sahel (below) windows during DJF from 2006 to 2015. Only occurrences that last for at least six consecutive hours are considered. Wind vectors are averaged over 2 grid points in both the x and y directions. No event of LLC Class-2 has been extracted in the Sahel area during DJF, hence the empty plot.

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Figure B2: Vertical profile of the mean potential temperature (solid line) for the lowest 2 km of the atmosphere, computed in Guinea and Sahel for all the 'one-hours' preceding LLC Class-2 occurrence and all the 'one-hours' after its disappearance. The shaded regions represent the 10th and 90th interpercentile range. No event of LLC Class-2 has been extracted in the Sahel area during DJF, hence the empty plot.

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