Editor

The authors have done an excellent job of responding to the reviewer comments. I am looking forward to seeing the final version of the revised manuscript.

Reviewer #2 raised some valuable points about the chosen metrics in this study and how those metrics compare with others previously used. The authors have done a great job of responding to the reviewer. I would encourage the authors to include some of this information (perhaps focusing on the general response and the response to comment #10) in the supplemental material of their revised manuscript. These sorts of discussions are useful for future studies, and I would hate to see that information get lost.

We thank the editor for the positive assessment of our manuscript and our responses to the reviewers comments. We agree that it would be potentially helpful to have some of our answers on model independence and the skill tests documented in the supplement. We have, therefore, extended section S3 in the supplement of the revised paper to also include a discussion about the potential circularity between calibration of the performance shape parameter and the subsequent skill tests, which both draw on future model information. For a summary on the different approaches used to establish model independence we have added a new section to the supplement (section S4 in the revised manuscript), where we also show an alternative "family tree" clustering based on the model error correlation distance.

Reviewer 1

Summary

The authors present a methodology for weighting CMIP6 models based on several performance metrics as well as on their independence from each other. This provides narrower bounds on future global mean temperature changes than in the unweighted ensemble, primarily by down-weighting the highly sensitive models that happen to have poor performance with respect to two reanalysis products and/or are closely related to other models. I found the paper to be nicely motivated, well organized and supported, and a useful contribution to the literature. There are a few areas that I think need to be clarified, and so I recommend minor revisions.

We thank the reviewer for the positive assessment and for the comments on our paper. Please find our answers to the comments highlighted in bold below.

Major Comments

* Figure 1 and the discussion around lines 241-242: the terminology of 0% to 100% trend-based seems too ambiguous to me and should just be written out explicitly. Couldn't the terms that are included just be stated explicitly in the figure? The figure doesn't really stand on its own, since one has to refer to these lines to know what exactly is meant by these. Additionally, it is not clear what the intermediate values (33%,50%, 66%) correspond to exactly. Upon multiple readings, I still cannot understand what is meant by these percentages at all, and I'm not completely sure what is actually meant by "50% tasTREND and 50% anomaly- and variance-based diagnostics" that forms the basis of the remaining analysis. Please clarify.

Thank you for pointing this out. The reviewer is correct, our notation in the original manuscript was ambiguous. What we are doing in our analysis is splitting 5 diagnostics into two parts: 1) tasTREND, 2) tasANOM, tasSTD, psIANOM, psISTD. Each of the categories in figure 1 relates to the relative importance of tasTREND compared to the other diagnostics, i.e.:

- 0% tasTREND + (25% tasANOM + 25% tasSTD + 25% psIANOM + 25% psISTD) [termed 'not-trend based' in the manuscript]
- 33% tasTREND + (17% tasANOM + 17% tasSTD + 17% psIANOM + 17% psISTD)
- 50% tasTREND + (13% tasANOM + 13% tasSTD + 13% psIANOM + 13% psISTD)
- 66% tasTREND + (8% tasANOM + 8% tasSTD + 8% psIANOM + 8% psISTD)
- 100% tasTREND + (0% tasANOM + 0% tasSTD + 0% psIANOM + 0% psISTD) [termed 'only tasTREND based' in the manuscript]

(values not summing up to 100% is due to rounding)

We have adjusted the paragraph in question as well as figure 1 in order to make this clearer (see figure 1 and line 259f in the revised manuscript).

* Discussion of Figure 2 around line 270: Should one have intuitively expected this from the math? I cannot seem to rationalize why using a model that is close to the CMIP6 MME to weigh CMIP6 would pull the CMIP6 MME mean away from the pseudo-observational "truth". This seems like a deficiency in the weighting. Shouldn't the weighting be resilient to this and do very little "harm" in this case?

Again, thank you for pointing this out. We did not mean to say that cases in which the perfect model is close to the unweighted MME *necessarily* lead to a decrease in skill and there are several examples where this is not the case (e.g., for pseudo observations from CanESM2 or

IPSL-CM5A-MR; see figure S2 in the revised manuscript). It is crucial, however, to point out that when we write 'close to the truth' we mean close to the truth in the evaluation periods (2041-60 or 2081-00). These periods are not used to inform the weighting and it is possible (in a pure model world as well as in the real world) that the information drawn from the past does not lead to a skill increase in the future if the constraint from the past is unrelated to the future projection. We have adapted our discussion of this topic to be clearer (see lines 291-308 in the revised manuscript).

In addition, skill might be dependent on the emission path. Looking at the time series plots using IPSL-CM5A-LR as pseudo-observations (figure S2 in the revised manuscript), for example, we see a slight downward shift of the distributions for SSP1-2.6 as well as SSP5-8.5. For the former, this leads to an increase in skill while it reduces skill for the latter. We have added a short discussion on this topic to the revised manuscript in lines 309-313.

We have also added additional information about the skill for each CMIP5 model used as pseudo-observation to figure S2 in the revised manuscript. Finally, we note that figures 2, 4, S2, and S4 have been updated in accordance with a comment from reviewer 2 (see last paragraph of our answer to their comment 10). For each CMIP5 pseudo-observation we now exclude the direct CMIP6 predecessors (if existing) from the calculation (see line 236-237 and table S5 in the revised manuscript).

* Figure 4: The combined and performance-only weights are shown, but not the in-dependence weights. Is there a reason for this? Is it worth also showing the ECS or TCR from these models on this plot, so that one could see that higher ECS/TCR models tend to be down-weighted? I assume this is correct, to the extent that models that warm the most over the 21st Century have high ECS/TCR, but I don't recall the authors coming out and saying it. Modifying this figure in this way could be a compact way of making that point.

We had originally decided against showing independence weights to avoid the readers being overwhelmed by the figure (and because they could be inferred from the difference between combined weights and performance weights). Also, in the original figure we had shown the weights relative to the median weight, so that the distance of a model with, e.g., twice the equal weight would show at the same distance from '1' (equal weighting) as a model with ½ of the weight (see also your last minor comment). However, we realise that this might be slightly harder to interpret so we have changed it in the revised manuscript.

We now show normalised weights for all three cases: independence, performance, and combined. In addition we now indicate TCR by coloring the labels accordingly (Figure 4 in the revised manuscript) and we have added a table containing all values to the supplement (Table S2).

* Figure 4: I'm surprised to see several well-regarded models having relatively low performance weights (UKESM, HadGEM, CanESM, CESM), whereas some models that are typically poor performers seem to do well here (GISS, FGOALS, INM-CM). Any comment? Is it possible that your performance metrics are too restrictive (just involving tas and psl, two fields that may not adequately discriminate models with good vs bad moist physics that governs feedback and ECS), allowing poor performing models to get high weights?

The reviewer is right, several typically well-regarded models receive rather low weights in our scheme. However, we point out that most of the models mentioned as examples have very high TCR. Based on our analysis (and other studies, see, e.g., Tokarska et al., 2020, Nijsse et al., 2020) these very high warming models are less likely and therefore they are down-weighted. In some cases (UKESM, HadGEM, CanESM) the main reason is the obvious mismatch between the observed and simulated warming over the course of the 20th century, which the modeling groups acknowledge in their technical description papers of the models.

It is indeed possible that our particular diagnostics choice leads to typically less well-regarded models receiving relatively high weights. This means that according to our chosen diagnostics

they are performing well compared to other models. It is possible that we would need to include more or other diagnostics to downweight models which have, e.g., bad moist physics, since the weighting method does not include knowledge about specific parameterizations. This point highlights the importance of careful diagnostics choices and the fact that the weighting is always aimed at a particular target and diagnostics choice. The weighting is not supposed to tease out which model is best in every case, and depending on the target and diagnostics choice the models receiving the highest or lowest weights will be different. This does not mean models receiving low weights in this case are bad models in general, as the reviewer realized some low weight models in our case are well regarded models and considered good models in general. But it means that based on their performance in simulating historical warming trends they are considered less likely here.

Minor Comments

*line 61: should be "model's"*line 78: should be "method's"

Done.

*Line 250: I don't see where the 10-20% statement comes from. By my eye, the medians range from near 0% to slightly larger than 25%.

The reviewer is correct, we changed this.

*Figure 1: titles should be "leave-one-out"

We changed the caption so this is no longer applicable.

*Figure 2 caption: should be "which"

Done.

*Figure 2: To clarify, the similarity between pseudo-obs and MME is only assessed over the "Diagnostic period" right? (Side-note: "diagnostic period" only appears in the figure and is not discussed in the text.) By my eye, MPI looks closer to the MME than does CanESM, so I'm a bit confused here. Is the reason because similarity in the evolution of GMST only one of the several metrics employed, and MPI does worse in the ones that cannot be gleaned from this figure?

We now introduce the terms diagnostic period in the main text of the revised manuscript (lines 215). Regarding the second point: the reviewer is correct in assuming that the performance of the models in the diagnostics that inform the weighting can not be inferred from figure 2 in general. We have added a sentence to the caption of figure 2 to make that clear.

*Line 309: "allows us" or "allows one"; also, it seems like some reference to all the performance metrics work done by Gleckler et al seems appropriate here. I believe they also advocate for comparing against multiple observational datasets.

This sentence does no longer exist but we have added a reference to Gleckler et al. (2008) in line 108 in the revised manuscript, where we motivate the usage of more than one observational dataset.

*Line 314: I don't see the motivation for these 3 groupings. Is it in any way objective?

This paragraph no longer exists in the revised manuscript.

*Figure 6: too small to read, suggest stacking the two panels vertically rather than placing them next to each other horizontally

Done.

*Line 334: should be "model's"

Done.

*I don't think the average reader should be expected to know how to interpret a figure like Figure 5. Only the meaning of the colors are explained in the caption. What does the rest signify?

We have added additional description to figure 5 and now provide a more detailed description of the clustering approach in the supplement (section S6 in the revised manuscript).

*Line 391 "The weighting also largely reconciles CMIP6 with 5": what is this referring to specifically, and is there a figure in particular being referenced?

We were referring to the fact that the constrained CMIP6 TCR is closer to the CMIP5 TCR range from, e.g., the IPCC AR5 (1°C-2.5°C). However, this sentence was slightly misplaced here and is no longer included in the revised manuscript.

*Figure 4: Are all weights less than or equal to 1 in absolute units, and only exceed when expressed relative to equal weighting as is done in the figure? Otherwise I'm a little confused about why a model would have a weight in excess of 1. How exactly is wi used? weighted avg of X = sum(wi*Xi)/sum(wi)?

We now show normalised weights for all three cases: independence, performance, and combined. See also our answer to your major point regarding figure 4 above.

Reviewer 2

Summary

Some models are more consistent with historical observations than others. In climate projection, it makes intuitive sense to give more weight to the models that are more consistent with observed climate shifts and less weight to models that are less consistent with observed climate shifts.

But how?

This paper reports on a method of assigning model weights that relies on two distinct distance measures: the distance of models from observations and the distance of models from other models. The method requires the specification of two parameters that determine how each of these distances are turned into model weights. The method for determining the parameter associated with inter-model distance is poorly explained(see specific comment 8 below). The method for determining the parameter associated with distance from observations is also poorly explained, but for many experiments, involves future-time-pseudo-observations from the future states that are the objective of the prediction (see comment 10 below). In other words, the tuning method appears to render the tests of the method to be of the "in-sample" variety. To weaken the degree of "in-sampleness" an additional test is performed using CMIP5 runs. However, since one expects many of the CMIP6 models to be closely related to the CMIP5 models, there are strong reasons to believe that this test is not truly "out-of-sample" either.

Even with the use of "future-time-pseudo-observations" in the tuning procedure, the improvements from this weighting scheme seem very modest in comparison with, for example, those obtained in Abramowitz and Bishop (2015, J. Climate) – (using a a method that solely required historical observations for the weights). The revised paper should include some attempt to compare/contrast/explain the Abramowitz and Bishop results.

A superficially appealing feature of the method is that it gives more weight to models that are both skillful and statistically independent of other models. However, this independence is just described in terms of inter-model distance and not in terms of the independence of the model error. Is there some unstated proof that increased inter-model distance equates to increased model error independence? (It seems easy to think of counter examples). As demonstrated in Bishop and Abramowitz (2013), it is the independence of the error of the individual models comprising an ensemble forecast (as measured by inter-model forecast error correlation) that increases the predictive power of the ensemble. The revised paper needs to address the issue of the relationship or lack of relationship between inter-model distance and model error independence.

After applying the method to the CMIP6 ensemble members, the authors find reduced warming relative to the simple sample means of CMIP6 ensembles for the high and low CO2 concentration scenarios considered. However, any confidence in this prediction must be strongly tempered by the "in sample" circular- nature of the testing and tuning procedures used by this method. My overall recommendation would be that the paper be returned to the authors to address the specific comments below and to include results from experiments in which only historical observations (or model-based-historical-pseudo-observations) were used to determine the weights. This constitutes major revision.

We thank the reviewer for the critical assessment of our manuscript. The reviewer raises several important questions in the general comments above. Most of them we address in our answers to the specific comments as summarised below. In addition, we discuss the rationale behind our model independence metric in the following:

- Calculation of the independence shape parameter: see comment 8
- Calculation of the performance shape parameter: see comment 10

- Out of sample skill tests: see comment 10
- Skill improvement and comparison with Abramowitz and Bishop (2015): see comment 11; in addition we have added several references to the approach used therein in the revised manuscript.
- Model distance versus model error independence: see below

Model-model distance and model error correlation

The weighting method we apply in our study separates between a model's performance and independence. For establishing either measure, different metrics have been used in the past (see line 145 in the revised manuscript). In the case of independence, one could, for example, argue that it should be based on our knowledge of a model's inner workings (such as shared components, parameterizations or heritage with other models). However, this information is not always easily accessible and is, in addition, hard to quantify. Therefore, we here use an output-based definition of independence: given a generalised distance metric (based on the climatology of two variables) we define independence as a model distance to all other models in the ensemble. This is equivalent to the distance of the models' errors:

$$e_i - e_j = (m_i - obs) - (m_j - obs) = m_i - m_j$$

where e is the model error, m is the model, and obs the observation.

This approach has the advantage that it does not rely on observations, which are often geographically sparse and restricted in time. It, therefore, allows, in theory, establishing model independence based on hundreds of years of pre-industrial control runs or based on variables which do not have reliable global observations, such as precipitation.

Here we use surface air temperature and surface pressure as the basis for our estimate of independence. This follows the work of Merrifield et al. (2020), who show that using these two variables allow a clear separation of initial-condition members of the same model as well as closely related models on the one side and independent models on the other side (see, e.g., figure 5 in Merrifield et al., 2020). In addition, in our manuscript we show qualitative results of our independence classification as a model dependence tree in figure 5 and discuss several clusters where the "inner workings" are known (line 389-395 in the revised manuscript). As a further test we insert artificial new models into the ensemble (see figure 6 and related discussions). This allows us to investigate the change in independence weight based on the relation of the inserted model to the rest of the multi-model ensemble.

Bishop and Abramowitz (2013) follow a different approach that is based on the assumption that independent models have uncorrelated error time series. This approach can not directly be applied to our framework since we base our weighting on time-aggregated (mean, standard deviation, trend) spatially resolved fields. The main question the reviewer seems to pose, therefore is: Do the two approaches deliver fundamentally different results?

To test this we assume that the concept of error independence also holds for time-averaged spatial fields. We apply an independence weighting based on the spatial correlation of model errors and contrast the results with our original results (based on model distances). S_{ij} in equation (1) then becomes the matrix of model error correlation distances:

$$S_{ij} = 1 - CORR_{spatial}(m_i - obs, m_j - obs)$$

Figure R1 below shows the models "family tree" equivalent to figure 5 in the manuscript based on these correlation distances. While the grouping of models is mostly the same as in figure 5,

there are also some obvious differences. The difference between the closest related models (e.g., UKESM1-0-LL and HadGEM3-GC1-LL) and the maximum distance between any two clusters of models is considerably larger. Several models have changed to a different cluster (e.g., NorESM2-MM or AWI-CM-1-1-MR). Without a detailed analysis, however, we can not make any clear statements on which clustering is "more correct".



Figure R1: Similar to figure 5 in the revised manuscript but based on error correlation distances instead of model-model distances. Note that for this case we do not use any area weighting.

Based on the general similarity of the two trees, we do not expect the change in the independence metric to have a major influence on the results. In a second step we, therefore, look at the weighted distributions based on independence weights using these error correlation distances. The results are presented in figure R2 below. Compared to figure 8 in the revised manuscript there are only minimal differences. This at least shows that there are no strong disagreements between the approaches. One reason for the similarity is certainly also the fact that the weighting is dominated to a large degree by the performance weighting and, in particular, by the low weights of some of the strong warming models.

In summary we, therefore, argue that either approach might be appropriate to use, and the main conclusions in our manuscript are the same for an independence matric based on correlation. For simplicity we, therefore, prefer to continue using our original metric basing independence directly on model-model distances which does not require observations and thus eliminates one potential source of uncertainty. We have, however, added section S4 to the supplement discussing or method to estimate model independence in the context of other approaches.



Figure R2: Similar to figure 8 in the revised manuscript but with the independence weighting based on error correlation distances instead of model-model distances. Note that for this case we do not use any area weighting in the independence weighting calculation.

Specific comments

1. Line 16. Consider explaining what TCR is in the abstract to appeal to a broader audience.

Indeed our study aims at a quite general audience and therefore focuses mainly on projections of future global warming which are widely known. In the revised manuscript we no longer mention TCR in the abstract.

2. Line 31. Do you mean model uncertainty, unknown model climate error, unknown model-climate-sensitivity-to-CO2 error or model climate differences? We know what the model is, and we can determine its climate past, present and future by running it. We can also determine the differences between the climates of different models. Given the limitations of the spatio-temporal distribution of observations, the uncertain thing is the actual climate both past, present, and future, is it not?

Model uncertainty here refers to the error of both present and future climate. In particular to its bias, since for climate projections we are concerned with correctly estimating distributions of trajectories, rather than individual trajectories like for weather and climate prediction.

"Model uncertainty" has become a standard piece of terminology in this subfield, following its popularization by Hawkins and Sutton (2009). It is also mentioned as "structural" uncertainty or error, referring to the structure of the model (which is assumed to be different between different climate models, hence the "model" label). We have updated the paragraph in question to make that more clear (lines 30-34 in the revised manuscript).

3. Line 35. Lorenz, the father of chaos theory, argued that while the accuracy of weather forecasts was limited to a few weeks the climate of a system was not sensitive to specified initial conditions and could be known provided the forcing on the system was known. I guess "climate" in the sense of Lorenz refers to the statistical description of the attractor of the chaotic system. When you refer to "internal variability" do you just mean slow modes of the model's chaotic attractor that might possibly be confused with a change in the mean of the model's attractor if the ensemble size was too small?

"Internal variability" indeed refers to initial condition sensitivity; the terminology has become standard in the climate literature following papers like Hawkins and Sutton (2009) or Deser et

al. (2012). Here "climate" refers to the statistical description of the attractor of the system which these models attempt to represent - including the atmosphere but also the ocean, ice, and land surface. Particularly for the ocean, coupled models and the real earth's coupled system show variations on timescales of, at the very least, multiple years (e.g., due to ENSO) that depend on the initial conditions. Recently some efforts have sought to identify predictability on the order of decades, though if this exists it is assumed (here and generally) to be small.

Because GCMs are expensive to run and have unknown but expected long timescales before ensemble variance that properly samples the climatology is achieved, CMIP models are not at the point where many of them have enough ensemble members to adequately sample the attractor (in contrast to weather prediction, where that is currently achievable and in fact often achieved). With a small number of ensemble members and long timescales, internal variability is convolved with forced responses. These can be isolated with "large ensembles" (of several tens of simulations differing only by initial conditions) but the CMIP ensemble includes many models which are expected to differ in their bias, some of which also include multiple realizations from the same model, which are expected to differ among each other only in terms of their "internal variability" or due to sampling. We have added some discussion to the paragraph in question (lines 34-40 in the revised manuscript).

Deser, C., Phillips, A. S., Bourdette, V., & Teng, H. (2012). Uncertainty in climate change projections: the role of internal variability. Climate Dynamics, 38(3–4), 527–546. https://doi.org/10.1007/s00382-010-0977-x

4. Line 102: I'm guessing you are referring to Section 3.2 of Brunner et al., 2019. Is that correct? If so, please state this in the text. Your wording suggested that you had estimated an observation error variance. However, on reading Section 3.2 of Brunner et al., 2019, I'm now guessing that you are referring to how your derived weights change depending on which subset of all observations you use. Are you suggesting that the reason for your weights changing is because the observations have different errors? Can you rule out the possibility that your weighting scheme isn't just over-fitting each individual observational data set? In any case, the revised paper needs to clarify whether in fact you are referring to the size of the change in weights associated with using differing observational data sets. Also, the observed values are known. They are not uncertain. The errors of the observed values are unknown. It is the observational error that is uncertain.

Thank you for pointing this out, the wording was unclear in the original manuscript. Indeed, it has been pointed out in the literature that using different observational datasets can lead to diverging results in some cases (e.g., Gleckler et al. 2008, Lorenz et al. 2018, Brunner et al. 2019) due to differences in the datasets. We referred to these differences in the observational datasets as observational uncertainty but no longer do so in the revised manuscript.

What we are concerned with here is bias in the observational datasets, which are a central challenge in climate science. In the presence of such biases, it is not unexpected that the results of the weighting change based on the datasets used. To get a reference that is as robust as possible, we are using a combination of two observational datasets (ERA5 and MERRA2) to calculate the model-observation distances and further the performance weights. The datasets are combined by taking the center of the observational spread at each grid cell (following Brunner et al. 2019 who also discuss other approaches; see their section 3.2 as well as section S2 and figure S3 in their supplement). We have clarified that and added additional information to the section in question in the revised manuscript (lines 103-110).

Gleckler, P. J., Taylor, K. E., & Doutriaux, C. (2008). Performance metrics for climate models. Journal of Geophysical Research Atmospheres, 113(6), 1–20. https://doi.org/10.1029/2007JD008972

5. Line 145. "We want to..." If there was a hypothetical user of the climate projection that only cared about temperature trend and not about year-to-year variability, might you not be doing them a

disservice by down-weighting members that have an excellent temperature trend but poor inter-annual variability? Consider changing to "We choose to..."

The reviewer is correct in pointing out that the selection of diagnostics for establishing the models performance weights should depend on the target in question. In our study we look at temperature change in two time periods as a target, which is closely related to the temperature trend. Therefore, the temperature trend is indeed a powerful diagnostic.

However, it also is strongly influenced by internal variability (i.e., it differs quite strongly between initial-condition members of the same model) which is not desirable for a good diagnostic as we argue in line 171 of the revised manuscript: "Ideally, a performance weight is reflective of underlying model properties and does not depend on which ensemble member is chosen to represent that model (i.e., on internal variability). tasTREND does not fulfil this requirement: the spread within one model is the same order of magnitude as the spread among different models."

We therefore use "a balanced combination of climate system features (i.e., diagnostics) relevant for the target to inform the weighting to minimise the risk for skill decreases. This guards against the possibility of a model "accidentally" fitting observations for a single diagnostic while being far away from them in several others (and hence possibly not providing a skilful projection of the target variable)." (line 454 of the revised manuscript)

In this sense we argue that even if a user is only interested in a model simulating future temperature trend correctly, it might still be important to also include other diagnostics. This can help to avoid weighting a model highly because it "accidentally" matches the observations in a given historical period due to, e.g., internal variability.

6. Line 147-149. Equations should be added to precisely describe these observation derived quantities – perhaps in an appendix or supplementary material.

We have now added a mathematical description of the diagnostic calculation to the supplement (section S2) and reference it in the revised manuscript in line 160.

7. Line 170. You must state what was used as a proxy for a perfect model. I would think that the derived sigma_D must be related to the ensemble variance of the model states around the time averaged state. That quantity will depend on the model will it not? Please clarify.

We have adjusted our description of the shape parameter calculation in the revised manuscript in order to make this more clear. In the revised manuscript we now refer to the iterative test used to the performance shape parameter as parameter calibration (lines 182-191). In addition we have added additional information including a schematic of the calibration test to the supplement (section S3).

8. Line 183. I looked at Section 2.3 of Brunner et al., 2019 for an explanation but Brunner et al. (2019) just directs the reader to Lorenz et al., 2018. Your work needs to be reproducible. When referring to another paper for a key explanation, you must give very specific information about where in the paper the explanation resides (e.g. a section number) to ensure reproducibility. You have not done this.

The reviewer rightfully points out that we should have been more clear in referencing this important information. The calculation of the independence shape parameter and reasoning behind it is described in detail in the supplement of Brunner et al. (2019; section S3.1), which we now explicitly mention. In addition we now provide a summary as well as a discussion of the chosen value in the context of our study in the supplement of the revised manuscript (see line 200 and supplement section S5 in the revised manuscript).

9. Line 191. The method used to evaluate performance given here seems almost identical to that given in Abramowitz and Bishop (2015) but no reference is given to this paper or others that may have used this approach before. Such literature is relevant and should be cited.

Thank you for pointing this out. We have added several references to the relevant literature which used similar approaches before (see line 206 in the revised manuscript).

10. Line 200-205. Here, we learn that sigma_D weights are determined in part from information from a place that is inaccessible in reality: the future. Only model futures are accessible. By line 205 we learn that the model future states (rather than observations) are, in fact, an integral part of choosing the weights. This is a significant departure from many other observation-based methods for improving ensemble forecasts and projections. The use of future time observations in the training causes all of the associated tests to be "in-sample" tests – dramatically reducing their trustworthiness. Since the CMIP5 models belong to the same general class of human produced climate simulators they can barely be considered "out-of-sample". Please comment on the limitations of this approach. In addition, you have not clarified how the method of tuning for future states interacts with the method to determine sigma D referred to on line 170 (see previous comment).

The reviewer rightfully points out that there is some influence from the future model states included in the weights via the performance parameter calibration. However, there also seems to be some misunderstanding regarding our approach. We adapted the sections in question to make it more clear in the revised manuscript.

The model performance weights are proportional to each model's generalised distance (a combination of 5 diagnostics) to the observations (D_i) as given in the numerator of equation (1). The proportionality constant is the performance shape parameter sigma_D, which translates these distances into the weights. It is indeed established using the target period, i.e., the future model states. The weighting for the ensemble is then calibrated as a whole using this single parameter, and it is not the case that the weight of each model is calibrated individually through its historical simulation.

Crucially, this means that the weighting is still dominated by the comparison of models to the observations only. Consider, for example, a case where the diagnostics are really poorly chosen: this could be because they are dominated by (random) internal variability or because they do not have any physical relationship to the target. The weighting then would not have any skill, regardless of the sigma_D parameter.

As, for example, Sanderson et al. (2017) state, selecting sigma_D only based on historical information might lead to overconfident results as a more skillful representation of the base state does not necessarily translate to a more skillful representation of the future. Selecting sigma_D only based on historical information would a priori assume that the chosen metric is relevant for the projection. One way of approaching the problem might be to apply the method on the historical and then test the result in a perfect model test, potentially adjusting the method in an iterative approach to maximise skill.

In our weighting approach we already include such a perfect model test in the calculation of the weights in order to avoid overconfident results. To avoid confusion between the setting of the parameter and the subsequent testing of method skill we have changed the terminology in our manuscript and refer to the former as *parameter calibration* to separate it from the later perfect model tests which are used to calculate the skill of the weighting. In addition we have added a section in the supplement detailing and visualising this parameter calibration (section S3 and figure S1 in the revised manuscript).

Finally, addressing the question of the relationship between the calibration of the performance shape parameter and the subsequent testing of the skill of the method, we would argue that the circularity is quite limited. There are several reasons for this:

- As we point out above, the weighting is, to a large degree, based on the model's distance to historical observations, with future observations only influencing them via sigma_D, which is a single value constant across all models, over time, and all metrics.
- The parameter calibration does not aim at maximising (mean) skill, but rather ensures that the results are not overconfident. Take the example of poorly chosen diagnostics again: in such a case, any separation into better or worse models would be overconfident as it would be based on pure chance. During the parameter calibration this would become obvious and sigma_D would be relaxed to a large value (in order to avoid this overconfidence) leading to an approximation of equal weighting. Subsequently testing the skill of the method can still be insightful to estimate the actual increase in skill (or the lack thereof in the case of badly chosen diagnostics).
- We use two different model pools to draw the perfect models from in our investigation of the method's skill. The first one is based on CMIP6 data, and one could therefore argue that it has a stronger potential circularity as the same models have been used to calibrate sigma_D. However, this test is mainly used to investigate the relative differences between different combinations of diagnostics and to select the best performing one (see figure 1 and related discussion). Since any remaining circularity is the same for all cases shown in figure 1, a comparison between them should still be valid. We have adapted the abstract as well as section 3.1 to make that more clear.
- For the second test, we use CMIP5 models, which have not been used in the parameter calibration, as perfect models. Here, another potential issue arises: several CMIP6 models are related to CMIP5 models and are therefore not independent. However, about eight years of additional model development lie between the two generations. In addition, it has been noted that several CMIP6 models have a much higher climate sensitivity and are, hence, quite different from their predecessors (at least in their response to anthropogenic forcing, which dominates the future period used for the perfect model test).

To further increase the independence between the CMIP5 and CMIP6 ensembles, we now exclude directly-related models from the perfect model test in the revised manuscript. So, for example, when weighting based on the CMIP5 model HadGEM2-ES we exclude the CMIP6 models HadGEM3-GC31-LL and UKESM1-0-LL from the evaluation. A list of CMIP6 models excluded for each CMIP5 model can be found in table S5 in the supplement and we have added some discussion about this topic in section S3 of the supplement.

11. Line 266-280. Here we learn that the method is very prone to creating decreased skill relative to the multi-model unweighted mean. This negative result is in contrast to the positive results found in Abramowitz and Bishop (2015) using the method of Bishop and Abramowitz (2013).

Thank you for pointing this out, this was not expressed clearly in the original manuscript. In fact, the method produces a median skill increase of about 12-22% when using CMIP5 models as pseudo observations (see figure 3a in the revised manuscript). Nonetheless, it is correct that there can be a decrease in skill from the unweighted to the weighted multi-model ensemble based on our skill metric when using some CMIP5 models as pseudo-observations. However, these instances are limited to only a few (about 15% across SSPs and target periods) cases. We have revised the paragraph in question to make this more clear (line 314-322 in the revised manuscript).

We note that the change in skill also depends on the skill metric used and the target it is applied to. Here, our target is 20-year mean, global mean temperature change from 1995-2014 to two future periods (2041-60 and 2081-00). As a skill metric, we use the continuous ranked probability skill score (CRPSS), a measure for ensemble forecast quality. Note that this does not only evaluate the distance between the (un-) weighted mean and the reference but also considers the full distribution.

Reduced global warming from CMIP6 projections when weighting models by performance and independence

Lukas Brunner¹, Angeline G. Pendergrass^{2,1*}, Flavio Lehner^{1*}, Anna L. Merrifield¹, Ruth Lorenz¹, and Reto Knutti¹

¹Institute for Atmospheric and Climate Science, ETH Zurich, Zurich, Switzerland

²National Center for Atmospheric Research, Boulder, CO, USA

*Now at: Department of Earth and Atmospheric Sciences, Cornell University, Ithaca, NY, USA

Correspondence: Lukas Brunner (lukas.brunner@env.ethz.ch)

Abstract. The sixth Coupled Model Intercomparison Project (CMIP6) constitutes the latest update on expected future climate change based on a new generation of climate models. To extract reliable estimates of future warming and related uncertainties from these models, the spread in their projections is often translated into probabilistic estimates such as mean and likely range. Here, we use a model weighting approach, which accounts for <u>a model's the models</u>' historical performance based on several

- 5 diagnostics as well as possible-model inter-dependence within the CMIP6 ensemble, to calculate constrained distributions of global mean temperature change. We investigate the skill of our approach in a perfect model test, where we remove each CMIP6 model from the ensemble in turn, use it as pseudo-observation use previous-generation CMIP5 models as pseudo-observations in the historical period, and evaluate the weighted CMIP6 ensemble against it in the future. This is complemented by a second perfect model test drawing on the previous-generation CMIP5 models as-. The performance of the so weighted distribution in
- 10 matching the pseudo-observations in the future is then evaluated and we find a mean increase in skill of about 17 % compared to the unweighted distribution. In addition, we show that our independence diagnostics metric correctly clusters models known to be similar based on a CMIP6 "family tree", which enables applying a weighting based on the degree of inter-model dependence. We then apply the weighting approach, based on two observational estimates (ERA5 and MERRA2), to constrain CMIP6 projections in weak (SSP1-2.6) and strong (SSP5-8.5) climate change scenarios. Our results show a reduction in pro-
- 15 jected mean warming for both scenarios because some CMIP6 models with high future warming receive systematically lower performance weights. The mean of end-of-century warming (2081-2100 relative to 1995-2014) for SSP5-8.5 with weighting is 3.7 °C, compared to 4.1 °C without weighting; the likely (66 %) uncertainty range is 3.1 °C to 4.6 °C, a decrease in spread of 13 %. For SSP1-2.6, weighted end-of-century warming is 1 °C (0.7 °C to 1.4 °C). Applying the weighting to estimates of Transient Climate Response (TCR) yields 1.9 °C (1.6 °C to 2.1 °C a reduction in the likely uncertainty range of 46 %), which
- 20 is consistent with estimates from previous model generations and other lines of evidencea reduction of -0.2 °C in the mean and -24% in the likely range compared to the unweighted case.

1 Introduction

Projections of future climate by Earth System Models provide a crucial source of information for adaptation planing, mitigation decisions, and the scientific community alike. Many of these climate model projections are coordinated and provided within the
frame of the Coupled Model Intercomparison Projects (CMIPs), which are now in phase 6 (Eyring et al., 2016). A typical way of communicating information from such multi-model ensembles (MMEs) is by combining them into probabilistic distributions, such as through a best estimate and uncertainty rangean uncertainty range or a probabilistic distribution. In doing so it is important to make sure that the different sources of uncertainty are identified, discussed, and accounted for, to provide reliable information without being overconfident. Typically In climate science typically three main sources of uncertainty are identified

30 in MMEs: (i) uncertainty in future emissions, (ii) internal variability of the climate system, and (iii) model response uncertainty (e.g., Hawkins and Sutton, 2009; Knutti et al., 2010).

Uncertainty due to future emissions can easily be isolated by making projections conditional on scenarios such as the Shared Socioeconomic Pathways (SSPs) in CMIP6 (O'Neill et al., 2014) or the Representative Concentration Pathways (RCPs) in CMIP5 (van Vuuren et al., 2011). The other two sources of uncertainty are harder to quantify since reliably separating

- 35 them is often challenging (e.g., Kay et al., 2015; Maher et al., 2019). Model uncertainty arises due to different responses and feedbacks of (sometimes also referred to as structural uncertainty or response uncertainty) is used here to describe the differing responses of climate models to a given radiative forcing, leading to different estimates of mean warming or Transient Climate Response (TCR) (e.g., Forster et al., 2013) forcing due to their structural differences following the definition by Hawkins and Sutton (2009). Such different responses to the same forcing can emerge, among other things, due to different
- 40 processes and feedbacks as well as due to the parametrisations used in the different models (e.g., Zelinka et al., 2020). Internal variability, finally, here refers to a model's sensitivity to the initial conditions as captured by initial-condition ensemble members (e.g., Deser et al., 2012). In this sense, it stems from the chaotic behavior of the climate system at different time scales and is highly dependent on the variable of interest as well as the period and region averaged overconsidered. While, for example, uncertainty in global mean temperature is mainly dominated by differences between models, regional temperature trends are
- 45 considerably more dependent on internal variabilityas can be estimated from. Recently, efforts have been made to use so-called Single Model Initial-condition Large Ensembles (SMILEs) (Lehner et al., 2020; Maher et al., 2019; Merrifield et al., 2019)to investigate internal variability in the climate projections more comprehensively (e.g., Kay et al., 2015; Maher et al., 2019; Lehner et al., 201

Depending on the composition of the investigated MME, uncertainty estimates often fail to reflect that included models are not always independent from each other. In the development process of climate models, ideas, code and even full components

50 are shared between institutions or models might be branched from each other in order to investigate specific questions. This can lead to some models (or model components) being copied more often, resulting in an over-representation of their respective internal variability or sensitivity to forcing (Bishop and Abramowitz, 2013; Boé, 2018; Boé and Terray, 2015)(Masson and Knutti, 2011; Bis The CMIP MMEs in particular have not been designed with the aim of including only independent models and are therefore sometimes referred to as "ensembles of opportunity" (e.g., Tebaldi and Knutti, 2007) incorporating as many models as possible.

- 55 When calculating probabilities based on such MMEs it is therefore important to account for model inter-dependence in order to accurately translate model spread into estimates of mean change and related uncertainties (Knutti, 2010; Knutti et al., 2010). In addition, not all models represent the aspects of the climate system relevant to a given question equally well. To account for that, a variety of different approaches have been used to weight, sub-select, or constrain models based on their historical performance. This has been done both regionally and globally as well as for a range of different target metrics such as end-of-century
- 60 temperature change or TCR (see, e.g., Brunner et al., 2020b; Eyring et al., 2019; Knutti et al., 2017a, for an overview)Transient Climate Response (TCR) (for an overview see, e.g., Knutti et al., 2017a; Eyring et al., 2019; Brunner et al., 2020b). Global mean temperature increase in particular is one of the most widely discussed effects of continuing climate change and the main focus of many public and political discussions. With the release of the new generation of CMIP6 models, this discussion has been sparked yet again, as several CMIP6 models show stronger warming than most of the earlier-generation CMIP5 models

65 (Forster et al., 2020; Zelinka et al., 2020; Swart et al., 2019; Gettelman et al., 2019; Voldoire et al., 2019; Golaz et al., 2019; Andrews et a This raises the question of whether these models are accurate representations of the climate system and what that means for the interpretation of the historical climate record and the expected change due to future anthropogenic emissions.

Here, we use the Climate model Weighting by Independence and Performance (ClimWIP) method (e.g., Merrifield et al., 2019; Brunner of weight models in the CMIP6 MME. Weights are based on (i) each models model's performance in simulating historical proper-

- 70 ties of the climate system such as horizontally resolved anomaly, variability, and trend fields, and (ii) its independence from the other models in the ensemble, estimated based on shared biases of climatology. In contrast to many other methods, which constrain model projections based on only one observable quantity, such as the warming trend (e.g., Giorgi and Mearns, 2002; Ribes et al., 201' ClimWIP is based on multiple diagnostics, representing different aspects of the climate system. These diagnostics are chosen to evaluate a model's performance in simulating observed climatology, variability, and trend patterns. Note that, in contrast to
- 75 other approaches such as emergent constraint-based methods, some of these diagnostics might not be highly correlated with the target metric (however, it is still important that they are physically relevant – to avoid introducing noise without useful information in the weighting). Combining a range of relevant diagnostics is less prone to overconfidence, since the risk of up-weighting a model because it "accidentally" fits observations for one diagnostic, while being far away from them in several others is greatly reduced. In turn, methods which are based on such a basket of diagnostics have been found to generally lead
- to weaker constraints (Sanderson et al., 2017; Brunner et al., 2020b), as the effect of the weighting typically weakens when adding more diagnostics (Lorenz et al., 2018).

ClimWIP has already been used to create estimates of regional change and related uncertainties for a range of different variables such as Arctic sea ice (Knutti et al., 2017b), Antarctic ozone concentrations (Amos et al., 2020), North American maximum temperature (Lorenz et al., 2018) and European temperature and precipitation (Merrifield et al., 2019; Brunner et al., 2019).

85 (Brunner et al., 2019; Merrifield et al., 2020). Recently, Liang et al. (2020) have used an adaptation of the method to constrain changes in global temperature using global mean temperature trend as single diagnostic for both the performance and independence weighting. Here, we focus on investigating the ClimWIP methods method's performance in weighting global mean temperature changes when informed by different a range of diagnostics. To assess the robustness of these choices, we perform an out-of-sample perfect model test using CMIP5 and CMIP6 as pseudo-observations. Based on these results, we select a com-

- 90 bination of diagnostics which capture not only a model's transient warming but also its ability to reproduce historical patterns in climatology and variability fields in order to increase the robustness of the weighting scheme and minimize the risk of skill decreases due to the weighting. This approach is particularly important for users interested in the "worst case" rather than in mean changes. We also look into the inter-dependencies among the models, showing the ability of our diagnostics in clustering models with known shared components using a "family tree" (Masson and Knutti, 2011; Knutti et al., 2013) and further the
- 95 skill of the independence weighting to account for this. We then calculate combined performance-independence weights based on two reanalysis products in order to also account for the uncertainty in the observational record. Finally, we apply these weights to provide constrained distributions of future warming and CTRTCR.

2 Data and Methods

2.1 Model data

- 100 The analysis is based on all currently available CMIP6 models which provide surface air temperature (tas) and sea level pressure (psl) for the historical, SSP1-2.6, and SSP5-8.5 experiments. We use all available ensemble members, which is a total of 129 runs from 33 models (see table \$3.54 in the supplementary material for a full list including references). We use models post-processed within the ETH Zurich CMIP6 next generation archive, which provides additional quality checks and re-grids models onto a common 2.5°×2.5° latitude-longitude grid, using second order conservative remapping (see Brunner et al., 2020a, for details). In addition, we use the first-one member of all CMIP5 models providing the same variables and the
- corresponding experiments (historical, RCP2.6, RCP8.5) which is a total of 27 models (see table \$4.55 for a full list).

2.2 Reanalysis data

To represent historical observations in tas and psl, we use two reanalysis products: ERA5 (C3S, 2017) and MERRA2 (Gelaro et al., 2017; Gi (GMAO, 2015a, b; Gelaro et al., 2017). Both products are regridded to a $2.5^{\circ} \times 2.5^{\circ}$ latitude-longitude grid using second order

- 110 conservative remapping and are evaluated in the period 1980-2014. Within the framework of the model weighting, they are combined to provide an estimate of observational uncertainty (see Brunner et al., 2019, for details)We use a combination of these two observational datasets following the results of Lorenz et al. (2018) and Brunner et al. (2019), who show that using individual datasets separately can lead to diverging results in some cases. It has been argued that that combining multiple datasets (e.g., by using their full range or their mean) yields more stable results (Gleckler et al., 2008; Brunner et al., 2019).
- 115 Here we use the mean of ERA5 and MERRA2 at each grid point as reference equivalent to Brunner et al. (2019). Finally, we also compare our results to globally averaged merged temperatures from the Berkley Earth Surface Temperature (BEST) data set (Cowtan, 2019).

2.3 Model weighting scheme

120

We use an updated version of the ClimWIP method described in Merrifield et al. (2019) and Brunner et al. (2019) Brunner et al. (2019) and Merrifield et al. (2020), which is based on earlier work by Lorenz et al. (2018), Knutti et al. (2017b), Sanderson et al. (2015b), and Sanderson et al. (2015a); it can be downloaded at: https://github.com/lukasbrunner/ClimWIP.git. It assigns a weight w_i to each model m_i that accounts for both model performance as well as independence,

$$w_i = \frac{e^{-\left(\frac{D_i}{\sigma_D}\right)^2}}{1 + \sum_{j \neq i}^M e^{-\left(\frac{S_{ij}}{\sigma_S}\right)^2}},\tag{1}$$

where D_i and S_{ij} are the generalised distances of model m_i to the observations and to model m_j , respectively. The shape parameters σ_D and σ_S set the strength of the weighting, effectively determining the point at which a model is considered to be 125 "close" to the observations or to another model (c.f., section 2.5).

This updated version of ClimWIP assigns the same weight to each initial-condition ensemble member of a model, which is adjusted by the number of ensemble members (see the revised version of Merrifield et al., 2019, for a detailed discussion) (see Merrifield et To illustrate this additional step in the weighting method, consider a single performance diagnostic d. d is calculated for each model and ensemble member separately, hence $d = d_k^k$ with *i* representing individual models , and *k* running over all ensemble

130

members K_i of model m_i (in CMIP6, from one to 50 members in CMIP6). For each model m_i , the mean diagnostic d'_i is,

$$d_i' = \frac{\sum_{k=1}^{K} d_i^k}{K_i} \underbrace{\text{for all } i.}_{(2)}$$

 d'_i is then used to calculate the generalised distance D_i and further the performance weight w_i via (1). A detailed description of this processing chain can be found in section S2 in the supplement. An analogous process is used for distances between

135

models. This setup allows a consistent comparison of model fields to each other and to observations in the presence of internal variability and, in particular, also enables the use of variance-based diagnostics. In addition, it ensures a consistent estimate of the performance shape parameter σ_D in the perfect model test calibration (see section 2.5), based on the average weight per model; in previous work, in contrast, it the calibration was based on only one ensemble member per model.

Weighting target and diagnostics 2.4

- We apply the weighting to projections of annual mean, global mean temperature change from two SSPs, representing weak 140 (SSP1-2.6) and strong (SSP5-8.5) climate change scenarios. Changes in two 20-year target periods representing mid-century (2041-2060) and end-of-century (2081-2100) conditions are compared to a 1995-2014 baseline. In addition, we weight TCR values from all available models obtained from an update of the data set described in Tokarska et al. (2020). The weights are calculated from global, horizontally-resolved diagnostics based on annual mean data in the 35-year period 1980-2014. We use
- 145 different diagnostics for the calculation of the independence and performance parts of the weighting, as proposed in the revised version of Merrifield et al. (2019)Merrifield et al. (2020).

The goal of the independence weighting is to identify structural similarities between models (such as shared offsets or similar spatial patterns) which are interpreted to be indications of inter-dependence arising from, e.g., shared components or parametrisations. In the past, combinations of horizontally-resolved regional temperature, precipitation, and sea level pressure fields,

- 150 have typically been used (e.g., Brunner et al., 2019; Sanderson et al., 2017; Knutti et al., 2013; Boé, 2018; Lorenz et al., 2018). Following (e.g., Knutti et al., 2013; Sanderson et al., 2017; Boé, 2018; Lorenz et al., 2018; Brunner et al., 2019). Building on the work of Merrifield et al. (2019)Merrifield et al. (2020), we use a combination of two global, climatology-based diagnostics, the spatial pattern of climatological temperature (tasCLIM) and sea level pressure (pslCLIM), that as similar diagnostics were found to work well for clustering CMIP5-generation models known to be similar. This definition of independence does not
- 155 Beside our approach, several other methods to tackle this issue of model dependence exist. Among them are approaches which use other metrics to establish model independence (e.g., Pennell and Reichler, 2011; Bishop and Abramowitz, 2013; Boé, 2018), which select a more independent sub-set of the original ensemble (e.g., Leduc et al., 2016; Herger et al., 2018a), or even treat model similarity as an indication for robustness and give models which are closer to the multi model mean more weight (e.g., Giorgi and Mearns, 2002; Tegegne et al., 2019). Neither of these definitions of independence hold in a purely strictly sta-
- 160 tistical sense (Annan and Hargreaves, 2017), but we still stress that it is important to account for different degrees of model inter-dependencies as well-good as possible when developing probabilistic estimates from an "ensemble of opportunity" such as CMIP6. We validate this approach in section 4.2 of the results Additional discussion about our method to calculate model independence in the context of other approaches can be found in section S4 of the supplement.
- The performance weighting, in turn, allocates more weight to models which better represent the observed behavior of the climate system as measured by the diagnostics, while down-weighting models with large discrepancies from the observations. We use multiple diagnostics to limit overconfidence in the case where a model fits the observations well in one diagnostic by chance, while being far away from them in several others. For example, we want to avoid giving heavy weight to a model based solely on its representation of the temperature trend if its year-to-year variability differs strongly from observed yearto-year variability. The performance weights are based on five global, horizontally-resolved diagnostics: temperature anomaly
- 170 (tasANOM; calculated from tasCLIM by removing the global mean), temperature variability (tasSTD), psIANOM, and psISTD as well as temperature trend (tasTREND). A detailed description of the diagnostic calculation can be found in section S2 in the supplement. We use anomalies instead of climatologies in the performance weight in order to avoid punishing models for absolute bias in global-mean temperature and pressure, because these are not correlated with projected warming (Flato et al., 2013; Giorgi and Coppola, 2010). This can be different for regional cases, where, e.g., absolute temperature biases have
- 175 been shown to be important for constraining projections of Arctic sea ice extent (Knutti et al., 2017b) or European summer temperatures (Selten et al., 2020).

One aim of our study is to find an optimal combination of diagnostics that successfully constrains projections for our target quantity (global temperature change) while avoiding overconfidence or susceptibility to uncertainty from internal variability. For example, tasTREND is a powerful diagnostic because of its clear physical relationship to and high correlation with pro-

180 jected warming (e.g., Tokarska et al., 2020; Nijsse et al., 2020)(e.g., Nijsse et al., 2020; Tokarska et al., 2020). However, while it has the highest correlation to the target of all investigated diagnostics, it also has the largest uncertainty due to internal vari-

ability (i.e., spread of tasTREND across ensemble members of the same model). Ideally, a performance weight is reflective of underlying model properties and does not depend on which ensemble member is chosen to represent that model(i. e., on internal variability)... tasTREND does not fulfil this requirement: the spread within one model is the same order of magnitude as the spread among different models. To find a compromise, we divide our diagnostics into two groups: trend-based diag-

185

as the spread among different models. To find a compromise, we divide our diagnostics into two groups: trend-based diagnostics (tasTREND) and not-trend based diagnostics (tasANOM, tasSTD, psIANOM, and psISTD). Different combinations of these two groups (ranging from only not-trend based to only tasTREND) are evaluated in section 3.1 and the best performing combination is selected for the remainder of the study.

2.5 **Calculation Estimation** of the shape parameters

- 190 The shape parameters σ_D and σ_S are two constants which determine the width of the Gaussian weighting functions -for all models. As such they are responsible for translating the generalised distances into weights. In case of the performance weighting, small values of σ_D lead to very aggressive weighting with a few models receiving all the weight, while large values lead to more equal weighting. It is important to note that, while σ_D sets this "strength" of the weighting, the rank of a model (i.e., where it lies on the scale from best to worst) is purely based on its generalised distance to the observations. To estimate a
- 195 performance shape parameter σ_D that weights models based on their historical performance without being overconfident, we use a calibration approach based on the perfect model test detailed in Knutti et al. (2017b) in Knutti et al. (2017b) and detailed in section S3 in the supplement. In short, the test calibration selects the smallest σ_D value (hence the strongest weighting) for which 80 % of perfect models "perfect models" fall within the 10-90 percentile range of the weighted distribution -in the target period. Smaller σ_D values lead to less models fulfilling this criterion and hence to too narrow, overconfident projections. Note
- 200 that methods that simply maximize correlation of the weighted mean to the target in a perfect model test often tend to pick small values of σ_D that result in projections that are overconfident in the sense that the uncertainty ranges are too small (Knutti et al., 2017b). A similar issue arises for methods which estimate σ_D based only on historical information as better performance in the base state does not necessarily lead to a more skill representation of the future, e.g., if the chosen diagnostics are not relevant for the target (Sanderson and Wehner, 2017).
- 205 The independence weighting has a subtle but fundamentally different dependence on its shape parameter σ_S : small values lead to equal weighting, as all models are considered to be independent, but so do large values, as all models are considered to be *de*pendent. Hence, the effect of the independence weighting is strongest if the shape parameter is chosen such that it identifies clusters of models as similar (down-weighting them) while still correctly identifying models which are far from each other as independent (hence giving them relatively more weight)(see revised version of Merrifield et al., 2019, for a more detailed discussion includin
- 210 . For a detailed discussion including SMILEs see Merrifield et al. (2020). To estimate σ_S , we use the information from models with more than one ensemble member. We know that Simply put, we know that initial-condition ensemble members are copies of the same model that differ only due to internal variability, and therefore we have a priori some information about the correct independence weighting, distances that must be considered "close" by σ_S . The method for calculating σ_S is described in detail in section 3 of the supplement of Brunner et al. (2019). Here, we arrive at a value of $\sigma_S = 0.54$, which we use throughout
- 215 the manuscript. It is worth noting that σ_S is based only on historical model information, and is therefore independent from

observations or the selected target period or scenario. Following the method described in detail in Brunner et al. (2019), we arrive at a value of $\sigma_S = 0.54$, which we use throughout the manuscriptand scenario. Additional discussion of the selected σ_S value in the context of the multi-model ensemble used in this study can be found in the supplement (section S5).

2.6 Validation of the performance weighting

- To investigate the skill of ClimWIP in weighting CMIP6 global mean temperature change and the effect of the different diagnostic combinations (different relative importance of tasTREND), we apply a perfect model test (Abramowitz and Bishop, 2015; Boé and Terra As a skill measure, we use the continuous ranked probability skill score (CRPSS), a measure for ensemble forecast quality, defined as the relative error between the distribution of weighted models and a reference (Hersbach, 2000). Here, we define the CRPSS as relative use the relative CRPSS change between the unweighted and weighted cases (in %), with positive values
- 225 indicating a skill increase. The CRPSS is calculated separately for both SSPs and future time periods, since we expect to find different skill for different projected climate states.

The first perfect model test is based only on the CMIP6 MME and focuses on evaluating the performance weighting only focuses on the relative skill differences when applying performance weights based on different combinations of diagnostics (results are presented in section 3.1). We explain its implementation based on an example perfect model m_i with only one

- ensemble member for simplicity here: (i) the model m_j is taken as pseudo-observation and removed from the CMIP6 MME; (ii) the output from m_j during the historical diagnostic period (1980-2014) is used to calculate the performance diagnostics for the remaining models $(d'_{i\neq j})$; (iii) the generalised model-"observation" distances $(D_{i\neq j})$ and the performance weights $(w_{i\neq j})$ are calculated and applied to the MME (excluding m_j); (iv) the CRPSS is calculated in the target periods using the future projections of m_j as reference. This is done iteratively, using each model in CMIP6 MME in turn as pseudo-observation.
- For perfect models with more than one ensemble member (m_j^k) , all members are removed from the ensemble in (i), $d'_{i\neq j}$ is calculated for each member separately in (ii) and then averaged, and the CRPSS is also calculated for each ensemble member in (iv) and averaged.

We note that a similar perfect model test is also an integral part of ClimWIP as it is used to estimate This approach is structurally similar to the one used to calibrate the performance shape parameter σ_D as integral part of ClimWIP (described in

- section 2.5), which introduces a small amount of circularity in this test. However, it is still valuable to investigate the skill of the weighting method using this test to (i) the metric and aim of this perfect model test are quite different. It is used to show the potential for an increase in skill through a skill increase through the performance weighting, as well as the risk of a decrease $\frac{1}{7}$ (ii) cross-check the based on the selected σ_D calculation, and (iii) compare different fractions of trend-versus not-trend-based diagnostics, in order and to establish the most skilful combination of diagnostics.
- The second perfect model test (section 3.2) is conceptually equivalent similar, but pseudo-observations are now drawn from CMIP5 instead of CMIP6. This test has the advantages that we can always use the full CMIP6 MME (without having to remove any models) and that the the perfect models have not been used to estimate σ_D and can be considered independent, at least in a methodological sense. Note that they are not necessarily independent in a model sense. Even though one might argue that also the CMIP5 pseudo-observations are not fully out-of-sample as several CMIP6 models descend from are related to

- 250 CMIP5 models and might be structurally similar to their predecessors, which was the case for the CMIP5 and 3 generations (Knutti et al., 2013). However, there are also considerable differences between CMIP5 and 6 that arise from many years of additional model development, a longer observational record to tune calribrate to, and differing spatial resolutions. In addition, the emission scenarios that force CMIP5 and 6 in the future (RCPs and SSPs, respectively) result in slightly different radiative forcings (Forster et al., 2020) ; determining how these scenario families differ is currently an active area of researchand several
- 255 CMIP6 have been shown to lead to considerably more warming than most CMIP5 models. We do not discuss these similarities and differences between the model generations in detail here; instead we simply use CMIP5 simply as a source of additional for pseudo-observations to evaluate the skill of ClimWIP for in weighting the CMIP6 MMEto improve the fit to a given. To avoid cases with the highest potential of remaining dependence between generations we exclude CMIP6 models which are direct predecessors of the respective CMIP5 model used as pseudo observations (see table S5 for a list).

260 2.7 Validation of the independence weighting

To validate that the information in the diagnostics chosen for the independence weighting (tasCLIM and pslCLIM) can identify models known to be similar, we use a hierarchical clustering approach based on Müllner (2011) and implemented in the Python SciPy package (www.scipy.org). We use the linkage function with the average method applied to the horizontally-resolved distance fields between each pair of models (see section S6 in the supplement for more details). This approach is conceptually similar to the work from Masson and Knutti (2011) and Knutti et al. (2013) and follows their example of showing similarity as model "family trees". The hierarchical clustering is *not* used in the model weighting itself; we use it here only to show that qualitative information about model similarity can be inferred from model output using the two chosen diagnostics and to compare it to the results from the independence weighting.

- The independence weighting (denominator in equation (1)) quantifies the similarity information extracted from the pairwise distance fields via the independence shape parameter (σ_S ; see section 2.5). The independence weighting estimates where two models fall on the spectrum from completely independent to completely redundant and weights them accordingly. In order to test this approach, we successively add artificial "new" models into the CMIP6 MME: for an example model with two members $(m_j^1 \text{ and } m_j^2)$, we remove the first member and add it as additional model (m_{M+1}) . In an idealized case, where all models are perfectly independent from each other and all ensemble members of a model are identical, we would expect the weight of the member that remains (m_j^2) to go down by a factor 1/2, while the weight of all other models would stay the same. However, in
- a real MME, where there is internal variability and complex model inter-dependencies exist, we would not necessarily expect such simple behaviour; several other models might also be (rightfully) affected by adding such a duplicate while the effect on the m_j^2 would be smaller (see section 4.2)



Figure 1. Continuous ranked probability skill score (CRPSS) relative to the unweighted ensemble for the performance weighting based on a leave-one-out perfect model test with CMIP6 for (a) mid-century and (b) end-of-century temperature change relative to 1995-2014. The x-axis shows different combinations of the two diagnostic groups (see section 2.4) ranging from only not-trend based (0 % tasTREND) to only trend-based (100 % tasTREND). Values not summing to 100 % is due to rounding in the labels only.

3 Evaluation of the weighting in the perfect model test

280 3.1 Leave-one-out perfect model test with CMIP6

We start by calculating the performance weights in the diagnostic period (1980-2014) in a pure model world and without using the independence weighting. In this first step we focus on the evaluation of the performance weighting relative skill differences when using different combinations of diagnosticsand on calculating the ideal performance shape parameters (σ_D). Figure 1 shows the distribution of the CRPSS (with positive values indicating an increase in projection skill due to the weighting and vice versa; see section 2.6) evaluated for two the mid- and end-of-century target periods, the two SSPs, and for different combinations of diagnostics range from only not-trend based (0% tasTREND ; using only tasANOM, tasSTD , psIANOM, and psISTD + 25% tasANOM + 25% tasSTD + 25% psIANOM + 25% psISTD = 100%) to only tasTREND based (100% tasTREND). Overall, all diagnostic combinations tend to increase median skill compared to the unweighted projections, but there is a considerable range of CRPSS values and they can be negative. In evaluating the different cases we hence focus on two important aspects of the CRPSS distribution: (i) the median as best estimate of expected relative skill change and (ii) the 5th and 25th percentiles in particular if they are negative. Negative CRPSS values indicate a worsening of the projections compared to the unweighted case. Since the goal of the weighting is to improve the projections based on performance and dependence of the models, the risk of negative CRPSSs should be minimised.

We find the σ_D -values to be correctly <u>chosen calibrated</u> by the method in order to limit the risk for a strong skill decrease (CRPSS is close to zero or positive for the 25th percentile in almost all cases). For the mid-century period, the median skill



Figure 2. Time series of temperature change (relative to 1995-2014) for unweighted (gray) and weighted (colored) CMIP6 mean (lines) and likely (66 %) range (shading) as well as the CMIP5 models serving as pseudo-observations (dashed lines). Shown are the cases wich-which lead to (a) the largest decrease in skill (CMIP5 pseudo-observation: CanEMS2MIROC-ESM) and (b) to the largest increase (MPI-ESM-LR) for SSP5-8.5 in the end-of-century target period. Note that no inference on the performance of the CMIP5 models can be drawn from this figure. Diagnostic period refers to the 1980-2014 period, which informs the weights; the target periods to 2041-2060 and 2081-2100.

increases by about 10% to 20% across both SSPs and all up to 25% depending on SSP and combination of diagnostics. The magnitude of potential negative CRPSSs in a "worst-case" scenario (5th percentile), however, is better constrained using a balanced combination of diagnostics (e.g., 50% tasTREND). In the end-of-century period, the median skill is more variable (mainly due to the selected performance shape parameters σ_D ; see table S1), with combinations that include both trend and not-trend diagnostics again performing best.

300

305

Using 50 % tasTREND and 50 % anomaly- and variance-based diagnostics (tasANOM, tasSTD, pslANOM, about 13 % tasANOM, 13 % tasSTD, 13 % pslANOM, and 13 % pslSTD) optimises the combination of median CRPSS increases and avoidance of possible negative CRPSSs; we therefore use this combination to calculate the weights for the rest of the analysis. Note that the two SSPs and time periods have slightly different σ_D values (ranging from 0.35 to 0.58; table S1), leading to slightly differing weights even though the historical information is the same. This arises from differences in confidence when applying the method for different targets. However, since the σ_D values are found to be so similar we use the mean value from the two SSPs and time periods in the following for simplicity, hence $\sigma_D = 0.43$. This does not have a strong influence on the results but simplifies their presentation and interpretation.

3.2 Perfect model test using CMIP5 as pseudo-observations

310 We now use each of the 27 CMIP5 models in turn as pseudo-observation and include both the performance and independence parts of the method. For all considerations in this section, we use the CMIP5 merged historical and RCP runs corresponding to the CMIP6 historical and SSP runs, i.e., RCP2.6 to SSP1-2.6 and RCP8.5 to SSP5-8.5. This allows an evaluation of the skill of the <u>full</u> weighting method applied to the full CMIP6 MME in the future. Figure 2 shows two cases selected to lead to the largest decrease (figure 2a) and increase (figure 2b) in the CRPSS for SSP5-8.5 in the end-of-century period when applying

- 315 the weights. The figures reveal This reveals an important feature of the weighting: if the unweighted MME is already close to the "truth" the risk for a skill decrease is highest (constraining methods in general: there is a risk that the information from the historical period might not lead to a skill increase in the future. In the case shown in figure 2a). In other words, using the CMIP5 model CanESM2, which happens to be close to the unweighted CMIP6 MME mean , as weighting based on pseudoobservations to weigh CMIP6 tends to pull the CMIP6 MME mean from MIROC-ESM shifts the distribution downwards, while
- 320 projections from MIROC-ESM end up warming more than the unweighted mean in the future. This reflects the possibility that information drawn from real historical observations might not lead to an increase in projection skill in some cases. Here cases of decreasing skill appear for about 15% of pseudo-observations.

The largest skill increases, in turn, often comes from pseudo-observations rather far away from the pseudo-observational "truth". In the reverse case unweighted mean. It seems that, if the "truth" is pseudo-observations behave very different from the

- 325 MME mean e.g., the CMIP5 model MPI-ESM-LR being rather different from the CMIP6 MME mean –, the potential for a skill increase is highest (figure 2b).model ensemble in the historical period, there is a good chance that they will continue to do so in the future. One explanation for this could be a systematic difference between the models in the ensemble and the pseudo observation due to, e.g., a missing feedback or component. An important cautionary takeaway is thus to not only maximise median mean skill increase when setting up the method, as the cases with highest skill might come from rather "unrealistic"
- 330 pseudo-observations (i.e., the ones on the tails of the model distribution, like). This is illustrated in figure 2-and figure S1S5 in the supplement (e.g., using the CMIP5 GFDL or GISS models as pseudo observations). However, in many cases we do not necessarily expect the real climate to follow such an extreme trajectory but rather be closer to the unweighed unweighted MME mean (in part because real observations tend to be used in model development and tuning). It is thus important to use a balanced set of multiple diagnostics and not only optimise for maximal correlation in choosing σ_D , which might make
- 335 the highest possible skill increases unattainable, but maybe more importantly guard against even more substantial skill decreases.

Finally, it is important to note that the skill of the weighting for a given pseudo-observation also depends on the target. In isolated cases that can mean that the weighting leads to an increase in skill for one SSP while it leads to a decrease in the other (e.g., IPSL-CM5A-LR as pseudo-observation) or to an increase in one time period and to a decrease in the other (e.g.,

340 <u>CSIRO-Mk3-6-0</u>). An overview of the weighting based on each of the 27 CMIP5 models can be found in figure <u>S1-S5</u> in the supplement.

To look into the skill change more quantitatively, figure 3a shows the skill distribution of weighting CMIP6 to predict each of the pseudo-observations drawn from CMIP5 for both target time periods and scenarios. We note again that for each CMIP5 pseudo-observation directly related CMIP6 models are excluded (see table S5 for a list). Compared to the leave-one-out perfect

model test with CMIP6 shown in figure 1 the increase in median CRPSS is lower and the risk for negative CRPSSs is slightly higher. This is not unexpected for a test sample , which has not been used for training (i.e., the estimation of the σ_D -value) and which is structurally different from CMIP6 in several aspects (such as forcing scheme and maximum amount of warming).



Figure 3. (a) Similar to figure 1 but using 27 CMIP5 models as pseudo-observations and showing only the 50 % tasTREND case. (b) Map of median of the CRPSS values relative to the unweighted ensemble for 2041-2060 under SSP5-8.5

But the setup still achieves a median CRPSS increase of about 10% to 20%12% to 22%, with the risk of for a skill reduction being mostly confined to less than 25\%, clearly showing confined to about 15\% of cases and to a maximum decrease of about 25%. This clearly shows that ClimWIP can be used to provide reliable estimates of future global temperature change and related uncertainties from the CMIP6 MME.

350

Finally, we consider the question of whether there are regional patterns in the skill change by investigating a map of median CRPSSs for SSP5-8.5 in the mid-century period in figure 3b (see figure $\frac{82-56}{10}$ in the supplement for the other cases). Note that each CMIP6 model is still assigned only one weight, but the CRPSS is calculated at each grid point separately. The skill

- 355 increases almost everywhere with the northern hemisphere having a slightly higher amplitude. A notable exception is the North Atlantic, where weighting leads to a slight decrease in median skill. Indeed, this is the only region where the unweighted CMIP6 mean underestimates the warming from CMIP5. Weighting the CMIP6 ensemble leads to a slight strengthening of the underestimation in this region, while it reduces the difference almost everywhere else.
- In summary, weighting CMIP6 in a perfect model test using five different diagnostics to establish model performance and two diagnostics for independence shows an increase in a clear increase in median skill compared to the unweighted distribution for the vast majority of cases and consistent over both investigated scenarios and time periods. Looking into the geographical distribution reveals an increase in skill almost everywhere, with some decreases found in the Southern Ocean, particularly in SSP1-2.6 (figure <u>\$256</u>). Importantly, skill increases almost everywhere over land, thus benefiting assessments of climate impacts and adaptation where people are affected most directly.

365 4 Weighting CMIP6 projections of future warming based on observations

So far we have selected a combination of diagnostics, which leads to the highest increase in median skill while minimising the risk for a skill decrease based on an out-of-sample perfect model test with CMIP6 in section 3.1. We also argued that we use the same shape parameters (which determine the strength of the weighting) for all cases, namely $\sigma_S = 0.54$ for independence and $\sigma_D = 0.43$ for performance. In section 3.2 we then evaluated this setup by using 27 pseudo-observations drawn from the

370 CMIP5 MME. In this section we now calculate weights for CMIP6 based on observed climate and validate the effect of the independence weighting.

We use observational surface air temperature and sea level pressure estimates from the ERA5 and MERRA2 reanalyses to calculate the performance diagnostics (tasANOM, tasSTD, tasTREND, psIANOM, psISTD). The combination of two reanalysis products allows to account for observational uncertainty, which has been found to be important for robust weighting

375 in earlier work by Brunner et al. (2019) and Lorenz et al. (2018). As independence diagnostics we continue to use model-model distances in tasCLIM and pslCLIM.

4.1 Calculation of weights for CMIP6

Figure 4 shows the combined performance and independence weights assigned to each CMIP6 model by ClimWIP when applied to the target of global temperature change. Three general regimes can be identified: (i) models which represent historical

- 380 observations better than average receive relative weights mostly between In addition also the individual performance and independence weights are shown. All three cases are individually normalised. Applying the combined weight, about half of the models receive more weight than in a simple arithmetic mean and about half receive less. The best performing model, GFDL-ESM4, has about four times more influence than it would have without weighting (about 0.13 compared to 0.03 in the case with equal weighting). The three lowest performing models, MIROC-ES2L, CanESM5, and HadGEM3-GC31-LL, in turn
- 385 receive less than 1 and 2 (with a maximum of about 4), (ii)models which represent historical observations slightly less well, but can still be considered skillful representations of the climate system, receive relative weights mostly between 1 and 0.5, and (iii)models which can be considered less skillful based on their past performance receive weights of less than 0.2. /20 of the equal weighting (about 0.001).

Indeed, several recent studies have found that models which show more future warming per unit of greenhouse gas are less

likely based on comparison with past observations (e.g., Jiménez-de-la Cuesta and Mauritsen, 2019; Nijsse et al., 2020; Tokarska et al., 2020; Consistent with their findings models with high TCR receive very low performance (and combined) weights (label colours in figure 4). Among the five lowest ranking models four have a TCR above 2.5 °C and all models with TCR above 2.5 °C receive less then equal weight. The eight highest ranking models, in turn, have TCR values ranging from 1.5 °C to 2.5 °C and lie, therefore, rather in the middle of the CMIP6 TCR range. See table S2 in the supplement for a summary of all model weights

395 and TCR values.

In addition to the combined weighting, figure 4 also shows the pure performance weights. The relative differences independence and performance weights separately. We discuss model independence in more detail in the next section. For the model



Figure 4. Combined independence-performance weights for each CMIP6 model (line with dots) and was well as pure performance weights (squares) relative to equal weighing and pure independence weights (triangles). Weights smaller than 0.2 times All three cases are individually normalised and the equal weighting are only each model would receive in a normal arithmetic mean is shown as their approximate combined weight for reference (fractions in the right bottom cornerdashed line). The labels are coloured by each models TCR value: $> 2.5^{\circ}$ C - red, $> 2^{\circ}$ C - yellow, $> 1.5^{\circ}$ C - green, and $\le 1.5^{\circ}$ C - blue. The number of ensemble members per model is shown in brackets after the model name in the x-axis labels.

performance weighting, the relative difference to the combined weights are weighting (i.e., the influence of the independence weighting) is mostly below 50 %, with the MIROC model family being one notable exception. Both MIROC models are very independent, which shifts MIROC6 from a below-average model (based on the pure performance weight; black-square in figure 4) to an above-average model in the combined weight (black-dot) effectively more than doubling its performance weight. For MIROC-ES2L the scaling due to independence is similarly high(not visible in figure 4), but its total weight is still dominated by the very low performance weight. In the next section we investigate if these independence weights indeed correctly

represent the complex model inter-dependencies in the CMIP6 MME and down-weight models which are highly dependent on

405

400

4.2 Validation of the independence weighting

other models appropriately.

To test if model inter-dependence can correctly be inferred from model output in general, we first take a quantitative approach, somewhat different to the model (independence) weighting itself. Focusing on the independence weights in figure 4 one can broadly distinguish three cases: (i) relatively independent models, (ii) clusters of models which are quite dependent, and



Figure 5. Model "family tree "-for all 33 CMIP6 models used in this study similar to Knutti et al. (2013). Based Models branching further to the left are more dependent, models branching further to the right are more independent. The analysis is based on global, horizontally resolved tasCLIM and pslCLIM in the period 1980-2014. The independence shape parameter σ_S is indicated as dashed vertical line, an estimation of internal variability as grey shading. Labels with the same colour indicate models with obvious dependencies such as shared components or same origin (models with no clear dependencies are labelled in black). Weak relations such as remote "ancestors" are not eolored together (e.g., BCC-CSM2-MR and CESM2).



Figure 6. Similar to figure 4 but removing one variant initial-condition ensemble member from (a) MIROC6 and (b) MPI-ESM1-2-HR and adding it as separate model when calculating the independence weights (the "new" model is not shown in the plot). Models with obvious dependencies to the "new" model have bold labels (same as in equivalent to figure 5) have bold labels. The change in the combined weight relative to the original weight is shown as blue bars using the right axis.

410 (iii) models for which the independence weighting does not really influence the weighting. To visualise and discuss these cases somewhat quantitatively, we show a CMIP6 model family tree similar to the work by Masson and Knutti (2011) and Knutti et al. (2013).

Using the same two diagnostics, namely horizontally resolved global temperature and sea level pressure climatologies (from 1980-2014) we apply a hierarchical clustering approach (section 2.7). Figure 5 shows the resulting <u>"family tree" family tree</u> of

- 415 CMIP6 models similar to the work by Masson and Knutti (2011) and Knutti et al. (2013). Models In this tree models which are closely related branch further to the left, while very independent model clusters branch further to the right. The mean distance between two initial-condition members of the same as an estimation for the internal variability in the generalised distance is indicated as grey shading. Model which have a distance similar to this value (e.g., the two CanESM5 model versions) are basically indistinguishable. The independence shape parameter used through the manuscript ($\sigma_S = 0.54$) is shown as dashed
- 420 <u>vertical line</u>.

A comprehensive investigation of the complex inter-dependencies within the multi-model ensemble in use and further between models from the same institution or of similar origin is beyond the scope of this study and will be subject of future work. Here we limit ourselves to pointing out several base features of the output-based clustering, which serve as indications that it is skilful in identifying inter-dependent models. The labels of models with the same origin or with known shared compo-

425 nents are marked in the same colour , as this is in figure 5. These two factors are the most objective measure for a priori model dependence we have. The information about the model components is taken from each models model's description page on the ES-DOC explorer (https://es-doc.org/cmip6/) as listed in table S3-S4 in the supplement.

Figure 5 clearly shows that clustering models based on the selected diagnostics performs well: models with shared components or with the same origin (indicated by the same colour) are always grouped together. Looking into a bit more detail we find, for

- 430 example, that closely related models such as low and high resolution versions (MPI-ESM-2-LR and MPI-ESM-2-HR; CNRM-CM6-1 and CNRM-CM6-1-HR) or versions with only one differing component (CESM2 and CESM2-WACCM; INM-CM5-0 and INM-CM4-8; both differing only in the atmosphere) are detected as being very similar. Both MIROC models, which have been identified as very independent based on figure 4are again, in turn, are found to be very far away from each other and even further away from all other models in the CMIP6 MME.
- To investigate if the independence weighting correctly identifies and weights models based on their degree of inter-dependence translates model distance into weights we now look at two models as examples: one model that performs well and is relatively independent (MIROC6) and another that also performs well but is more dependent (MPI-ESM1-2-HR). Each has multiple ensemble members; we remove one member from each and add it to the MME as an additional model as detailed in section 2.7. In the first case (figure 6a; MIROC6 which is among the least dependent models), the original weight is reduced by almost
- 440 1/2, which is close to what we would expect in the idealised case. All other models are unaffected by adding a duplicate of MIROC6, even the other model from the same center, MIROC-ES2L which differs in atmospheric resolution and cumulus treatment (Tatebe et al., 2019; Hajima et al., 2019). Based on the "family tree" shown in figure 5 this behaviour is not surprising: the two MIROC models are not only identified as the most independent models in the CMIP6 MME but also as very independent

from each <u>other</u>. While some of the components and parameterizations are similar, updates in parameterizations and in the tuning of the parameters appear to be sufficient here to create a model that behaves quite differently.

The second case (figure 6b; MPI-ESM1-2-HR which is among the most dependent models) shows a very different picture. The strongest effect on the original weight is found for the copied model itself, which is reduced by about 0.820%, but also several other models are affected: MPI-ESM1-2-LR (reduced by 0.86), AWI-CM-1-1-MR (0.9), NESM3 (0.93), MRI-ESM2-0 (0.94), and CAMS-CSM1-0 (0.94). Looking into the these models in more detail, we conclude that the inter-dependencies
detected by our method can be traced to shared components in most cases: MPI-ESM1-2-LR is just the low resolution version of MPI-ESM1-2-HR (run with a T63 atmosphere instead of T127 and a 1.5° ocean instead of 0.4°), AWI-CM-1-1-MR and

- NESM3 share the atmospheric (ECHAM6.3) and similar land (JSBACH3.x) components, and CAMS-CSM1-0 shares a similar atmospheric (ECHAM5) component, while MRI-ESM2-0 does not have any obvious dependencies. Information about the models can be found in their reference publications (Mauritsen et al., 2019; Gutjahr et al., 2019; Semmler et al., 2019; Yang et al., 2020; Chen et al., 2019; Yukimoto et al., 2019) and on the ES-DOC explorer, which provides detailed information about
- all model used in this study. The links to each models information page can be found in table $\frac{83}{54}$ in the supplementary material.

4.3 Applying weights to CMIP6 temperature projections and TCR

- Figure 7 shows a timeseries of unweighted and weighted projections based on a weak (SSP1-2.6) and strong (SSP5-8.5) climate
 change scenario. For both scenarios a clear shift in the mean towards less warming is visible, which is also reflected in the upper uncertainty bound. Notably, however, the lower bound hardly changes, leading to a reduction in projection uncertainty in total. This becomes even clearer when investigating the two 20-year periods, reflecting mid- and end-of-century conditions (figure 8a and table \$2\$3).
- Based on these results, warming exceeding 5 °C by the end of the century is very unlikely even under the strongest climate change scenario SSP5-8.5. The mean warming for this case is shifted downward to about 3.7 °C and the 66 % (likely) and 90 % ranges are reduced by 12 % and 30 %, respectively. For SSP1-2.6 in the end-of-century period as well as both SSPs in the midcentury period, reductions in the mean warming of about 0.1 °C to 0.2 °C are found. The likely range is reduced by about 30% - 20% to 30 % in these three cases. A summary of all-weights and warming values for all models as well as all statistics can be found in tabletables S2 and S3 in the supplement. Recent studies that use historical temperature trend as an observational
- 470 constraint for future warming lead to similar conclusions, with lower constrained warming compared to unconstrained (both in the mean and upper percentiles of the distributions) (e.g., Tokarska et al., 2020; Nijsse et al., 2020)(e.g., Nijsse et al., 2020); Tokarska et al., 2020; Nijsse et al., 2020); Tokarska et al., 2020; Nijsse et al., 2020; Nijsee et al., 2020; Nijsee et al., 2020; Nijsee et al., 2

To investigate the influence of remaining internal variability in our combination of diagnostics on the weighting, we also perform a bootstrap test. Selecting only one random member per model (for models with more than one ensemble member) we calculate weights and the corresponding unweighted and weighted temperature change distributions. This is repeated 100

475 times, providing uncertainty estimates for both the unweighted and weighted percentiles. The mean values of the weighted percentiles taken over all 100 bootstrap samples are very similar to the values from the weighting based on the full MME (including all ensemble members; see figure \$3\$57) confirming the robustness of our approach.



Figure 7. Timeseries of temperature change (relative to 1995-2014) for unweighted (gray) and weighted (colored) CMIP6 mean (lines) and likely (66%) range (shading). Three observational datasets are also shown in black; note that BEST is not used to inform the weighting and is only shown for comparison here.

480

490

We also apply weights to TCR estimates in figure 8b. For four models included in the weighting of temperature change we do not yet have all information available to estimate TCR (FGOALS-g3, CanESM5-CanOE, FIO-ESM-2-0, MCM-UA-1-0); these are omitted in figure 8b. For the remaining 29 models we find a finding an unweighted mean TCR value of about 2 °C with a likely range of $\frac{1.6 \circ \text{C}}{1.6 \circ \text{C}} \frac{1.6 \circ \text{C}}{1.6 \circ \text{C}}$ to 1.9 °C (1.6 °C to 2.1 °C 1.6 °C to 2.2 °C), a reduction of 46%-36% in the likely range. These values are consistent with recent studies based on emergent constraints which estimate the likely range of TCR to be $\frac{1.5 \circ \text{C}}{1.3 \circ \text{C}}$ to $2.2 \circ \text{C}}{1.3 \circ \text{C}}$ to $2.1 \circ \text{C}$ (Nijsse et al., 2020) and 1.2 °C to 2.0 °C (Tokarska et al., 2020) and they are also very similar to the range of 1.5° to 2.2° from Sherwood et al. (2020) who combine multiple lines of evidence. They are also consistent but substantially more narrow than 485 the likely range from the fifth assessment report of the IPCC (IPCC, 2013) based on CMIP5: 1 °C to 2.5 °C.

Figure 8b clearly shows that almost all models with higher than equal weights lie within the likely range, and only one model lies above it (KACE-1-0-GFIO-ESM-2-0). This is a strong indication that TCR values beyond about $2.5 \,^{\circ}\text{C}$ are unlikely when weighting based on several diagnostics and when accounting for model independence. The weighting also largely reconciles CMIP6 with 5 by giving less weight to some of the models in CMIP6 that warm most strongly.





5 Discussion and Conclusions

We have used the Climate model Weighting by Independence and Performance (ClimWIP) method to constrain projections of future global temperature change from the CMIP6 multi-model ensemble. Based on a leave-one-out perfect model test, a combination of five global, horizontally-resolved diagnostic fields (anomaly, variance, and trend of surface air temperature and anomaly and variance of sea level pressure) was selected to inform the performance weighting. The skill of weighting based on this selection was tested and confirmed in a second perfect model test using CMIP5 models as pseudo-observations. Our results clearly show the usefulness of this weighting approach in translating model spread into reliable estimates of future changes and in particular into uncertainties that are consistent with observations of present day climate and observed trends.

500

495

We also discussed the remaining risk for decreasing skill compared to the raw distribution which is a crucial question in all weighting or constraining methods. We show the importance of using a balanced combination of climate system features (i.e., diagnostics) relevant for the target to inform the weighting to minimise the risk for skill decreases. This guards against the possibility of a model "accidentally" fitting observations for a single diagnostic while being far away from them in several others (and hence possibly not providing a skilful projection of the target variable).

505 weighting, showing that models get correctly down-weighted based on an estimate of dependence derived from their output. To inform the independence weighting we used two global, horizontally resolved fields (climatology of surface air temperature

By adding copies of existing models into the CMIP6 multi-model ensemble we verified the effect of the independence

and sea level pressure) which we showed to allow a clear clustering of models with obvious inter-dependencies using a CMIP6 "family tree".

From these tests we conclude that ClimWIP is skilful in weighting global mean temperature change from CMIP6 using the

- selected setup. We hence use it to calculate weights for each CMIP6 model and apply them in order to obtain probabilistic es-510 timates of future changes. Compared to the unweighted case these results clearly show that the CMIP6 models which lead to the highest warming are less probable, confirming earlier studies (e.g., Tokarska et al., 2020; Nijsse et al., 2020)(e.g., Nijsse et al., 2020); Sherv We find a weighted mean global temperature change (relative to 1995-2014) of 3.7 °C with a likely (66 %) range of 3.1 °C to 4.6 °C by the end of the century when following SSP5-8.5. With ambitious climate mitigation (SSP1-2.6) a weighted mean
- change of 1 °C (likely range: 0.7 °C to 1.4 °C) is projected for the same period. 515

On the policy level, this highlights the need for quick and decisive climate action to achieve the Paris climate targets. For climate modeling on the other hand, this approach demonstrates the potential to narrow the uncertainties in CMIP6 projections, particular on the upper bound. The large investments in climate model development have so far not led to reduced model spread in the raw ensemble, but the use of climatological information and emergent transient constraints has the potential to provide

more robust projections with reduced uncertainties, that at the same time are more consistent with observed trends, thus 520 maximizing the value of climate model information for impacts and adaptation.

Code availability. The ClimWIP model weighting package is available under a GPLv3 at https://github.com/lukasbrunner/ClimWIP.git

Author contributions. LB, ALM, and RK were involved in conceiving the study. LB did the analysis and created the plots substantially supported by AGP. LB wrote the manuscript with contributions from all authors. The ClimWIP package was implemented by LB and RL; 525 AGP wrote the script used to create tables S4 and S6.

Competing interests. The authors declare that they have no conflict of interest.

530

for helpful discussions and comments on the manuscript. The EUCP project This work was carried out in the frame of the EUCP project which is funded by the European Commission through the Horizon 2020 Programme for Research and Innovation: Grant Agreement 776613. Ruth Lorenz was funded and Anna L. Merrifield was co-funded by the European Union's Horizon 2020 Research and Innovation program: Grant Agreement 641816 (CRESCENDO). Flavio Lehner is Flavio Lehner was supported by a Swiss NSF Ambizione Fellowship (Project PZ00P2_174128). This material is partly based upon work supported by the National Center for Atmospheric Research, which is a major facility sponsored by the National Science Foundation (NSF) under Cooperative Agreement No. 1947282, and by the Regional and Global Model Analysis (RGMA) component of the Earth and Environmental System Modeling Program of the U.S. Department of Energy's Office

Acknowledgements. The authors thank Martin B. Stolpe for providing the TCR values as well as Martin B. Stolpe and Katarzyna B. Tokarska

- 535 of Biological & Environmental Research (BER) via NSF IA 1844590. This study was generated using Copernicus Climate Change Service Information 2020 from ERA5. The authors thank NASA for providing MERRA2 and Berkeley Earth for providing BEST. We acknowledge the World Climate Research Programme, which, through its Working Group on Coupled Modelling, coordinated and promoted CMIP5 and 6. We thank the climate modeling groups for producing and making available their model output, the Earth System Grid Federation (ESGF) for archiving the data and providing access, and the multiple funding agencies who support CMIP5 and 6 and ESGF. A list of all CMIP6
- 540 runs and their references can be found in table S6 in the supplement. We thank all contributors to the numerous Python open source packages which were crucial for this work, in particular the xarray project Python project xarray (http://xarray.pydata.org). The authors thank two anonymous reviewers for their helpful comments on our work.

References

Abramowitz, G. and Bishop, C. H.: Climate model dependence and the ensemble dependence transformation of CMIP projections, J. Clim., 28, 2332–2348, https://doi.org/10.1175/JCLI-D-14-00364.1, 2015.

- Abramowitz, G., Herger, N., Gutmann, E., Hammerling, D., Knutti, R., Leduc, M., Lorenz, R., Pincus, R., and Schmidt, G. A.: ESD Reviews: Model dependence in multi-model climate ensembles: weighting, sub-selection and out-of-sample testing, Earth Syst. Dyn., 10, 91–105, https://doi.org/10.5194/esd-10-91-2019, https://www.earth-syst-dynam.net/10/91/2019/, 2019.
 - Amos, M., Young, P. J., Hosking, J. S., Lamarque, J.-F., Abraham, N. L., Akiyoshi, H., Archibald, A. T., Bekki, S., Deushi, M., Jöckel,
- 550 P., Kinnison, D., Kirner, O., Kunze, M., Marchand, M., Plummer, D. A., Saint-Martin, D., Sudo, K., Tilmes, S., and Yamashita, Y.: Projecting ozone hole recovery using an ensemble of chemistry-climate models weighted by model performance and independence, Atmospheric Chemistry and Physics Discussions, 2020, 1–26, https://doi.org/10.5194/acp-2020-86, https://www.atmos-chem-phys-discuss. net/acp-2020-86/, 2020.

Andrews, T., Andrews, M. B., Bodas-Salcedo, A., Jones, G. S., Kuhlbrodt, T., Manners, J., Menary, M. B., Ridley, J., Ringer, M. A., Sellar,

- A. A., Senior, C. A., and Tang, Y.: Forcings, Feedbacks, and Climate Sensitivity in HadGEM3-GC3.1 and UKESM1, Journal of Advances in Modeling Earth Systems, 11, 4377–4394, https://doi.org/10.1029/2019MS001866, 2019.
 - Annan, J. D. and Hargreaves, J. C.: On the meaning of independence in climate science, Earth System Dynamics, 8, 211–224, https://doi.org/10.5194/esd-8-211-2017, 2017.
- Bishop, C. H. and Abramowitz, G.: Climate model dependence and the replicate Earth paradigm, Climate Dynamics, 41, 885–900, https://doi.org/10.1007/s00382-012-1610-y, 2013.
 - Boé, J.: Interdependency in Multimodel Climate Projections: Component Replication and Result Similarity, Geophysical Research Letters, 45, 2771–2779, https://doi.org/10.1002/2017GL076829, http://doi.wiley.com/10.1002/2017GL076829, 2018.
 - Boé, J. and Terray, L.: Can metric-based approaches really improve multi-model climate projections? The case of summer temperature change in France, Climate Dynamics, 45, 1913–1928, https://doi.org/10.1007/s00382-014-2445-5, 2015.
- 565 Brunner, L., Lorenz, R., Zumwald, M., and Knutti, R.: Quantifying uncertainty in European climate projections using combined performanceindependence weighting, Environmental Research Letters, 14, 124010, https://doi.org/10.1088/1748-9326/ab492f, http://dx.doi.org/10. 1038/ngeo3017, 2019.

Brunner, L., McSweeney, C., Ballinger, A. P., Hegerl, G. C., Befort, D. J., O'Reilly, C., Benassi, M., Booth, B., Harris, G., Lowe, J., Coppola, E., Nogherotto, R., Knutti, R., Lenderink, G., de Vries, H., Qasmi, S., Ribes, A., Stocchi, P., and Undorf, S.: Comparing methods to constrain future European climate projections using a consistent framework, J. Clim., pp. 1–62, https://doi.org/10.1175/jcli-d-19-0953.1, 2020b.

Chen, X., Guo, Z., Zhou, T., Li, J., Rong, X., Xin, Y., Chen, H., and Su, J.: Climate Sensitivity and Feedbacks of a New Coupled Model CAMS-CSM to Idealized CO 2 Forcing: A Comparison with CMIP5 Models, Journal of Meteorological Research, 33, 31–45, https://doi.org/10.1007/s13351-019-8074-5, 2019.

Brunner, L., Hauser, M., Lorenz, R., and Beyerle, U.: The ETH Zurich CMIP6 next generation archive : technical documentation, https://doi.org/10.5281/zenodo.3734128, 2020a.

C3S: ERA5: Fifth generation of ECMWF atmospheric reanalyses of the global climate, https://doi.org/10.24381/cds.f17050d7, accessed: 26.3.2020, 2017.

Cowtan, K.: The Climate Data Guide: Global surface temperatures: BEST: Berkeley Earth Surface Temperatures, https://climatedataguide.
 ucar.edu/climate-data/global-surface-temperatures-best-berkeley-earth-surface-temperatures, last modified 09 Sep 2019, 2019.

- Deser, C., Phillips, A., Bourdette, V., and Teng, H.: Uncertainty in climate change projections: the role of internal variability, Clim. Dyn., 38, 527–546, https://doi.org/10.1007/s00382-010-0977-x, http://link.springer.com/10.1007/s00382-010-0977-x, 2012.
- Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., and Taylor, K. E.: Overview of the Coupled Model Intercomparison Project Phase 6 (CMIP6) experimental design and organization, Geoscientific Model Development, 9, 1937–1958, https://doi.org/10.5194/gmd-9-1937-2016, https://www.geosci-model-dev.net/9/1937/2016/, 2016.
- Eyring, V., Cox, P. M., Flato, G. M., Gleckler, P. J., Abramowitz, G., Caldwell, P., Collins, W. D., Gier, B. K., Hall, A. D., Hoffman, F. M., Hurtt, G. C., Jahn, A., Jones, C. D., Klein, S. A., Krasting, J. P., Kwiatkowski, L., Lorenz, R., Maloney, E., Meehl, G. A., Pendergrass, A. G., Pincus, R., Ruane, A. C., Russell, J. L., Sanderson, B. M., Santer, B. D., Sherwood, S. C., Simpson, I. R., Stouffer, R. J., and Williamson, M. S.: Taking climate model evaluation to the next level, Nature Climate Change, 9, 102–110, https://doi.org/10.1038/s41558-018-0355-y, http://dx.doi.org/10.1038/s41558-018-0355-y, 2019.
- Flato, G., Marotzke, J., Abiodun, B., Braconnot, P., Chou, S., Collins, W., Cox, P., Driouech, F., Emori, S., Eyring, V., Forest, C., Gleckler, P., Guilyardi, E., Jakob, C., Kattsov, V., Reason, C., and Rummukainen, M.: Evaluation of Climate Models, in: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assess- ment Report of the Intergovernmental Panel on Climate Change, edited by Stocker, T., Qin, D., Plattner, G.-K., Tignor, M., Allen, S., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P., Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- Forster, P. M., Andrews, T., Good, P., Gregory, J. M., Jackson, L. S., and Zelinka, M.: Evaluating adjusted forcing and model spread for historical and future scenarios in the CMIP5 generation of climate models, Journal of Geophysical Research Atmospheres, 118, 1139– 1150, https://doi.org/10.1002/jgrd.50174, 2013.
- Forster, P. M., Maycock, A. C., McKenna, C. M., and Smith, C. J.: Latest climate models confirm need for urgent mitigation, Nature Climate
 Change, 10, 7–10, https://doi.org/10.1038/s41558-019-0660-0, 2020.
- Gelaro, R., McCarty, W., Suárez, M. J., Todling, R., Molod, A., Takacs, L., Randles, C. A., Darmenov, A., Bosilovich, M. G., Reichle, R., Wargan, K., Coy, L., Cullather, R., Draper, C., Akella, S., Buchard, V., Conaty, A., da Silva, A. M., Gu, W., Kim, G. K., Koster, R., Lucchesi, R., Merkova, D., Nielsen, J. E., Partyka, G., Pawson, S., Putman, W., Rienecker, M., Schubert, S. D., Sienkiewicz, M., and Zhao, B.: The modern-era retrospective analysis for research and applications, version 2 (MERRA-2), Journal of Climate, 30, 5419–5454, https://doi.org/10.1175/JCLI-D-16-0758.1, 2017.
- Gettelman, A., Hannay, C., Bacmeister, J. T., Neale, R. B., Pendergrass, A. G., Danabasoglu, G., Lamarque, J., Fasullo, J. T., Bailey, D. A., Lawrence, D. M., and Mills, M. J.: High Climate Sensitivity in the Community Earth System Model Version 2 (CESM2), Geophysical Research Letters, 46, 8329–8337, https://doi.org/10.1029/2019GL083978, https://onlinelibrary.wiley.com/doi/abs/10.1029/2019GL083978, 2019.
- 610 Giorgi, F. and Coppola, E.: Does the model regional bias affect the projected regional climate change? An analysis of global model projections: A letter, Climatic Change, 100, 787–795, https://doi.org/10.1007/s10584-010-9864-z, 2010.
 - Giorgi, F. and Mearns, L. O.: Calculation of average, uncertainty range, and reliability of regional climate changes from AOGCM simulations via the "Reliability Ensemble Averaging" (REA) method, Journal of Climate, 15, 1141–1158, https://doi.org/10.1175/1520-0442(2002)015<1141:COAURA>2.0.CO;2, 2002.
- 615 Gleckler, P. J., Taylor, K. E., and Doutriaux, C.: Performance metrics for climate models, J. Geophys. Res. Atmos., 113, 1–20, https://doi.org/10.1029/2007JD008972, 2008.

GMAO: MERRA-2 tavg1_2d_slv_Nx: 2d,1-Hourly,Time-Averaged,Single-Level,Assimilation,Single-Level Diagnostics V5.12.4, https://disc.gsfc.nasa.gov/api/jobs/results/5e7b68e9ed720b5795af914a, accessed: 25.3.2020, 2015a.

GMAO: MERRA-2 statD_2d_slv_Nx: 2d,Daily,Aggregated Statistics,Single-Level,Assimilation,Single-Level Diagnostics V5.12.4, https:
 //disc.gsfc.nasa.gov/api/jobs/results/5e7b648f4900ab500326d17e, accessed: 25.3.2020, 2015b.

- Golaz, J. C., Caldwell, P. M., Van Roekel, L. P., Petersen, M. R., Tang, Q., Wolfe, J. D., Abeshu, G., Anantharaj, V., Asay-Davis, X. S., Bader, D. C., Baldwin, S. A., Bisht, G., Bogenschutz, P. A., Branstetter, M., Brunke, M. A., Brus, S. R., Burrows, S. M., Cameron-Smith, P. J., Donahue, A. S., Deakin, M., Easter, R. C., Evans, K. J., Feng, Y., Flanner, M., Foucar, J. G., Fyke, J. G., Griffin, B. M., Hannay, C., Harrop, B. E., Hoffman, M. J., Hunke, E. C., Jacob, R. L., Jacobsen, D. W., Jeffery, N., Jones, P. W., Keen, N. D., Klein, S. A., Larson,
- V. E., Leung, L. R., Li, H. Y., Lin, W., Lipscomb, W. H., Ma, P. L., Mahajan, S., Maltrud, M. E., Mametjanov, A., McClean, J. L., McCoy, R. B., Neale, R. B., Price, S. F., Qian, Y., Rasch, P. J., Reeves Eyre, J. E., Riley, W. J., Ringler, T. D., Roberts, A. F., Roesler, E. L., Salinger, A. G., Shaheen, Z., Shi, X., Singh, B., Tang, J., Taylor, M. A., Thornton, P. E., Turner, A. K., Veneziani, M., Wan, H., Wang, H., Wang, S., Williams, D. N., Wolfram, P. J., Worley, P. H., Xie, S., Yang, Y., Yoon, J. H., Zelinka, M. D., Zender, C. S., Zeng, X., Zhang, C., Zhang, K., Zhang, Y., Zheng, X., Zhou, T., and Zhu, Q.: The DOE E3SM Coupled Model Version 1: Overview and Evaluation at Standard
 Resolution, Journal of Advances in Modeling Earth Systems, 11, 2089–2129, https://doi.org/10.1029/2018MS001603, 2019.
- Gutjahr, O., Putrasahan, D., Lohmann, K., Jungclaus, J. H., Von Storch, J. S., Brüggemann, N., Haak, H., and Stössel, A.: Max Planck Institute Earth System Model (MPI-ESM1.2) for the High-Resolution Model Intercomparison Project (HighResMIP), Geoscientific Model Development, 12, 3241–3281, https://doi.org/10.5194/gmd-12-3241-2019, 2019.
- Hajima, T., Watanabe, M., Yamamoto, A., Tatebe, H., Noguchi, M., Abe, M., Ohgaito, R., Ito, A., Yamazaki, D., Okajima, H., Ito, A.,
 Takata, K., Ogochi, K., Watanabe, S., and Kawamiya, M.: Description of the MIROC-ES2L Earth system model and evaluation of its climate-biogeochemical processes and feedbacks, Geoscientific Model Development Discussions, 5, 1–73, https://doi.org/10.5194/gmd-

2019-275, 2019.

- Hawkins, E. and Sutton, R.: The Potential to Narrow Uncertainty in Regional Climate Predictions, Bulletin of the American Meteorological Society, 90, 1095–1108, https://doi.org/10.1175/2009BAMS2607.1, http://journals.ametsoc.org/doi/10.1175/2009BAMS2607.1, 2009.
- 640 Herger, N., Abramowitz, G., Knutti, R., Angélil, O., Lehmann, K., and Sanderson, B. M.: Selecting a climate model subset to optimise key ensemble properties, Earth Syst. Dyn., 9, 135–151, https://doi.org/10.5194/esd-9-135-2018, 2018a.
 - Herger, N., Angélil, O., Abramowitz, G., Donat, M., Stone, D., and Lehmann, K.: Calibrating Climate Model Ensembles for Assessing Extremes in a Changing Climate, J. Geophys. Res. Atmos., 123, 5988–6004, https://doi.org/10.1029/2018JD028549, 2018b.
 - Hersbach, H.: Decomposition of the Continuous Ranked Probability Score for Ensemble Prediction Systems, Weather and Forecast-
- ing, 15, 559–570, https://doi.org/10.1175/1520-0434(2000)015<0559:DOTCRP>2.0.CO;2, http://journals.ametsoc.org/doi/abs/10.1175/
 1520-0434%282000%29015%3C0559%3ADOTCRP%3E2.0.CO%3B2, 2000.
 - IPCC: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change [Stocker,, 9, Cambridge University Press, 2013.
- Jiménez-de-la Cuesta, D. and Mauritsen, T.: Emergent constraints on Earth's transient and equilibrium response to doubled CO2 from post-1970s global warming, Nature Geoscience, 2015, https://doi.org/10.1038/s41561-019-0463-y, 2019.
 - Kay, J. E., Deser, C., Phillips, A., Mai, A., Hannay, C., Strand, G., Arblaster, J. M., Bates, S. C., Danabasoglu, G., Edwards, J., Holland, M., Kushner, P., Lamarque, J. F., Lawrence, D., Lindsay, K., Middleton, A., Munoz, E., Neale, R., Oleson, K., Polvani, L., and Vertenstein, M.: The community earth system model (CESM) large ensemble project : A community resource for studying climate change in the

presence of internal climate variability, Bulletin of the American Meteorological Society, 96, 1333–1349, https://doi.org/10.1175/BAMS-

655 D-13-00255.1, 2015.

- Knutti, R.: The end of model democracy?, Clim. Change, 102, 395–404, https://doi.org/10.1007/s10584-010-9800-2, 2010.
- Knutti, R., Furrer, R., Tebaldi, C., Cermak, J., and Meehl, G. A.: Challenges in combining projections from multiple climate models, Journal of Climate, 23, 2739–2758, https://doi.org/10.1175/2009JCLI3361.1, 2010.
- Knutti, R., Masson, D., and Gettelman, A.: Climate model genealogy: Generation CMIP5 and how we got there, Geophysical Research
 Letters, 40, 1194–1199, https://doi.org/10.1002/grl.50256, 2013.
- Knutti, R., Rugenstein, M. A., and Hegerl, G. C.: Beyond equilibrium climate sensitivity, Nature Geoscience, 10, 727–736, https://doi.org/10.1038/NGEO3017, http://dx.doi.org/10.1038/ngeo3017, 2017a.
- Knutti, R., Sedláček, J., Sanderson, B. M., Lorenz, R., Fischer, E. M., and Eyring, V.: A climate model projection weighting scheme accounting for performance and interdependence, Geophysical Research Letters, 44, 1909–1918, https://doi.org/10.1002/2016GL072012, http://doi.wiley.com/10.1002/2016GL072012, 2017b.
 - Leduc, M., Laprise, R., de Elía, R., and Šeparović, L.: Is institutional democracy a good proxy for model independence?, J. Clim., 29, 8301–8316, https://doi.org/10.1175/JCLI-D-15-0761.1, 2016.
 - Lehner, F., Deser, C., Maher, N., Marotzke, J., Fischer, E., Brunner, L., Knutti, R., and Hawkins, E.: Partitioning climate projection uncertainty with multiple Large Ensembles and CMIP5/6, Earth System Dynamics Discussions, pp. 1–28, https://doi.org/10.5194/esd-2019-93, 2020.
- 670 Liang, Y., Gillett, N. P., and Monahan, A. H.: Climate Model Projections of 21st Century Global Warming Constrained Using the Observed Warming Trend, Geophys. Res. Lett., 47, 1–10, https://doi.org/10.1029/2019GL086757, 2020.
 - Lorenz, R., Herger, N., Sedláček, J., Eyring, V., Fischer, E. M., and Knutti, R.: Prospects and Caveats of Weighting Climate Models for Summer Maximum Temperature Projections Over North America, Journal of Geophysical Research: Atmospheres, 123, 4509–4526, https://doi.org/10.1029/2017JD027992, http://doi.wiley.com/10.1029/2017JD027992, 2018.
- 675 Maher, N., Milinski, S., Suarez-Gutierrez, L., Botzet, M., Dobrynin, M., Kornblueh, L., Kröger, J., Takano, Y., Ghosh, R., Hedemann, C., Li, C., Li, H., Manzini, E., Notz, D., Putrasahan, D., Boysen, L., Claussen, M., Ilyina, T., Olonscheck, D., Raddatz, T., Stevens, B., and Marotzke, J.: The Max Planck Institute Grand Ensemble: Enabling the Exploration of Climate System Variability, Journal of Advances in Modeling Earth Systems, https://doi.org/10.1029/2019MS001639, 2019.

Masson, D. and Knutti, R.: Climate model genealogy, Geophysical Research Letters, 38, 1-4, https://doi.org/10.1029/2011GL046864, 2011.

- 680 Mauritsen, T., Bader, J., Becker, T., Behrens, J., Bittner, M., Brokopf, R., Brovkin, V., Claussen, M., Crueger, T., Esch, M., Fast, I., Fiedler, S., Fläschner, D., Gayler, V., Giorgetta, M., Goll, D. S., Haak, H., Hagemann, S., Hedemann, C., Hohenegger, C., Ilyina, T., Jahns, T., Jimenéz-de-la Cuesta, D., Jungclaus, J., Kleinen, T., Kloster, S., Kracher, D., Kinne, S., Kleberg, D., Lasslop, G., Kornblueh, L., Marotzke, J., Matei, D., Meraner, K., Mikolajewicz, U., Modali, K., Möbis, B., Müller, W. A., Nabel, J. E., Nam, C. C., Notz, D., Nyawira, S. S., Paulsen, H., Peters, K., Pincus, R., Pohlmann, H., Pongratz, J., Popp, M., Raddatz, T. J., Rast, S., Redler, R., Reick, C. H., Rohrschneider,
- 685 T., Schemann, V., Schmidt, H., Schnur, R., Schulzweida, U., Six, K. D., Stein, L., Stemmler, I., Stevens, B., von Storch, J. S., Tian, F., Voigt, A., Vrese, P., Wieners, K. H., Wilkenskjeld, S., Winkler, A., and Roeckner, E.: Developments in the MPI-M Earth System Model version 1.2 (MPI-ESM1.2) and Its Response to Increasing CO2, Journal of Advances in Modeling Earth Systems, 11, 998–1038, https://doi.org/10.1029/2018MS001400, 2019.
- Merrifield, A. L., Brunner, L., Lorenz, R., and Knutti, R.: Weighting scheme to incorporate large ensembles in multi-model ensemble projections, Earth System Dynamics, https://doi.org/10.5194/esd-2019-69, https://doi.org/10.5194/esd-2019-69, 2019.

27

Merrifield, A. L., Brunner, L., Lorenz, R., Medhaug, I., and Knutti, R.: An investigation of weighting schemes suitable for incorporating large ensembles into multi-model ensembles, Earth System Dynamics, 11, 807–834, https://doi.org/10.5194/esd-11-807-2020, https://esd. copernicus.org/articles/11/807/2020/, 2020.

Müllner, D.: Modern hierarchical, agglomerative clustering algorithms, pp. 1–29, http://arxiv.org/abs/1109.2378, 2011.

- 695 Nijsse, F. J. M. M., Cox, P. M., and Williamson, M. S.: An emergent constraint on Transient Climate Response from simulated historical warming in CMIP6 models, Earth System Dynamics, pp. 1–14, https://doi.org/10.5194/esd-2019-86, https://doi.org/10.5194/esd-2019-86, 2020.
- O'Neill, B. C., Kriegler, E., Riahi, K., Ebi, K. L., Hallegatte, S., Carter, T. R., Mathur, R., and van Vuuren, D. P.: A new scenario framework for climate change research: the concept of shared socioeconomic pathways, Climatic Change, 122, 387–400, https://doi.org/10.1007/s10584-013-0905-2, http://link.springer.com/10.1007/s10584-013-0905-2, 2014.
 - Pennell, C. and Reichler, T.: On the Effective Number of Climate Models, J. Clim., 24, 2358–2367, https://doi.org/10.1175/2010JCLI3814.1, http://journals.ametsoc.org/doi/abs/10.1175/2010JCLI3814.1, 2011.
 - Ribes, A., Zwiers, F. W., Azaïs, J. M., and Naveau, P.: A new statistical approach to climate change detection and attribution, Climate Dynamics, 48, 367–386, https://doi.org/10.1007/s00382-016-3079-6, 2017.
- 705 Sanderson, B. and Wehner, M.: Appendix B. Model Weighting Strategy, Forth Natl. Clim. Assess., 1, 436–442, https://doi.org/10.7930/J06T0JS3, 2017.
 - Sanderson, B. M., Knutti, R., and Caldwell, P.: A representative democracy to reduce interdependency in a multimodel ensemble, Journal of Climate, 28, 5171–5194, https://doi.org/10.1175/JCLI-D-14-00362.1, 2015a.

Sanderson, B. M., Knutti, R., and Caldwell, P.: Addressing interdependency in a multimodel ensemble by interpolation of model properties,
 Journal of Climate, 28, 5150–5170, https://doi.org/10.1175/JCLI-D-14-00361.1, 2015b.

- Sanderson, B. M., Wehner, M., and Knutti, R.: Skill and independence weighting for multi-model assessments, Geoscientific Model Development, 10, 2379–2395, https://doi.org/10.5194/gmd-10-2379-2017, 2017.
- Selten, F. M., Bintanja, R., Vautard, R., and van den Hurk, B. J.: Future continental summer warming constrained by the present-day seasonal cycle of surface hydrology, Scientific Reports, 10, 1–7, https://doi.org/10.1038/s41598-020-61721-9, http://dx.doi.org/10.1038/s41598-020-61721-9, http://dx.doi.000-61721-9, http://dx.doi.000-61721-9, ht

715

s41598-020-61721-9, 2020.

- Semmler, T., Danilov, S., Gierz, P., Goessling, H., Hegewald, J., Hinrichs, C., Koldunov, N. V., Khosravi, N., Mu, L., and Rackow, T.: Simulations for CMIP6 with the AWI climate model AWI-CM-1-1, Earth and Space Science Open Archive, p. 48, https://doi.org/10.1002/essoar.10501538.1, 2019.
- Sherwood, S., Webb, M. J., Annan, J. D., Armour, K. C., Forster, P. M., Hargreaves, J. C., Hegerl, G., Klein, S. A., Marvel, K. D., Rohling, E. J., Watanabe, M., Andrews, T., Braconnot, P., Bretherton, C. S., Foster, G. L., Hausfather, Z., von der Heydt, A. S., Knutti, R., Mauritsen,
- T., Norris, J. R., Proistosescu, C., Rugenstein, M., Schmidt, G. A., Tokarska, K. B., and Zelinka, M. D.: An assessment of Earth's climate sensitivity using multiple lines of evidence, Reviews of Geophysics, pp. 1–92, https://doi.org/10.1029/2019rg000678, 2020.
 - Swart, N. C., Cole, J. N., Kharin, V. V., Lazare, M., Scinocca, J. F., Gillett, N. P., Anstey, J., Arora, V., Christian, J. R., Hanna, S., Jiao, Y., Lee,W. G., Majaess, F., Saenko, O. A., Seiler, C., Seinen, C., Shao, A., Sigmond, M., Solheim, L., Von Salzen, K., Yang, D., and Winter, B.: The
- 725 Canadian Earth System Model version 5 (CanESM5.0.3), Geoscientific Model Development, 12, 4823–4873, https://doi.org/10.5194/gmd-12-4823-2019, 2019.
 - Tatebe, H., Ogura, T., Nitta, T., Komuro, Y., Ogochi, K., Takemura, T., Sudo, K., Sekiguchi, M., Abe, M., Saito, F., Chikira, M., Watanabe, S., Mori, M., Hirota, N., Kawatani, Y., Mochizuki, T., Yoshimura, K., Takata, K., O'Ishi, R., Yamazaki, D., Suzuki, T., Kurogi, M., Kataoka,

T., Watanabe, M., and Kimoto, M.: Description and basic evaluation of simulated mean state, internal variability, and climate sensitivity in MIROC6, Geoscientific Model Development, 12, 2727–2765, https://doi.org/10.5194/gmd-12-2727-2019, 2019.

Tebaldi, C. and Knutti, R.: The use of the multi-model ensemble in probabilistic climate projections, Philosophical Transactions of the Royal Society A: Mathematical, Physical and Engineering Sciences, 365, 2053–2075, https://doi.org/10.1098/rsta.2007.2076, http://rsta.royalsocietypublishing.org/cgi/doi/10.1098/rsta.2007.2076, 2007.

730

- Tegegne, G., Kim, Y.-o., and Lee, J.-k.: Spatiotemporal reliability ensemble averaging of multi-model simulations, Geophys. Res. Lett., p.
 2019GL083053, https://doi.org/10.1029/2019GL083053, https://onlinelibrary.wiley.com/doi/abs/10.1029/2019GL083053, 2019.
- Tokarska, K. B., Stolpe, M. B., Sippel, S., Fischer, E. M., Smith, C. J., Lehner, F., and Knutti, R.: Past warming trend constrains future warming in CMIP6 models, Science Advances, 6, eaaz9549, https://doi.org/10.1126/sciadv.aaz9549, https://advances.sciencemag.org/ lookup/doi/10.1126/sciadv.aaz9549, 2020.
 - van Vuuren, D. P., Edmonds, J., Kainuma, M., Riahi, K., Thomson, A., Hibbard, K., Hurtt, G. C., Kram, T., Krey, V., Lamarque, J. F., Masui,
- 740 T., Meinshausen, M., Nakicenovic, N., Smith, S. J., and Rose, S. K.: The representative concentration pathways: An overview, Climatic Change, 109, 5–31, https://doi.org/10.1007/s10584-011-0148-z, 2011.
 - Voldoire, A., Saint-Martin, D., Sénési, S., Decharme, B., Alias, A., Chevallier, M., Colin, J., Guérémy, J., Michou, M., Moine, M., Nabat, P.,
 Roehrig, R., Salas y Mélia, D., Séférian, R., Valcke, S., Beau, I., Belamari, S., Berthet, S., Cassou, C., Cattiaux, J., Deshayes, J., Douville,
 H., Ethé, C., Franchistéguy, L., Geoffroy, O., Lévy, C., Madec, G., Meurdesoif, Y., Msadek, R., Ribes, A., Sanchez-Gomez, E., Terray, L.,
- 745 and Waldman, R.: Evaluation of CMIP6 DECK Experiments With CNRM-CM6-1, Journal of Advances in Modeling Earth Systems, 11, 2177–2213, https://doi.org/10.1029/2019MS001683, https://onlinelibrary.wiley.com/doi/abs/10.1029/2019MS001683, 2019.
 - Yang, Y.-M., Wang, B., Cao, J., Ma, L., and Li, J.: Improved historical simulation by enhancing moist physical parameterizations in the climate system model NESM3.0, Climate Dynamics, 54, 3819–3840, https://doi.org/10.1007/s00382-020-05209-2, https://doi.org/10.1007/ s00382-020-05209-2, 2020.
- Yukimoto, S., Kawai, H., Koshiro, T., Oshima, N., Yoshida, K., Urakawa, S., Tsujino, H., Deushi, M., Tanaka, T., Hosaka, M., Yabu, S., Yoshimura, H., Shindo, E., Mizuta, R., Obata, A., Adachi, Y., and Ishii, M.: The meteorological research institute Earth system model version 2.0, MRI-ESM2.0: Description and basic evaluation of the physical component, Journal of the Meteorological Society of Japan, 97, 931–965, https://doi.org/10.2151/jmsj.2019-051, 2019.
- Zelinka, M. D., Myers, T. A., McCoy, D. T., Po-Chedley, S., Caldwell, P. M., Ceppi, P., Klein, S. A., and Taylor, K. E.: Causes
 of Higher Climate Sensitivity in CMIP6 Models, Geophysical Research Letters, 47, 1–12, https://doi.org/10.1029/2019GL085782, https://onlinelibrary.wiley.com/doi/abs/10.1029/2019GL085782, 2020.