

Authors' responses to Bridget Scanlon's Interactive Comment:

We thank Bridget Scanlon for her comments and support for the general intent of the paper. We are encouraged that, apart from minor comments, the methods, results and findings of the paper are considered robust and useful. We appreciate that the appended comments are helpful and intended to improve the manuscript.

This response is also provided in the form of a pdf supplement.

We provide the following responses to specific comments:

Abstract:

'In the abstract the authors refer to ENSO as seasonal; however, I think of ENSO as more interannual with 3 – 5 yr timescales'

We agree and propose simply to omit the word 'seasonal'. The original reason for the inclusion of this adjective was the known (inter-annual) periodicity in the control that ENSO exerts on seasonal precipitation (i.e. heavy rains, drought).

L40 – L42:

'The authors refer to baseflow from groundwater sustaining rivers and wetlands being fundamentally important, especially in semiarid and arid regions; however, one should recognize that in many semiarid regions surface water recharges groundwater.'

This characterisation is misrepresentative and a regretted oversight on our part; we propose to amend the wording as follows:

"Baseflow from groundwater sustains rivers and wetlands in the absence of rainfall and is therefore fundamentally important to the ecology of semi-arid and arid regions in particular..."

to be replaced by:

"Bidirectional flows between surface water and groundwater are fundamentally important to the ecology of semi-arid and arid regions (drylands), where surface water often recharges groundwater and baseflow from groundwater can sustain rivers and wetlands in the absence of rainfall"

L45:

'The authors suggest that future management of freshwater resources will be a critical issue linked to climate change but I think it is already a critical issue because of climate extremes (droughts and floods)'

Agreed – we propose to amend this sentence:

“Future climate change in which anthropogenic emissions of greenhouse gases transform patterns of natural variability, together with substantial socio-economic change, predicate that management of freshwater resources will become a critical task ((Famiglietti, 2014).”

to read:

“Climate change in which anthropogenic emissions of greenhouse gases transform patterns of natural variability, together with substantial socio-economic change, predicates that management of freshwater resources has and will increasingly become a critical task (Famiglietti, 2014).”

L63 – 65:

‘In contrast to the statement from Wada et al., 2014; recent reports suggest that water use has been stable or decreasing in past decades in the U.S. and China. Zhou, F., et al. (2020), Deceleration of China’s human water use and its key drivers, Proceedings of the National Academy of Sciences, 117(14), 7702-7711. https://www.usgs.gov/special-topic/water-science-school/science/trends-water-useunited-states-1950-2015?qt-science_center_objects=0#qt-science_center_objects’

We appreciate this insight and helpful suggestion. Zhou et al. (2020) specifically refer to a slowing in the *rate* of the increase in water use in China (Zhou et al., 2020) whereas the USGS online report clearly highlights a decline in total freshwater withdrawals in the US. To avoid confusion, we have revised the text here to refer simply to the estimate that relates to global groundwater withdrawals so that it is amended to read:

“In the context of an ~85% increase in global groundwater abstraction from 1979 to 2010 (Wada et al., 2014), an understanding...”

L175-180:

‘GRACE uses a baseline from 2004 – 2009; however, I think it would be better to calculate anomalies for your data based on your entire record. It is not clear why you don’t combine JPL and CSR mascons? JPL relies on models to process the GRACE data whereas CSR does not. The combination should be more robust.’

We appreciate these comments and, with respect to the suggested baseline, have conducted the analysis as proposed. Our TWS anomalies are calculated based on the entire record which means they have been adjusted to a time mean baseline taken over the whole period of GRACE data. This text on lines 175 to 178 is intended to convey that methodology:

“GRACE Δ TWS is not a time-invariant measure (Wahr et al., 1998) and in the standard datasets all anomalies are given with respect to a baseline which is the mean over the period January 2004 to December 2009 (JPL NASA, 2019). Here, the completed available

GRACE Δ TWS time series is examined with respect to climate anomalies in the same timeframe. Consequently, the JPL-Mascon Δ TWS dataset has been rescaled with respect to a time-mean taken over the whole period of GRACE operation (08.2002 – 07.2016), which is the study reference period (SRP) (JPL NASA, 2019)."

To clarify we will amend this passage to:

"GRACE Δ TWS is not a time-invariant measure (Wahr et al., 1998) and in the standard datasets all anomalies are given with respect to a baseline which is the mean over the period January 2004 to December 2009 (JPL NASA, 2019). However, we examine the completed available GRACE Δ TWS time series with respect to climate anomalies over the consistent timeframe of the entire series. Consequently, the employed JPL-Mascon Δ TWS dataset has been rescaled with respect to a time-mean taken over the whole period of GRACE operation (08.2002 – 07.2016), which is the study reference period (SRP) (JPL NASA, 2019)."

On the use of CSR Mascons, we did consider this and we decided that for this particular use of the JPL Mascons, the strong correlation (rank correlation coefficients mostly >0.9) between the CSR Mascons and JPL Mascons Δ TWS time series, identified by Scanlon et al. (2016) (Scanlon et al., 2016), was sufficient to provide a robust result using JPL Mascons alone. Specifically, we relied on the following arguments:

- Since the data are adjusted with respect to the time mean baseline for the entire series, we are not concerned with long-term trends which are the major source of divergence between the JPL and CSR datasets.
- With the exception of California Central Valley, nine of the analysed aquifer systems are large ($>500 \text{ km}^2$) and 4 are medium scale ($>100 \text{ km}^2$), meaning that the correlation between JPL and CSR Mascon datasets is robust due to the scale of the aquifers studied.
- Our study is concerned with Tropical and Sub-Tropical latitudes where the resolution of JPL Mascons ($\sim 300 \text{ km}$) is equivalent to that of the raw GRACE data, minimising concerns of processing and again improving correlation between the two Mascon datasets.
- We consider that the greatest source of uncertainty in the GRACE data for this study is derived from the use of one or more LSM's to derive Δ GWS from Δ TWS, as discussed in the paper (L90-93, L188-190)

L237:

'It would be good to use a consistent time period for calculating all anomalies throughout the paper.'

We agree and this is indeed what we have done in the paper. To make this point as clear as possible, we will amend the current text:

“PCPA = monthly precipitation anomalies with respect to the time-mean baseline for the study reference period (SRP – 2002-16)”

to read:

“PCPA = monthly precipitation anomalies with the respect to the consistently applied study reference period time-mean baseline, 2002 – 2016.”

L310

‘I think it is important to indicate the uncertainties in these global datasets. For example, the water table depth map developed by Fan differs markedly from that developed by the British Geological Survey for Africa’

We agree that it is important to indicate the uncertainties in these global datasets and propose to amend the text as set out below. The divergence between the Fan and BGS datasets for Africa is primarily one of scaling in which Fan shows 5 divisions between 0 m and 25 m depth whereas the BGS shows only 2 divisions (Fan et al., 2013; MacDonald et al., 2012). Within the constraints of the BGS scaling, the two datasets apparently agree reasonably well for most of the continent. For this study, the relevant exception is the Karoo Basin where the two datasets disagree; BGS data show some shallower areas of water-table depth that do not appear in the Fan dataset. We recognise that the response of the Karoo Basin to recharge is to some extent anomalous in this study but we do not explore the reasons in any detail.

Proposed amendment:

L318 on: ‘However, K is based on permeability mapping from hydrolithologies that have a standard deviation of ~2 orders of magnitude (Gleeson et al., 2011) and this variance underlies the uncertainty in each of these datasets used. WTD is 30 arc-second (~1km) resolution dataset compiled from available observational data extended by modelled interpolation with both these data sources being subject to considerable sampling bias and model uncertainty respectively (Fan et al., 2013). All of these datasets are global and derived from combinations of observations and modelled data.’

L423:

‘Long and Mahler (2013) applied the analysis to karst aquifers, which are similar to surface water drainage systems. These systems differ markedly from many aquifer systems.’

The original and current intent of this reference (Long and Mahler, 2013) is to illustrate the use of transfer functions in hydraulic analysis and efficiencies that have been derived in the use of metrics. However, we do not wish there to be any confusion arising from this contextual use of the reference and propose to omit it and to amend L420-426:

“HM ultimately derives from the physical properties of the saturated portion of the aquifer system (Townley, 1995) and system memory as measured by Eq. (5) is

representative of the physical properties of an aquifer system and its climate. Long and Mahler (2013), for example, used a total of 16 metrics to describe particular North American Karst aquifer Systems in their IRF (θ in Eq. (5)). In contrast, von Asmuth and Knotters (2004) used 4 parameters to describe groundwater dynamics in their transfer function that they argue represents a more accurate description of the physical system than previously used parametric methods.”

to read:

“HM ultimately derives from the physical properties of the saturated portion of the aquifer system (Townley, 1995) and system memory as measured by Eq. (5) is representative of the physical properties of an aquifer system and its climate. Von Asmuth and Knotters (2004) use 4 parameters to describe groundwater dynamics in their transfer function (θ in Eq. (5)) that they argue represents a more accurate description of the physical system than previously used parametric methods (von Asmuth and Knotters, 2004).”

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