

# Authors' Response to Referee 1 (ESDD esd-2018-71)

March 6, 2019

*The large disagreement of projections of future net land-atmosphere CO<sub>2</sub> flux in Earth-system models is the biggest uncertainty in future climate projections (Arora et al., 2013; Friedlingstein et al., 2013). To tackle this issue, the application of emergent constraints (EC) to different carbon-cycle and ecosystem processes to reduce the range of the future land-sink estimates has become increasingly popular (Cox et al., 2013; Wenzel et al., 2014; Mystakidis et al., 2016; Wenzel et al., 2016). In this study Winkler et al. discuss the reasoning behind the application of EC in Earth-system modelling. They point to potential limitations, such as the need to accurately measure the predictor and to find a robust relationship between predictor and predictand, and how that might change over time. They then use the sensitivity of Leaf-Area Index (LAI) to CO<sub>2</sub> and temperature to constrain future estimates of Gross Primary Productivity in the Northern High-Latitudes. In my opinion, the theoretical examination of the EC framework, sources of uncertainty and its limitations is particularly noteworthy and useful for the community (discussion around Figures 1, 4 and 6). I find the manuscript in the present form rather strenuous to read, without a fluid structure, several repetitions and sometimes omissions and inconsistencies that generate confusion. This can easily be improved during the revision: my suggestion would be to have a complete conceptual part discussing uncertainties and complications of the EC method before moving to the analysis of LAI data. There are, however, other points of this study that I find more problematic, and that need consideration before I can recommend its publication. I first describe my general concerns, and then include more specific comments for your consideration.*

We thank the reviewer for her/his detailed and very constructive review of our manuscript. We appreciate that the reviewer finds our study particularly noteworthy and useful for the community, but we also notice the reviewer's concerns. All revisions done in response to the reviewer's comments improved the structure and overall readability of the manuscript.

## 1 General Comments

*1.1 The introduction delves into the assumptions underlying the EC, different studies using EC to constrain the carbon-cycle sensitivity to global change and their limitations and uncertainties. I find that the introduction is missing a motivation statement that explains: (i) the need for the conceptual study presented here; (ii) why did the authors focused on the relationship between LAI and GPP (more on this below); (iii) the rationale behind the choice of trying to constrain  $\Delta GPP$  in the NHL only, since models that do well at simulating the effect of boreal/temperate ecosystem CO<sub>2</sub> fluxes do not necessarily constrain better the global terrestrial sink (Schimel et al. (2015), Figure 3).*

We agree, that the introduction lacks a clear motivation why a conceptual approach to the EC method is needed. In recent years, many studies have been published applying the EC method to constrain essential entities of the Earth system. The method will become even more popular with the upcoming CMIP6 model simulations. However, the literature is missing a detailed description of the applicability and limitations of the EC method, resulting in a somewhat arbitrary application and methodological inconsistencies among various studies. To account for that, this conceptual study is needed, which elaborates on the behavior of the EC method (i); There is no specific reason why we based our conceptual study on the relationship between LAI and GPP. We adduced these variables to build a case study in order to scrutinize the EC method. In theory, the results are qualitatively transmissible to other sets of predictors and predictands (ii); We focused our analysis on the northern high latitudes (NHL) because of two reasons. First, ecosystems in NHL are barely influenced by human land use. Thus, the changes of vegetation greenness

are natural responses to the forcing rather than agricultural artifacts. Second, independent remote sensors (AVHRR, MODIS) yield comparable greening sensitivities for the NHL, although the MODIS time series is yet too short to derive a statistical robust estimate (iii). We edited the introduction in the revised manuscript to account for these comments.

*1.2 The description of Winkler et al. (P2, L25 - P3, L2) is partly (but with less detail) described in the methods. I suggest mentioning here just the relevant aspects of their study.*

In the revised manuscript, we added a description of the methodology in Winkler et al. (2019), especially, we make abundantly clear how  $\omega$  is derived and describe its characteristics for models and observations. Please see also responses to comment 1.5 and 2.26.

*1.3 However, from this paragraph, it seems that one of the main conclusions of this manuscript is also an outcome of Winkler et al. (2018) – I mean the values of  $3.4 \pm 0.2 \text{ Pg C yr}^{-1}$  which are then presented again in the results section. This leaves me wondering to which extent is this study original, compared to that in revision in Nature Communications. It's important that the authors clarify this, at least in their reply to the comments.*

Winkler et al. (2019) present constraints on projected future plant productivity in NHL using greening sensitivity as well as independent observational resources such as ground-measurements of  $\text{CO}_2$  and atmospheric  $\text{CO}_2$  inversion products. The study in hand focuses on the concept of EC, its applicability and limitations, building on the EC presented in Winkler et al. (2019). We agree that presenting the  $\Delta\text{GPP}$  constraint of  $3.4 \pm 0.2 \text{ Pg C yr}^{-1}$  as key result in both studies, is problematic. We address this issue in the revised manuscript. Please see also our response to comment 2.4.

*1.4 In the Methods section, the authors state that they "revisit the study of Winkler et al. (2018)" and "largely follow the methodology detailed in Winkler et al. (2018)". However, the reviewers (and potential readers) do not have access to this study to evaluate the methodology in detail nor to understand what exactly is being revisited. Moreover, that companion paper is not yet accepted for publication. Therefore, the authors should at least describe the methodology in more detail.*

Yes, we acknowledge that access to the companion article is needed to better comprehend the methodology in this study. The article by Winkler et al. (2019) is now published and openly available on the website of *Nature Communications* (<https://rdcu.be/bpELU>). We included a more comprehensive methods section in the revised manuscript. Please see also responses to comments 1.2 and 1.5.

*1.5 This is especially the case for the calculation of  $\omega$ , which is then used for a big part of the analysis of  $\text{LAI}_{\text{max}}$  drivers. You explain that a PCA is performed on both variables ( $\text{CO}_2$  and  $\text{GDD0}$ ) to derive a proxy time-series that summarizes the evolution of both variables. The PCA is indeed suitable for such type of analysis and is probably better than multiple linear regressions used in other studies (e.g. Zhu et al. (2016)). However, the authors give very little information about this crucial step of the analysis: is the PCA performed at pixel level, or for the large-scale aggregated values? What components do they retain from the PCA (I'm assuming only PC1 is retained)? What fraction of the variance does it explain? How does it relate to  $\text{GDD0}$  and  $\text{CO}_2$ ? How does it vary over time? Here, a plot showing  $\omega$  over time would be very helpful. Moreover, the authors should keep in mind that  $\omega$  does not "represent the overall forcing" (P9, L8-9), but only  $\text{CO}_2$  and temperature.*

We agree, that a more detailed description of the derivation of  $\omega$  needs to be provided. PCA was performed on large-scale aggregated values as well as on pixel level to investigate on spatial variations. We only retain the first principal component (denoted  $\omega$ ), which explains a large fraction of the variance, ranging approximately from 70% to 90% in models and observations (for more details see Table R1-1, also included in Supplementary Information in the companion article). Figure R1-1 depicts the temporal development of  $\text{CO}_2$  and  $\text{GDD0}$  as well as the principal component  $\omega$  for observations. This figure, with some modifications, has been included in the appendix of the revised manuscript. Yes, we acknowledge that  $\text{CO}_2$  and  $\text{GDD0}$  do not represent the overall historical forcing, but we assume that these are the main drivers causing observed changes in the NHL region. Please see also response to comment 1.2.

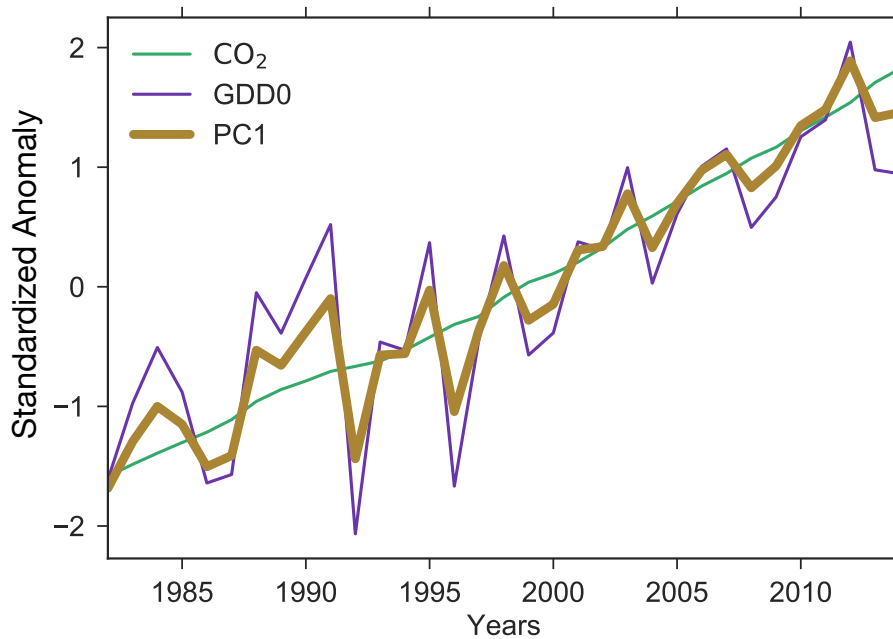


Figure R1- 1: Standardized temporal anomalies of annual averaged atmospheric CO<sub>2</sub> concentration, area-weighted averaged GDD0 for NHL, and their leading principal component  $\omega$  in observations.

Table R1- 1: Summary data for Principal Component Analysis and LAI<sub>max</sub> sensitivity estimation.

Model	Explained variance by $\omega$	LAI <sub>max</sub> sensitivity to $\omega$ , (m <sup>2</sup> m <sup>-2</sup> unit $\omega$ )	Correlation coefficient
MIROC-ESM	0.89	0.049 ± 3.3e-3	0.93
CESM1-BGC	0.83	0.014 ± 1.4e-3	0.86
GFDL-ESM2M	0.64	0.022 ± 3.2e-3	0.76
CanESM2	0.91	0.013 ± 1.0e-3	0.91
HadGEM2-ES	0.94	0.075 ± 3.5e-3	0.97
MPI-ESM-LR	0.77	0.028 ± 1.8e-3	0.94
NorESM1-ME	0.84	0.0088 ± 0.8e-3	0.88
Observations	0.9	0.045 ± 6.4e-3	0.78

1.6 The authors correctly state that one requirement of the EC method is that "a physically (or physiologically) based correlation between inter-model variations in an observable entity of the contemporary climate system (predictor) and a projected variable (predictand)" (P2, L26-27) exists. I find it, therefore, striking, that the authors do not discuss in any way why should LAI be used as a predictor of the CO<sub>2</sub> fertilization effect on GPP, and whether the linearity between the two variables in ESMs holds true for observations. Experimental CO<sub>2</sub> enrichment studies did not find a direct effect between CO<sub>2</sub> fertilization and increase in LAI (e.g. Körner et al., 2005) and LAI seems to increase non-linearly with increasing CO<sub>2</sub> (Norby et al. (2005)). Moreover, Norby et al. (2010) found strong influence of nutrient availability/limitation (not simulated in most CMIP5 ESMs) in the CO<sub>2</sub> fertilization effect on ecosystem productivity, possibly because of mycorrhizal effect (Terrer et al. (2016)). ? have also shown that under increasing CO<sub>2</sub>, allocation of carbon to leaves decreased, rather than increasing (as implicitly assumed here), which was not well simulated by DGVMs. The link between CO<sub>2</sub> fertilization, LAI and GPP is further complicated by how models simulate mortality and disturbances.

The link between LAI, GPP, and elevated CO<sub>2</sub> concentration is a complicated subject matter, as the referee thoroughly describes. In terms of *in-situ* measurements, there is no clear picture emerging. Körner et al. (2005) finds no significant coupling between elevated CO<sub>2</sub> and increased LAI in a Swiss forest site for a study period of four years. Norby et al. (2005) analyzed measurements from four different FACE experiments in the northern mid-latitudes (USA and Italy). They detect a non-linear relationship between increasing CO<sub>2</sub> and LAI. However, non-linearity is to be expected for such a sharp increase of CO<sub>2</sub> concentration (quasi-instant forcing of 174 ppm) and is not

comparable to the real-world response (annual forcing of 2-3 ppm). Please see also our response to comment 2.14 for a more detailed discussion on the assumption of linearity in the relationship between LAI and CO<sub>2</sub> for the last decades. Norby et al. (2005) also report, that their analysis suggested that at low LAI, elevated CO<sub>2</sub> was causing structural changes and substantial increase in absorbed photosynthetic active radiation, in general agreement with satellite measurements of low LAI regions, especially in NHL. De Kauwe et al. (2014) analyzed measurements from two FACE experiments located in North America (North Carolina and Tennessee, USA). They find that specific leaf area (SLA, the ratio of leaf area to leaf mass) decreased, but report a general increase in LAI as response to elevated CO<sub>2</sub>.

In general, conclusions drawn from FACE experiments, owed to their setup, are not representative for long-term observed changes on ecosystem-scale. We agree with the referee, that the current manuscript lacks an in-depth discussion on the causal link between predictor and predictand. However, this aspect is discussed in more detail in the companion paper by Winkler et al. (2019) and illustrated in Supplementary Figure 1 - *Schematic of the Emergent Constraint concept*. In the revised manuscript we discuss in more detail the causal link between predictor and predictand. Please see also our response to comment 2.41.

*1.7 I understand that the authors have a stronger background on earth-system modelling and I would not expect them to make a full case on the relationships between CO<sub>2</sub> fertilization, LAI and GPP. However, since they describe so well the need for a physical basis to the EC, they need to explain the choice of LAI as a predictor of future GPP (i.e. evidence for a mechanistic link), and whether the land-surface models composing the ESMs are able or not to correctly simulate the relevant processes for this relationship (see also Kolby Smith et al. (2016)). In the current version of the manuscript, the authors do not make a strong case for their choice, and there is limited evidence (mostly from model-based studies to the best of my knowledge) to suggest that LAI sensitivity to CO<sub>2</sub> can be a suitable predictor of future GPP. The authors could, for example, combine their analysis of LAI<sub>max</sub> sensitivity to CO<sub>2</sub> and temperature with GPP changes estimated from observation-based datasets (e.g. FLUXCOM).*

The referee makes a very good proposal in analyzing other observation-based datasets to corroborate the EC estimate. We already conducted such analysis and is part of Winkler et al. (2019). Amongst other data resources, we analyzed all available FLUXCOM datasets of upscaled eddy covariance flux measurements for NHL GPP. However, these datasets were designed not to capture long-term changes as well as interannual variability, and thus, cannot be applied for a temporal analysis (e.g. Anav et al., 2015). But one can build on the spatial information to investigate the correlation between LAI<sub>max</sub> and GPP. Using the climatologic mean of the recommended ensemble median of all FLUXCOM datasets and two independent sets of satellite observed LAI, we find a striking linear relationship for the northern high latitudes (Figure R1-2a and b). This tight linear relation between the two variables over a wide range of values suggests that changes in GPP also result in changes in LAI<sub>max</sub>. In general, model simulations and large-scale observational datasets clearly indicate that LAI sensitivity to CO<sub>2</sub> ( $\omega$  for temperature-limited ecosystems) is a suitable predictor of GPP for increasing CO<sub>2</sub> forcing in the NHL.

Further, we assess the relationship between changes in GPP and LAI exclusively using *in-situ* flux measurements, although these records are yet to short for a statistically robust analysis. We selected the longest FluxNet time series existing for the NHL, Hyytiala, Finland (61.8474° N, 24.2948° E, 1996 - 2014). We took the surrounding pixels of the long-term but rather coarse (AVHRR, 1/12°, 1982 - 2016) as well as short-term but higher resolution (MODIS, 500m, 2000 - 2016) satellite observations of LAI. We find contemporary trends in GPP and LAI, but the linear relation between the *in-situ* measured GPP and long-term AVHRR satellite datasets is rather weak due to the coarse resolution. Thus, to match the flux tower footprint, we have to make recourse to high resolution satellite observations of MODIS. MODIS LAI and AVHRR LAI (both analyzed in our study) have strong correlation and the latest AVHRR LAI datasets were developed by referencing to MODIS LAI (Zhu et al., 2013). For the MODIS time-series we find a much stronger relationship to the flux measurements and therefore confirms the tight connection between changes in GPP and LAI for the site in Hyytiala, Finland (Figure R1-2c). However, the overlapping period

of MODIS and FluxNet is yet too short to derive reliable estimates. [Anav et al. \(2015\)](#) also analyzed other eddy covariance flux measurement sites and find a general agreement on increasing GPP.

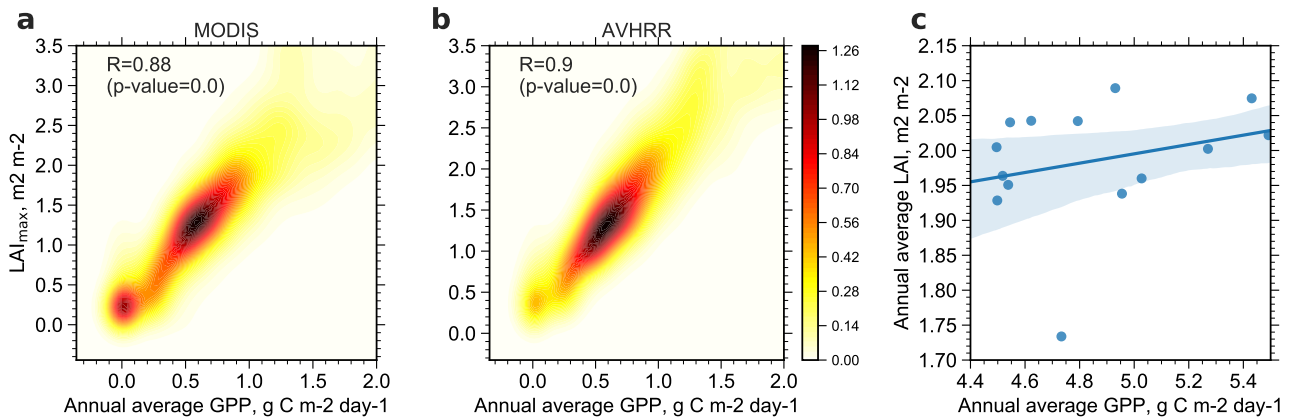


Figure R1- 2: Strong correlation in the climatologic mean in observational datasets between  $LAI_{max}$  derived from two independent satellite sensors, MODIS (a) and AVHRR (b), and the ensemble median annual average GPP from the FLUXCOM ensemble for the northern high latitudes. Color density indicates the probability distribution estimated using Gaussian kernel. c, Contemporary trends in the longest *in-situ* GPP flux measurement record in the NHL and the study site surrounding pixels of high resolution LAI satellite observations. The blue line shows the best linear fit and the shading shows the 95% confidence interval.

## 2 Specific comments:

### 2.1 P1, L 2: "promising results" of what?

This sentence has been rewritten to be more specific.

### 2.2 P1, L3: What do you mean by "difficult to measure variable [...] at a potential future"? If you are trying to estimate a future state of a variable, it is by definition non-measurable?

The statement 'difficult-to-measure' only refers to 'variable' and not to 'a potential future'. We rewrote this paragraph to avoid misunderstandings. Please see also response to comment 2.15

### 2.3 P1, L7: "greening sensitivity to the CO<sub>2</sub> forcing" ... but also temperature, right? (Methods).

We investigate both types of sensitivity, so, the greening response to rising CO<sub>2</sub> as well as to the combined signal of rising CO<sub>2</sub> and temperature (GDD0). The latter approach is necessary for temperature-limited ecosystems and is only applied in the analysis focusing on NHL.

### 2.4 P1, L18: Is the value of the GPP enhancement from this study or from Winkler et al. (in revision)?

This result is a subject in both studies, however, we discuss it with different perspectives. In the revised manuscript, we define the focus of this study more precisely to avoid such misunderstandings. Please see also response to comment 1.3.

### 2.5 P2, L4: "can have substantial uncertainties" ? remove can. They have.

We agree. The sentence was modified.

### 2.6 P2, L8: I'd move the "large-scale climate modes" to the paragraph about natural variability a few lines below

The sentence should give an overview of the range of aspects which are underrepresented in ESMs, from local short time-scale extreme events to long-term large-scale climatic modes. Hence, we prefer not to modify this section.

2.7 P2, L12: *"aims is to explore" ? "aims to explore"*

We corrected the sentence.

2.8 P2, L21: *"namely, AS a method. . ."*

We agree, the sentence reads better now.

2.9 P2, L24: *In theory, could another relationship (non-linear) be used?*

Yes, in theory, a non-linear relationship between predictor and predictand in an multi-model ensemble is conceivable, but this requires a reasonable process-based justification. For instance, there are attempts to establish an EC using an exponential relationship between historical warming (predictor) and equilibrated temperature increase for a doubling of atmospheric CO<sub>2</sub> (predictand, equilibrium climate sensitivity, ECS). This approach implies that models with strong historical warming should predict a disproportional high ECS. To build a reliable EC, one has to identify the process causing this disproportionality in the model ensemble.

2.10 P2, L27: *what do you mean by difficult to observe? Cox et al. (2013) used two variables that are relatively well observed (CO<sub>2</sub> growth rate and tropical temperature).*

Cox et al. (2013) used variations in the observables CO<sub>2</sub> growth rate and tropical temperature to constrain land carbon storage in the tropics, the latter being the variable that is difficult to observe.

2.11 P2, L32: *What do you mean by "confirmed"?*

The relationship between snow-albedo feedback strength of the current seasonal cycle and projected feedback to long-term warming has been detected in the CMIP3 ensemble and also exists in the CMIP5 ensemble. If an "Emergent Relationship" is independent of the analyzed model ensemble, it is considered as 'confirmed' in the EC literature.

2.12 P3, L17: *"2xCO<sub>2</sub> world": you mean in model simulations, not in CO<sub>2</sub> enrichment experiments, right?*

Yes, '2xCO<sub>2</sub> world' refers to model simulations. We rewrote this section to be more specific.

2.13 P6, L2: *Here you mention that you also use precipitation to derive  $\omega$ , however later you mention only CO<sub>2</sub> and GDD0 were used. If you don't use, can you justify the exclusion of precipitation (non-significant trends? Non-significant effects?)*

Precipitation and temperature are used to derive climatic regimes (see Figure 2 in the manuscript). For each climatic regime, we derive the greening sensitivity to CO<sub>2</sub>. First, only the sensitivities to rising atmospheric CO<sub>2</sub> concentration are calculated to obtain comparability between the different climatic regimes, vegetation classes, and latitudinal bands (see Figure 2 in the manuscript). Then, we focus on the northern high latitudes, where we also have to take temperature into account and derive the greening sensitivity to  $\omega$ , the combined signal of CO<sub>2</sub> and GDD0. We rewrote this passage to be more clear on that matter.

2.14 P7, L4-5: *can you provide any lines of evidence to justify the assumption (non-model based).*

The increase of observed CO<sub>2</sub> concentration (annual average) throughout the satellite era can be considered as quasi-linear (Figure R1-3). Our analyses of remote sensing datasets of LAI from different sources (AVHRR, MODIS) also suggest linearly increasing trends. This finding was also reported by several preceding studies (Zhu et al., 2016; Mao et al., 2016; Forkel et al., 2016; Mahowald et al., 2016). Forkel et al. (2016, Fig. 1 - *Amplification of plant activity in the northern biosphere*) analyzed several observational datasets for the northern ecosystems of the last 30 to 40 years and report evidence for a linearly changing system. The bottom line is, there is no observational indication of a non-linear relationship between LAI and CO<sub>2</sub>, at least not for the CO<sub>2</sub> forcing of the last decades (from ~340 to ~410 ppm.)

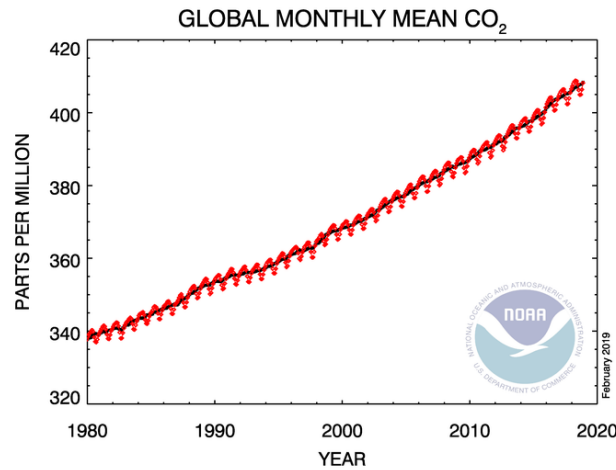


Figure R1- 3: A quasi-linear increase in observations of global monthly mean CO<sub>2</sub> concentrations since 1980; image taken from [https://www.esrl.noaa.gov/gmd/ccgg/trends/gl\\_full.html](https://www.esrl.noaa.gov/gmd/ccgg/trends/gl_full.html), February 14, 2019.

*2.15 P8, L4-5: What do you mean by "difficult to measure"? It's already repeated 2 times before.*

The concept of Emergent Constraints is to constrain an entity (predictand) of the Earth system that is not-directly or not-at-all observable (e.g. at a potential future state). This can be achieved by using an observable that is physically connected to the predictand. We understand the confusion about the term 'difficult-to-measure' with regard to projected estimates of GPP. To be more clear on that matter, we modified the terminology in the revised manuscript. Please see also response to comment 2.2.

*2.16 P8, L6-9: What evidence do you provide for this? CO<sub>2</sub> enrichment experiments contradict this assumption.*

Please see responses to comments 1.6 and 2.41.

*2.17 P8, L15: "large area" ? "large-scale"?*

We agree.

*2.18 P8, L16-32: This is somewhat confusing since up until now you mention that you will analyse NHL. Please reformulate before in other to make clear that first you look at global values, and then focus on NHL (and provide justification to do so).*

First, we present the observable on global scale aggregated for different climatic regimes, vegetation types, and latitudinal bands. Then, we show that LAI is only a meaningful predictor for changes in GPP in the northern high latitudes. We rewrote this section and provide better justification for our approach.

*2.19 P8, L19-21: How much does GDD0 contribute to  $\omega$  in the tropics? Can the low sensitivities in the tropics be due to your choice of temperature variable? I do not expect GDD0 to be a relevant temperature variable in the tropical band. . .*

Yes, we agree, GDD0 is not a relevant temperature variable for the tropical regions. We only consider GDD0 in the NHL, as part of  $\omega$ . When deriving sensitivities for global comparison (e.g. comparing tropical, mid-latitude, and high-latitude sensitivities; Figure 2 in the manuscript), we only consider the signal of rising CO<sub>2</sub> concentration and neglect temperature. Thereafter, we focus on the NHL, because there we obtain a clear LAI signal, e.g. not being distorted by human land use. We derive LAI sensitivity to  $\omega$ , so, also accounting for temperature and its variations, an important aspect for temperature-limited ecosystems. Also LAI in the tropics is quite sensitive to temperature variations, particularly to anomalies associated to ENSO. Thus, for a study focused on the tropics one should also consider temperature in estimating LAI sensitivities.

2.20 P9, L2-3: *Indeed, but perhaps this is because of your inadequate choice of predictor for temperature (GDD0, rather than annual T, or some other metric)?*

No, this is not the case. Please see response to comment 2.19. We only take temperature into account when focusing on NHL.

2.21 P9, L8-9: *not the overall forcing, just two components of the forcing (CO<sub>2</sub> and temperature). Please show the time-series of  $\omega$ .*

Please see our response to comment 1.5, Table R1-1 and the time-series of  $\omega$  in Figure R1-1.

2.22 P9, L17: *"all pixels": of the globe, or just NHL?*

In the manuscript, we state "We focus further analyses on the NHL region [...]" (P9, L2). Hence, we only show global comparison of LAI sensitivities to CO<sub>2</sub> in section 3.1 and, thereafter, we concentrate on LAI sensitivity to  $\omega$  in NHL. We modified this section to be more precise.

2.23 P9, L26-29: *Where do you show the corresponding increase in plant productivity? Where can I see that the distribution is approximately the same for the two variables? And if you have this data, where do you get GPP from, models or observations? Can you plot the GPP distribution for the same choice of pixels?*

We use CMIP5 model output to show that the distribution of pixels with significant changes of the predictor (LAI sensitivity to  $\omega$ , historical simulation) and the predictand (GPP, 1pctCO<sub>2</sub>) are approximately the same. Figure R1-4 compares respective distributions for all CMIP5 models. All models, except HadGEM2-ES, confirm that the pixels that show significant historical LAI sensitivity to  $\omega$  are approximately also the pixels showing significant changes in GPP for 2×CO<sub>2</sub>, resulting in similar distributions. Note, that the variables LAI and GPP had to be normalized for comparison in the same figure. This analysis is corroborating our statement in the manuscript, that averaging the equally distributed estimates does not affect the predictor-predictand relationship in the model ensemble (P9, L29-30). Also, the results shown in Anav et al. (2013, 2015) indicate spatial correlation of increasing GPP and LAI.

Long-term and large-scale changes in GPP still cannot be obtained from observations. Upscaled FluxNet measurements (i.e. FLUXCOM datasets) also rely on statistical models (e.g. neuronal networks) and are designed not to capture long-term changes (e.g. Anav et al., 2015). Thus, these datasets can only be applied for certain types of analyses, e.g. spatial patterns or natural variability. Please see also our response to comment 1.7. For completeness, we include a modified version of Figure R1-4 in the appendix of the manuscript.

2.24 P10, L3-8: *Is this also valid for ESM outputs?*

Yes, this statement is also valid for ESM simulation output. We discuss this aspect in the manuscript (P10, L27-30).

2.25 P10, L19: *What do you mean by "LAI<sub>max</sub> sensitivity cannot be accurately estimated irrespective of the window length"?*

This statement refers to Figure 4 in the manuscript. Figure 4 shows LAI<sub>max</sub> sensitivity to  $\omega$  for the historical period from 1860 to 2005 for different moving window lengths (15yr, 30yr, and 40yr). In the decades around the turn of the 20th century, LAI<sub>max</sub> sensitivity to  $\omega$  is fluctuating from negative to positive numbers for all window lengths. This is, because CO<sub>2</sub> forcing is low, and thus, natural variability dominates. Under these circumstances, LAI<sub>max</sub> sensitivity to  $\omega$  cannot be accurately estimated.

2.26 P10, L20-21: *Do you mean the signal to noise ratio of  $\omega$ ? Unfortunately you don't show the time-series, so it's hard to follow.*

When CO<sub>2</sub> forcing is low, natural variability (*noise*) is dominating and influencing the estimation of LAI<sub>max</sub> sensitivity to  $\omega$ . But, when CO<sub>2</sub> forcing grows stronger, the LAI response (*signal*) is exceeding the noise and LAI<sub>max</sub> sensitivity to the forcing can be estimated. Please see also our responses to comment 1.5 and 2.25., Table R1-1 and the time-series of  $\omega$  in Figure R1-1.



2.27 P10, L23-26: *But, in theory, that's the aim of the EC method. Do you mean that before considering using a given EC, one should evaluate the stability of the sensitivities?*

This section is refereeing to the comparability of sensitivities in window length and location between observations and models. In other words, the observed and modeled predictors have to be obtained from the same point in time (level of CO<sub>2</sub> forcing) and comparable temporal window lengths, so, all predictors have to be representative for the same state of the system. Yes, besides evaluating the stability of the predictor, one has to evaluate the comparability of predictors. So, the aim of the EC method is to use these predictors to constrain an entity of interest (predictand) at a different state (forcing) of the system.

2.28 P10, L29-30: *It's not really shown in Figure 4.*

We argue, that Figure 4 in the manuscript clearly shows that LAI<sub>max</sub> sensitivity estimation becomes more stable with strengthening forcing and increasing window length. Please compare different colored lines representing different window lengths and variability for different points in time, i.e. CO<sub>2</sub> concentration.

2.29 P11, 4-6: *Very good way to pose the question. But can you answer this in a pure model world? I'm not fully convinced.*

In this section, we show that the EC method can be applied also when the underlying relationship between predictor and predictand is changing with increasing forcing (e.g. from linear to non-linear). Predictions of future GPP are based on our current understanding of the system. We expect that saturation will occur with increasing CO<sub>2</sub>. In spite of this non-linear response, we illustrate that the EC relationship in the model ensemble can remain linear. From observations only, we cannot obtain ecosystem-scale estimates of GPP increase for a high CO<sub>2</sub>-world. So, yes, we can and must answer this question in a pure model world.

2.30 P11, L8-10: *Before (and after) you always use 7 models. It's not clear which model set is being used for which analysis. Are you using only 3 models to constrain future GPP changes? This does not seem consistent with Figure 6.*

Yes, we agree, this is confusing. In general, we use as many models as possible for the EC analysis (here, 7 models). In Figure 4 and 5 in the manuscript, we only show 3 of the 7 models, because all models show qualitatively the same. We selected these three models, because they span the full range of GPP predictions (CESM1-BGC: lowest estimate, HadGEM2-ES: highest estimate, and MIROC-ESM: closest to EC estimate). We generated two additional figures (shown in the appendix of the revised manuscript) which display the results of the other 4 models analogous to Figure 4 and 5 in the manuscript.

2.31 P11, L11-18: *Not that surprising since all models are based in some way or another in the Farquhar photosynthesis model, which for the ppm ranges of 1xCO<sub>2</sub> and 2xCO<sub>2</sub> can possibly be approximated by a linear function, and in DGVMs the allocations schemes to leaves are strongly coupled to GPP (e.g. models don't simulate well non-structural carbon reserves, or changes in allocation)? Also, if models prescribe fixed LAI<sub>max</sub> (as some do), then this will strongly depend on the chosen model parametrization.*

Yes, we agree, it is not surprising that all models show saturation at higher CO<sub>2</sub> levels. However, here we make the point, that despite the expected non-linearity of the predictor-predictand relationship at higher CO<sub>2</sub> levels, the inter-model relationship in the ensemble space can remain linear. This is a somewhat counter-intuitive aspect of the EC method and essential for its interpretation.

2.32 P11, L18: *Why not call it simply "thought experiments" or "conceptual experiments", for non-german readers?*

*Gedankenexperiment* is an universal scientific term such as the German word *Ansatz*.

2.33 P11, L21: *What do you mean by LAI? Annual values? Growing-season average? And why not LAI<sub>max</sub>?*

We intended to simplify the terminology in the *Gedankenexperiment*. Since the time dimension does not play a role in this conceptual framework, LAI expressed as annual average, growing season average, or annual maximum has no meaning. However, we acknowledge that the changed terminology can cause some confusion. Therefore, we stick to LAI<sub>max</sub> in the revised manuscript.

2.34 P11, L24: ". . . responses" ? add something like "of GPP to CO<sub>2</sub> and of LAI to GPP" for clarity.

We rewrote the sentence for clarity.

2.35 P11, L26 - P12, L5: *Why did you choose Scenario 3? Scenario 4 in Figure A2 is much more plausible (GPP saturating for high levels of CO<sub>2</sub> because of basic physiology (Farquhar)).*

We chose Scenario 3 to highlight the interplay of linear and non-linear relationships between forcing, predictor, and predictand. But we agree with the referee that Scenario 4 is the most plausible, which we also discuss in the manuscript (P12, L9-11).

2.36 P12, L9-10: *"timing of saturation": where can we see this?*

Figure 4 and Table 2 in the manuscript illustrate that the CMIP5 models show saturation of the relationship between  $\Delta\text{LAI}_{\text{max}}$  and  $\Delta\text{GPP}$  with increasing CO<sub>2</sub> forcing. The slopes in Figure 4 (detailed estimates in Table 2) reveal that the strength and 'timing' of saturation (i.e. at what level of CO<sub>2</sub> concentration) differs among the models. In the revised manuscript, we implemented more accurate description and references to tables and figures.

2.37 P12, L20: *what LAI<sub>max</sub> are you referring to here? I assume you used AVHRR, since you explained (well) why MODIS is not suitable. But you need to clarify.*

Yes, we used AVHRR data. We added more details to this section in the manuscript.

2.38 P12, L24-26: *in the model world. You need to discuss whether observations support this.*

Please see response to general comment 1.6.

2.39 P12, L26: *I assume you mean "LAI<sub>max</sub> sensitivities" to  $\omega$ . Is this simulated  $\omega$  or  $\omega$  from observations? Over which period? If it is simulated  $\omega$  you need to show how  $\omega$  from historical simulations compares with  $\omega$  from observations.*

LAI<sub>max</sub> sensitivity to  $\omega$  is calculated for observations and each model separately for approximately the same time period. Please see Table R1-1 for more details. This approach enables an accurate comparison between the simulated and observed predictor variables. Also, more details on  $\omega$  can be looked up in the supplementary information for Winkler et al. (2019).

2.40 P13, L14: *do models simulate compositional changes in these simulations? I.e. do they all include dynamic vegetation changes?*

Yes, most of the models include dynamic vegetation. In the revised manuscript, we include a short description of the representation of dynamic vegetation in CMIP5 models. In general, the historical and idealized model setups of the CMIP5 land components are comprehensively explained in several studies, such as Wenzel et al. (2014); Mahowald et al. (2016); Arora et al. (2013); Winkler et al. (2019). This is why we refrain from providing a detailed overview of the CMIP5 models in this study.

2.41 P13, L34: *But observations seem to point out that climate change (warming and drying) probably cancels out the CO<sub>2</sub> fertilization effect (Peñuelas et al., 2017), because of processes not well simulated by CMIP5 models - climate extremes, particularly heatwaves, mortality, disturbance and further reinforced by nutrient limitations (also not simulated by most CMIP5 models).*

The referee addresses one of the key problems in current climate and carbon cycle research. On the one hand, we expect that CO<sub>2</sub> fertilization is causing enhanced plant growth based on our physiological understanding. Many studies find evidence for this expectation. Especially, the Global Carbon Project suggests that ~30% of the anthropogenic CO<sub>2</sub> emissions are taken up the terrestrial biosphere (so, current land sink is ~ 12 Pg C yr<sup>-1</sup>, [Quéré et al., 2018](#)). On the other hand, observations (esp. on local scale) suggest that the net carbon uptake by plants for a higher CO<sub>2</sub> world is not changing due compensating effects (e.g. [Peñuelas et al., 2017](#)). Obviously, there is a paradox in place: Where do the 12 Pg carbon go every year, if plants do not take up more carbon with rising CO<sub>2</sub>? Future research needs to address this issue in more depth. For the NHL, however, we find robust observational evidence ([Keeling et al., 1996](#); [Myneni et al., 1997](#); [Graven et al., 2013](#); [Forkel et al., 2016](#); [Winkler et al., 2019](#)) that carbon uptake by plants is increasing, which is the baseline for the study in hand.

*2.42 P14, L8-9: Can you provide references for this?*

We assume the comment is referring to P15. There are several studies indicating that greening in the high northern latitudes is caused by indirect drivers associated to increasing CO<sub>2</sub>, such as warming and CO<sub>2</sub> fertilization ([Myneni et al., 1997](#); [Forkel et al., 2016](#); [Zhu et al., 2016](#)). At high LAI regions, GPP might also increase due to CO<sub>2</sub> fertilization without an enhancement of LAI. In rural areas, the observed greening is mainly caused by direct drivers such irrigation, application of fertilizers, and double cropping as shown recently by [Chen et al. \(2019\)](#). We added the references in the manuscript.

*2.43 P14, L2: is this an original result from this manuscript or from Winkler et al. in revision?*

We assume the comment is referring to P16. As we explained in response to comment 2.4, this result is subject in both studies, but discussed with different perspectives.

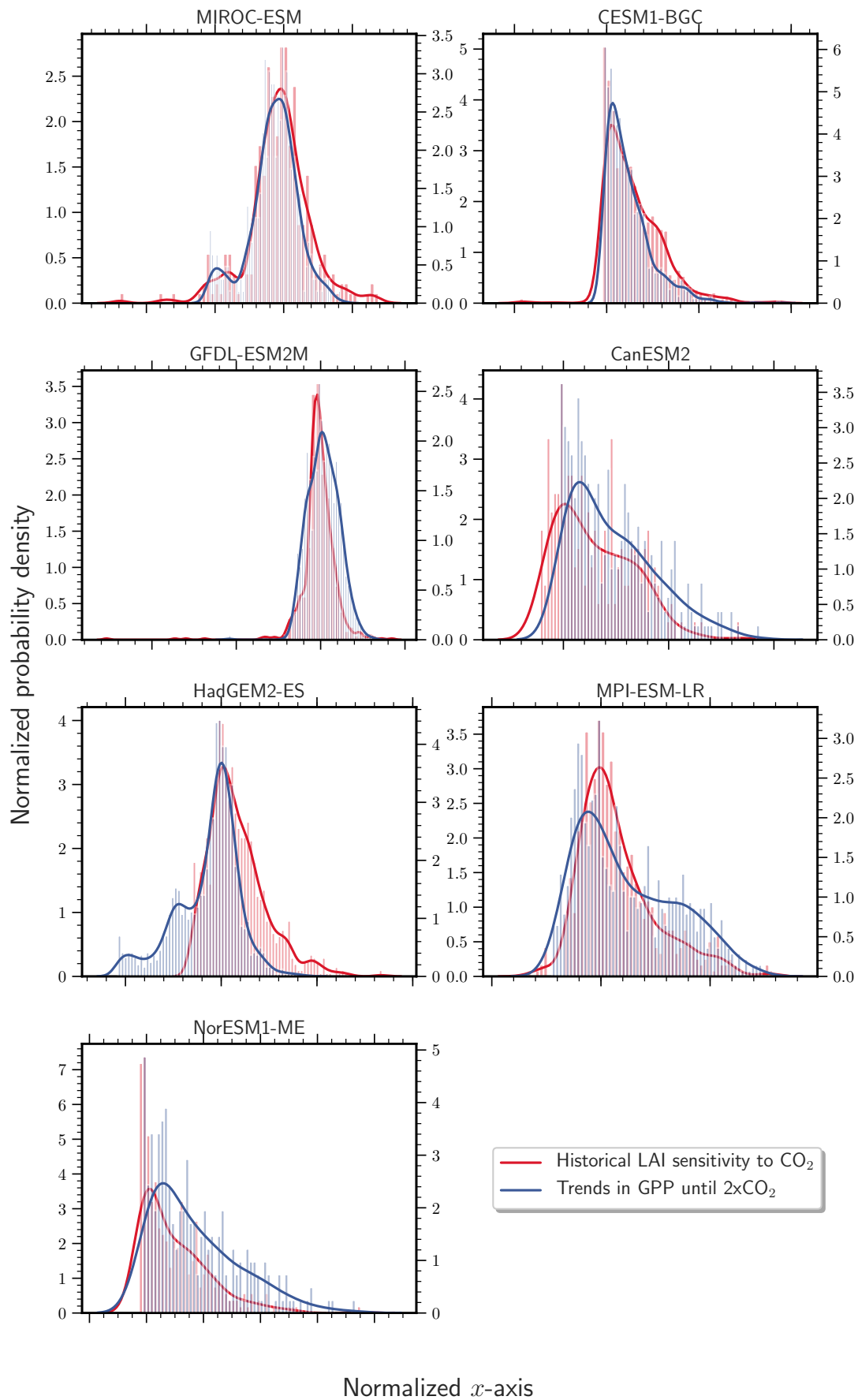


Figure R1- 4: Similar pixel distributions of LAI sensitivity to  $\omega$  (red, historical simulations) and temporal trends in GPP (blue, 1pctCO<sub>2</sub>, until 2xCO<sub>2</sub>) for NHL. All CMIP5 models are shown. Only significant pixels are included (Mann-Kendall test,  $p < 0.1$ ). To obtain comparability, the  $x$ -axis was normalized and has only qualitative meaning.

## References

- Anav, A., Friedlingstein, P., Beer, C., Ciais, P., Harper, A., Jones, C., Murray-Tortarolo, G., Papale, D., Parazoo, N. C., Peylin, P., Piao, S., Sitch, S., Viovy, N., Wiltshire, A., and Zhao, M. (2015). Spatiotemporal patterns of terrestrial gross primary production: A review. *Reviews of Geophysics*, 53(3):2015RG000483.
- Anav, A., Friedlingstein, P., Kidston, M., Bopp, L., Ciais, P., Cox, P., Jones, C., Jung, M., Myneni, R., and Zhu, Z. (2013). Evaluating the Land and Ocean Components of the Global Carbon Cycle in the CMIP5 Earth System Models. *Journal of Climate*, 26:6801–6843.
- Arora, V. K., Boer, G. J., Friedlingstein, P., Eby, M., Jones, C. D., Christian, J. R., Bonan, G., Bopp, L., Brovkin, V., Cadule, P., Hajima, T., Ilyina, T., Lindsay, K., Tjiputra, J. F., and Wu, T. (2013). Carbon–Concentration and Carbon–Climate Feedbacks in CMIP5 Earth System Models. *Journal of Climate*, 26:5289–5314.
- Chen, C., Park, T., Wang, X., Piao, S., Xu, B., Chaturvedi, R. K., Fuchs, R., Brovkin, V., Ciais, P., Fensholt, R., Tømmervik, H., Bala, G., Zhu, Z., Nemani, R. R., and Myneni, R. B. (2019). China and India lead in greening of the world through land-use management. *Nature Sustainability*, 2(2):122.
- Cox, P. M., Pearson, D., Booth, B. B., Friedlingstein, P., Huntingford, C., Jones, C. D., and Luke, C. M. (2013). Sensitivity of tropical carbon to climate change constrained by carbon dioxide variability. *Nature*, 494(7437):341–344.
- De Kauwe, M. G., Medlyn, B. E., Zaehle, S., Walker, A. P., Dietze, M. C., Wang, Y.-P., Luo, Y., Jain, A. K., El-Masri, B., Hickler, T., Wårlind, D., Weng, E., Parton, W. J., Thornton, P. E., Wang, S., Prentice, I. C., Asao, S., Smith, B., McCarthy, H. R., Iversen, C. M., Hanson, P. J., Warren, J. M., Oren, R., and Norby, R. J. (2014). Where does the carbon go? A model–data intercomparison of vegetation carbon allocation and turnover processes at two temperate forest free-air CO<sub>2</sub> enrichment sites. *New Phytologist*, 203(3):883–899.
- Forkel, M., Carvalhais, N., Rödenbeck, C., Keeling, R., Heimann, M., Thonicke, K., Zaehle, S., and Reichstein, M. (2016). Enhanced seasonal CO<sub>2</sub> exchange caused by amplified plant productivity in northern ecosystems. *Science*, 351(6274):696–699.
- Friedlingstein, P., Meinshausen, M., Arora, V. K., Jones, C. D., Anav, A., Liddicoat, S. K., and Knutti, R. (2013). Uncertainties in CMIP5 Climate Projections due to Carbon Cycle Feedbacks. *Journal of Climate*, 27:511–526.
- Graven, H. D., Keeling, R. F., Piper, S. C., Patra, P. K., Stephens, B. B., Wofsy, S. C., Welp, L. R., Sweeney, C., Tans, P. P., Kelley, J. J., Daube, B. C., Kort, E. A., Santoni, G. W., and Bent, J. D. (2013). Enhanced Seasonal Exchange of CO<sub>2</sub> by Northern Ecosystems Since 1960. *Science*, 341(6150):1085–1089.
- Keeling, C. D., Chin, J. F. S., and Whorf, T. P. (1996). Increased activity of northern vegetation inferred from atmospheric CO<sub>2</sub> measurements. *Nature; London*, 382(6587):146–149.
- Kolby Smith, W., Reed, S. C., Cleveland, C. C., Ballantyne, A. P., Anderegg, W. R. L., Wieder, W. R., Liu, Y. Y., and Running, S. W. (2016). Large divergence of satellite and Earth system model estimates of global terrestrial CO<sub>2</sub> fertilization. *Nature Climate Change*, 6(3):306–310.
- Körner, C., Asshoff, R., Bignucolo, O., Hättenschwiler, S., Keel, S. G., Peláez-Riedl, S., Pepin, S., Siegwolf, R. T. W., and Zotz, G. (2005). Carbon Flux and Growth in Mature Deciduous Forest Trees Exposed to Elevated CO<sub>2</sub>. *Science*, 309(5739):1360–1362.
- Mahowald, N., Lo, F., Zheng, Y., Harrison, L., Funk, C., Lombardozzi, D., and Goodale, C. (2016). Projections of leaf area index in earth system models. *Earth Syst. Dynam.*, 7:211–229.
- Mao, J., Ribes, A., Yan, B., Shi, X., Thornton, P. E., Séférian, R., Ciais, P., Myneni, R. B., Douville, H., Piao, S., Zhu, Z., Dickinson, R. E., Dai, Y., Ricciuto, D. M., Jin, M., Hoffman, F. M., Wang, B., Huang, M., and Lian, X. (2016). Human-induced greening of the northern extratropical land surface. *Nature Climate Change*, 6(10):959–963.

- Myneni, R., Keeling, C. D., Tucker, C. J., Asrar, G., and Nemani, R. R. (1997). Increased plant growth in the northern high latitudes from 1981 to 1991. *Nature*, 386:698–702.
- Mystakidis, S., Davin, E. L., Gruber, N., and Seneviratne, S. I. (2016). Constraining future terrestrial carbon cycle projections using observation-based water and carbon flux estimates. *Global Change Biology*, 22(6):2198–2215.
- Norby, R. J., DeLucia, E. H., Gielen, B., Calfapietra, C., Giardina, C. P., King, J. S., Ledford, J., McCarthy, H. R., Moore, D. J. P., Ceulemans, R., Angelis, P. D., Finzi, A. C., Karnosky, D. F., Kubiske, M. E., Lukac, M., Pregitzer, K. S., Scarascia-Mugnozza, G. E., Schlesinger, W. H., and Oren, R. (2005). Forest response to elevated CO<sub>2</sub> is conserved across a broad range of productivity. *Proceedings of the National Academy of Sciences*, 102(50):18052–18056.
- Peñuelas, J., Ciais, P., Canadell, J. G., Janssens, I. A., Fernández-Martínez, M., Carnicer, J., Obersteiner, M., Piao, S., Vautard, R., and Sardans, J. (2017). Shifting from a fertilization-dominated to a warming-dominated period. *Nature Ecology & Evolution*, 1(10):1438.
- Quéré, C. L., Andrew, R. M., Friedlingstein, P., Sitch, S., Pongratz, J., Manning, A. C., Korsbakken, J. I., Peters, G. P., Canadell, J. G., Jackson, R. B., Boden, T. A., Tans, P. P., Andrews, O. D., Arora, V. K., Bakker, D. C. E., Barbero, L., Becker, M., Betts, R. A., Bopp, L., Chevallier, F., Chini, L. P., Ciais, P., Cosca, C. E., Cross, J., Currie, K., Gasser, T., Harris, I., Hauck, J., Haverd, V., Houghton, R. A., Hunt, C. W., Hurtt, G., Ilyina, T., Jain, A. K., Kato, E., Kautz, M., Keeling, R. F., Klein Goldewijk, K., Körtzinger, A., Landschützer, P., Lefèvre, N., Lenton, A., Lienert, S., Lima, I., Lombardozzi, D., Metzl, N., Millero, F., Monteiro, P. M. S., Munro, D. R., Nabel, J. E. M. S., Nakaoka, S.-i., Nojiri, Y., Padin, X. A., Peregón, A., Pfeil, B., Pierrot, D., Poulter, B., Rehder, G., Reimer, J., Rödenbeck, C., Schwinger, J., Séférian, R., Skjelvan, I., Stocker, B. D., Tian, H., Tilbrook, B., Tubiello, F. N., Laan-Luijkx, I. T. v. d., Werf, G. R. v. d., Heuven, S. v., Viovy, N., Vuichard, N., Walker, A. P., Watson, A. J., Wiltshire, A. J., Zaehle, S., and Zhu, D. (2018). Global Carbon Budget 2017. *Earth System Science Data*, 10(1):405–448.
- Schimel, D., Stephens, B. B., and Fisher, J. B. (2015). Effect of increasing CO<sub>2</sub> on the terrestrial carbon cycle. *Proceedings of the National Academy of Sciences*, 112(2):436–441.
- Terrer, C., Vicca, S., Hungate, B. A., Phillips, R. P., and Prentice, I. C. (2016). Mycorrhizal association as a primary control of the CO<sub>2</sub> fertilization effect. *Science*, 353(6294):72–74.
- Wenzel, S., Cox, P. M., Eyring, V., and Friedlingstein, P. (2014). Emergent constraints on climate-carbon cycle feedbacks in the CMIP5 Earth system models. *Journal of Geophysical Research: Biogeosciences*, 119(5):794–807.
- Wenzel, S., Cox, P. M., Eyring, V., and Friedlingstein, P. (2016). Projected land photosynthesis constrained by changes in the seasonal cycle of atmospheric CO<sub>2</sub>. *Nature*, 538(7626):499–501.
- Winkler, A. J., Myneni, R. B., Alexandrov, G. A., and Brovkin, V. (2019). Earth system models underestimate carbon fixation by plants in the high latitudes. *Nature Communications*, 10(1):885.
- Zhu, Z., Bi, J., Pan, Y., Ganguly, S., Anav, A., Xu, L., Samanta, A., Piao, S., Nemani, R. R., and Myneni, R. B. (2013). Global Data Sets of Vegetation Leaf Area Index (LAI)<sub>3g</sub> and Fraction of Photosynthetically Active Radiation (FPAR)<sub>3g</sub> Derived from Global Inventory Modeling and Mapping Studies (GIMMS) Normalized Difference Vegetation Index (NDVI<sub>3g</sub>) for the Period 1981 to 2011. *Remote Sensing*, 5(2):927–948.
- Zhu, Z., Piao, S., Myneni, R. B., Huang, M., Zeng, Z., Canadell, J. G., Ciais, P., Sitch, S., Friedlingstein, P., Arneeth, A., Cao, C., Cheng, L., Kato, E., Koven, C., Li, Y., Lian, X., Liu, Y., Liu, R., Mao, J., Pan, Y., Peng, S., Peñuelas, J., Poulter, B., Pugh, T. A. M., Stocker, B. D., Viovy, N., Wang, X., Wang, Y., Xiao, Z., Yang, H., Zaehle, S., and Zeng, N. (2016). Greening of the Earth and its drivers. *Nature Climate Change*, 6(8):791–795.