Two drastically different climate states on an Earth-like terra-planet

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Abstract. We study an Earth-like terra-planet with an overland recycling mechanism bringing fresh water back from higher latitudes to the lower latitudes. By performing model simulations for such a planet we find two drastically different climate states for the same set of boundary conditions and parameter values: A Cold and Wet (CW) state (present-day Earth-like climate) with

10 dominant low-latitude precipitation and, a Hot and Dry (HD) state with only high-latitude precipitation. We notice that for perpetual equinox conditions, both climate states are stable below a certain threshold value of background soil albedo while above the threshold only the CW state is stable. Starting from the HD state and increasing background soil albedo above the threshold causes an abrupt shift from the HD state to the CW state resulting in a sudden cooling of about 35°C globally which is of the order of the temperature difference between the present-day and the Snowball Earth state. In contrast to the Snowball

15 Earth instability, we find that the sudden cooling in our study is driven by the cloud albedo feedback rather than the snow-albedo feedback. Also, when albedo in the CW state is reduced back to zero the terra-planet does not display a closed hysteresis. This is due to the high cloud cover in the CW state hiding the surface from solar irradiation. As a result, this reduction of background surface albedo has only a minor effect on the top of the atmosphere radiation balance, thereby making it impossible to heat the planet sufficiently strongly to switch back to the HD state. Additional simulations point to a similar abrupt transition from HD

20 state to the CW state for non-zero obliquity which is the only stable state in this configuration. Our study also has implications for the habitability of Earth-like terra-planets. At the inner edge of the habitable zone, the higher cloud cover in the CW state cools the planet and may prevent the onset of a runaway greenhouse state. At the outer edge, the resupply of water at lower latitudes stabilizes the greenhouse effect and keeps the planet in the HD state and may prevent water from getting trapped at higher latitudes in frozen form. Overall, the existence of bi-stability in the presence of an overland recycling mechanism hints at

25 the possibility of a wider habitable zone for Earth-like terra-planets at lower obliquities.

1 Introduction

Recent advancements in observational astrophysics like with the Kepler mission led to the discovery of a vast number of potentially habitable planets (Kopparapu et al. 2014). Habitable planets are planets which can maintain liquid water on their surface (Hart, 1978; Kasting et al., 1993; Kopparapu et al., 2014). The habitability of a planet is influenced by several factors

30 like stellar flux, orbit and planetary properties (Schulze-Makuch et al., 2011). In the case of Earth-like planets, it has been shown that the width of the habitable zone strongly depends on the water cycle and the carbonate-silicate cycle (Kasting et al., 1993; Kopparapu et al., 2013; Zsom et al., 2013). This is because, these cycles control the atmospheric concentration of water vapour, carbon-di-oxide and surface pressure of a planet. In our study, we focus on the role of the water cycle.

The water cycle on a planet strongly depends on the amount of surface water present which is controlled by a complex

35 interplay between processes like atmospheric circulation, precipitation and cloud formation. Depending on the amount of surface water, habitable planets fall into two classes: aqua-planets – planets covered with global oceans – and terra-planets – planets with

vast deserts or vegetated surfaces and a limited amount of water (Dune, Herbert, 1965; Abe et al., 2005, 2011, Leconte et al., 2013). In the present study, we investigate the climate of an Earth-like terra-planet for low obliquities.

Over the recent years, there has been a growing interest in the study of terra-planets due to three reasons: First, the absence of oceans on a terra-planet helps to isolate the effects of land surface processes and thus aids in better understanding of the land-atmosphere coupling (Aleina et al., 2013; Rochetin et al., 2014; Becker and Stevens, 2015). Second, terra-planets with optically thin atmospheres (like present day Earth's atmosphere) can maintain their inner edge of the habitable zone much closer to their parent star compared to aqua-planets due to their higher surface albedo (Abe et al., 2011; Zsom et al., 2013). Third, terra-planets at lower obliquities can support a wider habitable zone compared to that of aqua-planets (Abe et al., 2011). The reason is limited atmospheric access to water in the case of terra-planets results in a dry climate with water confined to the higher latitudes for lower obliquities (Abe et al., 2005; Abe et al., 2011; Leconte et al., 2013). The dry atmosphere of terra-planets limits the escape of hydrogen molecules and shows a higher resistance to the runaway greenhouse effect. Also, the dry atmosphere inhibits the formation of clouds, ice and snow and thus helps the planet to resist complete freezing (Abe et al., 2011; Leconte et al., 2013). However, the habitable areas on such a dry planet are confined to the edges/bottom of frozen ice caps (Leconte et al., 2013; Zsom et al., 2013). Whether at such edges liquid water can exist sufficiently permanent for life to evolve and persist is unclear (Zsom et al., 2013). Recycling mechanisms similar to those which occur on the present day Earth like ocean circulation and surface runoff must exist to maintain a long-lasting liquid water inventory. Leconte et al., 2013 argue that on dry planets, mechanisms like gravity driven ice flows and geothermal flux can maintain sufficient amounts of liquid water at the edges/bottom of large ice caps. The liquid water thus formed can eventually flow back to the lower latitudes to be re-available for evaporation. However, there exists no climate modelling study on an Earth-like terra-planet which actually implements either implicitly or explicitly such a recycling mechanism bringing fresh water back from higher to lower latitudes. In our study, we consider an Earth-like terra-planet with an unlimited subsurface water reservoir as a way to mimic the recycling mechanism. Even though the water reservoir in our case is unlimited, it is not similar to an aqua-planet because it includes additional resistances which restrict atmospheric access to water (i.e. soil and plant resistances along with aerodynamic resistance).

Evidence from paleo-climate modelling studies reveals that planets within the habitable zone can support multiple climate states but all of these states may not satisfy the stable liquid water on their surface. Like our Earth can exist in two different climate states – the present day warm state with abundant surface liquid water and the Snowball Earth state with no surface liquid water (Marotzke and Botzet, 2007; Voigt and Marotzke, 2010; Voigt et al., 2011). Recent studies on the habitability of aqua-planets also indicate that the presence of multi-stability can further complicate the interpretation of habitability of Earth-like planets (Linsenmeier et al., 2015; Boschi et al., 2013; Lucarini et al., 2013). With our study we demonstrate that under certain conditions Earth-like terra-planets exhibit two drastically different climate states and both these climate states can support habitable areas with long-lived surface liquid water.

It should be noted that the present paper is mainly descriptive in nature and is not meant to explain the mechanisms for the emergence of the two climate states and the way the transition between them happens. This is still under investigation. The paper is organised as follows: Sect. 2 describes our model, our terra-planet configuration and gives an overview on the simulations performed for this study. Sect. 3 discusses the two drastically different terra-planet climate states at perpetual equinox conditions. Sect. 4 describes the transition between the two climate states. In Sect. 5 hysteretic behaviour is discussed. Sect. 6 explores the role of the snow albedo feedback for the emergence of the different climate states. Sect. 7 is about the terra-planet climate at present day obliquity. Sect. 8 draws some general conclusions from our study, in particular for the habitability of terra-planets.

2 Model and simulation setup

2.1 Model

We use the ICOsahedral Non-hydrostatic (ICON) General Circulation Model jointly developed by the MPI for Meteorology and the German Weather Service (DWD) and run it in terra-planet configuration i.e. with a single globally extended continent. The model has a horizontal resolution of R2B04 (~ 160 Km) and 47 layers in the vertical. The atmosphere model uses a non-hydrostatic dynamical core on an icosahedral-triangular Arakawa C-grid (Zangl et al., 2015), and the model atmospheric physics is similar to ECHAM6 physics (Stevens et al., 2013). The radiative transfer calculations are based on the Rapid Radiative Transfer Model (Mlawer et al., 1997; Iacono et al., 2008). Convection is parameterized by the mass flux scheme of Tiedtke (1989) with modifications to penetrative convection by Nordeng (1994). Cloud cover is calculated based on relative humidity (Lohmann and Roeckner, 1996). For a complete description of the model physics and parameterizations, the reader is referred to Stevens et al. (2013). The ICON version used in this study inherits the land physics from ECHAM5 (Roeckner et al., 2004) extended by a layered soil hydrology (Hagemann and Stacke, 2015).

2.2 Terra-planet configuration

The Terra-planet configuration is designed to be highly symmetric with no orography and glaciers. The rotation rate and the solar constant are the same as for the present day Earth. We consider two situations: perpetual equinox (zero obliquity) and present day Earth's obliquity (23.5°). Background concentration of CH_4, N_2O and aerosols are fixed to zero, while water vapor is prognostic and CO_2 concentration in the atmosphere is fixed to 348 ppmv. The ozone distribution is assumed to be zonally uniform and meridionally symmetric. Land surface properties are assumed to be homogenous: Leaf Area Index (LAI) =3, surface roughness=0.05m. The albedo of snow over the land surface ranges between 0.4 - 0.8 depending on the surface temperature. We maintain the ground water table at constant depth of 1.2 m to mimic the recycling of water from higher to lower latitudes. The root depth and soil depth are both set to 6 m such that the roots always have access to the subsurface reservoir. By prescribing this subsurface water level thereby, the water cycle in the terra-planet configuration is not closed i.e., the water drawn out of the subsurface reservoir by evapotranspiration in the low-latitudes gets accumulated near the higher latitudes but it does not return back to the low-latitudes. Our terra-planet setup closely resembles that of recent 3-dimensional studies (Abe et al., 2011; Leconte et al., 2013) on Earth-like terra-planets except that those studies did not consider an overland water recycling pathway bringing back water from higher to lower latitudes that we mimic by the prescribed subsurface water reservoir.

2.3 Simulations

We study the effect of background surface albedo (α) on the climate in a series of simulations with α varying from 0.07 to 0.24 at perpetual equinox (0°) (Z7 to Z24 simulations in Table 1) and at present day conditions (23.5°) (S7 to S24 simulations in Table 1). All these simulations start from the same initial atmospheric state with a homogenous temperature, 290 K and moisture content, 25 kg m^{-2}. Due to the small thermal inertia of the land surface and the atmosphere, the simulations reach a steady state within 10 years. We continue the simulations for additional 30 years of which, the last 10 years are used in the analysis of the mean climate. To see the transition between different climate states at the perpetual equinox condition more clearly, we perform an additional simulation (TRANS). First, we simulate the planet in the stable Hot and Dry (HD) state for 30 years for $\alpha = 0.14$ and then increase α abruptly to $\alpha = 0.14 + 0.01$ corresponding to the Cold and Wet (CW) state and continue the simulation for another 30 years. Then, to check for hysteresis, we continue the TRANS simulation by switching the albedo back to 0.14 and continued lowering α stepwise until zero. Additionally, to test the sensitivity of the climate states to model

parameterizations, we performed simulations with a different convection scheme at perpetual equinox conditions (T7 to T24 simulations). Finally, to investigate the role of snow-albedo feedback on the terra-planet climate states, the terra-planet is simulated with dark snow (DS simulations in Table 1) (i.e. we assume snow albedo to be same as background surface albedo). The details of all terra-planet simulations are listed in Table 1.

3 Drastically different climate states

Figure 1 shows the time evolution of global mean surface temperature and precipitation for different terra-planet simulations at zero obliquity (Z7 to Z24 simulations). We notice that the terra-planet exists in two drastically different climate states: a Hot and Dry (HD) state for α < 0.15 and a Cold and Wet (CW) state for α ≥ 0.15. This is different to findings in previous studies (Abe et al., 2005, 2011; Leconte et al., 2013) on low obliquity terra-planets who found only the HD state. The mean climate in the two states is remarkably different (Fig. 2). The annual mean surface temperature in the CW state is below freezing point almost everywhere, except in the low-latitudes where it is around 10°C. High cloud cover (Fig. 2e) and a high planetary albedo (Fig. 2f) lower the surface absorption of incoming radiation and result in very low temperatures in the CW state. On the other hand, in the HD state, the global mean surface temperature is about 35°C higher than in the CW state.

Besides temperature, the other striking difference between the two states is the location of precipitation bands on the planet. In the HD state, no precipitation occurs in the low-latitude region between 40°S to 40°N. Only latitudes higher than 40° receive some amount of rainfall (Fig. 2b) (compare Abe et al., 2005, 2011; Leconte et al., 2013). In contrast, in the CW state, the precipitation and cloud cover are mainly concentrated in the lower latitudes with a banded structure similar to the present day equatorial Inter-Tropical Convergence Zone (ITCZ). The reason for the absence of CW state in previous studies is the lack of an effective mechanism that can recycle water trapped at the higher latitudes in the form of snow and ice back to the lower latitudes.

Figures 2c and d show the distribution of water on the planet. In the HD state, the very high temperatures in the low-latitudes raise the water vapor saturation limit and the moisture-holding capacity of the atmosphere allowing the planet to store a substantial amount of its water in the atmosphere (Fig. 2c) leaving the upper soil layers comparably very dry (Fig. 2d). Further, supressed precipitation at lower latitudes along with large export of moisture away from the low-latitudes result in water to be mainly present at higher latitudes in the HD state. On the other hand, in the CW state, due to the low annual mean temperatures, the planet can only hold little water in its atmosphere (Fig. 2c). This, together with continuous precipitation at lower latitudes, keeps the upper soil layers in the lower latitudes always wet (Fig. 2d).

The mean circulation pattern for the two climate states is shown in Fig. 3a, b. We notice that in both the states, the circulation pattern resembles the present-day hemispheric three cell structure. But when comparing the two states, the width and intensity of the circulation are very different: The Hadley cell is narrower and less intense in the CW state compared to a wider and a more vigorous circulation in the HD state. Neutrally stable atmospheric conditions in the HD state require larger mass flux to transport away heat and to stabilise the equatorial temperatures and hence support a more vigorous circulation (Held and Hou 1980; Caballero et al. 2008; Mitchell et al., 2008). Additionally, we notice that the circulation centre of the Hadley cell is very different in the two states (depicted by black arrows in Fig. 3 a,b). The HD state has its circulation centre much closer to the surface at about 750 hPa, while in the CW state the centre is much higher at 500 hPa.

Observations and modelling studies in the literature also report intense low centred circulations for present day Earth's climate in the tropical eastern pacific during the northern hemisphere summer time (Zhang et al. 2004, 2009; Nolan et al., 2007), for a Snowball Earth state (Pierrehumbert, 2005) and for Titan (Mitchell et al., 2006, 2009). These circulations have the tendency to export larger amounts of moisture out of the lower latitudes compared to high centred circulations. We also notice this increased transport of moisture away from the low-latitudes with the more vigorous low centred Hadley cell in the HD state

compared to that in the CW state (Fig. 3c). The reason is, in the CW state the existence of deep convection allows moisture to reach much higher elevations before getting exported to higher latitudes (Nolan et al., 2007). At higher elevations, more moisture can condense and be lost as precipitation due to lower temperatures and much less remains for being exported away. Instead, in the HD state the lack of precipitation allows more moisture to be exported.

5   Figure 4 shows the northward transport of energy (transport due to latent and sensible energy) by atmospheric circulation in the two climate states. For the CW state, latent energy transport is equatorward in the low-latitudes and poleward in the mid-latitudes. In contrast, in the HD state latent energy is exported poleward at all latitudes (latent energy curves in Fig. 4). The reason for the opposite sign in the low-latitudes is that in the CW state (as in the case of present day Earth) intense precipitation dries out the atmosphere and thus keeps most of the water within the lower latitudes limiting its poleward export (Fig. 3c) –
10 hence the net flow is always equatorward. Instead, in the HD state, by the absence of precipitation, moisture is retained in the air so that it is exported poleward.

The sensible energy transport in the HD state is considerably larger as compared to that in the CW state despite of a lower equator to pole temperature gradient. This implies that the atmospheric circulation is more efficient in transporting the sensible energy to the higher latitudes. Overall, the net northward transport of energy is dominated by the sensible energy transport and is
15 enhanced in the case of HD state. Further, the larger northward transport of energy in the HD state results in a smaller equator to pole temperature gradient compared to the CW state.

4 Transition to the CW state

Starting from the HD state ($\alpha = 0.14$), we abruptly increase the albedo to $\alpha = 0.14 + 0.01$ (simulation TRANS) and thereby initiate a shift to the CW state. The full transition from the HD state to the CW state takes about 4 years. The changes in annual
20 mean surface temperature, precipitation and snow cover during the course of transition are shown in Fig. 5 while changes in mean circulation are shown in Fig. 6. One can distinguish three transitional stages:

Stage 1: After the abrupt increase in α, small precipitation clusters appear in the low-latitude region and the terra-planet starts to cool down due to an increase in cloud cover and planetary albedo. Snow cover at higher latitudes starts to increase slowly. The mean meridional circulation structure changes – small circulation cells appear very close to the equator. This happens for a
25 period of one year.

Stage 2: The precipitation clusters aggregate into a band of deep convection around the equator accompanied by a sharp increase in precipitation. At this point, we still see precipitation bands even at higher latitudes around 50°. However, the precipitation intensity at higher latitudes is smaller (around 6 mm day^{-1} lower) compared to the precipitation at the equator. Surface temperature drastically decreases and the circulation structure is now associated with two cells, a shallow cell close to the
30 equator and a deep cell slightly away from the equator.

Stage 3: Precipitation bands at the higher latitudes start moving equatorward and precipitation intensity decreases compared to the previous stage. Surface temperature decreases further. Snow cover further increases and starts moving towards the equator. The circulation is now less intense with circulation centre around 500hpa.

Finally, the CW state is reached. Precipitation only occurs at low latitudes. Surface temperature is below freezing
35 almost everywhere on the planet except at the low latitudes. Snow cover reaches down to the 40° latitude.

5 Hysteresis

To further study the bifurcation structure, we investigated the hysteretic behaviour. Figure 7 shows the global mean surface temperature plotted as a function of α. The spontaneous transition from the HD to the CW state and the associated abrupt cooling

is seen as path 1. Starting from the CW state and lowering back α stepwise below the threshold value until zero does not lead back to the HD state. The reason is that the high cloud cover present in the CW state hides the surface from solar irradiation. Therefore, a reduction of α has only a minor effect on the top of the atmosphere radiation balance so that thereby it is impossible to heat the planet sufficiently strongly to switch back to the HD state. This is true even when repeating these simulations taking

5   snow albedo equal to background albedo (experiments DS; no figure shown). Thus, the planet remains in the CW state indicating that under the chosen conditions for the terra-planet setup the hysteresis is not closed (Fig. 7).

6 Does snow albedo feedback play a role for the emergence of the two climate states?

Changes in snow cover can lead to multiple climate states in the Earth system with drastically different global mean surface temperatures like the present day Earth and the Snowball Earth (Budyko 1969; Sellers 1969). In this case, the large temperature

10   difference between the two states is caused by the positive snow-albedo feedback. In our study, we also notice such huge differences in global mean surface temperatures between the two terra-planet climate states (Sect. 3). In order to test whether the snow albedo feedback is responsible for the huge temperature difference in our study, we performed additional simulations where we changed the snow albedo to be equal to the background albedo of soil (simulations DS in Table 1). With the darker snow, we still find the spontaneous transition between the two states. The existence of the bifurcation even in the simulations

15   with dark snow implies that the snow-albedo feedback is not the cause for the existence of the two states. However, the snow albedo feedback does enhance the drastic temperature change in the bright snow simulations by around 12°C compared to the dark snow simulations but is not solely responsible for the temperature difference between the two states. It seems the main contribution to the huge temperature difference comes from the positive cloud feedback. This can be inferred by looking at the planetary albedo curves in the DS experiments which only represent contributions from clouds and background soil. (Fig. 8).

20   ## 7 Terra-planet with seasonality

So far, we considered a planet without seasonality because obliquity was set to zero. Next, we investigate how the climate on a terra-planet changes with a seasonal orbit taking present day Earth's obliquity of 23.5° (S simulations in Table 1). We find that with seasonality the HD state disappears and the terra-planet always stays in the CW state (Fig. 9). This is contrary to previous studies on terra-planets which show that terra-planets always exist in a warm state for obliquities lower than 30° (Abe et al.,

25   2005, 2011). We notice that, in the presence of the seasonal cycle the meridional structure of insolation changes drastically resulting in a weaker meridional temperature gradient compared to the situation without seasonal cycle (Fig. 9a). This avoids overheating and drying of the lower latitudes (Fig. 9d) and stabilises the planet in the CW state.

8 Discussion and Conclusions

So far terra-planets have been investigated for a wide range of planetary properties like mass, rotation rate, atmospheric

30   composition and orbit. Here, we focus on climate simulations of a terra-planet with low obliquity (< 30°), flat orography, and otherwise Earth like conditions. An important difference to previous studies concerns the treatment of atmospheric access to water. Past studies on possible climates of terra-planets prescribed a limited water inventory, which for low obliquities leads to trapping of the water at high latitudes and stabilizing the planet in a state with no precipitation at low latitudes, similar to what we call in our study 'Hot and Dry' state. By contrast, in the present study we assume an unlimited subsurface water reservoir,

35   which is still different from an aqua planet configuration (having also an unlimited water reservoir), because soil and plant resistances restrict the atmospheric access to surface water in addition to the restriction from aerodynamic resistance. For such an Earth-like terra-planet with restricted water access we find two drastically different climate states, a Hot and Dry (HD) state

characterized by a hot climate with precipitation confined to higher latitudes and a Cold and Wet (CW) state which is closer to present-day Earth's climate with precipitation mainly occurring at lower latitudes and an intense cycling of water there. Compared to the other studies mentioned above, we find this additional CW state only, because we implicitly assume for our terra-planet a mechanism restoring water from the higher latitudes to lower latitudes, refilling sufficiently the subsurface water reservoir to maintain the very active low latitude water cycle in this state. The difference in global mean temperature between these two climate states is 35°C (with the same boundary conditions), which is in the same order of magnitude as the temperature difference between present day and the Snowball Earth climate (Pierrehumbert et al., 2011; Micheels and Montenari, 2008; Fairchild and Kennedy, 2007; Hoffman and Schrag, 2002).

Similar to the abrupt transition to a Snowball Earth state, for perpetual equinox conditions one of our terra-planet simulations also shows an abrupt transition, namely from the HD state to the CW state. These two states exist for low background surface albedo α, while for high α only the CW state is possible. The abrupt transition occurs when in the HD state α is increased beyond a particular threshold value. While the transition to the Snowball Earth is driven by the snow albedo feedback, the transition in our study is mostly driven by the cloud albedo feedback, accompanied by huge changes in the global hydrological cycle, an increase in planetary albedo, and a complex reorganization of the circulation patterns. Moreover, we notice that in our setup the terra-planet exhibits an open hysteresis: even with background surface albedo reduced to zero it does not return back to the HD state. Additional terra-planet simulations with an obliquity as the real Earth all result in a CW state, hinting at another possible bifurcation from the HD to the CW state when the obliquity is increased to non-zero values, but further analysis is needed to confirm this.

Concerning the global water cycle, the HD and CW states share one important similarity. The planetary boundary layer at lower latitudes is extremely dry in the HD state with relative humidity of about 15%. Clearly, trade winds can maintain such a dry boundary layer only if the water supply at the surface is sufficiently limited. However, without any evapotranspiration at the surface the Hadley Cell would dry out completely, loosing the greenhouse effect that sustains the high temperatures of the HD state. Furthermore, in the CW state, to keep the rain along the equator one also needs a considerable water supply in the lower latitudes. In summary, both climate states are associated with a strong atmospheric transport of water from lower to higher latitudes, which has to be balanced by evapotranspiration at the lower latitudes. Therefore, allowing for both states to be potentially realized under the same boundary conditions, on a real planet, mechanisms must exist that can continuously restore water back to lower latitudes. Such mechanisms exist on each conceivable planet with low obliquity and water in the lower latitudes. For present day Earth this happens via the oceans. For our terra-planet the water is stored in frozen form at higher latitudes like in the past glacial states of our Earth. The resupply of water may happen via processes like melting of glaciers, transport of ice by gravity flows and melting at the bottom of large ice caps due to high pressure and geothermal heat flux. This provides liquid water, which is brought back to the lower latitudes by rivers (Abe et al., 2010; Leconte et al., 2013). Note that the huge differences in climate between these two states is primarily a consequence of the completely different functioning of the global hydrological cycle, so that the assumed recycling mechanism can be considered as an additional degree of freedom to the internal dynamics extending the range of possible terra-planet climates.

Our findings may have some relevance for estimates of the habitable zone for such Earth-like terra-planets. In the HD state liquid water is confined to the mid-latitudes (40°-50°) in both hemispheres, whereas in the CW state, there are sufficient precipitation and high enough temperatures for permanent liquid water at low latitudes (40°S - 40°N). Thus in both climate states life can potentially persist. At the outer edge of the habitable zone, the assumed resupply of water from higher to lower latitudes stabilizes the greenhouse effect, keeps the planet in the HD state and may prevent a situation with all water accumulated at the higher latitudes in the frozen form. At the inner edge of the habitable zone, by this resupply the planet can maintain precipitation

and high cloud cover at the equator in the CW state. Thereby, the planetary albedo is increased which cools the planet and may prevent the runaway greenhouse state with all the water well mixed in the atmosphere in the gas phase. On the whole, our study thus suggests that the presence of a mechanism which can recycle fresh water from the higher latitudes back to the lower latitudes and results in the two drastically different climate states may extend the habitable zone of Earth-like terra-planets at

5   lower obliquities. But substantiation of this hypothesis would need further simulations.

Appendix A: Sensitivity of the climate states to model convection scheme

To confirm that the two climate states of the terra-planet are not an artifact of convective parameterizations, we performed an additional simulation with a different convection scheme. By default, the model uses the Nordeng convection scheme (Nordeng 1994). We have replaced the default settings and simulated terra-planet simulations with the Tiedtke convection scheme (Tiedtke

10   1989) and we find the two drastically different climate states irrespective of the convection scheme (Fig. A1).

Competing interests. The authors declare that they have no conflict of interest.

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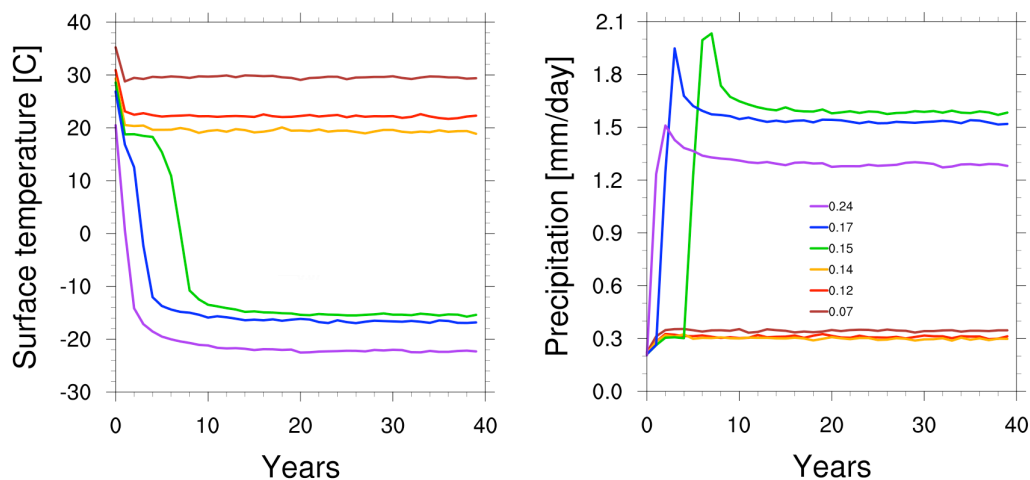
Simulations	Background soil albedo	Obliquity	Convection scheme	Snow albedo
Z7 – Z24	0.07, 0.12, 0.14, 0.15, 0.17, 0.24	0	Nordeng	Dynamic
S7 – S24	0.07, 0.12, 0.14, 0.15, 0.17, 0.24	23.5	Nordeng	Dynamic
TRANS	0.14 – 0.15 – 0.14 – 0.12 – 0.07- 0.00	0	Nordeng	Dynamic
T7 – T24	0.07, 0.12, 0.14, 0.15, 0.17, 0.24	0	Tiedtke	Dynamic
DS	0.07, 0.12, 0.14, 0.15, 0.17, 0.24	0	Nordeng	Constant at background soil albedo

Table 1. Summary of simulations performed in this study.

5 **Figure 1: Time series of global mean surface temperature (°C) and precipitation (mm day^{-1}) for different background soil albedo values in the Z7- Z24 simulations.**

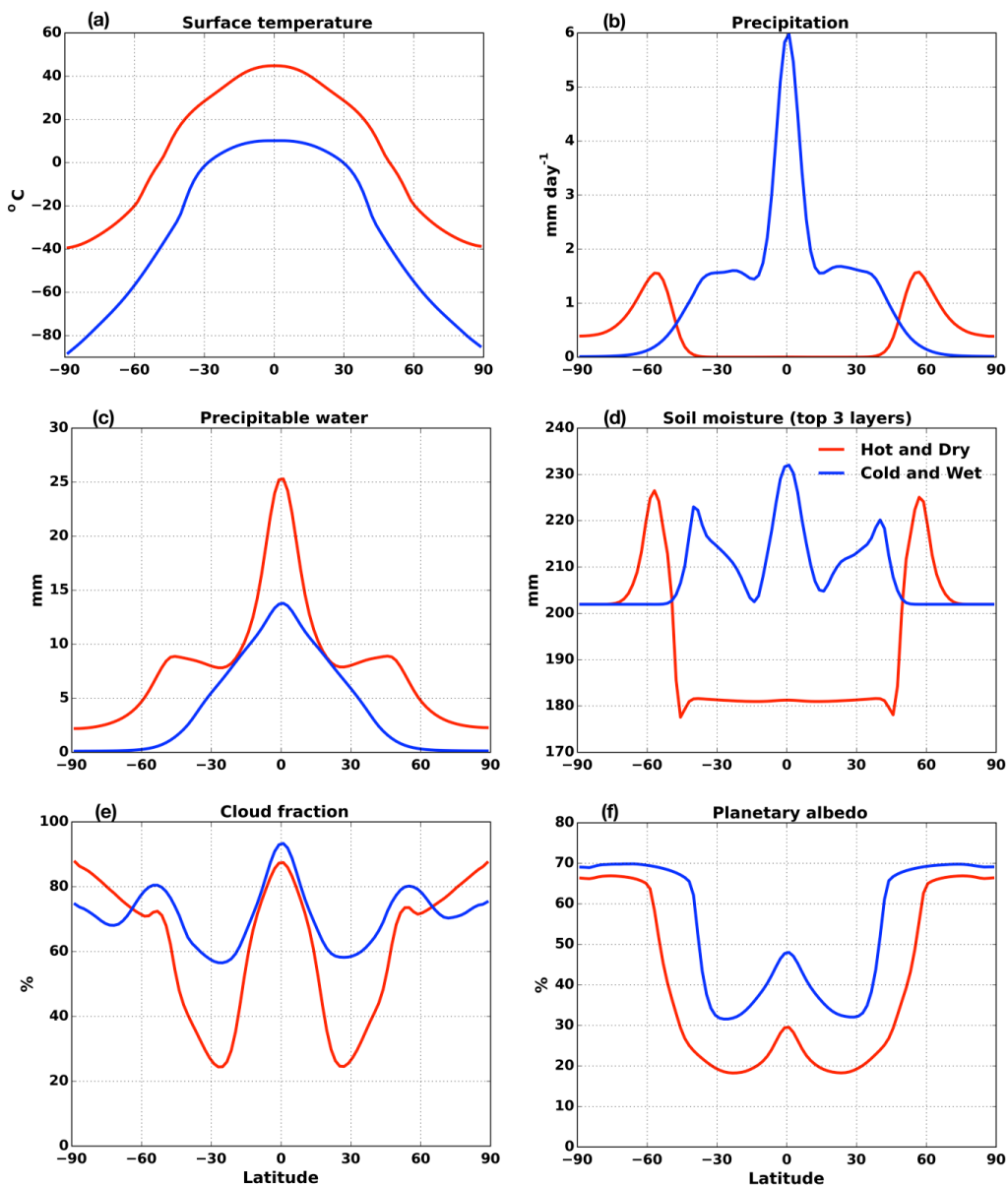



Figure 2: Annual mean meridional profiles of (a) surface temperature, (b) precipitation, (c) precipitable water, (d) soil moisture (averaged over the top three layers), (e) cloud fraction and (f) planetary albedo for the two terra-planet states: HD (α = 0.14) and CW (α = 0.15) in the simulations Z14 and Z15 averaged over a period of ten years.

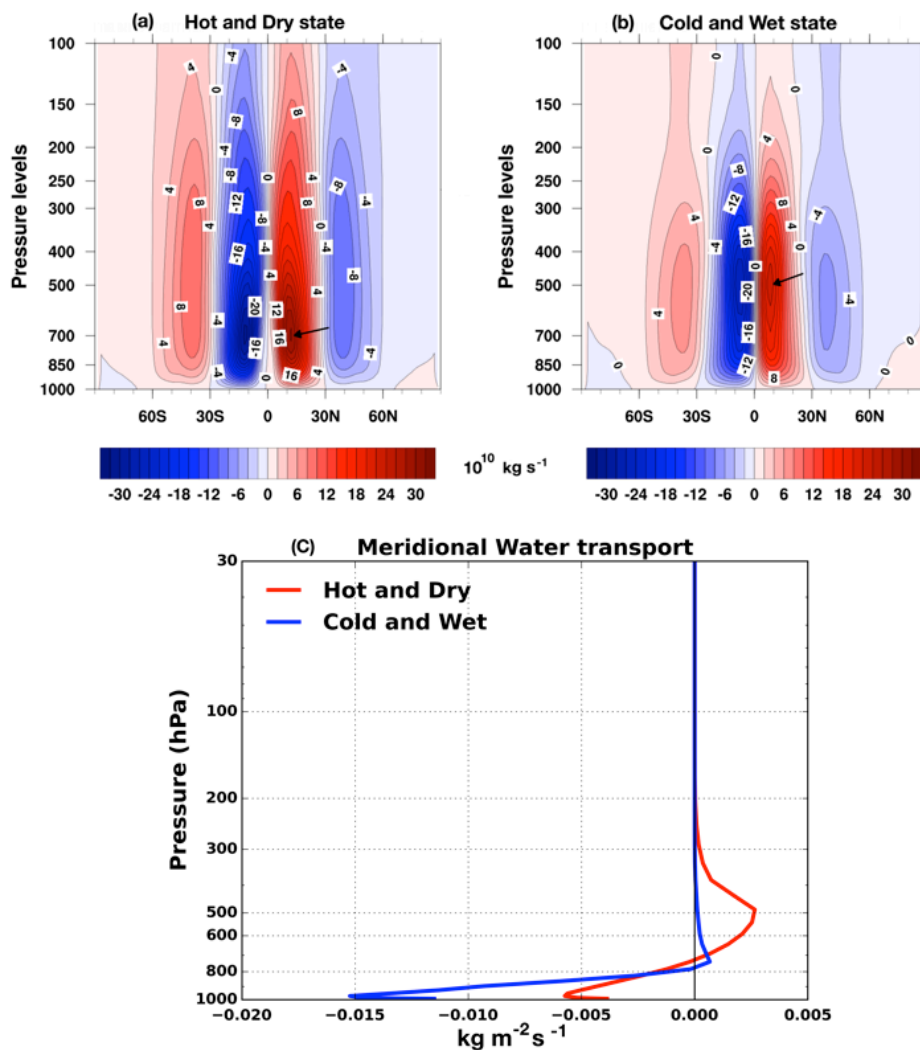



Figure 3: Annual mean behavior of (a) and (b) meridional stream function in 10^{10} kg s^{-1}, (c) vertical profiles of meridional transport of water at 10°N for the two terra-planet states in the simulations Z14 and Z15 averaged over a period of ten years. The black arrows in (a) and (b) denote the northern Hadley circulation centre.

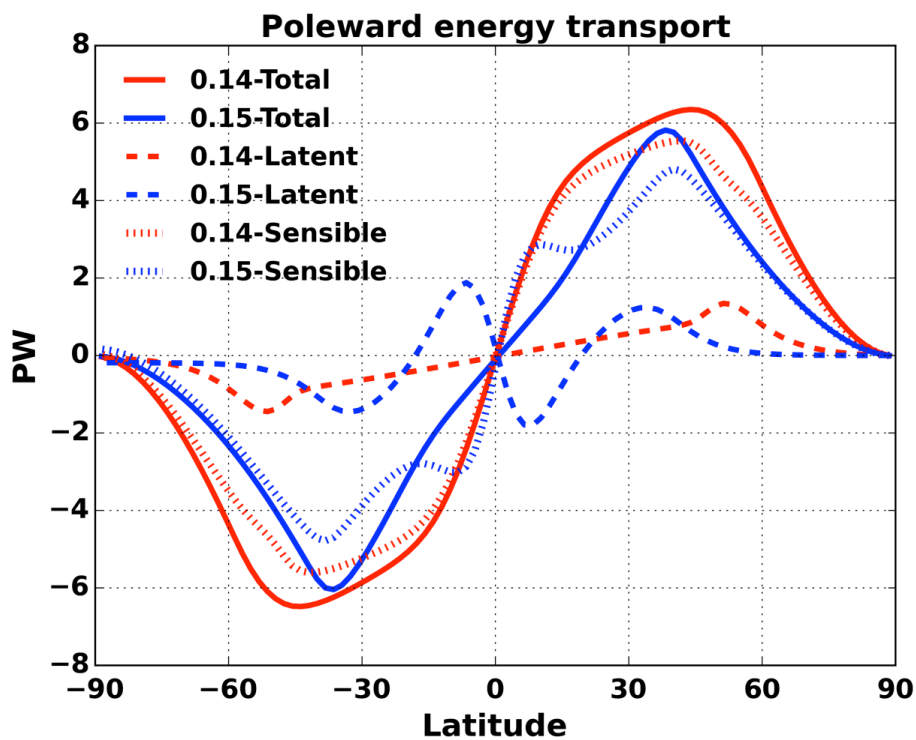



Figure 4: Annual mean behavior of northward transport of total, latent and sensible energy in peta watts (PW) for the two terra-planet states in the simulations Z14 and Z15 averaged over a period of ten years.

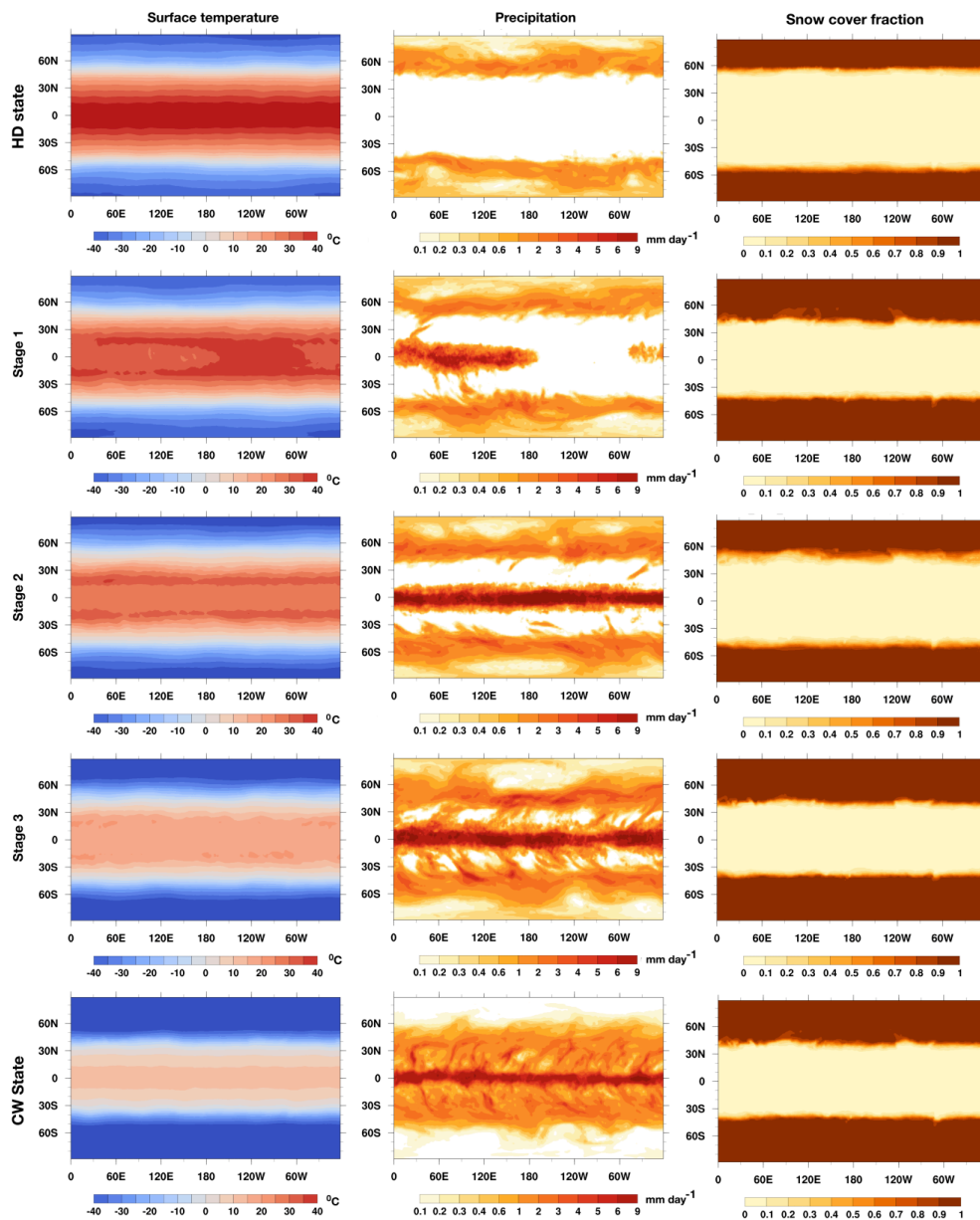



Figure 5: Annual mean surface temperature (left panels), precipitation (centre panels) and snow cover fraction (right panels) for different stages of the transition (TRANS simulation) from the HD state to the CW state.

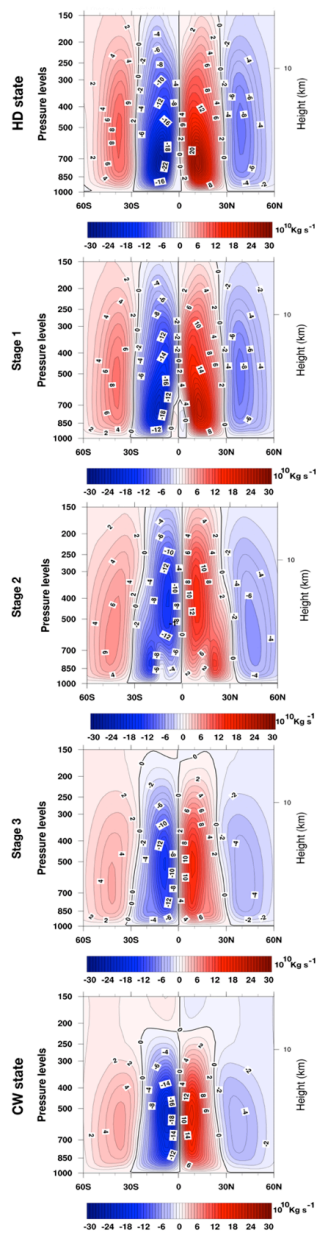



Figure 6: Annual mean meridional stream function in 10^{10} kg s^{-1} for different stages of the transition (TRANS simulation) from the HD state to the CW state.

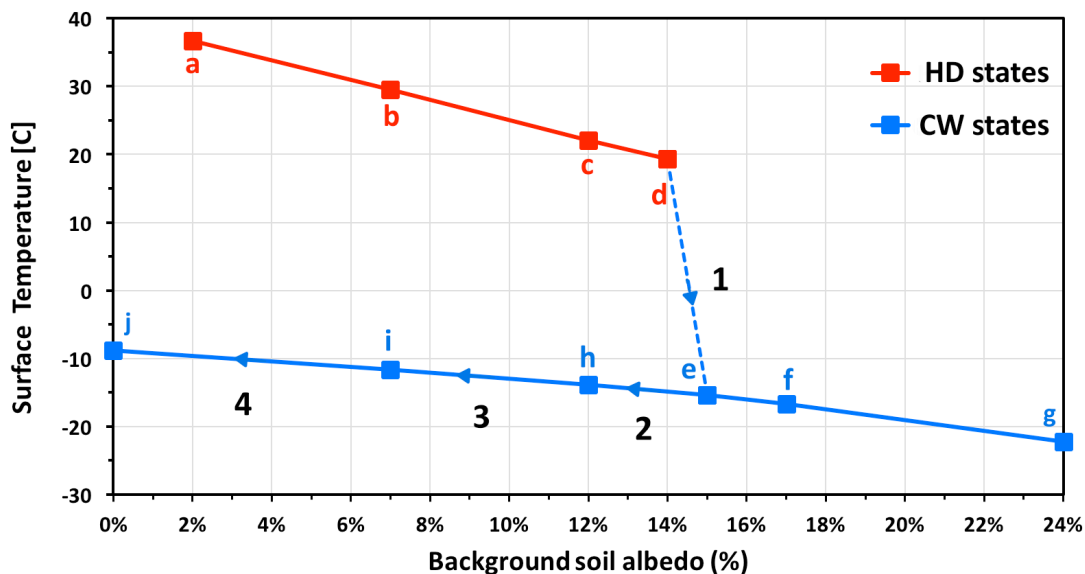



Figure 7: Global mean surface temperature as a function of background soil albedo (α). Simulated HD states are denoted by points a-d and simulated CW states by e-j. Path 1 denotes the spontaneous transition from HD to CW when increasing α and paths 2, 3 and 4 denote the reverse state development upon stepwise lowering of background albedo starting from the threshold value until zero. Obviously, the hysteresis doesn't close for the terra-planet at zero obliquity considered here.

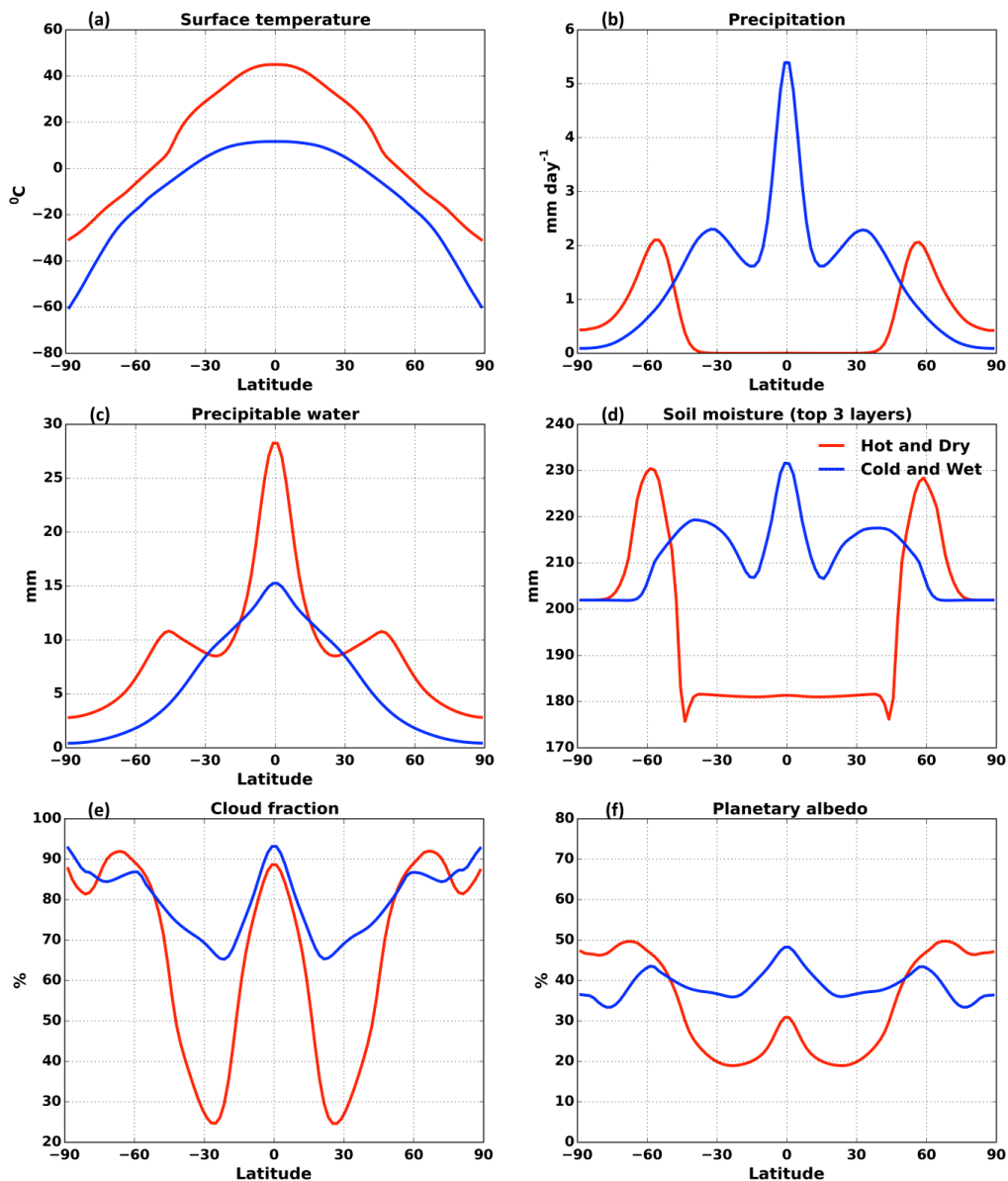



Figure 8: Annual mean meridional profiles of (a) surface temperature, (b) precipitation, (c) precipitable water, (d) soil moisture (averaged over the top three layers), (e) cloud fraction and (f) planetary albedo for the two terra-planet states averaged over a period of ten years similar to the Fig. 2 but with dark snow (snow albedo is same as background soil albedo – DS simulation).

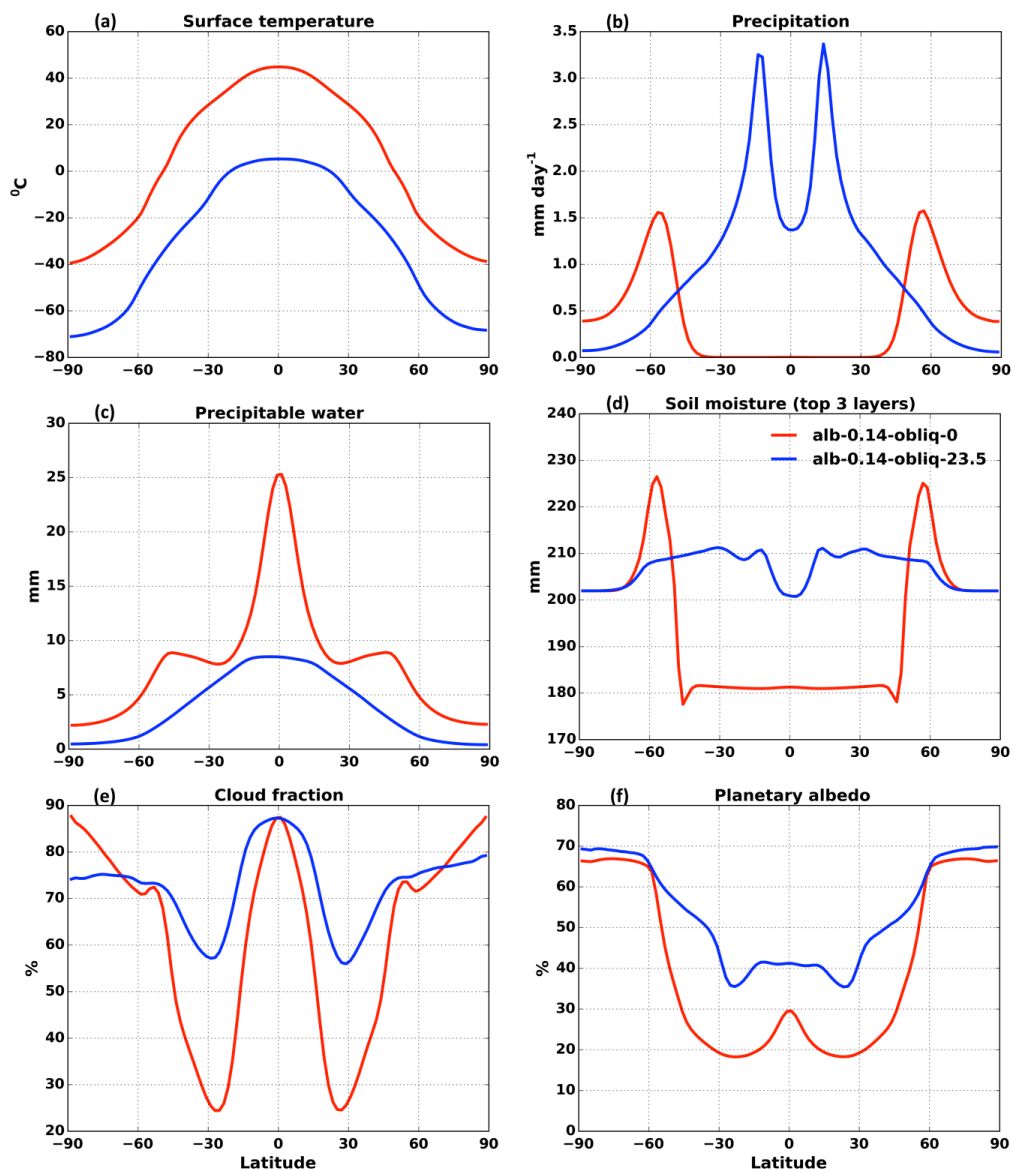



Figure 9: Annual mean meridional profiles of (a) surface temperature, (b) precipitation for the two terra-planet states with the same albedo of 0.14 but obliquities: 0° (red lines – Z14 simulation in Table 1) and 23.5° (blue line – S14 simulation in Table 1) averaged over a period of ten years.

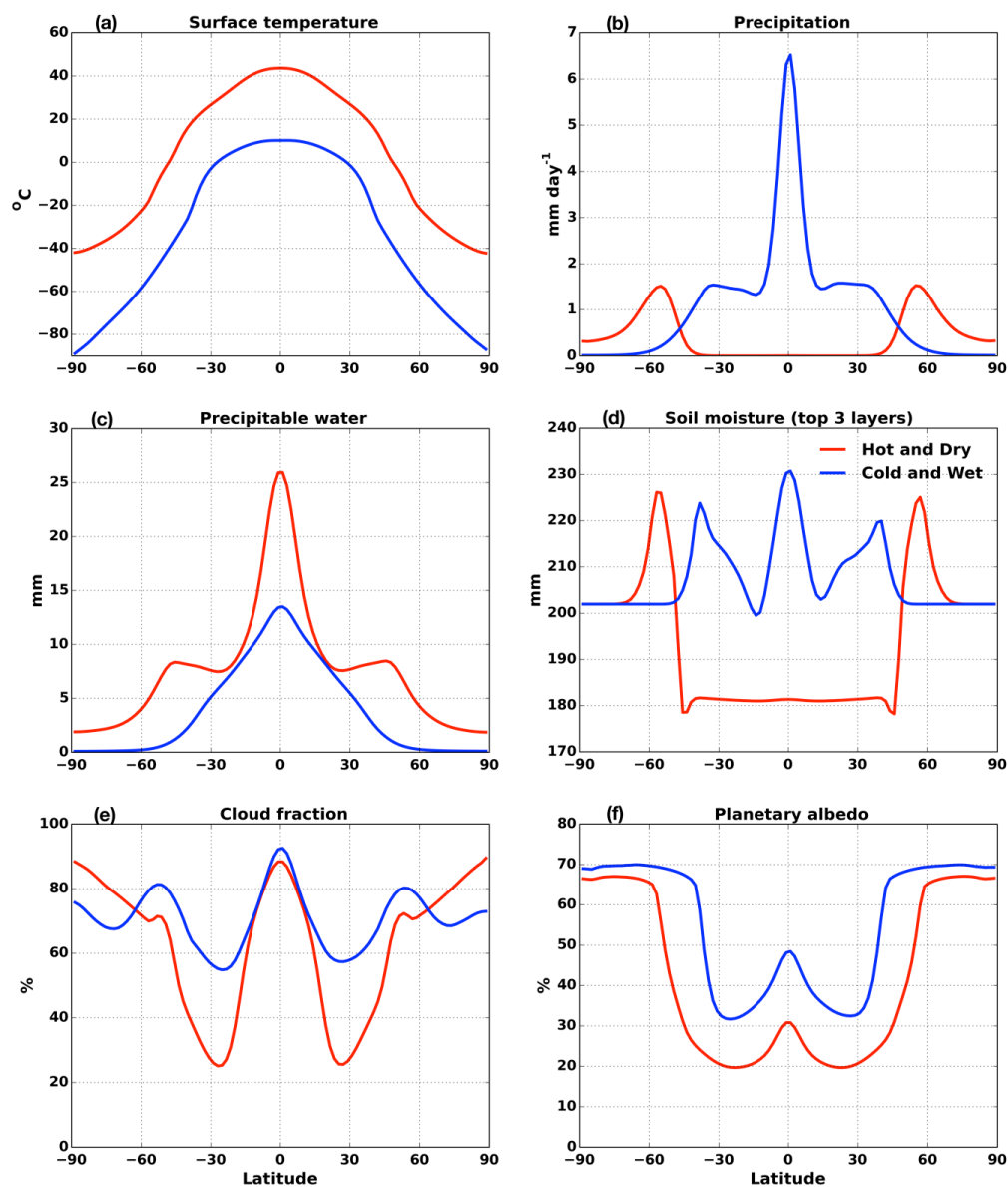



Figure A1: Annual mean meridional profiles of (a) surface temperature, (b) precipitation, (c) precipitable water, (d) soil moisture (averaged over the top three layers), (e) cloud fraction and (f) planetary albedo for the two terra-planet states averaged over a period of ten years similar to the Fig. 2 but with Tiedtke convection scheme plotted using T simulations in Table 1.