Regional scaling of annual mean precipitation and water availability with global temperature change

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Abstract. Changes in regional water availability belong to the most crucial potential impacts of anthropogenic climate change, but are highly uncertain. It is thus of key importance for stakeholders to assess the possible implications of different global temperature thresholds on these quantities. Using a subset of climate model simulations from the 5th phase of the Coupled Modeling Intercomparison Project (CMIP5), we derive here the sensitivity of regional changes in precipitation and precipitation minus evapotranspiration to global temperature changes. The simulations span the full range of available emissions scenarios and the sensitivities are derived using a modified pattern scaling approach. The applied approach assumes linear relationships on global temperature changes while thoroughly addressing associated uncertainties via resampling methods. This allows us to assess the full distribution of the simulations in a probabilistic sense. Northern high latitude regions display robust responses towards a wetting, while subtropical regions display a tendency towards drying but with a large range of responses. Even though both internal variability and the scenario choice play an important role in the overall spread of the simulations, the uncertainty stemming from the climate model choice usually accounts for about half of the total uncertainty in most regions. We additionally assess the implications of limiting global mean temperature warming to values below (i) $2K$ or (ii) $1.5K$ (as stated within the 2015 Paris Agreement). We show that opting for the $1.5K$-target might just slightly influence the mean response, but could substantially reduce the risk of experiencing extreme changes in regional water availability.

1 Introduction

Assessing regional changes in mean-annual precipitation, P, and precipitation minus evapotranspiration, $P - E$ (often also referred to as water availability), in the context of on-going global warming is of high relevance for a wide-range of socio-economic sectors. Regional differences in P and $P - E$ pose important challenges to farmers, water resources managers, stakeholders and decision-makers and a comprehensive, easily accessible communication and visualization of complex climate model output is necessary to allow for targeted adaptation and mitigation strategies.

The public and political debate on climate change is usually limited to a debate about global temperature change, which is, however, an abstract measure and does not enable end-users to infer direct implications for regional to local climate change, especially also with respect to hydroclimatological variables (Victor and Kennel, 2014; Seneviratne et al., 2016). However,

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due to its omnipresence in popular climate communication, global mean temperature T could be used as a general measure of climate change and thereby enable a different communication of regional climate impacts to the public: 'The regional change of a climate variable as a function of global warming'. Many studies use approaches following this guideline, with one of the most common techniques used being summarized as 'pattern scaling'.

5   In this study we follow the tradition of pattern scaling, but introduce a more rigorous, probabilistic assessment of the underlying uncertainties. Common pattern scaling approaches originally have the goal to use a spatial response pattern in a certain variable (e.g. regional temperature, precipitation) that is derived from observational or (usually) from climate model data with respect to global mean temperature or CO_2-changes in order to create a large number of additional scenarios (Santer et al., 1990; Mitchell, 2003). In this study we use the "large number of additional scenarios" created by the utilized pattern scaling technique to estimate the uncertainty distribution of the response pattern in a probabilistic approach. Pattern scaling approaches have been employed in a large number of studies (see e.g. Tebaldi and Arblaster (2014) for an overview), but many common approaches to estimate the response pattern are also subject to an ongoing debate (Tebaldi and Arblaster, 2014; Herger et al., 2015; Kravitz et al., 2016). To estimate the spatial response pattern in mean-annual P and $P - E$, we adapt here a technique based on the assumption that the scaling relationship between local temperature at each gridpoint and global mean temperature is linear and that the resulting maps of regression slopes could be used as the response pattern (Solomon et al., 2009). Following a more empirical approach without a priori assumptions on the relationship of regional variables on global temperature, it was recently shown that these findings also hold for extreme temperatures and extreme precipitation, mostly independent of emission scenarios (Seneviratne et al., 2016). This approach was further applied and extended in Wartenburger et al. (2017) by using a comprehensive set of hydroclimatological variables, including both mean-annual P and $P - E$. The assessment presented in this work builds upon the analysis presented in Wartenburger et al. (2017) by utilizing a similar data collection to quantify the associated response pattern.

The scaling relationship between global mean P and global warming has also been analysed in previous studies (Andrews et al., 2009; Frieler et al., 2011; Pendergrass and Hartmann, 2012; Fischer et al., 2014; Pendergrass et al., 2015). At global scales, mean precipitation scales positively with global temperature increase (Knutti et al., 2016), but the associated scaling coefficient is still subject to an ongoing debate and might not necessarily follow a linear relationship (Good et al., 2016). It was further shown that the magnitude of the scaling relationship depends on the emission scenario (Andrews et al., 2009; Frieler et al., 2011; Pendergrass and Hartmann, 2012; Pendergrass et al., 2015), whereas the scaling relationship of extreme precipitation is independent of the emission scenario (Pendergrass et al., 2015; Seneviratne et al., 2016). The scaling relationship between global mean $P - E$ and global warming was, to our knowledge, only assessed in Wartenburger et al. (2017) and more research is needed to evaluate the full range of potential impacts of regional water availability change.

We aim here to develop a methodological framework in order to assess regional changes in mean-annual P and $P - E$ over global land areas with respect to global warming by using a representative subset of climate models and considering different emission scenarios. We further account for the internal variability of each projection by considering the year-to-year variability of P and $P - E$. This enables us to generate conservative estimates of the uncertainty distribution of the scaling coefficient for P and $P - E$ at every gridpoint and within specific regions.

Another issue addressed within this study is related to the implications of different global warming-degree targets on regional P and $P - E$. At the United Nations Climate Change Conference held in Paris in 2015 (COP21), most nations agreed to limit the increase in global mean temperature to values "well below $2K$" and to ideally not surpass a warming of $1.5K$ above pre-industrial conditions. Thereby, previous goals to limit global warming to "only" $2K$ global warming are significantly intensified. However, this raises the question of potential implications and differences between these "warming-degree targets" with respect to changes in many other climate variables besides the (rather abstract) value of global mean temperature and especially at regional scales (Seneviratne et al., 2016; Schleussner et al., 2016; Guiot and Cramer, 2016; James et al., 2017). The framework developed within this study allows us to directly assess regional changes in P and $P - E$ in the context of these warming-degree targets, thereby providing important and useful information to decision makers, farmers, water resources managers, stakeholders and the general public within a specific region.

First, we introduce the climate model data that is utilized within this study before describing the methodological approach that is used to estimate the uncertainty distribution of the scaling coefficients of P and $P - E$ with respect to global warming (Sec. 2). We provide in the following illustrations of the median and the range of the scaling coefficients (Section 3). Next, we comprehensively assess the uncertainty that is stemming from the choice of emission scenario (Section 3.1) and how other sources of uncertainty contribute to the total uncertainty (Section 3.2). We further apply the new framework to analyse changes between different warming degree targets (Section 4) and summarize and discuss our results also within the context of previous assessments (Sec. 5).

2   Scaling - Data and Methodology

The Coupled Model Intercomparison Project, version 5 (CMIP5) ensemble (Taylor et al., 2012) includes climate model projections forced by four Representative Concentration Pathway (RCP) emissions scenarios (Moss et al., 2010). These scenarios correspond to their relative radiative forcings reached by the end of the 21st century with respect to the preindustrial period: 2.6, 4.5, 6.0, and $8.5Wm^{-2}$ (from hereon referred to as RCP2.6, RCP4.5, RCP6.0 and RCP8.5). We use a total of 14 climate models selected based on prerequisites provided in Fischer et al. (2014), which are only one model from each modeling center and thereof the newest with the highest resolution. Please note that not all climate models provide data for all emission scenarios (see Table 1). We consider a time period of 120 years beginning in 1980 and ending in 2099 comprising historical simulations for the first 26 years which emerge into simulations of the respective emission scenarios from 2006 onwards. The first 20 years (1980-1999) are used as a common baseline period and values in mean annual P and $P - E$ are assessed in relative terms [%] with respect to the baseline period. Since we focus here on global land areas, please note that in case $P - E < 0$ for all models and scenarios, those grid points were neglected.

For each model m and each emission scenario s within the 100-year period 2000-2099 (yr), the relative values of precipitation, $P_{m,s,yr}$, and precipitation minus evapotranspiration, $(P - E)_{m,s,yr}$, are regressed at each gridpoint (or averaged

over a certain region) against mean-annual global temperature, $T_{m,s,yr}$). We use an ordinary least squares fit to estimate the parameters of the linear equation:

$$P'_{m,s,yr} = r_{m,s} \cdot T_{m,s,yr} + I_{m,s} \tag{1}$$

with $r_{m,s}$ denoting the regression slope and $I_{m,s}$ the intercept (and likewise for $(P-E)_{m,s,yr}$). The slope itself provides us with an estimate of the regional scaling coefficient of P against global changes in T.

Given the annual residuals $R_{m,s,yr} = P_{m,s,yr} - P'_{m,s,yr}$, the uncertainty of the regression slope $r_{m,s}$ is assessed by resampling years yr' of the residuals ($R_{m,s,yr'}$) and fitting the regression slope against the new pairs ($T_{m,s,yr}, P'_{m,s,yr} + R_{m,s,yr'}$). Repeating this approach 1000-times at each gridpoint (or within each specific region) provides us with a comprehensive uncertainty measure $\epsilon_{m,s}$ of each model- and scenario-specific regression slope $r_{m,s}$ for both P and $P-E$. We like to point out that the uncertainty estimated through resampling residuals results in very similar results as computing the uncertainty through using different realisations of a single model (this is shown for CSIRO-Mk3-6-0 in the supplementary information). We further note, that scaling coefficients in low latitude regions represent the response within a larger area compared to those in high latitude regions. To also test the validity of the linearity assumption, we assessed (i) if the Kolmogorov-Smirnov test does not reject the null hypothesis that the annual residuals $R_{m,s,yr}$ are normally distributed and (ii) if there is no significant lag-1yr autocorrelation of all $R_{m,s,yr}$. Following this approach, in the majority of world regions and for most models, the linearity assumption is potentially valid. However, please note for the following sections, that in many hyper-arid regions both tests fail for the majority of models.

This approach allows us to distinguish between three different sources of uncertainty. As illustrated in the conceptual Fig. 1, these are (i) internal variability, σ_i, representing the uncertainty stemming from interannual variability for each model under each scenario, (ii) the model uncertainty, σ_m, related to the uncertainty across all models and for a specific scenario (e.g. in terms of variance: $\sigma_m = 1/n \sum_{m=1}^{n}(r_{m,rx} - \mu_r)^2$, with $n=14$ models and μ_r denoting the average of all scaling coefficient) and (iii) the scenario uncertainty, σ_s, related to the uncertainty across all scenarios and for a specific model or the multimodel mean (e.g. in terms of variance: $\sigma_s = 1/n \sum_{rx=1}^{n}(r_{m,rx} - \mu_r)^2$, with $n=4$ scenarios and μ_r denoting the average of all scaling coefficients). The total uncertainty denotes the uncertainty across all models, all scenarios and considering the internal uncertainty. Please note, that this approach of attributing uncertainties is very simplistic and neglects any potential relationship between the individual sources of uncertainty, but is suitable and useful to provide a general measure of the underlying uncertainty sources.

3   Scaling - Results

Considering the total uncertainty across all models and scenarios and by additionally including the internal variability, we are able to estimate the uncertainty distribution of the regional scaling coefficient of P ($P-E$) against globally-averaged T. Displayed in Fig. 2 are the median and the 10th and 90th quantile of the uncertainty distribution of the scaling coefficient

for both P and $P-E$ at each gridpoint. The median scaling coefficient shows positive values, and hence an increase in both P and $P-E$ with increasing T, in most parts of the northern high latitudes and Asia, but also in eastern Africa for both P and $P-E$. Negative values, and hence a decrease in both P and $P-E$ with increasing T, are found in the Mediterranean region, southern Africa, Australia and in parts of West Africa, as well as Central and South America. Comparing the 10th and

5   90th quantiles of the uncertainty distribution shows the range of possible scaling coefficients. This range does, in most regions and especially for $P-E$, include the zero coefficient, which means that the probability of experiencing a scaling response of different sign compared to the median response, is, at least, 10% or higher. The range is further generally much larger for $P-E$, pointing towards overall higher uncertainties in the estimation of the scaling relationship. The range is especially large in most subtropical regions (e.g. Sahara, Arab peninsula, India, Australia, etc.). Regions showing a significant increase of P

10   and $P-E$ with global warming are located mainly in the northern high latitudes.

Fig. 3 summarizes these findings by qualitatively showing the probability of experiencing either a positive or negative scaling response in P with respect to global warming. A very likely increase ($90-100\%$ probability) in regional P with ongoing global warming is hence found only within regions of the northern high latitudes, whereas a likely increase ($66-100\%$ probability) is located also in many parts of Asia and North America and to a minor extent also in some regions of South America and

15   Africa. A likely decrease is located in most parts of the Mediterranean region, southern Africa, northeastern South America, Central America and along the Australian coastal regions. A decrease that is very likely is only found in South Africa. Most other regions show either uncertain or no change. Fig. 3 also illustrates selected quantiles of the uncertainty distribution of the scaling coefficient of P as a function of global temperature increase for a comprehensive subset of SREX regions (Seneviratne et al., 2012) as outlined in the map (see also Table 2 for more information). Very certain responses within the SREX-regions

20   are only found for those in the northern high latitudes (ALA, NAS, NEU), while most other regions show a large spread of the uncertainty distribution (especially e.g. in NEB, NAU).

Similarly for $P-E$, Fig. 4 displays a very likely increase ($90-100\%$ probability) in regional $P-E$ with ongoing global warming for an even smaller portion of land in the northern high latitudes, whereas a likely increase ($66-100\%$ probability) is located throughout the northern high latitudes and similarly to P in many parts of Asia and North America and to a minor

25   extent also in some regions in South America and Africa. A likely decrease is located in parts of the Mediterranean region, southern Africa, northeastern South America, Central America and some parts of Australia. A very likely decrease is only found for single gridpoints, primarily in Central America. Most other (and when compared to P an even higher number of) regions show either uncertain or no change. Fig. 4 also illustrates selected quantiles of the uncertainty distribution of the scaling coefficient of $P-E$ as a function of global temperature increase for the same set of SREX regions as shown in Fig. 3. Very

30   certain responses are again only found in the northern high latitudes (ALA, NAS) and southern Asia (SAS), while most other regions show a very large and even larger spread of the uncertainty distribution when compared to estimates of P.

Please note that the results for both P and $P-E$ are neither substantially influenced by the unequal number of available models per scenario nor individual climate models of the ensemble that potentially exhibit a large hydroclimatological drift (Liepert and Previdi, 2012; Liepert and Lo, 2013) (refer to the supplementary material for more information).

3.1 Scenario uncertainty

The probability of experiencing an increase or decrease in regional P and $P - E$ with global warming depends on the emission scenario. At global scales, mean precipitation scaling was shown to depend on the emission scenario (Andrews et al., 2009; Frieler et al., 2011; Pendergrass and Hartmann, 2012; Pendergrass et al., 2015), whereas the scaling of extreme precipitation is independent of the emission scenario (Pendergrass et al., 2015; Seneviratne et al., 2016). Here we assess the relationship of regional changes in P and $P - E$ on the emission scenario by analysing the uncertainty distributions of the scaling coefficient for each scenario individually. A conceptual representation of the probability of the scaling coefficient being postive/negative is displayed for P in Fig. 5 and for $P - E$ in Fig. 6 (similar to the total uncertainty as shown in Fig. 3 and Fig. 4). In general, the fraction of regions showing either likely or very likely changes is increasing with the emission scenario for both P and $P - E$, pointing towards a larger uncertainty in the estimation of the scaling coefficient in case the climate change forcing is weak (RCP2.6, RCP4.5). Further, regions showing very likely changes are more common under high emission scenarios (RCP6.0, RCP8.5). The drying response in the Mediterranean region is e.g. not evident when considering the RCP2.6 scenario alone, whereas a very likely decrease is found in the RCP8.5 scenario. In fact, individual grid points of the Mediterranean region (e.g. in central Spain) even show a likely increase in P and $P - E$ in the RCP2.6 scenario. On the other hand, a likely drying response in parts of central and northen Australia found in RCP2.6 disappears for higher emission scenarios for P, or even turns into a wetting response for $P - E$. Robust signals are, again, found in most parts of the northern high latitudes, showing a (very) likely increase across all emission scenarios. A (very) likely decrease across all scenarios is further found for parts of southern Africa and parts of the Amazon region. Also note here that the overall conclusions for both P and $P - E$ are not influenced by the unequal number of available models per scenario (see supplementary information).

A more detailed look on the underlying uncertainty distributions for P ($P - E$) within each SREX-region is provided in Fig. 7 (Fig. 8). It is clearly evident that the uncertainty is largest for low emission scenarios throughout all regions and in most cases lowest for the RCP8.5 scenario (with overall larger uncertainties in $P - E$, please note the different y-axis scales). Additionally, the uncertainty distribution of the higher emission scenarios is usually found within the uncertainty range of the low emission scenarios, pointing towards a more certain signal. However, there are often huge differences regarding the median response and the location and shape of the uncertainty distribution of a particular emission scenario with respect to other emission scenarios. This is especially evident when comparing low to high emission scenarios. Most prominently, for many regions (e.g. WNA, CAM, MED, WAS, CAS, please see Table 2 for more information on the acronyms) the uncertainty distribution of the RCP6.0 or RCP8.5 scenario is located mainly within the lowest tercile of the RCP2.6 scenario, leading to a dryer response in P ($P - E$) with global warming for high emission scenarios. This finding is, however, reversed in a few other regions (especially NAU and for $P - E$, to a certain extent also in ALA, EAF, SEA). The shapes of the uncertainty distributions for both P and $P - E$ are also different between regions and emission scenarios. While the distributions for the low emission scenarios are, in most cases, unimodal, there are bimodal distributions in some (e.g. for P: NEU, WAS, CAS, for $P - E$: MED, EAS, SEA) and even multimodal distributions in a few other regions (e.g. for P: AMZ, NAS, for $P - E$: CNA, WAF, EAF) for the higher emissions scenerios (especially RCP8.5). Please note, however, that not all models provide data for the RCP2.6 and

RCP6.0 emission scenarios, which might also cause differences between those and the other scenarios (see Table 1 for more information).

3.2 Sources of uncertainty

Besides the scenario uncertainty we also introduced two other sources of uncertainty in Sec. 2, the internal variability and the model uncertainty which contribute to the total uncertainty. Here we assess the fraction of uncertainty which each source contributes to the total uncertainty. We follow the approach of Hawkins and Sutton (2009), which was also adapted in Orlowsky and Seneviratne (2013). Therefore we compare (i) the average over the variances of the uncertainty distributions of each model under each emission scenario (internal uncertainty, , σ_i), (ii) the average of the variances of scenario-specific uncertainty distributions of each model (model uncertainty, , σ_m) and (iii) the variance of the averages of all uncertainty distributions within a specific scenario (scenario uncertainty, σ_s). Even though this approach of attributing uncertainties is very simplistic (see section 2), it provides basic information on the composition of different uncertainty sources within the total uncertainty. The percentage of the total uncertainty that stems from a particular source is illustrated for both P and $P - E$ in Fig. 9. For all SREX regions, there is generally no single source of uncertainty contributing more than 80% to the total uncertainty. However, in most regions the largest source of uncertainty stems from model uncertainty, which is contributing up to ca. 3/4 of the total uncertainty in some regions and is especially large in most northern high latitude regions (CGI, NEU, NAS, except ALA), possibly related to the relatively large and certain scaling response in these regions. Internal variability contributes between $20 - 40\%$ to the total uncertainty, with highest values found in particular for several regions adjacent to the Pacific Ocean (e.g. WNA, SEA, SAU), probably related to previously assessed modes of inter-annual variability, such as the Quasibiennial Oscillation, QBO (e.g. Labat et al., 2004). Internal variability seems to be rather low in other tropical to subtropical regions, such as AMZ, EAS and NAU for P and NEB, WAF and SAS for $P - E$. Scenario uncertainty contributes between ca. $5 - 30\%$ to the total uncertainty with those regions reaching highest values that have differing locations of the uncertainty distributions between low and high emission scenarios as shown in Fig. 7 and Fig. 8 (e.g. WNA, ENA, CAM, MED, WAS, CAS, and NAU). Differences in P between emission scenarios are further not solely caused by varying radiative forcing due to differing greenhouse gas emissions, but also due to differences in the black carbon forcing (Pendergrass and Hartmann, 2012). It is further interesting to note that scenario uncertainty is generally lower and internal variability generally larger for $P - E$ when compared to P. However, even though the scenario uncertainty is the overall weakest source of uncertainty in most regions, it is by no means negligible. Please note, again, that the scenario uncertainty interferes especially with the rather large model uncertainty and we do not account for such relationships in this approach.

4 Application: Assessing warming degree limits

As agreed in COP21, the increase in global mean temperature should be limited to values well below the previously set goal of $2K$, preferably to not more than $1.5K$ above pre-industrial conditions. However, global mean temperature is an abstract value and provides no information about direct implications at regional scales and with respect to other climate variables such

as regional, mean-annual P or $P - E$. The framework developed within this study enables us to directly assess the regional response of P and $P - E$ to these targets and to study differences between them. Using the uncertainty distributions of each SREX region and scaling them to either $1.5K$ or $2K$ (as illustrated in Fig. 10 for P and in Fig. 11 for $P - E$) allows us to study differences both in the median response as well as in the tails of the distribution. It is, however, naturally evident that

5   in regions with a weak median scaling response the difference between the warming degree targets is small regarding the median response itself (e.g. AMZ, CEU, WAS), whereas in regions with a stronger median scaling (e.g. ALA, NEU, NAS) an additional $0.5K$ warming could lead to substantial differences. Nonetheless, even though the difference in the median might be small, the difference in the tails of the uncertainty distributions are in most cases significant and stress an increased risk of experiencing strong changes in P and $P - E$. As an example for the Mediterranean region (MED): The median responses of

10   P to $1.5K$ global warming vs $2K$ global warming are not strongly different, while there are stronger differences at the tails, showing that the 1.5K limit would avoid a decrease of P of more than 20%, which can on the other hand not be excluded with the 2K limit. This behavior is even more evident for $P - E$ and also occurs in regions with almost no median response (e.g. CEU). Also the irregularity of the distribution further amplifies this behavior in some regions. In summary, opting for a low warming degree target (such as $1.5K$) might just slightly influence the mean response but could substantially reduce the risk

15   of experiencing more extreme changes in regional P and $P - E$.

5   Conclusions

We developed here a framework building upon the pattern scaling approach to assess regional changes in mean-annual P and $P - E$ with respect to global mean T-increase by utilizing a comprehensive subset of climate models and considering all available emission scenarios. We further took into account internal variability from each projection by accounting for the

20   year-to-year variability of P ($P - E$). This enabled us to assess a conservative estimate of the uncertainty distribution of the scaling coefficient of P ($P - E$) to global warming at every gridpoint or within SREX regions.

Analysing maps of the median response and the responses in the 10th and 90th quantile of the gridpoint-specific uncertainty distributions showed low uncertainties and positive scaling coefficients (thereby a certain increase in P and $P - E$ with global warming) within most northern high latitude regions. Slight decreases in the median response together with large uncertainties

25   (and thereby an uncertain decrease in P and $P - E$ with global warming) are found for most subtropical regions. Uncertainties are, however, larger for estimates of $P - E$ and do hence not permit robust conclusions for many regions. Our results support previous findings of hydroclimatological changes (Greve and Seneviratne, 2015), but provide a new, probabilistic and rigorous perspective on the assessment of uncertainties in regional hydroclimatological changes under conditions of ongoing global warming and extend the wealth of studies investigating pattern scaling approaches of climate variables (Tebaldi and Arblaster,

30   2014; Herger et al., 2015).

Assessing scenario-specific uncertainty distributions revealed strong regional differences between different emission scenarios. It is evident that weaker climate change signals within the low-emission scenarios (RCP2.6, RCP4.5) lead to high uncertainties in the estimation of scaling coefficients. A very likely change in regional P only emerges under high emission

scenarios (RCP6.0, RCP8.5) and is even less likely to occur for $P - E$. In some regions low emission scenarios further show likely opposite changes compared to changes identified in higher emission scenarios (both switching from a likely wetting to a very likely drying in parts of MED, or from a likely drying to no change in P or even a likely wetting in $P - E$ in NAU). A closer look a the uncertainty distributions shows large differences both in location and shape across regions and emission scenario. However, in most cases higher emission scenarios point towards a dryer response than low emission scenarios (with a few regions showing, however, the opposite behavior).

This led us to the analysis of the relative contribution of single sources of uncertainty to the overall uncertainty. It is shown that model uncertainty is largest in most regions, but is, however, in no case contributing more than 80% to the overall uncertainty. It is therefore important that both internal variability and scenario uncertainty are considered as well in order to get a complete picture of the total uncertainty. Comparing mean annual P and $P - E$ shows that scenario uncertainty is generally lower and internal variability generally larger for $P - E$.

We further assessed the implications of different warming degree limits on changes in regional P and $P - E$. At the COP21, most nations agreed to limit the increase in global mean temperature to values well below the previously set goal of $2K$ and to consider limiting warming to not more than $1.5K$ above pre-industrial conditions. Comparing these two targets reveals naturally little differences in the mean response in regions where the mean response is small anyway. However, since uncertainties are large, especially for $P - E$, there is a nonlinear increase in the risk of experiencing more extreme changes. Therefore, opting for a low warming degree target (such as $1.5K$) might just slightly influence the mean response but could substantially reduce the risk of experiencing extreme changes in regional P and $P - E$. This means that even though the discussion about the implications of $1.5K$ vs. $2K$ global warming might be moot for the mean response, it is, given the underlying, large uncertainties of climate projections, absolutely necessary to more closely investigate the potentially large increase in the risk of experiencing extreme change. This is especially important in order to enable robust decision making to ensure adequate development pathways and to avoid the risk of maladaptation; in the specific case of changes in mean-annual P and $P - E$ this is e.g. of high relevance for water resources managers and farmers.

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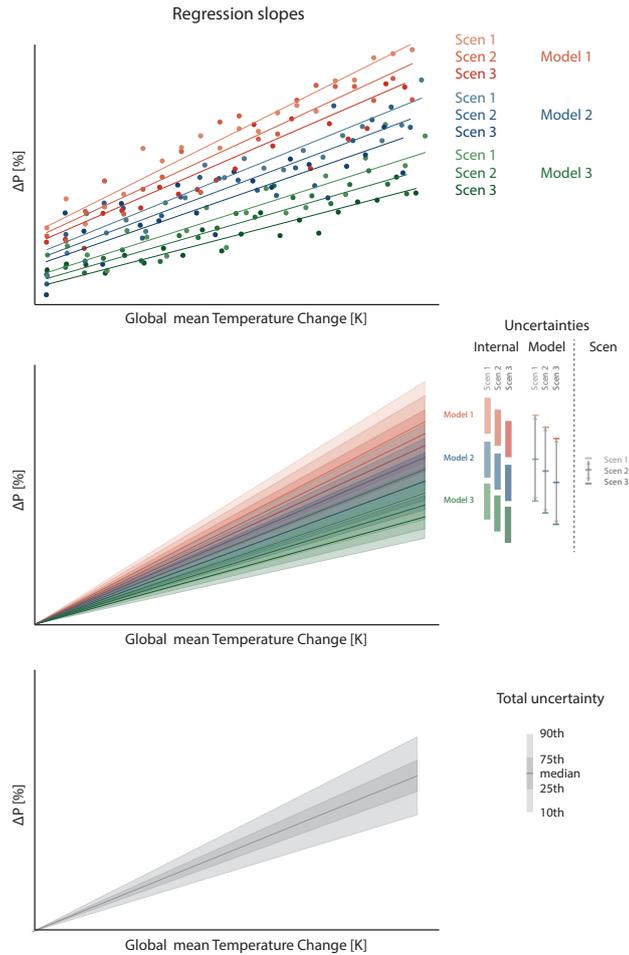
Figure 1. Conceptual illustration of deriving the uncertainty distribution of the scaling coefficient of P with respect to global warming from multiple models forced by multiple emissions scenarios. (Top) For each model under each scenario we regress the relative change in mean-annual P (ΔP, with respect to a baseline period) against global mean T. We thereby obtain the regression slope which is the scaling coefficient of P to global warming. The year-to-year variability further causes the estimate of the slope to be uncertain. We account for this uncertainty by numerically estimating the uncertainty distribution of each model- and scenario-specific regression slope through resampling the residuals in a bootstrapping approach. (Middle) This uncertainty is associated with every model run and represents the internal variability. The average of the uncertainties stemming from the range of all individual models within a certain scenario represents the model uncertainty and the uncertainty associated with the range of all scenario-specific multi-model means represents the scenario uncertainty. The uncertainty distribution is illustrated here as a function of global temperature increase. (Bottom) The Total uncertainty combines all sources of uncertainty and provides a conservative estimate of regional ΔP as a function of global warming, that can be used to assess either the median response or to study changes in any other quantile of the uncertainty distribution.
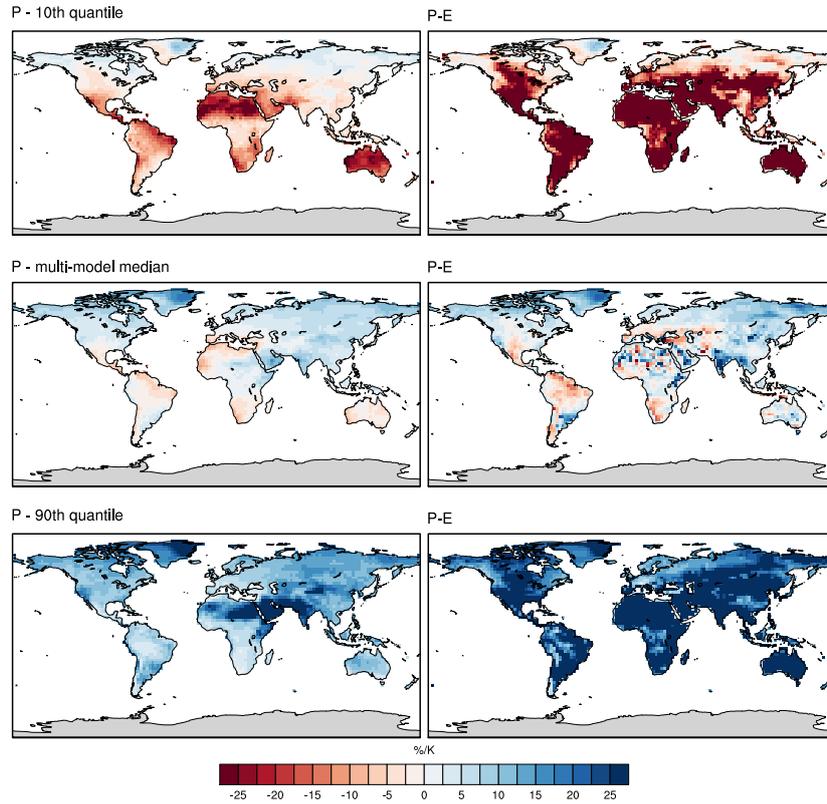
Figure 2. Median (middle), 10th (top) and 90th quantile (bottom) of the sensitivity of P (left) and $P - E$ (right) to changes in global mean temperature $[\%/K]$. A total of 14 CMIP5 models and all scenarios (RCP2.6, RCP4.5, RCP6.0, RCP8.5) are considered.
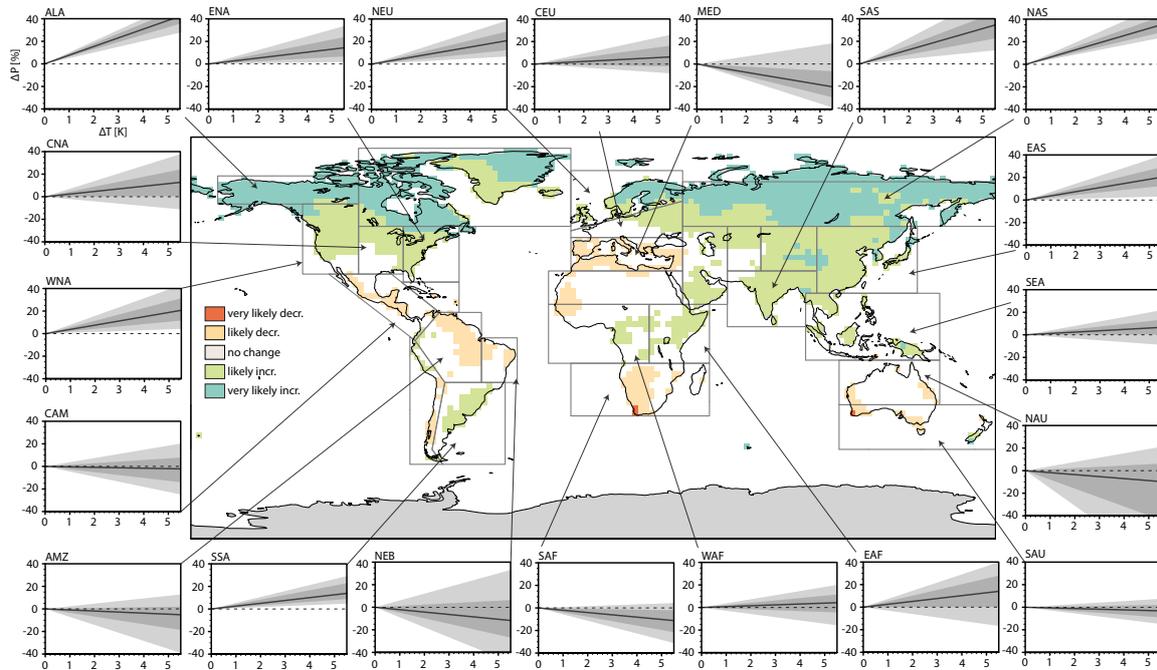
Figure 3. Conceptual summary of the probability that the slope of P is negatively/positively different from zero considering all climate models and all scenarios. Panel plots illustrates the uncertainty distribution of the sensitivity of P to global temperature change as a function of global mean temperature change averaged for each SREX regions outlined in the map (the shading in each panel plot corresponds to those illustrated in Fig. 1).
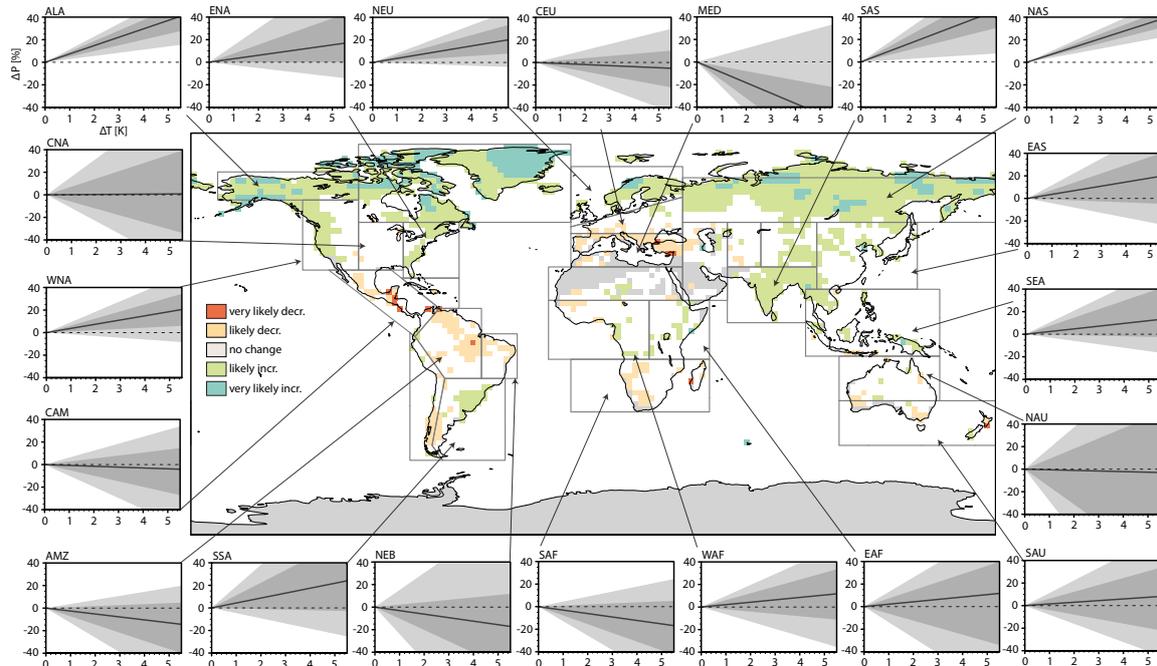
Figure 4. Conceptual summary of the probability that the slope of $P - E$ is negatively/positively different from zero considering all climate models and all scenarios. Panel plots show the uncertainty distribution of the sensitivity of $P - E$ to global temperature change as a function of global mean temperature change averaged for each SREX regions outlined in the map (the shading in each panel plot corresponds to those illustrated in Fig. 1).
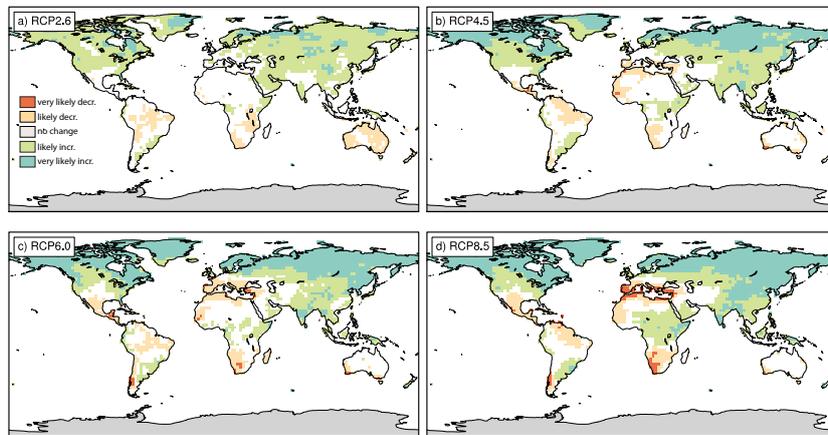
Figure 5. Conceptual summary of the probability that the slope of P is negatively/positively different from zero considering all climate models and a) the RCP2.6, b) the RCP4.5, c) RCP6.0 and d) RCP8.5 emission scenario only. See Fig. 3 for comparison.
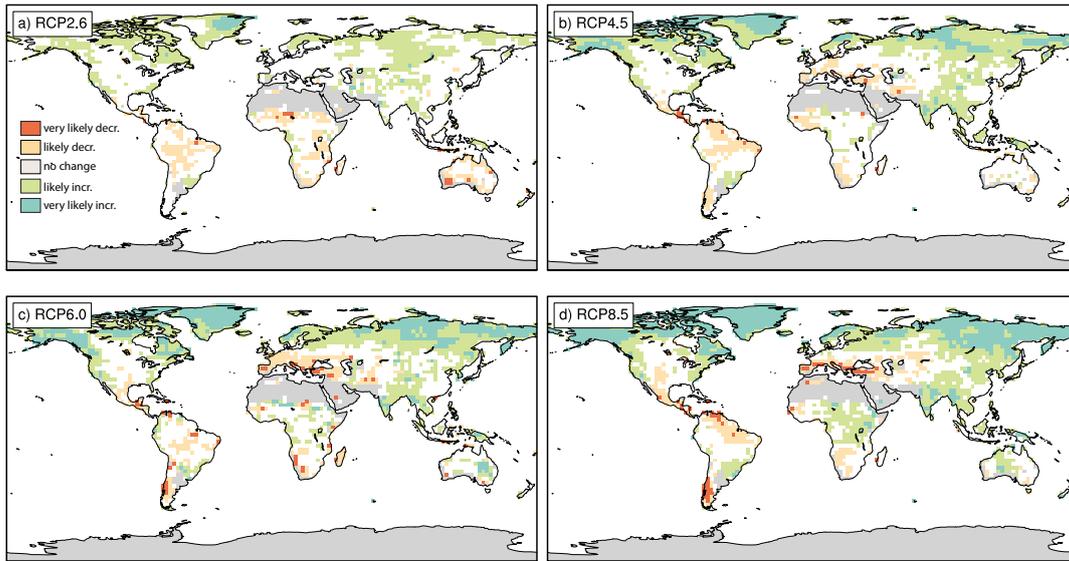
Figure 6. Conceptual summary of the probability that the slope of $P - E$ is negatively/positively different from zero considering all climate models and a) the RCP2.6, b) the RCP4.5, c) RCP6.0 and d) RCP8.5 emission scenario only. See Fig. 4 for comparison.
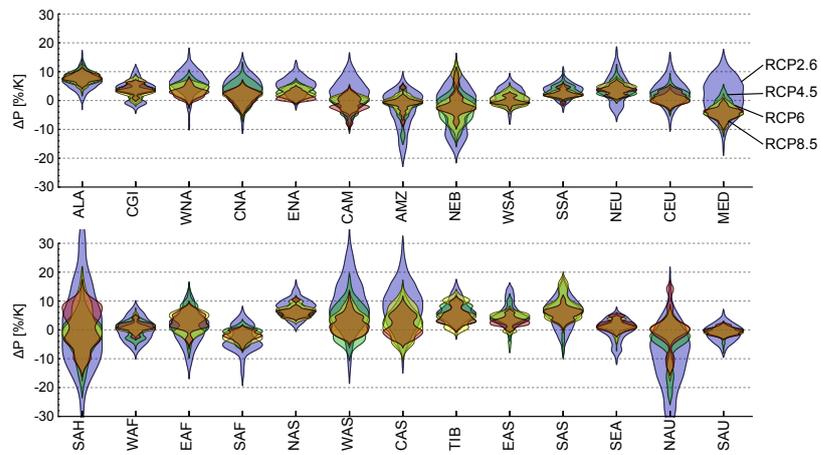
Figure 7. Uncertainty distributions (shown as violin plots) of the sensitivity of P to global mean temperature change for each emission scenario averaged over all SREX regions (as outlined in Table 2 and Fig. 3).
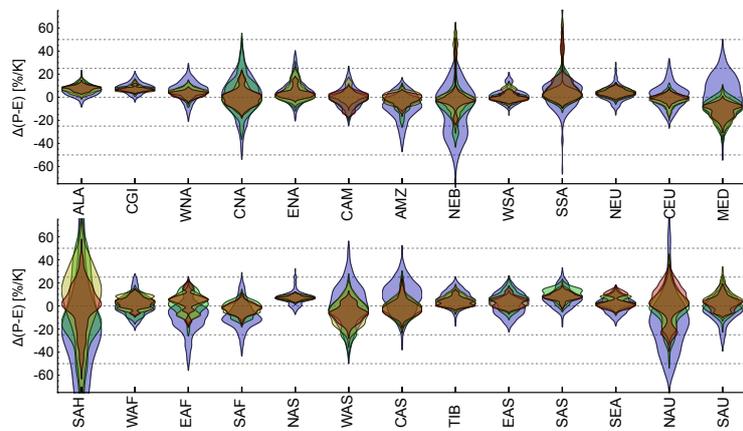
Figure 8. Uncertainty distributions (shown as violin plots) of the sensitivity of $P - E$ to global mean temperature change for each emission scenario averaged over all SREX regions (as outlined in Table 2 and Fig. 4). It is important to note that the data considered to estimate the area average is scarce in some regions (e.g. SAH). Please also refer to Fig. 7 for more information.
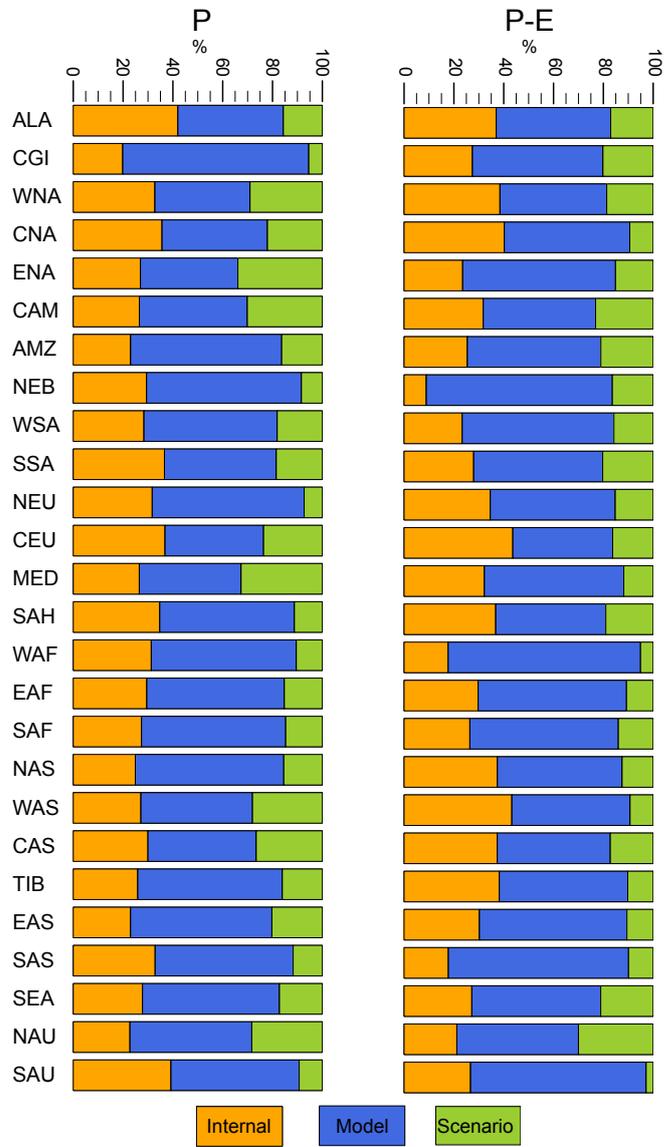
Figure 9. Sources of uncertainty in the sensitivity of P (left) and $P - E$ (right) to global mean temperature change averaged over each SREX regions as outlined in Fig. 3.
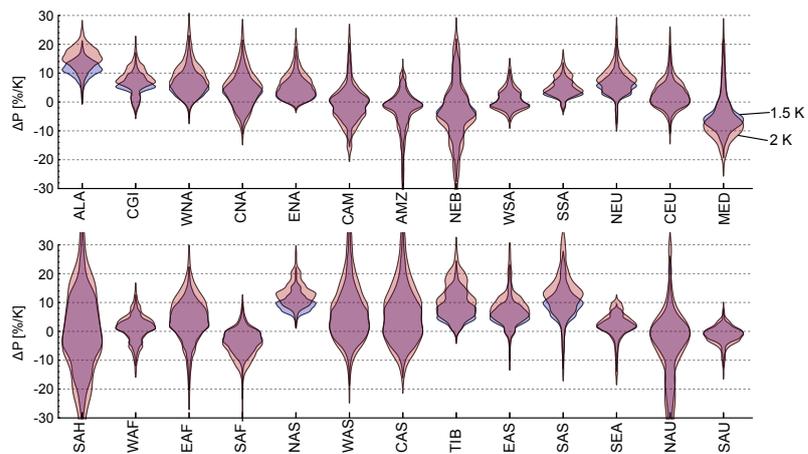
Figure 10. Uncertainty distributions (shown as violin plots) of the sensitivity of P to global mean temperature change for two different degrees of global mean temperature change, which correspond to the widely used warming degree limits of $1.5K$ and $2K$. The estimates are averaged over all SREX regions (as outlined in Fig. 3 and Table 2).
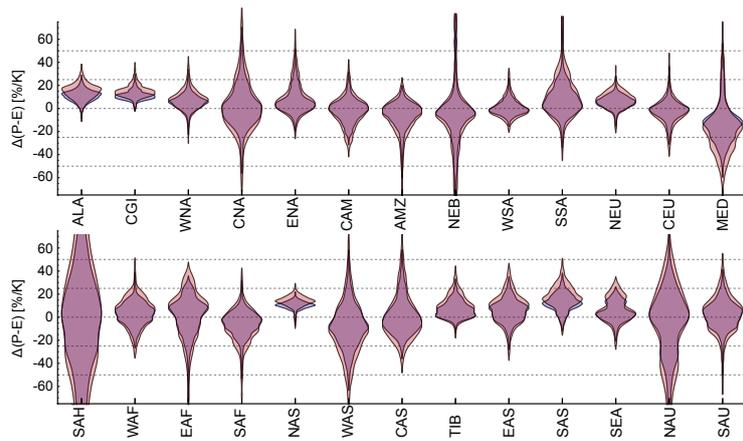
Figure 11. Uncertainty distributions (shown as violin plots) of the sensitivity of $P - E$ to global mean temperature change for two different degrees of global mean temperature change, which correspond to the widely used warming degree limits of $1.5K$ and $2K$. The estimates are averaged over all SREX regions (as outlined in Fig. 3 and Table 2). It is important to note that the data considered to estimate the area average is scarce in some regions (e.g. SAH). Please also refer to Fig. 10 for more information.

Table 1. List of models and availability under each emission scenario.

Model	RCP2.6	RCP4.5	RCP6.0	RCP8.5
ACCESS1-3		x		x
bcc-csm1-1	x	x	x	x
CanESM2	x	x		x
CESM1-BGC		x		x
CMCC-CMS		x		x
CNRM-CM5	x	x		x
CSIRO-Mk3-6-0	x	x	x	x
FGOALS-g2	x	x		x
GISS-E2-R	x	x	x	x
HadGEM2-ES	x	x		x
IPSL-CM5A-MR	x	x	x	x
MIROC5	x	x	x	x
MRI-GCM3	x	x	x	x
NorESM1-M	x	x	x	x

Table 2. List of Acronyms for all 26 SREX-regions (Seneviratne et al., 2012)

Region	SREX-Acronym
Alaska/NW Canada	ALA
Eastern Canada/Greenland/Iceland	CGI
Western North America	WNA
Central North America	CNA
Eastern North America	ENA
Central America/Mexico	CAM
Amazon	AMZ
NE Brazil	NEB
West Coast South America	WSA
South- Eastern South America	SSA
Northern Europe	NEU
Central Europe	CEU
Southern Europe/the Mediterranean	MED
Sahara	SAH
Western Africa	WAF
Eastern Africa	EAF
Southern Africa	SAF
Northern Asia	NAS
Western Asia	WAS
Central Asia	CAS
Tibetan Plateau	TIB
Eastern Asia	EAS
Southern Asia	SAS
Southeast Asia	SEA
Northern Australia	NAU
Southern Australia/New Zealand	SAU