

**Soil C management,
large-scale Earth
system modelling**

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Soil carbon management in large-scale Earth system modelling: implications for crop yields and nitrogen leaching

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We explore cropland management alternatives and the effect these can have on future C and N pools and fluxes using the land use-enabled dynamic vegetation model LPJ-GUESS. Simulated crop production, cropland carbon storage, carbon sequestration and nitrogen leaching from croplands are evaluated and discussed. Compared to the version of LPJ-GUESS that does not include land use dynamics, estimates of soil carbon stocks and nitrogen leaching from terrestrial to aquatic ecosystems were improved.

We explore trade-offs between important ecosystem services that can be provided from agricultural fields such as crop yields, retention of nitrogen and carbon storage. These trade-offs are evaluated for current land use and climate and further explored for future conditions within the two future climate change scenarios, RCP 2.6 and 8.5. Our results show that the potential for carbon sequestration due to typical cropland management practices such as no-till and cover-crops proposed in literature is not realised, globally or over larger climatic regions. Our results highlight important considerations to be made when modelling C–N interactions in agricultural ecosystems under future environmental change, and the effects these have on terrestrial biogeochemical cycles.

1 Introduction

Growing population along with rapidly changing dietary preferences pose a key economical and environmental challenge of this century (Gerland et al., 2014; Hertel, 2015). According to estimates made by the United Nations Food and Agriculture Organisation (FAO), food production will need to be doubled by 2050 in order to meet the global food demand (FAO, 2008). Since the beginning of the 20th century there has been an increase in crop yields and overall production, especially since the 1950s (Steffen et al., 2015), as a result of agricultural intensification driven by substantial

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advances in agricultural practices and technology, improved crop varieties and an increased application of N and P fertiliser (Evans, 1999; Spano et al., 2003). In addition, agricultural land area globally has expanded, with presently around 35 % of the total land surface being covered by cropland and pastures (Ramankutty et al., 2008). A further increase in cropland area by at most around 5 Mha (> 3 times the present area; Eitelberg et al., 2015) would be possible, but a range of societal and political pressures on land resources may limit the conversion of additional land area to agricultural production in many regions. Increasing yields on existing cropland would reduce pressure for further land conversion. Yield increases may be achieved through further development of high-yielding varieties or through further improvements in the efficiency of agricultural practices, the latter especially in regions where gaps between actual and potential yields are large (Licker et al., 2010; Mueller et al., 2014). The enhanced input of nitrogen (N) into ecosystems, jointly with other technical developments, has played a major role in the large increase in agricultural productivity over the last 50 years, often termed the “green revolution”. However, the associated environmental effects have often been detrimental, with negative impacts on biodiversity and water quality, and substantial emissions of N trace gases that affect air quality and climate, such as nitrous oxide (N_2O), a potent greenhouse gas (Galloway et al., 2004; Rockstrom et al., 2009; Tilman et al., 2002; Vitousek et al., 1997). A large fraction of the N_2O emitted to the atmosphere today originates from terrestrial sources, mostly from fertiliser use on agricultural soils (Zaehle et al., 2011; Park et al., 2012; Ciais et al., 2013). Fertiliser use also promotes nitrate leaching which causes eutrophication and algal blooms in watersheds and coastal seas, with follow-on effects such as loss of fish populations and recreational value, and health risks through contamination of drinking water (Cameron et al., 2013). Even in Europe, where environmental regulations are relatively advanced, around 70 % of the population live in areas where the levels of nitrate in drinking water either exceed the recommended value (ca. 20 % of the population) or have reached at least half this level (ca. 50 % of the population; Grizetti, 2011).

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Previous studies of the effects of agriculture on global biogeochemical cycles have typically focused on the largest immediate impacts, like the carbon losses following deforestation (e.g. Ciais et al., 2013; Houghton et al., 2012; Le Quéré et al., 2015). It is estimated that over the last 150–200 years, the conversion of natural to managed ecosystems, especially croplands, has released ca. 180 Pg carbon (C, current rate is $\sim 1 \text{ Pg C yr}^{-1}$) from the terrestrial biosphere to the atmosphere by disturbing soils and through the harvesting and burning of biomass (Le Quéré et al., 2014). This sum is equivalent to around a third of the anthropogenic CO_2 concentration in the atmosphere today. However, the land-use related carbon flux is one of the most uncertain terms in the global carbon budget (Ciais et al., 2013; Le Quéré et al., 2015), and studies with dynamic vegetation models (DVMs) incorporating representations of land-use change (LUC) have shown that the actual estimate is highly dependent on the management practices assumed in the model (Bondeau et al., 2007; Levis et al., 2014; Lindeskog et al., 2013; Pugh et al., 2015).

Available knowledge on the effects of interactions between nitrogen and carbon cycles in terrestrial ecosystems is largely based on simulations with DVMs representing potential natural vegetation (e.g., Thornton et al., 2009; Zaehle and Dalmonech, 2011; Smith et al., 2014). The results obtained with these models suggest that soil N processes governing plant available nitrogen can constrain vegetation growth and the strength of the terrestrial carbon sink (e.g., Zaehle et al., 2011; Wårlind et al., 2014). Only two global modelling frameworks have been put forward with both detailed cropland ecosystem functioning and coupled C–N cycling simulated in a consistent fashion (Arora, 2003; Drewniak et al., 2013). No study has applied such a model at global scale to investigate joint impacts of environmental change and land management on associated changes in agricultural yields, water pollution and carbon balance.

The production of food and the protection of the environment often require conflicting strategies. Compared to forests, agricultural lands have lower carbon sequestration rates and enhanced nitrogen leaching, and agricultural production is hence done at the expense of other ecosystem services that those lands might otherwise provide; inten-

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sifying agricultural production might have further detrimental effects on these ecosystems (Tilman et al., 2002). At the same time, the world's population needs to be fed, and a more nitrogen-intensive agricultural system with higher productivity may result in a lower overall area in use for agriculture, leaving aside a greater area that can be devoted to the provision of other ecosystem services. This debate, often termed land-sparing vs. land-sharing, is currently a matter of great scientific and political debate (Phalan et al., 2011). These trade-offs between agricultural production on the one hand and carbon sequestration and reduction of nitrogen leaching on the other have given rise to a number of mitigation strategies in agricultural practice that have only a limited impact on production but contribute to other ecosystem services.

From an ecosystem carbon-pool-size perspective, the largest effect on terrestrial carbon storage through agricultural management practices is in fact induced through harvest (Smith et al., 2012; Pugh et al., 2015), as this removes a considerable amount of biomass each year from the ecosystem. Harvest is, however, the fundamental purpose of croplands and pastures, and is rarely discussed as an element of a management strategy targeting enhanced environmental value. Management practices need to move into focus in global-scale modeling, and some of the key interventions are reviewed below. One management option related to harvest is residue-removal after harvest (Lal and Bruce, 1999). Removing residues for use in bio-fuel production is an appealing measure, as making multiple use of the existing croplands may be seen as a win-win situation (Lal, 2004b; Smith et al., 2012). However, not incorporating residues into soils results in their becoming drained of soil organic carbon (SOC); which retains water and nutrients and thus affects the soil fertility (Lal, 2004b; Smith et al., 2012). Another practice that is often debated is tillage (Lal, 2004a, 2008). Different forms of tillage have been used for centuries to promote the release of nutrients from the soil organic matter (SOM) for uptake by crops. However, the aeration of the soil associated with the mechanical disturbance of the soil profile increases heterotrophic respiration (R_h), and thus enhances soil C losses to the atmosphere (Chatskikh et al., 2009; Lal, 2004a). No-till management has gained popularity as a potential climate change mit-

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igation measure, as it can prevent management-related losses of soil C stocks (Lal, 2004a). But while no-till is expected to favour carbon retention by agricultural fields, the strength and persistence of any sink is debated (Lal, 2004a; Schlesinger, 2000). In a review on soil C sequestration, comparing conventional and no-till, Baker et al. (2007) found that the top soil in no-till treatments did contain more C, but the C density below the top soil layer in the conventionally tilled fields was higher, and there was no significant difference in overall C densities between the two treatments. Moreover, the conversion of N to plant-available forms is reduced in untilled soils and can thus lead to lower crop productivity, which could in the long run decrease the soil's ability to store water and nutrients because the reduced release is partly counterbalanced by a reduced input of new organic material (Lal, 2004a).

Timing of farmer practices such as sowing and harvest, fertilisation or irrigation are crucial for calculation of yields and – from the fertilisation perspective in particular – N leaching, but these practices are unlikely to affect carbon cycling on a global scale (Pugh et al., 2015). A major challenge is therefore to identify practices that increase soil C sequestration while maintaining or increasing current yields and at the same time without detrimental side-effects on ecosystems.

In this study we employ the land-use enabled version of a global DVM, LPJ-GUESS (Lindeskog et al., 2013; Olin et al., 2015), to explore and quantify the effectuality of alternative management strategies targeting mitigation of the negative effects of agriculture on carbon and nitrogen cycles. To this end, we extended the model to include N dynamics for crops, and the response of different N-application rates was evaluated on local to regional scale. Management options considered are tillage, cover-crops and manure application. We quantify management effects on soil carbon pools, yields and nitrogen losses through leaching from croplands and evaluate the model globally and for a representative range of climatic regions. In addition, the persistence and direction of these effects under future climate change scenarios are explored.

2 Materials and methods

2.1 LPJ-GUESS

LPJ-GUESS (Smith et al., 2014) is a DVM that simulates dynamic vegetation response to climate, atmospheric CO₂ levels ([CO₂]) and N input through competition for light, N, and water on a daily time step. Vegetation is represented by plant functional types (PFTs) that differ in their growth form, phenology, life-history strategy, distributional temperature limits and N requirements. C3 and C4 photosynthetic pathways are discriminated for grasses. Leaf-level net photosynthesis is calculated following a Farquhar-type approach, modified by Collatz et al. (1991, 1992) and scaled to the canopy following Haxeltine and Prentice (1996). Canopy conductance of water vapor and respiration of plant compartments other than leaves follows (Sitch et al., 2003). For potential natural vegetation, carbon allocation and stand dynamics, based on competition among age-classes of trees co-occurring in a number (here 5) of replicated patches in each grid cell, are modelled on a yearly time-step (Smith et al., 2001; Hickler et al., 2004). Disturbance by wildfire and other events such as storms are accounted for. Details of the representation of soil and plant physiological and growth processes are provided in Smith et al. (2001, 2014); Olin et al. (2015).

Soil C–N dynamics in LPJ-GUESS are based on the CENTURY model (Parton et al., 1993) in which SOM and litter are represented by 11 pools that differ in their C to N ratios (C : N), which are dynamic within prescribed limits (Smith et al., 2014). Mobilisation of mineral N is the result of heterotrophic decay and respiration which depends on the C : N and decay rates (K_d) of the SOM pools. Values of K_d are dynamic and vary between these pools, and are also modified by factors related to temperature and water content of the soil (Smith et al., 2014). Immobilisation of mineral N occurs when the C : N ratio of transferred SOM is larger than that of the receiving pool. Both organic and mineral N leaching are represented in the model and are related to percolation; for organic leaching there is also a dependency on soil silt and clay fractions.

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The present study uses the managed land version of the model (Lindeskog et al., 2013; Olin et al., 2015). Crops are represented in the model by crop functional types (CFTs), which differ in their temperature requirements for survival, heat requirements for growth and their C allocation patterns. Land-use and land-cover change are modelled on a yearly basis (Lindeskog et al., 2013) based on the externally supplied fractional area change within a grid-cell. Sowing dates are determined dynamically in the model based on the prevailing climate in the grid cell (Lindeskog et al., 2013; Waha et al., 2011), and crops are harvested when specified heat sum requirements are fulfilled. The crops are allowed to adapt to the local conditions by adjusting the heat sum requirements to the historic climate (Lindeskog et al., 2013), reflecting a difference in varieties of a given crop grown in different climatic zones.

The allocation of C and N for the CFTs in the C-N version of LPJ-GUESS (Olin et al., 2015) is done on a daily time-step. C allocation depends on the plant's development stage which in turn is based on temperature and day length, and follows Penning de Vries et al. (1989), as described in detail for winter and spring wheat in Olin et al. (2015), and for maize in Table A1. N requirements for the plant vary during the growing period. This is reflected in the model by applying fertilisers proportionally at different developmental stages; see Appendix for more information on the timing of N fertiliser application for different CFTs.

At present, the C-N version of LPJ-GUESS is limited to three CFTs that are based on wheat and maize growth characteristics: a C3 crop with dynamic selection between spring and autumn sowing (presented here as winter wheat, WW), a C3 crop with sowing enforced in spring (spring wheat, SW) and a C4 crop (maize, MA). Allocation for SW and WW is described in Olin et al. (2015), MA-specific allocation parameters are listed in Tables A1–A2. For comparison with yield data, we adopt these three types to represent the entire spectrum of crops grown globally. In particular, wheat and rapeseed that have spring and autumn sown varieties were simulated as WW, whereas other C3 crops (beans, rice, tubers, etc.) were modelled as SW, since these are typically spring sown. Sorghum and millet were modelled as MA.

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LPJ-GUESS has been evaluated against a range of experimental and observational data types, e.g. CO₂-fertilisation experiments (Olin et al., 2015; Smith et al., 2014), ecosystem dynamics (Smith et al., 2014), vegetation seasonality (Lindeskog et al., 2013) and C fluxes at various scales (Ahlström et al., 2012; Piao et al., 2013; Wraneby et al., 2008). In Olin et al. (2015), the growth response to N fertiliser application on site scale (under ambient and elevated CO₂) and over a larger region (western Europe) was evaluated.

2.1.1 Cropland management

The cropland management options implemented in LPJ-GUESS are sowing, irrigation, tillage, N-application, cover crops and residue management. The latter four options are relevant for this study and will be described below.

Tillage

Tillage is implemented using a tillage factor (f_T) which affects K_d for selected SOM pools on croplands. Two tillage routines were implemented: moderate tillage where f_T affects the surface microbial pool and humus, and the microbial and slow turnover pool of the soil; and full tillage in which K_d for the metabolic and structural surface pools and the passive and metabolic pools of the soil are also affected. The two tillage levels are not intended to represent different tillage practices, but rather to span uncertainties in the overall effect of tillage on soil respiration rates. The value of f_T (1.94) is taken from Chatskikh et al. (2009), and modifies K_d ($K'_d = f_T K_d$) throughout the year.

N application

Fertilisers are applied as mineral N (Olin et al., 2015). The timing of fertiliser applications in the model roughly coincides with the crucial developmental periods of plants being applied at the development stages (DS) 0, 0.5 and 0.9 (Olin et al., 2015) in CFT-specific amounts listed in Table A2.

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Here we have extended the available N fertiliser application management options to also include manure application in the first of the three events (DS = 0, sowing). The manure is derived from the mineral N application and is applied as an increase in the metabolic and structural SOM pools with a C : N of 30, which has been chosen to represent the C and N content in manure from sources ranging from poultry waste (C : N ~ 15) to straw-rich manure from livestock (C : N \gtrsim 40), (Nieder and Benbi, 2008). As the metabolic and structural SOM pools have different turnover (decomposition) rates the manure-derived N becomes available for an extended period in the soil.

Cover crops

Cover crops are intermediate crops that are grown in-between the main agricultural growing seasons. This can occur either as a fallow that stretches over the subsequent growing season or within a year (Follett, 2001). A common practice is to sow N-fixing plants such as legumes as cover crops, but grasses are also used. If the cover crop is not harvested but, for example, ploughed in, some of the captured or retained nutrients, as well as the carbon content of the crop biomass, are retained in the soil, enhancing nutrient availability.

In our implementation, cover crops are grown in-between two growing periods of the generic main crop used if the crop-free period is longer than 15 days. At the time of sowing of the subsequent main crop, the cover crop biomass is added to the soil litter pool. C and N allocation of the cover crop is done daily, with a leaf-to-root ratio that depends on the plant water status. In case of water stress, a functional balance response is introduced and allocation to roots increases relative to leaves. Cover crops are modelled as grasses, being “planted” with an initial C mass of 0.01 kg C m^{-2} and N mass that is based on the C : N_{min} value for grasses (C : N_{min} = 16). Symbiotic N-fixation, such as in legumes – common as cover crops in temperate latitudes – is not yet implemented.

Residue removal

A measure to increase the soil fertility and decrease the water loss, in particular in arid areas, is to leave the residues on the ground after harvest (Lal, 2004a; Smith et al., 2012). This practice is represented in our model by removing only a fraction (default set to 75 %) of the biomass remaining following harvest, thus leaving the rest as litter which enters the normal soil-decomposition calculations.

2.2 Experimental setups

Our study is divided into two parts. In the first part we test the ability of LPJ-GUESS to simulate present-day soil C and yield response to management by comparing simulated results with datasets of soil C in crop fields, potential C sequestration after a change in management, and global yield statistics. In the second part of the study, we investigate the effectuality of alternative crop management options described in Sect. 2.1.1 for mitigating climate change through increased carbon retention in cropland soils. The sensitivity of soil carbon sequestration to these management options is first studied for present-day climate conditions, assessing relative effects in different regions. Subsequently, we force the model with General Circulation Model (GCM)-simulated climate under a 21st century future climate projection to investigate combined effects of future changes in multiple ecosystem drivers on cropland ecosystem carbon balance.

For simulation over the recent historic period (1901–2006), gridded monthly mean observations from CRU (precipitation, air temperature and cloudiness, Mitchell and Jones, 2005) were used. For the future climate simulations, monthly climate data were adopted from four GCMs (CCSM4, Gent et al., 2011; MPI-ESM-LR, e.g. Stevens et al., 2013; IPSL-CM5A-LR, Dufresne et al., 2013; and HadGEM2-ES, Collins et al., 2011) from the CMIP5 data set (Taylor et al., 2011) and were bias corrected against CRU for monthly means over the period from 1961–1990, as described in (Ahlström et al., 2013). Climate data for the contrasting RCPs 2.6 and 8.5 radiative forcing projections

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(Moss et al., 2010) were selected based on the availability of projections of future N-fertilisation.

For all simulations soil C and N pools were initialised using a 500 year “spin-up” using atmospheric [CO₂] from the first historic year (1901 for the historical CRU-based simulations and 1850 for the CMIP5 simulations) combined with repeatedly cycled, detrended climate input using the first 30 years of the historic climate data set. For comparison with the CRU simulations described above, the spin-up in the CMIP5 simulations was set to 450 years, followed by a simulation for years 1850–1901 with dynamic climate but constant [CO₂] (using the [CO₂] for 1901).

N atmospheric deposition was provided as decadal-varying monthly averages from the ACCMIP data set (Lamarque et al., 2010) transformed to the resolution of the climate data following (Smith et al., 2014; Wårlind et al., 2014).

As N-fertiliser input for the croplands, data from (Zaehle et al., 2010a) were used for the historical time period starting from 1901 (CRU) and 1850 (CMIP5); for the future period (2006–2100), a dataset described in Stocker et al. (2013) was used, which expands on the data set from Zaehle et al. (2010a), and includes simulated future fertiliser applications from integrated assessment models (RCP 2.6, Bouwman et al., 2013; RCP 8.5 Riahi et al., 2011). In addition, a simulation using N-fertiliser information from AgGRID (Elliott et al., 2014) was performed for the comparison of yields with national statistics from the FAO. The AgGRID dataset provides a long-term mean N fertiliser input for each grid cell representing present day (approximately the year 2000). In these simulations the input from (Zaehle et al., 2010a) was used until 1990, subsequently switching over to AgGRID data.

Land cover information was adopted from (Hurtt et al., 2011), with the forested, rangeland and urban classes treated as natural land cover. During spin-up, cropland fraction was linearly increased from an assumed baseline of zero at 1750 to the first historic value (1901 for CRU and 1850 for CMIP5). The number of years for this transition (150 years for the CRU-based and 100 years for the CMIP5 simulations) was chosen to ensure that the soil C and N pools of the natural vegetation fraction of each

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mately 200 grid cells, evenly spread over the continents (Fig. B2). Some 200 of the grid cells were either in the boreal zone or not included in the climate dataset. Management practices were enabled from year 1990 and throughout the remaining simulation period. For the simulations using CRU climate input, the last 30 years of climate and [CO₂] (381 ppm), N deposition and fertiliser from the last year, were repeated until 2100, the end of the CMIP5 climate data set, in order to allow soil carbon and nitrogen pools to reach a new equilibrium after the management shift.

2.2.2 Management, global soil C and N leaching

The effect of the different management strategies considered (no-tillage, manure application, cover-crops and leaving residues) on simulated global crop yields, soil C pool size, and N leaching were tested in a factorial experiment where managements was turned on at the beginning of the simulation. The simulated yields, soil C and N-leaching were then compared with a baseline simulation (F_{std} , Table 2.1.1) with settings as in (Lindeskog et al., 2013; Smith et al., 2014; Olin et al., 2015).

To be able to compare our results with previous estimates of global soil C and N pools and N leaching from LPJ-GUESS (Smith et al., 2014), a simulation with potential natural vegetation (PNV) was also conducted. In addition, an optimised simulation set-up was created (F_{opt}), where the management from Table 1 that yielded the largest increase in soil carbon per grid cell was selected, for the CRU and CMIP5 simulations.

3 Results

3.1 Yield comparison

LPJ-GUESS wheat (C3) and maize (C4) yields were simulated using the gridded N-fertiliser dataset (Elliott et al., 2014) and compared to reported yields from FAO¹ for

¹FAOSTAT, <http://faostat3.fao.org/home/E>

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the years 1996–2005 (Fig. 1). The overall model agreement with reported wheat yields per country was good across all wheat producing countries, with a correlation coefficient of 0.73. Maize yields had a lower agreement (correlation coefficient 0.46), with simulated yields overestimating the observations for most countries that have a low maize production, (e.g., Mexico, China and many African countries; Fig. 1). However, with exception of China and Mexico, yields in high producing countries were captured well, including the largest producer, the USA.

The total simulated production (wet weight) of all agricultural crops (including cereals, tubers and pulses) of 2.7 Gt, was within 30 % of what is reported to the FAO, 3.5 Gt for the period 1996–2005 (cereals, 2.12; coarse grain, 0.93; roots & tubers, 0.28²; pulses, 0.06; oil crops, 0.11).

3.2 Simulated soil C and its response to management

Simulated soil C pools (0–1.5 m) for the selected grid cells (Sect. 2.2.1) were compared against data from soil cores from agricultural fields for the four climatic regions (Batjes, 2005). This comparison did not aim to reproduce observed C values at the individual field scale, as this would require to capture individual site meteorology as well as details on land-use history. Consequently, per-site comparison of simulated vs. observed soil C resulted in low correlations of 0.05–0.14, but the mean and spread over the climatic zones were captured by the model (Table 2).

In Fig. 2, the simulated mean soil C sequestration response to the three managements (no-till, manure and cover-crops) is compared to estimates of potential soil C sequestration from Stockmann et al. (2013) for the simulated climatic regions over the historic period (1990–2006). Besides the model's average regional response to the three management options, Fig. 2 illustrates how the soil C sequestration in response to the onset of management (here: in the year 1990, see Sect. 2.2.1) evolves over time. The simulated long term (100 years) mean soil C sequestration by using manure on tropical

²Corrected for moisture content, value from FAOSTAT, 0.68 Gt.

soils was ca. $0.001 \text{ kg C m}^{-2} \text{ yr}^{-1}$, declining to negligible levels by the end of the simulated period. For no-till, the long term mean C sequestration was $0.003 \text{ kg C m}^{-2} \text{ yr}^{-1}$ or higher for all treatments, and levelled off to ca. $0.002 \text{ kg C m}^{-2} \text{ yr}^{-1}$ by simulation year 2100. The highest mean C sequestration rates were found for manure in the humid temperate climatic regions ($0.006 \text{ kg C m}^{-2} \text{ yr}^{-1}$) and for cover-crops in the tropical humid regions ($0.008 \text{ kg C m}^{-2} \text{ yr}^{-1}$), in both cases levelling off to below $0.001 \text{ kg C m}^{-2} \text{ yr}^{-1}$ by the end of the simulation period.

3.3 Global responses to management

The simulated management options resulted in an increase in cropland soil C, for all climatic regions (Fig. 3), with the largest global increase, as expected, for the option in which the management that yielded the largest carbon sequestration in a given grid cell was chosen (F_{opt}). With the exception of no residue removal, the simulated management treatments reduced N leaching (expressed here as negative anomalies), with cover-crop resulting in the largest decline. Cover crops and no-residue removal had opposite effects on both yields and N leaching. The reduction in N leaching from cover crops ($\sim 15\%$) was accompanied by a decline in simulated global yields of 5%. The large negative effect of cover crops on simulated yields in the temperate humid climatic region is due to the implicit competition over the available N between the cover crop and the main crop, the low temperature makes the decomposition of the SOM slow and in turn the release of N more evenly spread throughout the year. The N retained in the system is locked in SOM, and not easily available for plant uptake, the opposite happens in the tropical regions and especially so for the humid tropics, where turnover of the SOM is relatively fast due to the prevailing warm and moist conditions. Leaving all the residues on the fields (no residue removal) was the only treatment that increased the modelled yields both globally and for all climatic regions, but with the environmental “cost” of an increase in N-leaching. The increase in both modelled yields and N leaching is obtained because N becomes available for plant uptake and transport over

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a longer period, and between the growing periods there is nothing growing that can take up the available nitrogen. In all treatments, the soil N pools were higher than for the standard simulation (Table 3), which is caused by the reduction in leaching and the incorporation of nitrogen in SOM.

In general, the soil C pools simulated with the managed land version of LPJ-GUESS were slightly larger than simulated with PNV (Table 3), which is due to higher C storage in pastures compared to the natural vegetation they have replaced (e.g. Central Asia and parts of the Great Plains of North America) and also in high-productivity croplands that receive high inputs of N fertilisers (e.g. Egypt and western China; results not shown here).

From the simulations of different cropland management options, the management combination that yielded the largest SOC stocks 1996–2005 was chosen for each grid cell (F_{opt}); the spatial patterns are shown in Fig. 4, F_{CC} and F_{NT} being the most dominant and with distinct differences with F_{CC} mostly in humid tropical areas and F_{NT} in subtropical and temperate regions.

Figure 5 depicts the evolution over time of the effects of implementing the different soil carbon sequestration managements for two future climate change, CO_2 and land-use change scenarios. The spread that can be seen around the simulations with CRU forcing in Fig. 5 originates from the GCM climate variability, which can be seen also during the historic period (Fig. B1b). In the scenarios of land-use change (Hurtt et al., 2011), there is a steady increase of cropland area globally, which is most extreme for RCP 2.6 (Fig. B1a). Differences between the RCP 2.6 and 8.5 cases regarding the effects of management are consistently seen only for cropland soil C storage, with values being higher for RCP8.5 compared to RCP 2.6. Manure and no-tillage did not affect calculated N leaching or yields under future conditions any more than for present-day forcing. The effect of cover crops and best carbon management for RCP 8.5 was an enhanced reduction of yields and enhanced N leaching compared to the standard model set-up.

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tions of soil N pools agree well with other estimates, the soil C pools are at the low end of generally reported global estimates (Ciais et al., 2013; Stockmann et al., 2013), which can partially be attributed to absence of wetland and permafrost processes in the current model setup (Miller and Smith, 2012; Tarnocai et al., 2009; Wania et al., 2009).

When making projections on global C pools, the information on land-use history is vital (Pongratz et al., 2014). In our simulations, the assumption that the main – natural to cropland – conversions started at 1750, could have resulted in overestimations of C stored in agricultural soils. When focusing on site data collected for croplands and grouped by four climate regions (Table 2), simulated average C pools in LPJ-GUESS were higher than observations, especially for temperate soils. It is to be expected that many of these sites, especially in temperate environments, would have been under land use for very long periods, which could well lead to lower C pools compared to our modelling assumptions. Moreover, we do not have comprehensive information on present management practices, or how these would have changed over time. Still, the among-grid-cell variations in C pools were similarly large in observed and simulated data (Table 2), suggesting that the model response is sound overall.

4.1.1 Carbon sequestration potential through cropland management

Global-scale modelling of the impacts of specific land-management options is in its infancy, but since a number of future climate and socioeconomic scenarios highlight the importance of land-based mitigation, and because of the multiple trade-offs that exist with other ecosystem services, they are of importance for future research and practical applications. Published estimates of the global carbon sequestration potential on existing cropland due to different types of cropland management range between 0.34–0.57 Pg C yr⁻¹ for present-day environmental conditions (Lal, 2004a). Our model estimate when implementing the best practice (from the perspective of maximising C storage) globally is roughly 20% of that value. Figure 3 shows a global increase of ca. 5% in soil C for the F_{opt} case (ranging from 4 to 9% between large regions),

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interactions between vegetation, soils and the abiotic environment. In a review of the potential for countries to fulfil emissions reduction obligations under the Kyoto protocol IPCC (1996); Schlesinger (2000) found only a small or even no potential for C sequestration in cropland soils, while (Powlson et al., 2014) argued that no-tillage over tillage enhances some important soil properties but has a small overall effect of total agricultural soil C.

4.2 Yields

Compared to other measures of global C flows, statistics on crop production and yields are relatively accessible, and encompass relatively long time-series, albeit with differing quality between individual countries. While yield is not a direct measure of the net primary productivity (NPP), it is a good proxy for trends and variability of carbon flows on croplands (Haberl et al., 2007) and thus relevant for the estimation of fluxes and pools on agricultural fields. From a food production perspective, Olin et al. (2015) showed that including C–N dynamics and fertiliser input significantly increased model performance compared to the C-only version of LPJ-GUESS (e.g. Rosenzweig et al., 2014) for yield modelling and responses of yields to environmental changes. This was expected, since the C-only version intentionally represents a situation not limited by nutrients. The data sets used in this study were either designed for crop modelling in the AgMIP project (Elliott et al., 2014), or for studying global flows of carbon and nitrogen (rather than yields) (Stocker et al., 2013; Zaehle et al., 2011). When using the former (Elliott et al., 2014), the model performance was significantly improved (an increase in model agreement with observed yields from (R^2) 0.25 to 0.53 for WW and from 0.1 to 0.25 for MA). However, since the AgMIP dataset lacks information on temporal variations and trends, it could not be applied to transient historical or future simulations of global yields, C and N flows. Previous studies with global models have simulated yields (e.g. PEGASUS Deryng et al., 2011, R^2 for WW = 0.22, MA = 0.39, and DayCent Stehfest et al., 2007 R^2 for WW = 0.66, MA = 0.67). Our results compare favourably with these studies for WW, but less so for MA. The C-N version of our model has not yet

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been evaluated and parameterised against observations of maize yields, and the lower degree of agreement with data was expected.

4.3 N leaching

Global estimates of N leaching from terrestrial ecosystems are uncertain (Gruber and Galloway, 2008), and the estimates with LPJ-GUESS fall well within the broad range of published annual global totals (Table 3). Only a few other global studies with DVMs (e.g. Smith et al., 2014; Stocker et al., 2013; Yang et al., 2009; Zaehle et al., 2010b) have reported N leached from terrestrial ecosystems. For models that included N-fertiliser applications, we estimated a range from 63 Tg N yr^{-1} (Yang et al., 2009)³ to 133 Tg N yr^{-1} (Stocker et al., 2013)⁴. None of these simulation studies accounted for croplands explicitly, non-harvested grasslands in Zaehle et al. (2010b) and harvested grasslands in Stocker et al. (2013) were used as proxies for croplands. Zaehle et al. (2010b), estimated the total N leached to aquatic ecosystems from terrestrial sources to be 86 Tg N yr^{-1} , out of which 57 Tg N yr^{-1} was attributed to agricultural ecosystems. These estimates for the entire land surface are considerably larger than the estimates provided here ($24\text{--}66 \text{ Tg N yr}^{-1}$ for the simulations including croplands, Table 3). Among the simulations performed here, the simulation without residue removal (F_{NR}) was the only one in which N leached from croplands was of comparable magnitude to the findings of Zaehle et al. (2010b). In our study fertilisers are applied at specific crop developmental stages with amounts that match the CFT specific demand (see Table A2), whereas in Zaehle et al. (2010b) three applications with equal amounts were spread using climate indicators defining the peak in the growing season. This could lead to higher leaching when fertiliser application is not timed to coincide with the peak of the

³Derived by scaling their average $0.47 \text{ g N m}^{-2} \text{ yr}^{-1}$ by the ice-free land area of $1.33 \times 10^{14} \text{ m}^2$, consistent with the estimates done elsewhere in this study.

⁴derived from the N_2O emissions of 0.8 Tg N yr^{-1} stemming from N-leaching and the constant fraction of leached N that is emitted as N_2O , 0.6% that is assumed in the study.

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found for e.g. maize and soybean in the US (Wilhelm et al., 2004) and millet in Niger (Bationo et al., 1993).

Manure (F_{MN}) application had minor effects on any of the investigated processes, both globally and in any particular climatic region (Fig. 3). The relatively low effect on soil C might be caused by the relatively small fraction of the total N applied at sowing (which is the time when manure was also applied), 8 % for WW and 11 % for SW and MA. In terms of yield, the relatively high C : N (30), might have reduced crop productivity slightly, since the manure-N will not be available for plant uptake at sowing, but will be released from the SOM during the growing season. Still, in some of the high producing regions (e.g. north-western Europe and parts of China), manure application was the most effective management for carbon sequestration (Fig. 4); these are all areas where the N application rates in the data set used here are high (Zaehle et al., 2011), and thus the amount of carbon added to the soil is relatively large.

By contrast with moderate tillage, complete absence of tillage resulted in enhanced soil C, with only small to moderate yield reduction, and a small reduction of N loss through leaching. Depending on the regional climate and N-fertiliser applications, reductions in crop productivity by up to 0.5 t ha^{-1} were also reported for maize and winter wheat grown in the USA in a recent meta-analysis, comparing tillage to no-tillage (Ogle et al., 2012). A larger effect on C sequestration (at similarly small to moderate effects on yields) was only found when optimising for carbon sequestration also resulted in a moderate reduction in yields while achieving a reduction in the modelled N-leaching by ca. 30 % (Fig. 3). Considering the high global demand for food today and in the future, a 5 % yield reduction may be difficult to motivate in exchange for a 5 % increase in soil C and reduced leaching. Avoiding the loss of food production would require either further intensification (likely resulting in enhanced N losses through leaching) or expansion of crop and pasture areas (potentially interfering with other ecosystem services). In this regard, regional differences are crucial to consider. Large vegetation carbon stocks in tropical forest ecosystems motivate the protection of these systems, limiting the further expansion of managed land in these ecosystems. Given that tropical areas

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tend also to have the largest yield gaps (Licker et al., 2010), a much better strategy in these regions is to invest in sustainable intensification of existing managed land.

The initial difference between F_{opt} and F_{CC} in Fig. 3, where F_{CC} had a positive effect on yields until mid 1960s, is due to the fact that in the model, the cover-crops are being sown with a finite initial carbon and nitrogen mass. This results in more available nitrogen in the fields with this management (basically a fertilisation via the seeds), despite the indirect competition for nitrogen between the cover-crop and the main crop which subsequently also results in a relatively larger nitrogen export through leaching. Cover-crops have been used to re-vitalise croplands, the results shown here implies that the model partly captures this, but the simulated indirect competition is too strong and further studies and model developments are needed to better represent cover-crop management. Also, as the cover-crop implementation does not include symbiotic N-fixation, the simulated reduction in yields with that management could very well have resulted in the opposite effect, but as was seen for N-leaching prior to 1960 and also for the no-residue removal, maybe also an increase in the relative N-leakage.

5 Conclusions

We have presented a global model analysis highlighting effects of alternative crop management strategies for a range of core ecosystem processes and the services derived from them, related to interactions of climate change and land use change.

Our large-scale approach based on the simplifying assumption of uniform management across regions does not faithfully represent actual conditions, but instead allows the influence of different management actions to be evaluated, and geographical difference to be highlighted.

Results demonstrate that effects of management on cropland can be beneficial for carbon and nutrient retention without risking (large) yield losses. Nevertheless, effects on soil carbon are small compared with extant stocks in natural and semi-natural ecosystem types and managed forests. While agricultural management can be tar-

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geted towards sustainable goals, from a climate change or carbon sink perspective avoided deforestation or reforestation constitutes a far more effective overall strategy for maintaining and enhancing global carbon sinks. However, enhanced carbon storage in agricultural soils could also be seen as a surrogate for enhanced soil structure and reduced erosion having additional (non-climate) environmental benefits.

Appendix: Allocation

In Olin et al. (2015) relationships between allocation to leaves (g_L), stem (g_{St}), root (g_R) and grains (g_Y) based on the allocation model of Penning de Vries et al. (1989) were established using a logistic growth function, a Richards curve (Richards, 1959), (Eq. A1):

$$f_i = a + \frac{b - a}{1 + e^{-c(DS-d)}} \quad (A1)$$

where f_i is the daily allocation of assimilates to a plant organ relative to e.g. the shoot, a is the asymptote when $DS \rightarrow 0$, b is the upper asymptote when $DS \rightarrow \infty$, c the growth rate, and d is the DS of maximum growth.

The relative relationships of daily assimilate allocation to the organs described with Eq. (A1):

$$f_1 = \frac{g_R}{g_R + g_L + g_{St}}, f_2 = \frac{g_L}{g_L + g_{St}}, f_3 = \frac{g_Y}{g_R + g_L + g_{St} + g_Y} \quad (A2)$$

And combining the equations in Eq. (A2) yields:

$$\begin{aligned} g_R &= f_1(1 - f_3) \\ g_L &= f_2(1 - f_1)(1 - f_3) \\ g_{St} &= (1 - f_2)(1 - f_1)(1 - f_3) \\ g_Y &= f_3 \end{aligned} \quad (A3)$$

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See Olin et al. (2015) for more details on how these relationships were derived.

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Table 2. Soil C pools (0–1.5 m) in four climatic regions, observed (Batjes, 2014) and simulated, with the 0.05 and 0.95 percentiles in parenthesis. In the last two columns, the correlation coefficients and the p values to demonstrate that there is a positive relationship between simulated and observed values are shown.

Climatic region	Soil C (kg C m^{-2})		Corr.	P
	Observed	Simulated		
temperate, dry	5.7 (1.3–12.0)	10.0 (1.4–27.4)	0.14	0.0479
temperate, humid	8.4 (1.9–21.2)	11.9 (5.2–26.2)	0.05	0.4686
tropical, dry	6.0 (1.7–12.8)	7.6 (2.5–16.7)	0.07	0.3427
tropical, humid	11.2 (2.5–28.7)	7.9 (3.7–16.0)	0.13	0.0504

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Table A1. The parameters for the factors f_1 , f_2 and f_3 in the carbon allocation scheme (Eq. A2) for spring wheat (SW), winter wheat (WW) and maize (MA).

parameter		SW	WW	MA
f_1 :	a	0.62	0.53	0.24
	b	-0.02	0	1.22
	c	5.80	7.63	18.10
	d	0.55	0.55	1.12
f_2 :	a	0.86	0.8	0.68
	b	0.19	0.20	-0.06
	c	28.65	13.99	12.48
	d	0.55	0.55	0.81
f_3 :	a	0	0	0
	b	1	1	1
	c	8.27	8.32	28.52
	d	1.10	1.15	1.03

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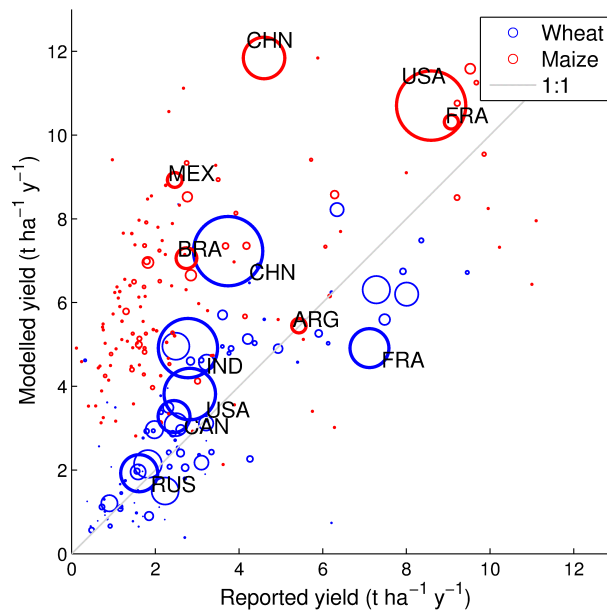


Figure 1. Per country comparison of simulated yields for WW (wheat) and MA (maize) against reported yields from FAO (1996–2005). Marker size indicates each country’s total production. The top 6 producer countries of both crops are labelled with an abbreviation: ARG, Argentina; BRA, Brazil; CAN, Canada; CHN, China; FRA, France; IND, India; MEX, Mexico; RUS, Russia; USA, United States.

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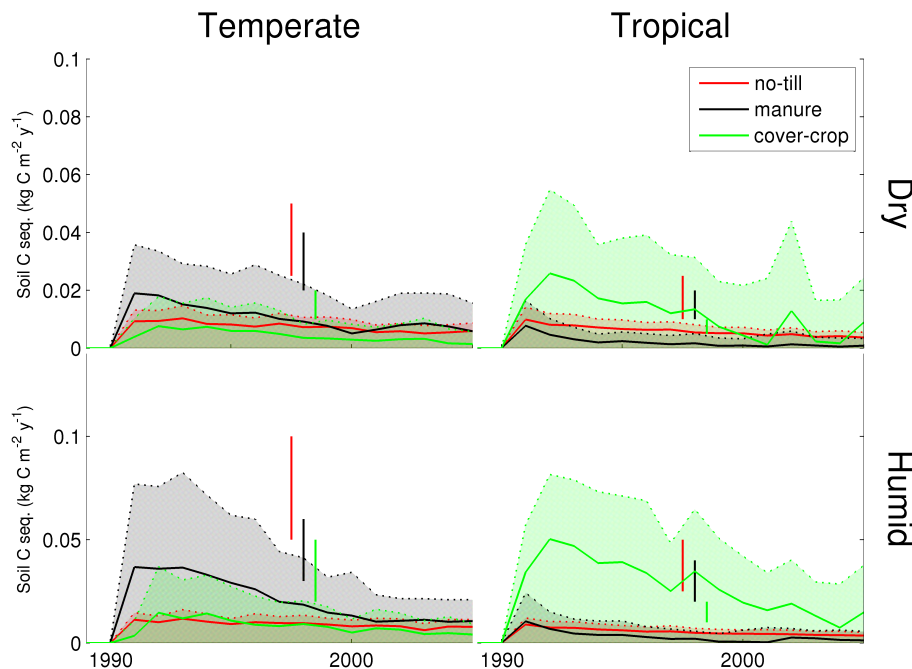


Figure 2. Simulated mean C sequestration following the implementation of management over the CRU historic period on agricultural soils averaged (thick lines) for the selected grid cells in the four climatic regions, compared to estimates (horizontal lines) from Stockmann et al. (2013). Dotted lines indicate the mean plus 2 standard deviations from all grid cells in each climatic region. The vertical lines do not represent specific years, but the potential over time to sequester C on cropland soils.

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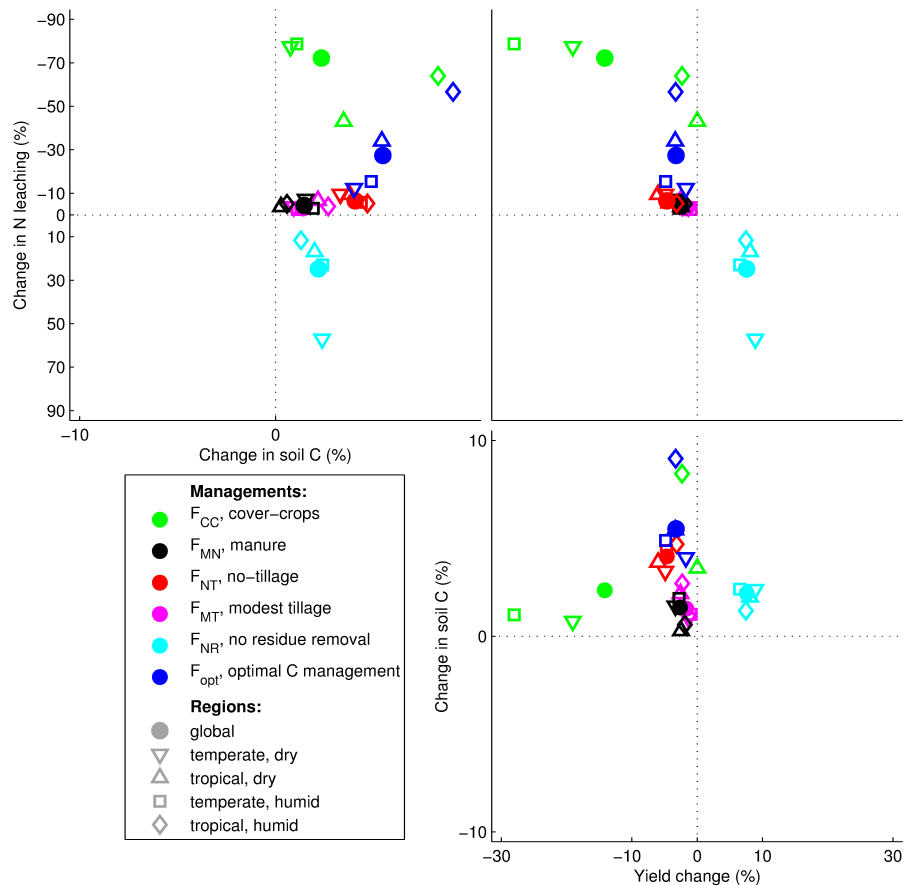


Figure 3. The simulated relative response (%) of soil carbon to management options (Table. 1) compared to the standard setup, averaged for 1996–2005 and displayed as the global response (filled symbol) and per climatic region. Note the reversed axes for N leaching (all axes display scales from reduced to enhanced ecosystem services).

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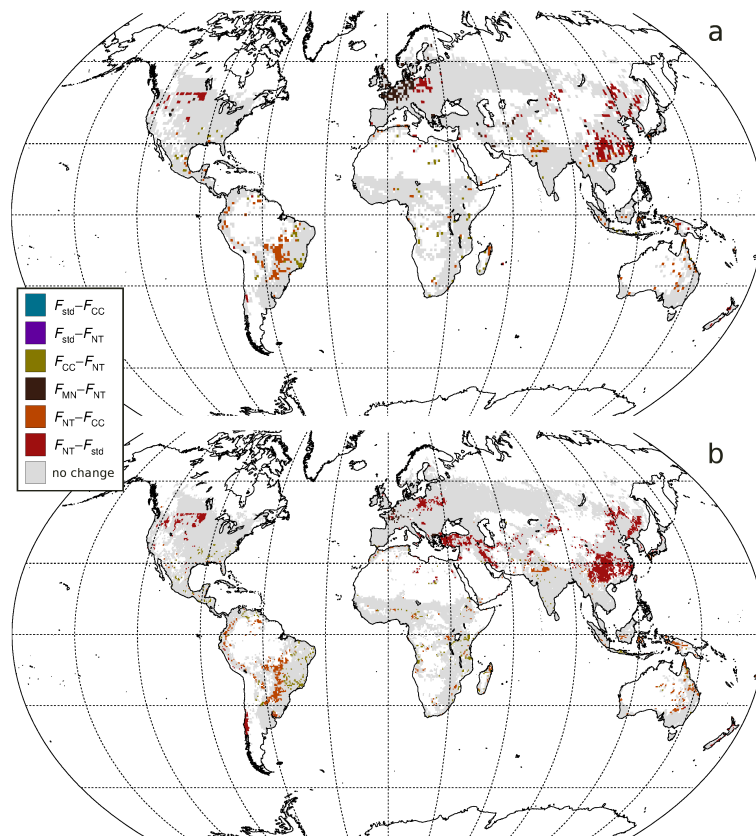


Figure 6. Grid cells where different management options resulted in the highest soil carbon in 2000 (Fig. 4) compared to 2050, **(a)** RCP 2.6 and **(b)** 8.5.

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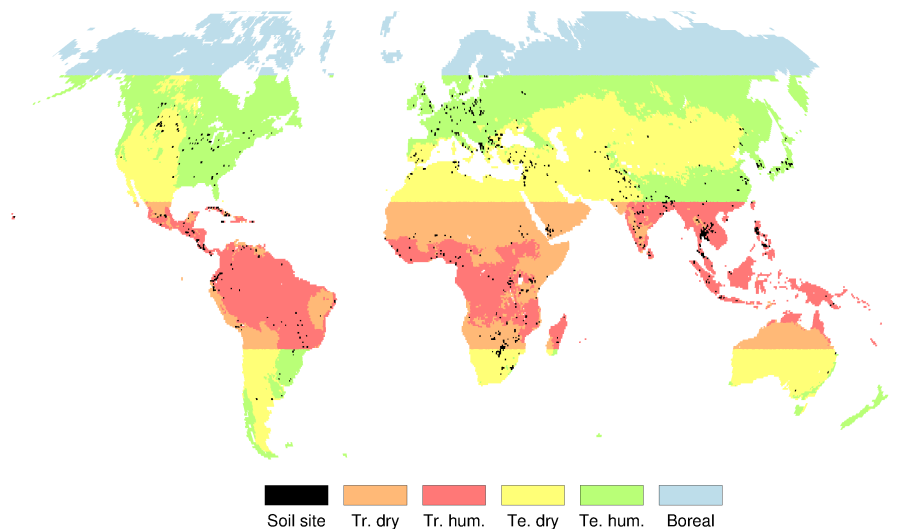


Figure B2. Climatic regions, as defined in Sect. 2.2.1 (Tropical dry, Tropical humid, Temperate dry, Temperate humid, Boreal), black pixels show the cropland soil column sites from Batjes (2005), used in the soil carbon comparison.

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