

1 **COMMENTS TO REVIEWERS**

2 **Reviewer 1**

3 1.- The discussion in the introduction on page 3 Line numbers 13-29 is not specific enough and especial
4 the claimed increase in extreme summer weather events and changed baroclinicity needs a much more
5 critical reflection.

6

7 Although we understand the criticism raised by the reviewer we must emphasize that this section is only
8 a small introduction to some key elements of the Arctic climate within the context of climate change.
9 Nevertheless, based on the comprehensive assessments provided by Tang et al (2013) and Cohen et
10 al. (2014), we have included in the new version of the manuscript some descriptions on the mechanisms
11 linking summer extreme events with climate change anomalies in the Arctic;

12 “According to Cohen et al. (2014) there are three major dynamical frameworks to propagate the
13 anomalous climate signals originated in the Arctic (namely changes in SIE and SCE) toward mid-
14 latitudes. These include: (1) changes in storm tracks; (2) changes in the characteristics of the jet stream;
15 and (3) anomalous planetary wave configurations triggered by regional changes in the tropospheric
16 circulation. Tang et al. (2013) compared the role played by both SIE and SCE in what concerns their
17 capacity to change atmospheric circulation and inducing extreme summer extremes in northern mid-
18 latitudes. These authors have found that despite the stronger decrease in SCE compared to SIE, the
19 latter provides a stronger response in terms of atmospheric circulation anomalies. Often related with
20 climatic extremes. Tang et al. (2013) provide evidence that the combined reductions of SIE and SCE
21 are associated to “widespread upper-level height increases, weaker upper-level zonal winds at high
22 latitudes, a more amplified upper-level pattern, and a general northward shift in the jet stream”.

23

24 2.- With respect to the presented figures the authors should discuss the origin of the obvious inter-
25 annual and decadal-scale changes?

26 We agree with the reviewer’s suggestion to provide some additional information on the mechanisms
27 responsible for the inter-annual and decadal-scale changes. Thus, we have added text in red in line 18,
28 page 7 describing the EOF analysis performed by Yu (2007) that is useful to account for the origin of the
29 decadal and interannual variability in both global evaporation and regions found in our study.

30 “...wind forcing playing a dominant role. According to Yu (2007) and after performing an EOF analysis
31 of Evaporation and its related variables (wind speed and air–sea humidity differences), the wind forcing

1 is mainly responsible for the decadal change through two mechanisms, one direct, “greater wind speed
2 induces more evaporation by carrying water vapor away from the evaporating surface to allow the air–
3 sea humidity gradients to be reestablished at a faster pace” and a second one indirect “the enhanced
4 surface wind strengthens the wind-driven subtropical gyre, which in turn drives a greater heat transport
5 by the western boundary currents, warms up SST along the paths of the currents and extensions, and
6 causes more evaporation by enlarging the air–sea humidity gradients” . The EOF analysis also showed
7 that the interannual variability of Evaporation occurred on similar time scales to those of the El Niño–
8 Southern Oscillation. The a), b) and c) panels in Figure 1 also show the...”

9 Then we have reminded this origin when describing our figures after line 21, page 7 in the previous
10 version of the manuscript

11 “Although important interannual variability superimposed to a pronounced decadal-scale variability
12 previously commented trends are significant in most of the grid points encircled, and are especially clear
13 for the Atlantic, Pacific and Mediterranean sources. Similar results were reached when evaporation
14 taken from ERA- 24 Interim was used (not shown)”.

15

16 3.- The manuscript needs a critical reading to eliminate some typing errors.

17 We have corrected the text

18

19 **REVIEWER 2**

20 1.- Page 5, Lines 24-29: "i)" appears to be stated incorrectly. Perhaps if the word "extension" is replaced
21 with "contraction", then it would be correct. But if this replacement is made, then "i)" would essentially
22 say the same thing as "ii)". As "i)" is currently stated, it is not consistent with Kapsch et al. (2013) (they
23 discuss high ice years (HIYs) on their page 746)

24 The reviewer is correct. Thanks for spotting this contradiction. We have removed option i)

25 “... Using a very different methodology Kapsch et al. (2013) in the paper entitled “Springtime
26 atmospheric energy transport and the control of Arctic summer sea-ice extent” demonstrated that i)
27 ~~enhanced water vapour and clouds in spring, together with the associated greenhouse effect, are~~
28 ~~related to the extension of sea ice during the summer; and ii)~~ in areas of summer ice retreat, a
29 significantly enhanced transport of humid air is evident during spring, producing increased cloudiness
30 and humidity resulting in an enhanced greenhouse effect.”

1 **Atmospheric moisture transport, the bridge between ocean** 2 **evaporation and Arctic ice melting**

3

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9

10 **Abstract**

11 Changes in the atmospheric moisture transport have been proposed as a vehicle for
12 interpreting any of the most significant changes in the Arctic region. The increasing moisture
13 over the Arctic during last decades it is not strongly associated with the evaporation that takes
14 place within the Arctic area itself, despite the fact that the sea-ice cover is decreasing. Such
15 increment is consistent is more dependent on the transport of moisture from the extratropical
16 regions to the Arctic that has increased in recent decades, and is expected to increase within a
17 warming climate. This increase could be due either to changes in circulation patterns which
18 have altered the moisture sources, or to changes in the intensity of the moisture sources
19 because of enhanced evaporation, or a combination of these two mechanisms. In this short
20 communication we focus on the assessing more objectively the strong link between ocean
21 evaporation trends and Arctic Sea ice melting. We will critically analyze several recent results
22 suggesting links between moisture transport and the extent of sea-ice in the Arctic, this being
23 one of the most distinct indicators of continuous climate change both in the Arctic and on a
24 global scale. To do this we will use a sophisticated Lagrangian approach to develop a more
25 robust framework on some of these previous disconnecting results, using new information and
26 insights. Results reached in this study seems to stress the connection between two climate
27 change indicators, namely an increase in evaporation over source regions (mainly the
28 Mediterranean Sea, the North Atlantic Ocean and the North Pacific Ocean in the paths of the
29 global western boundary currents and their extensions) and Arctic ice melting precursors.

30

1 **1 The outstanding role of Arctic climate within the global climate system**

2 The last IPCC Assessment Report has confirmed that the main components of the climate
3 system have been warming (atmosphere, oceans) or shrinking (cryosphere) since the 1970s, as
4 a result of global warming induced by the significant increment in concentration of
5 Greenhouse Gases of anthropogenic origin (AR5, IPCC, 2013). The so called hiatus in the
6 rise of global air temperature since the late 1990s is not observed in the relentless decadal
7 shift of temperature distributions in both hemispheres (Hansen et al., 2012) neither in the
8 frequency of extreme hot events over the continents (Seneviratne et al., 2014). The much
9 larger capacity of the oceans to store heat, in respect to the atmosphere, has played a
10 fundamental role storing the excessive heat retained in the climate system either in the Pacific
11 (Kosaka and Xie, 2013) or the Atlantic (Chen and Tung, 2014) oceans.

12 However, global warming is a very uneven phenomenon impossible to be encapsulated by a
13 single indicator relative to one subsystem, such as the global average of near surface
14 atmospheric temperature. The spatial pattern of observed temperature trends is very
15 asymmetrical and regionalized, with continents warming more than oceans, and with high
16 latitudes also presenting considerably higher warming rates than mid-latitude and tropical
17 regions. In particular, several authors have shown that the rise in Arctic near surface
18 temperature (AST) has been twice as large as the global average throughout most of the year
19 (e.g., Screen and Simmonds, 2010; Tang et al., 2014, Cohen et al., 2014). Additionally, the
20 evolution of the climate in the Arctic region is often associated to two important indicators;
21 the summer and autumn sea-ice-extent (SIE) and the spring and summer snow-cover extent
22 (SCE), both characterized by a very significant decline since the 1970s and widely recognized
23 as some of the most undeniable indicators of continuous climate change affecting the climate
24 system (Tang et al., 2014; IPCC, 2013).

25 Nevertheless, the opposite evolution of AST and SIE indices in recent decades emphasize that
26 both phenomena are not independent and, actually, are known to reinforce each other (Tang et
27 al., 2014), as changes in surface albedo (associated with melting snow and ice) tend to
28 enhance warming in the Arctic (Serreze and Francis, 2006) as shown in the recent review
29 paper Cohen et al. (2014). Nevertheless both indicators (AST and SIE) may also respond to
30 other mechanisms including changes in atmospheric circulation patterns (Graverson et al.,
31 2008), ocean circulation (Comiso et al., 2008), or changes in radiative fluxes associated to
32 cloud cover and water vapour content in the atmosphere (Schweiger et al., 2008; Kapsch et

1 al., 2013), though the absorption of the outgoing long-wave radiation from the surface by the
2 increased atmospheric moisture and then remitted toward the Arctic surface, resulting in the
3 surface warming and sea-ice decline (Kapsch et al., 2013). In particular, changes in the
4 atmospheric moisture have been proposed as a vehicle for interpreting the most significant
5 changes in the Arctic region either due to increase transport from middle latitudes (Lucarini
6 and Ragone, 2011; Zanhg et al., 2012) or via enhance local evaporation (Bintanja and Seltan,
7 2014). However, some of the recent studies showed that the evaporation from the Arctic
8 surface appears not to be an important moisture source (e.g., Graverson et al., 2008; Park et
9 al., 2015).

10

11 According to some authors, the recent rise on the incidence of summer extreme weather
12 events over northern hemisphere continental land masses (Coumou and Rahmstorf, 2012;
13 Seneviratne et al., 2014) is probably driven by the accelerated decline of summer SIE and
14 SCE observed in recent decades (Francis and Vavrus, 2012; Tang et al., 2014). According to
15 this hypothesis, the observed weakening of poleward temperature gradient triggered changes
16 in atmospheric circulation namely slower progression of Rossby waves (Francis and Vavrus,
17 2012) and the existence of a planetary-scale wave life cycle (Bagget and Lee, 2015) that is
18 highly amplified (blocking) despite a reduced meridional temperature gradient (consistent
19 with Francis and Vavrus, 2012). These mechanisms have favoured more persistent weather
20 conditions that are often associated to extreme weather events, such as the mega-heatwave in
21 Russia in 2010 (Barriopedro et al., 2011) or long drought in central USA (Coumou and
22 Rahmstorf, 2012). However, there is currently a wide debate on the nature of mechanism(s)
23 responsible for this increment of persistent weather patterns associated to such extreme
24 climatic events (Cohen et al., 2014), with some authors suggesting other drivers (albeit
25 equally exacerbated by global warming) such as the role of drying soils associated with earlier
26 SCE melting (Tang et al., 2014) or simply related to tropical extra-tropical interactions
27 (Palmer, 2014). According to Cohen et al. (2014) there are three major dynamical
28 frameworks to propagate the anomalous climate signals originated in the Arctic (namely
29 changes in SIE and SCE) toward mid-latitudes. These include: (1) changes in storm tracks;
30 (2) changes in the characteristics of the jet stream; and (3) anomalous planetary wave
31 configurations triggered by regional changes in the tropospheric circulation. Tang et al.
32 (2013) compared the role played by both SIE and SCE in what concerns their capacity to

1 change atmospheric circulation and inducing extreme summer extremes in northern mid-
2 latitudes. These authors have found that despite the stronger decrease in SCE compared to
3 SIE, the latter provides a stronger response in terms of atmospheric circulation anomalies.
4 Often related with climatic extremes, Tang et al. (2013) provide evidence that the combined
5 reductions of SIE and SCE are associated to “widespread upper-level height increases, weaker
6 upper-level zonal winds at high latitudes, a more amplified upper-level pattern, and a general
7 northward shift in the jet stream”.

8 Considering all the above reasons the Arctic sector emerges as the most sensitive region of
9 the climate system to the effects of global warming but it also represents an area where
10 current and future changes are bound to affect the climate at a much larger scale (Screen and
11 Simmonds, 2010; Tang et al., 2014, Cohen et al., 2014).

12

13 **2 Main mechanisms relating sea ice decline and increase moisture transport**

14 The atmospheric branch of the hydrological cycle plays a fundamental role establishing the
15 link between the Arctic system and the global climate. However, to the best of our
16 knowledge, this role has not been fully accounted objectively, although the transport of
17 moisture from the extratropical regions to the Arctic has increased in recent decades (Zhang et
18 al., 2012), and is expected to further increase under global warming, independently of the
19 climate change scenario considered (Kattsov et al., 2007). Some works try to explain extreme
20 events of atmospheric moisture transport to the Arctic throughout the occurrence of
21 atmospheric rivers (Woods et al., 2013) and Rossby wave breaking events (Liu and Barnes,
22 2015). The general increase of moisture could be due either to changes in circulation patterns
23 which have altered the location of the most important moisture sources, or result from
24 changes in the magnitude of the existing moisture sources as a consequence of enhanced
25 evaporation, or a combination of these two mechanisms (Gimeno et al., 2012; 2013).

26 Most studies of changes on moisture transport towards the Arctic Climate make use of one of
27 three possible techniques, namely (1) Eulerian approaches (e.g. Jakobson and Vihma, 2010),
28 which can be used to estimate the ratio of advected-to-recycled moisture and to calculate the
29 moisture transport between predetermined source and sink regions; (2) isotope analysis (e.g.,
30 Kurita, 2011), but neither this nor the Eulerian techniques are capable of a proper
31 geographical identification of the sources; or (3) more complex Lagrangian computational
32 techniques that are able to infer the sources of the precipitation that falls in a target region and

1 thus overcome the limitations of (1) and (2). An analysis of the performance of these
2 Lagrangian techniques and their advantages over Eulerian and isotope analysis was recently
3 given by Gimeno et al. (2012). Here we will critically analyze some of the previous
4 assessments that have established the link between moisture transport from mid-latitudes
5 towards the Arctic region and changes in Arctic SIE. In addition, we will use a sophisticated
6 Lagrangian approach to contrast these existing results using new information and insights.

7 In recent years a number of mechanisms have been put forward relating the strength of
8 moisture transport and Arctic SIE. These mechanisms vary significantly in the nature of their
9 main driver, including; i) hydrological, such as increments in Arctic river discharges (Zhang
10 et al., 2012) or increments in precipitation due to enhanced local evaporation due to less SIE
11 (Bintanja and Selten, 2014), ii) radiative, particularly through rises in cloud cover and water
12 vapour (Kapsch et al., 2013), iii) dynamical, namely more summer storms with unusual
13 characteristics crossing the Arctic, (Simmonds and Rudeva, 2012). Most likely these different
14 mechanisms coexist to a certain extent and are not necessarily mutually exclusive, for
15 instance the autumn and early positive trend in SCE (Estilow et al., 2015) can be closely
16 related to positive trends in Eurasian rivers (Yang et al., 2007). In particular, two of these
17 works (Zhang et al., 2012; Kapsch et al., 2013) provide novel insight on the role played by the
18 transport of moisture and the melting of sea ice or snow cover. Their main findings are
19 summarized below:

20 1. According to Zhang et al. (2012) in their work entitled "Enhanced poleward moisture
21 transport and amplified northern high-latitude wetting trend", the authors provide strong
22 evidence to support; i) that there is a trend in the net poleward atmospheric moisture transport
23 (AMT) towards the Eurasian Arctic river basins, ii) that this net AMT is captured in 98% of
24 the gauged climatological river discharges, iii) that the upward trend of 2.6% net AMT per
25 decade is in good agreement with the 1.8% increase per decade in the gauged discharges.

26 The increase in Arctic river discharge is a possible cause of ~~melting~~ sea-ice **melting** in
27 agreement with several studies realized over the Canadian Arctic region support these results
28 (e.g., Dean et al., 1994; Nghiem et al., 2014). Thus, AMT can be seen to have an important
29 role to play in this process. Nevertheless, Zhang et al. (2012) used a very simple analysis of
30 integrated moisture fluxes, in which they calculated moisture transport from predetermined
31 source and sink regions, and were unable to identify the moisture source regions directly.

32

1 2. Using a very different methodology Kapsch et al. (2013) in the paper entitled “Springtime
2 atmospheric energy transport and the control of Arctic summer sea-ice extent” demonstrated
3 that ~~i) enhanced water vapour and clouds in spring, together with the associated greenhouse
4 effect, are related to the extension of sea ice during the summer; and ii)~~ in areas of summer
5 ice retreat, a significantly enhanced transport of humid air is evident during spring, producing
6 increased cloudiness and humidity resulting in an enhanced greenhouse effect.

7 As for Kapsch et al. (2013), global balances of atmospheric moisture flux were used, which
8 allowed neither the identification of the moisture sources nor any assessment of their role in
9 the variability of the moisture transport.

10

11 **3 Identifying objectively the main sources of moisture for large Eurasian** 12 **rivers basins**

13 The analysis adopted here to discuss existing results is mostly based on the Lagrangian
14 particle dispersion model FLEXPART (Bintanja and Selten, 2014; Stohl and James, 2004),
15 using data from 1979 to 2013 obtained from the ERA-Interim reanalysis of the ECMWF (Dee
16 et al., 2011), which can be considered the state of the art reanalysis in terms of the
17 hydrological cycle (Trenberth et al., 2011; Lorenz and Kunstmann, 2012). The analysis will
18 be restricted to years after 1979 in order to avoid working with results obtained prior to the
19 incorporation of satellite data in the reanalysis. Using a horizontal resolution of 1° in latitude
20 and longitude and a resolution of 61 vertical levels, the algorithm tracks atmospheric moisture
21 along trajectories. A 3-D wind field moves a large number of so-called particles (air parcels)
22 resulting from the homogeneous division of the atmosphere. The specific humidity (q) and the
23 position (latitude, longitude and altitude) of all the particles are recorded at 6-hour intervals.
24 The model then calculates increases (e) and decreases (p) in moisture along each trajectory at
25 each time step by means of variations in (q) with respect to time i.e., $e-p = m \, dq/dt$. The
26 quantity ($E-P$) is calculated for a given area of interest by summing ($e-p$) for all particles
27 crossing a 1° grid column of the atmosphere, where E and P are the rates of evaporation and
28 precipitation, respectively. The particles are tracked and a database is created with values of
29 $E-P$ averaged and integrated over 10 days of transport, this being the average residence time
30 of water vapour in the atmosphere (Numaguti, 1999). The main sources of moisture for the
31 target area (in terms of when and where the air masses that reach the target area acquire or
32 lose moisture) are shown through the analysis of the 10-day integrated ($E-P$) field. For a

1 comprehensive review see Gimeno et al. (2012), which provides details of the limitations of
2 this Lagrangian approach, its uncertainty and significance, and its advantages and
3 disadvantages with respect to other methods of estimating moisture sources. For further
4 information on FLEXPART model see Stohl et al. (2004).

5

6 According to Zhang et al. (2012), temporal lags must be considered when linking AMT from
7 lower latitudes with snowpack accumulation and also between this and Arctic river
8 discharges. Thus, summer Arctic river discharge can be related to the result of the melting of
9 the snowpack that accumulated during the preceding months, while the AMT most related to
10 the summer river discharge corresponds to that resulting from snowpack accumulation during
11 the period October - March. We therefore choose this period to estimate the moisture sources
12 for the target region formed by the Ob, Yenisei and Lena rivers basins, as in the work of
13 Zhang et al. (2012). The central panel of figure Fig. 1 shows that the main moisture sources
14 are located over the Mediterranean Sea, and the smaller Caspian and Black Seas, as well as
15 the North Atlantic Ocean and to a somewhat lesser degree the North Pacific Ocean in the
16 paths of the global western boundary currents and their extensions. This result is striking
17 because these source regions seem to match those areas with the highest trend in terms of
18 evaporation in the past few decades.

19

20 **4 Trends in evaporation from main sources: possible consequences**

21 Using some of the best estimates of evaporation, namely those derived from the OAFflux data
22 (Yu and Weller, 2007), strong increasing trends can be seen in evaporation from the oceans
23 since 1978, with the upward trend being most pronounced during the 1990s. The spatial
24 distribution of these trends (Yu, 2007) shows that while the increase in evaporation has
25 occurred globally, it has primarily been observed during the hemispheric winter and is
26 strongest along the paths of the global western boundary currents and any inner Seas with
27 wind forcing playing a dominant role. According to Yu (2007) and after performing an EOF
28 analysis of Evaporation and its related variables (wind speed and air-sea humidity
29 differences), the wind forcing is mainly responsible for the decadal change through two
30 mechanisms, one direct, “greater wind speed induces more evaporation by carrying water
31 vapor away from the evaporating surface to allow the air-sea humidity gradients to be
32 reestablished at a faster pace” and a second one indirect “the enhanced surface wind

1 strengthens the wind-driven subtropical gyre, which in turn drives a greater heat transport by
2 the western boundary currents, warms up SST along the paths of the currents and extensions,
3 and causes more evaporation by enlarging the air–sea humidity gradients”. The EOF analysis
4 also showed that the interannual variability of Evaporation occurred on similar time scales to
5 those of the El Niño–Southern Oscillation. The a), b) and c) panels in Figure 1 also show the
6 evolution of the average evaporation derived from OAFLUX for the main moisture sources
7 for the Arctic river basins (those circled with a blue line and the entire Mediterranean basin
8 sea). Although important interannual variability superimposed to a pronounced decadal-scale
9 variability previously commented trends are significant in most of the grid points encircled,
10 and are especially clear for the Atlantic, Pacific and Mediterranean sources. Similar results
11 were reached when evaporation taken from ERA-Interim was used (not shown). The
12 differences in the composites of the moisture sources of the Arctic river basins between the
13 decade 2001-10 and the decade 1981-90 are also shown in Figure 1, with greenish colours
14 indicating regions where their contribution as a source intensified over these years. From
15 these results it seems clear that there is an enhanced moisture contribution from those
16 moisture regions where the evaporation increased.

17 We have repeated the procedure considering the region analyzed by Kapsch et al. (2013), i.e.
18 in this case, the late spring (April and May) moisture sources detected are related to the area
19 where the September sea-ice anomaly is encountered. Overall results (Figure 2) are quite
20 similar to those presented for the Arctic river basins, and the main moisture sources are also
21 placed in the paths of the global western boundary currents in both the North Atlantic and the
22 North Pacific Oceans (the main one in this case), and in the Mediterranean basins (more
23 moderated in this case).

24 In this regard the intensification of evaporation in these source regions could have a dual
25 effect on the reduction of September Arctic ice, through (1) intensification of summer river
26 discharge and (2) enhancement of the greenhouse effect due to an increase in cloudiness and
27 humidity over the ice-melting regions.

28

29 **Summary and conclusions**

30 We have made a critical assessment of the results obtained in two important recent works that
31 offer new understanding on the role played by the transport of moisture and the melting and
32 the melting of sea ice or snow cover (Zhang et al., 2012; Kapsch et al., 2013). The Lagrangian

1 analysis adopted in our approach seems to stress the connection between two climate change
2 indicators, namely an increase in evaporation over source regions and Arctic ice melting. We
3 are confident that our results provide the necessary link between these two realms and suggest
4 an intricate chain of events related to (1) positive trends in evaporation in specific ocean areas
5 that correspond to the main moisture source regions of Eurasian rivers, (2) upward trends in
6 atmospheric transport from these regions to the Arctic river basins/ regions where ice-melting
7 occurs, and (3) trends in river discharges/moisture and cloud cover. These developments merit
8 further and more comprehensive study in terms of their effects on present and future climates.

9

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11 programme, within the project ACPCA and by the Spanish MINECO and FEDER within the
12 project TRAMO.

13

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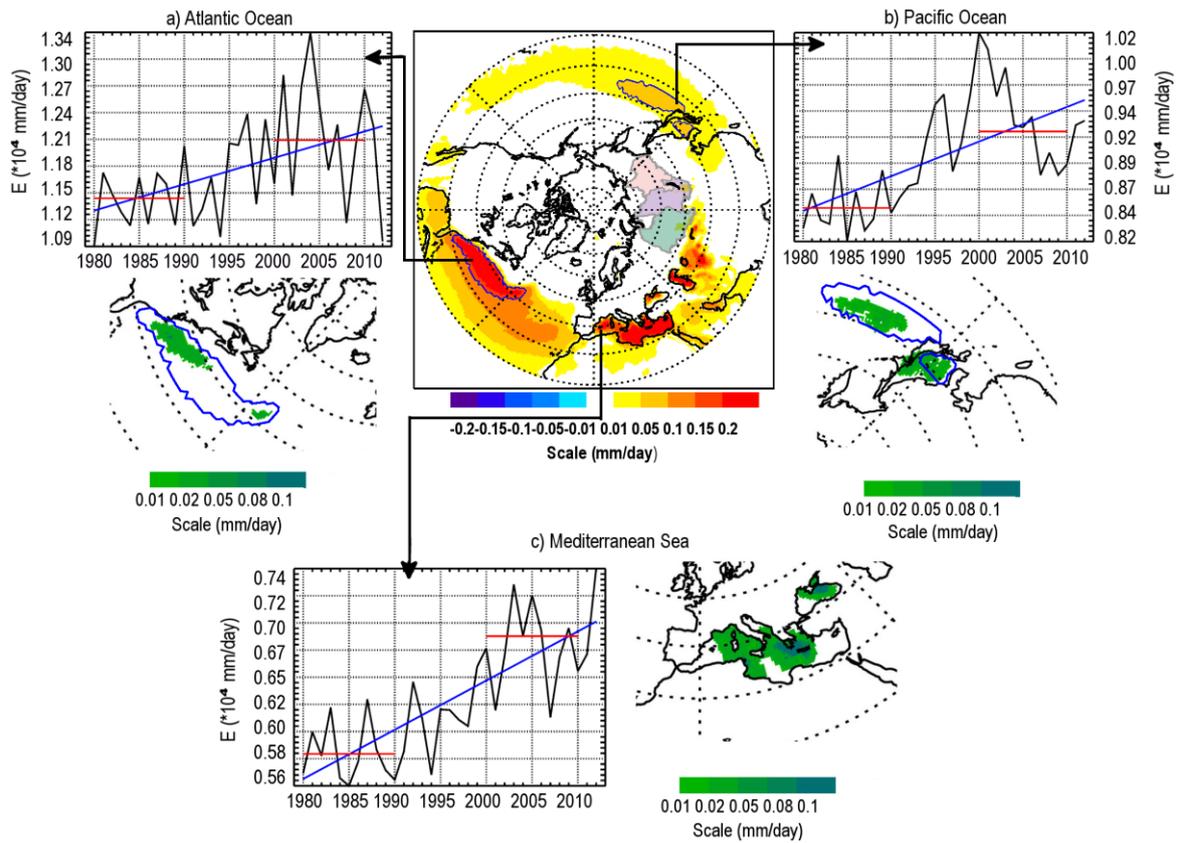
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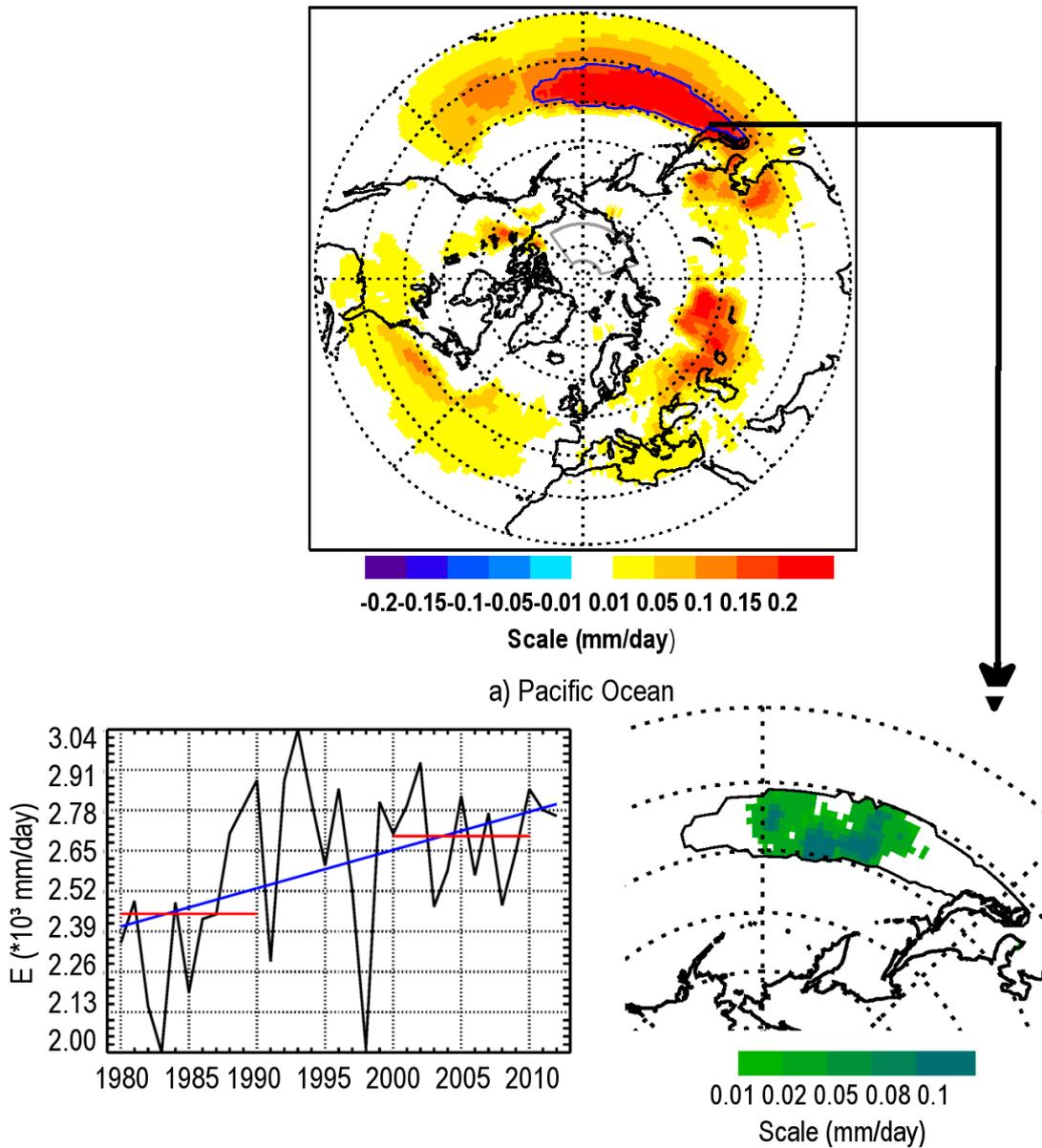
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2 Figure 1. (Central panel) Climatological October-March 10-day integrated (E-P) values
3 observed for the period 1979 – 2012, for all the particles bound for the Ob, Yenisei and Lena
4 rivers basins (green, purple and pink areas respectively indicate the basin area), determined
5 from backward tracking. Warm colours represent regions acting as moisture sources for the
6 tracked particles. Plots in green show the significant positive differences at the 95% level
7 after bootstrap test (1000 interactions) in the composites of the moisture sources of the Arctic
8 river basins between the decades 2001-10 and 1981-90. Temporal series show the evolution
9 of the average evaporation derived from OAFUX dataset for the main moisture sources for
10 the Arctic river basins (the Atlantic and Pacific sources, those circled with a blue line in the
11 central figure, and for the whole Mediterranean Sea basin). The blue lines are the linear trend
12 and the red ones denoted the 10-year periods used on composites.
13



1
 2 Figure 2. As Fig.1 but for the Kapsch area (115°-215° E ; 75°-85° N), denoted with the grey
 3 contour in the bottom panel.